

1 **Diffusive nitrous oxide (N₂O) fluxes across the sediment-water-**
2 **atmosphere interfaces in aquaculture shrimp ponds in a**
3 **subtropical estuary: Implications for climate warming**

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25 **ABSTRACT**

26 Emissions of the potent greenhouse gas nitrous oxide (N₂O) from aquaculture remain a
27 large knowledge gap in the global N₂O budget. The water column and the sediment of
28 aquaculture ponds present very different environmental conditions, but their relative
29 contributions to N₂O production and emission are poorly resolved. We sampled three
30 aquaculture ponds in the Min River Estuary in southeastern China monthly throughout
31 the farming season. Based on the dissolved N₂O concentrations within the water column
32 and in sediment porewater, we calculated the diffusive N₂O fluxes across the water-
33 atmosphere interface (WAI) and sediment-water interface (SWI). The diffusive N₂O flux
34 averaged 216.9 nmol m⁻² h⁻¹ across WAI and 16.0 nmol m⁻² h⁻¹ across SWI. The estimated
35 N₂O production rate under steady-state condition was 0.13 nmol L⁻¹ h⁻¹ in the water
36 column and 1.07 nmol L⁻¹ h⁻¹ in sediment porewater. Hence, the water column
37 compartment and the sediment compartment of the aquaculture ponds played different
38 roles in N₂O dynamics. Based on our data, it is calculated that China's coastal aquacultural
39 ponds would emit 0.2 Gg N₂O yr⁻¹, or less than 1% of all aquaculture N₂O emission in
40 China. Therefore, coastal shrimp aquaculture has a relative minor climate impact
41 compared to other aquaculture operations. Future studies should examine the role of N-
42 cycling functional genes on N₂O production and the mechanisms regulating N₂O emission
43 from aquaculture ecosystems.

44
45 **Keywords:** Nitrous oxide (N₂O) fluxes; Sediment-water interface (SWI); Water-

46 atmosphere interface (WAI); Nitrogen substrate; Temperature; Aquaculture ponds

47 **1. Introduction**

48 Nitrous oxide (N₂O) is an ozone-depleting greenhouse gas (GHG), contributing
49 substantially to radiative forcing and global climate change (Maavara et al., 2018; Quick
50 et al., 2019; Ravishankara et al., 2009). The Industrial Revolution and intensive
51 fertilization for farming have increased nitrogen availability significantly in both land
52 and water environments (Swaney et al., 2012; Howarth et al., 1996; Zhang et al., 2020),
53 which in turn has caused the increase in atmospheric N₂O. N₂O has increased to 335.3
54 ppbv this year (NOAA, 2022), exceeding the preindustrial concentration by ~24%, and
55 the agricultural sector contributes ~60% (4.3 Tg N yr⁻¹) of the global N₂O emissions from
56 anthropogenic activities (Tian et al., 2020; Webb et al., 2021). China has promised to be
57 carbon neutral by 2060 (Yang et al., 2022). To effectively mitigate the impact of climate
58 change, it is therefore crucial to better understand the N₂O biogeochemical processes in
59 various agriculture ecosystems.

60 Nitrous oxide can be produced in soil (or sediment) and water by two main microbial
61 processes, nitrification and denitrification (Audet et al., 2017; Li et al., 2021; Wrage et
62 al., 2005; Wu et al., 2021), with oxygen (O₂) and nitrogen (N) availabilities being
63 important controlling factors (Murray et al., 2015; Xiao et al., 2019a; Yang et al., 2020a).
64 Similar to crop fields, aquaculture systems (e.g., aquaculture ponds) receive heavy N
65 loadings and are hotspots for N₂O production (Hu et al., 2012; Yang et al., 2020a; Yuan
66 et al., 2019, 2021) and emission (Paudel et al., 2015; Ye et al., 2022; Yogev et al., 2018).
67 An aquaculture system contains two compartments—the water column and the sediment,

68 both of which can contribute to N₂O production but are characterized by different
69 reduction-oxidation conditions and microbial community compositions (Beaulieu et al.,
70 2015; Freing et al., 2012; Yuan et al., 2021). The anoxic sediment receives a large amount
71 of organic matter (OM) from animal feces and residual feeds (Chen et al., 2016; Yang et
72 al., 2021), where microbial mineralization of the OM fuels N₂O production (Lin and Lin,
73 2022; Yuan et al., 2021). On the other hand, the overlying water column tends to be better
74 aerated and may support aerobic microbial activities, such as nitrification.

75 Although efforts have been made to quantify N₂O emissions from aquaculture
76 systems, there is a paucity of information on the relative contributions of the two
77 compartments (Hu et al., 2012). To close the knowledge gap, we analysed the diffusive
78 N₂O fluxes across the SWI and the WAI of aquaculture ponds in southeastern China. The
79 objectives were to: (1) characterize the temporal variations in diffusive N₂O fluxes across
80 the sediment-water-atmosphere interfaces; (2) explore the environmental factors that
81 drive the temporal variations in N₂O fluxes; (3) compare the roles of sediment and water
82 in N₂O production and emission from the aquaculture ponds. We hypothesized that: (1)
83 N₂O fluxes would exhibit distinct temporal variations in response to differences in
84 environmental variables (e.g., temperature and nitrogen availability); (2) the water
85 column would contribute more than the sediment to the overall N₂O emission to air.

86 **2. Materials and methods**

87 *2.1. Study area*

88 Our research area is located in the Shanyutan wetland of the Min River Estuary

89 (MRE), southeastern China (Figure 1). The region has a humid subtropical monsoon
90 climate (Tong et al., 2010), with a mean annual temperature of 19.6 °C and an average
91 annual precipitation of 1,390 mm (Yang et al., 2020b). The wetland has three dominant
92 vegetation types: two with the native species *Cyperus malaccensis* and *Phragmites*
93 *australis*, and one with the invasive species *Spartina alterniflora*. During the past decades,
94 large extent of the MRE tidal saltmarshes (mainly dominated by *C. malaccensis* and *S.*
95 *alterniflora*) were converted to aquaculture ponds for shrimp (*Litopenaeus vannamei*)
96 due to rising demand for seafood (Stokal et al., 2021; Yang et al., 2020c). Shrimp
97 farming typically begins in May and ends in November, with one crop of shrimp
98 produced annually. For more details of the aquaculture operation, please see Yang et al.
99 (2020c). In this study, three aquaculture ponds of ~1.5 m deep were selected for monthly
100 sampling of water and sediment from April 2019 to January 2020, for a total of 10
101 sampling campaigns in each pond. In each pond, samples were collected from three sites:
102 one near the bank, one in the feeding zone, and one at the center of the pond.

103 2.2. Sediment collection and bulk properties

104 Three surface sediments (top 15 cm) were collected at each site using a steel cylinder
105 (5 cm in diameter), stored in sterile sample bags and transported in a cooler to the
106 laboratory within 4–6 hr. All samples were stored at 4 °C, and analyzed within 72 hr. In
107 the laboratory, the samples were analyzed for physicochemical properties of sediments,
108 dissolved N₂O concentrations and physicochemical properties of porewater, as explained
109 below.

110 Subsamples of sediment were freeze-dried, homogenized and ground to fine powder
111 to determine pH and salinity. Sediment pH was determined via a pH meter (Thermo
112 Fisher Scientific, Sunnyvale, California, USA) in a sediment-to-water ratio of 1:2.5 (w/v
113 with added deionized water). Sediment salinity (SAL) was measured by a Eutech
114 Instruments-Salt6 salinity meter (Thermo Fisher Scientific, San Francisco, California,
115 USA) in a sediment-to-water ratio 1:5 (w/v). Soil water content (SWC) and bulk density
116 (BD) were determined after drying fresh soil at 105 °C for 48 h (Percival and Lindsay,
117 1997; Yin et al., 2019); weight loss after drying was used to calculate sediment porosity
118 (POR) (Zhang et al., 2013). *In situ* sediment temperature (T_s) were measured by a
119 portable temperature meter (IQ150, IQ Scientific Instruments, Carlsbad, California,
120 USA).

121 2.3. N_2O concentration and dissolved chemicals in sediment porewater

122 The dissolved N_2O concentrations in sediment porewater was measured according
123 to Dutta et al. (2015). Briefly, a subsample of sediment (6 cm^3) was collected via a 10
124 mL syringe, transferred to a 55-mL glass serum vial, and then sealed using an open-
125 topped screw cap and a halobutyl rubber septum. The vial was shaken vigorously in an
126 oscillator (IS-RDD3, China) for 10 min to achieve gas equilibrium between the slurry
127 and the headspace. The concentration of headspace N_2O (approximately 10 mL) was then
128 analysed on a gas chromatograph (GC) with an electron capture detector (ECD)
129 (Shimadzu GC-2014, Kyoto, Japan). Three N_2O gas standards, namely 0.3, 0.4, and 1.0
130 ppm, were used in the calibration. The detection limit for N_2O was 0.02 ppm, and the

131 relative standard deviations of N₂O were ±4.5% in 24 h. The calculation of sediment
132 porewater dissolved N₂O concentration (nmol L⁻¹) followed the method of [Ding et al.](#)
133 (2005) and [Johnson et al. \(1990\)](#).

134 The rest of the bulk sediment was centrifugated at 5000 rpm for 10 min (Cence®
135 L550, [De Vittor et al., 2012](#)) to extract porewater. After being filtered through 0.45 µm
136 acetate fiber membranes (Biotrans™ nylon membranes), the porewater filtrates were
137 analyzed for the levels of nitrate-nitrogen (NO₃⁻-N) and ammonia-nitrogen (NH₄⁺-N) on
138 a flow injection analyzer (Skalar Analytical SAN⁺⁺, Netherlands), and Cl⁻ and SO₄²⁻ on
139 an ion chromatograph (Dionex 2100, Thermo Fisher Scientific, Sunnyvale, California,
140 USA).

141 *2.4. Water sample collection and analysis*

142 Water column samples were taken from the surface layer (~10 cm below the surface)
143 and the bottom layer (~5 cm above the sediment) using a 1.5-L organic glass hydrophores,
144 and then transferred into 150 mL polyethylene bottles and 55 mL pre-weighed serum
145 glass bottles. All water samples were preserved with saturated HgCl₂ solution (~0.5 mL)
146 ([Borges et al., 2018](#); [Marescaux et al., 2018](#)) and transported in a cooler to the laboratory
147 within 4–6 hr. Approximately 100 mL water sample was filtered through a 0.45 µm
148 acetate fiber membrane (Biotrans™ nylon membranes) and the filtrate was used to
149 analyze the levels of NO₃⁻-N and NH₄⁺-N, total dissolved nitrogen (TDN), Cl⁻ and SO₄²⁻.

150 Dissolved N₂O concentrations were determined by headspace equilibration and gas
151 chromatography ([Yu et al., 2013](#); [Musenze et al., 2014](#)). Briefly, nitrogen gas (N₂; >99.9%

152 purity) was injected into each serum glass bottle to displace a 25-mL headspace and the
153 bottle was shaken vigorously for 10 min (IS-RDD3, China) to attain air-water
154 equilibrium. After waiting for 0.5 hour, a 5 mL headspace sample was extracted and
155 injected into a gas chromatograph as explained earlier (see Section 2.3). The *in situ*
156 dissolved N₂O concentrations (nmol L⁻¹) were calculated according to Yu et al. (2013)
157 and Musenze et al. (2014).

158 2.5. Diffusive N₂O fluxes across the sediment-water interface (SWI)

159 The diffusive N₂O fluxes across SWI (F_{S-W} , nmol m⁻² h⁻¹; positive values indicate
160 N₂O fluxes from sediment to water) were calculated as (Gruca-Rokosz and Tomaszek,
161 2015; Tan et al., 2014):

$$162 \quad F_{S-W} = D_S \times \Delta C / \Delta Z = (D_w \times \text{POR}) \times (C_S - C_W) / \Delta Z \quad (\text{Eq. 1})$$

163 where D_S is the diffusion coefficient of N₂O in sediment (cm² s⁻¹); $\Delta C / \Delta Z$ is the gradient
164 for dissolved N₂O concentration with depth; POR is sediment porosity; C_S is dissolved
165 N₂O concentration in sediment porewater (nmol L⁻¹); C_W is the dissolved N₂O
166 concentration in overlying water (near the sediment surface) (nmol L⁻¹); Z is diffusion
167 distance (cm); D_w is the diffusion coefficient of N₂O in water (cm² s⁻¹), which was
168 calculated as:

$$169 \quad D_w = -6.0 \times 10^{-9} T_w^2 + 10^{-9} T_w - 3.0 \times 10^{-7} \quad (\text{Eq. 2})$$

170 where T_w is the temperature in overlying water (near the sediment surface) (°C).

171 2.6. Diffusive N₂O fluxes across the water-atmosphere interface (WAI)

172 The diffusive N₂O fluxes across SWI (F_{W-A} , nmol m⁻² h⁻¹; positive values indicate

173 N₂O fluxes from water to air) were calculated as (Musenze et al., 2014):

$$174 \quad F_{W-A} = [2.07 + (0.215 \times U_{10}^{1.7})](Sc/660)^{-n} \times (C_W - C_{eq}) \quad (\text{Eq. 3})$$

175 where C_W is the dissolved N₂O level (nmol L⁻¹) in the surface water; C_{eq} is the N₂O level
176 in water that is in equilibrium with air at the *in situ* temperature; U_{10} is the frictionless
177 wind speed (W_S ; m s⁻¹) at 10 m height above the water surface according to Crusius and
178 Wanninkhof (2003); Sc is the Schmidt number for N₂O (Wanninkhof, 1992); n is a
179 constant that varies between 0.50 ($W_S > 3$ m s⁻¹) and 0.66 (for $W_S \leq 3$ m s⁻¹) (Cole and
180 Caraco, 1998).

181 2.7. Auxiliary data

182 Meteorological parameters, including air pressure (A_P), wind speed (W_S) and air
183 temperature (T_A), were collected by an automated meteorological station on site. In each
184 sampling campaign, *in situ* water salinity (Sal), temperature (T_W) and dissolved oxygen
185 (DO) were measured by a Eutech Instruments-Salt6 salinity meter (Thermo Fisher
186 Scientific, San Francisco, California, USA), a temperature meter (IQ150, IQ Scientific
187 Instruments, Carlsbad, California, USA) and a multiparameter probe (550A YSI, USA),
188 respectively.

189 2.8. Statistical analysis

190 All data were checked for normality and homogeneity of variance before further
191 statistical analysis. One-way ANOVA was conducted in SPSS 22.0 (IBM, Armonk, NY,
192 USA) to explore the effects of sampling time on various environmental variables, N₂O
193 concentrations and diffusive N₂O fluxes. Spearman correlation analysis was performed

194 to examine the relationships between diffusive N₂O fluxes (or dissolved N₂O
195 concentrations) and environmental parameters, using R corrplot and Hmisc packages.
196 The extent to which environmental variables regulated the temporal variations in
197 diffusive N₂O fluxes (or dissolved N₂O concentrations) was analysed using Redundancy
198 Analysis (RDA) in CANOCO 5.0 (Microcomputer Power, Ithaca, USA). The
199 significance level was set at $p < 0.05$ for all analyses.

200 **3. Results**

201 *3.1. Physical and chemical characteristics*

202 The physicochemical properties of porewater and surface water are shown in [Figure](#)
203 [2](#). Air, water and sediment temperatures increased from April toward July and August,
204 then decreased toward January. Across all sampling months, the mean temperature, NO₃⁻
205 -N and NH₄⁺-N concentrations in sediment porewater were 20.25±0.59 °C, 0.22±0.03 mg
206 L⁻¹ and 0.33±0.04 mg L⁻¹, respectively, which were significantly lower than in the surface
207 water (25.36±1.01 °C, 1.19±0.20 mg L⁻¹ and 0.52±0.13 mg L⁻¹, respectively) ($p < 0.05$ or
208 < 0.01).

209 *3.2. Dissolved N₂O concentration in sediment porewater and water column*

210 Dissolved N₂O concentration varied significantly over time ($p < 0.01$; [Figure 3](#)). It
211 ranged from 2.11±0.18 to 21.68±2.60 nmol L⁻¹ in the surface water, 1.83±0.11 to
212 22.04±5.54 nmol L⁻¹ in the bottom water, and 5.02±0.63 to 40.48±5.31 nmol L⁻¹ in the
213 sediment porewater ([Figure 3](#)). The mean N₂O concentration was significantly higher in
214 sediment porewater (18.94±3.05 nmol L⁻¹), followed by bottom water (9.88±2.11 nmol

215 L⁻¹) and surface water (9.29±2.12 nmol L⁻¹) ($p<0.001$). The highest N₂O concentration
216 was observed in September in the sediment porewater and in July in the surface water
217 and bottom water (Figure 3).

218 3.3. N₂O fluxes across the sediment–water–atmosphere interfaces

219 The diffusive N₂O fluxes across the SWI were always positive (ranged 1.51–48.84
220 nmol m⁻² h⁻¹; Figure 4a), indicating net N₂O releases from the sediment to the overlying
221 water. The N₂O fluxes across the SWI varied significantly between months ($F_{df=9}=4.962$,
222 $p=0.001$) with considerably higher values from August to October (Figure 4a).

223 The diffusive N₂O fluxes across the WAI of the ponds showed significant temporal
224 variations ($F_{df=9}=32.227$, $p<0.001$; Figure 4b), with higher fluxes from June to August,
225 and lower fluxes from November to January (Figure 4b). Overall, the diffusive N₂O
226 fluxes across WAI ranged from 25.06 to 507.87 nmol m⁻² h⁻¹ (Figure 4b), indicating that
227 the aquaculture ponds were an N₂O emission source to the atmosphere.

228 Across all the sampling campaigns, the mean diffusive N₂O flux across WAI (216.85
229 ± 55.52 nmol m⁻² h⁻¹) was an order of magnitude higher than that across SWI (16.00 ±
230 4.60 nmol m⁻² h⁻¹) ($F_{df=1} = 38.319$ $p < 0.01$).

231 3.4. Environmental drivers of N₂O concentrations and fluxes

232 Spearman correlation analysis showed that N₂O fluxes across SWI (or porewater
233 N₂O concentrations) correlated positively with T_s ($p<0.001$) and NO₃⁻-N ($p<0.01$), but
234 negatively with sediment salinity ($p<0.01$), porewater Cl⁻ ($p<0.05$) and SO₄²⁻
235 concentrations ($p<0.01$) (Figure 5a). N₂O fluxes across WAI (or water column N₂O

236 concentrations) correlated positively with T_A , T_W , NO_3^- -N and TDN ($p < 0.01$ or $p < 0.001$),
237 but negatively with A_P , DO, pH, salinity, Cl^- and SO_4^{2-} concentrations ($p < 0.01$ or $p < 0.001$)
238 (Figure 5b).

239 Based on the results of RDA, T_S (explaining 57.2% of the variations) and NO_3^- -N
240 (16.9%) were the variables that explained most of the temporal variations in N_2O fluxes
241 across SWI (or porewater N_2O concentrations) (Figure 6a), whereas variations in N_2O
242 fluxes across WAI (or water column N_2O concentrations) were mostly explained by T_W ,
243 (64.6%), followed by NO_3^- -N (21.4%) and TDN (8.8%) (Figure 6b).

244 4. Discussion

245 Nitrous oxide emissions from agriculture have been a main focus in climate science
246 due to the increase in agricultural land use and application of fertilizer (Del Grosso et al.,
247 2008; Shcherbak et al., 2014). While the estimates of agricultural N_2O emissions are
248 reasonably well constrained at the global level, uncertainties persist at the regional and
249 local levels (Reay et al., 2012). Earlier studies also highlighted that N_2O emissions from
250 aquaculture systems remain a critical knowledge gap, especially considering the rapid
251 expansion of the aquaculture sector worldwide (Reay et al., 2012).

252 The whiteleg shrimp *Litopenaeus vannamei* is one of the main species farmed in
253 small-hold earthen ponds along the China's coast (BFMA, 2019). The sediment and the
254 water column of aquaculture ponds present very different reduction-oxidation
255 environments to drive the different N_2O production pathways, for example, via
256 incomplete denitrification in anoxic condition and nitrification in oxic condition, using

257 NO_3^- -N or NH_4^+ -N as substrate (Hu et al., 2012, 2013; Yuan et al., 2021; Yang et al.,
258 2020d). In contrast to the expectation that the sediment is a sink of nitrogenous substrates
259 from deposition of animal wastes and excess feed (Avnimelech and Ritvo, 2003;
260 Hargreaves et al., 1998), our data showed that both NO_3^- -N and NH_4^+ -N concentrations
261 were higher in the water column than in the sediment porewater (Figure 2). This perhaps
262 reflects the high efficiency of the shrimp to convert feed to biomass and its ability to
263 recycle nutrients within the water column (Avnimelech and Ritvo, 2003; Lacoste and
264 Gaertner-Mazouni, 2016; Zhang et al., 2016). The large increase in surface-water NO_3^- -
265 N concentration in July–October was likely caused by the increased use of feed and
266 increased feeding activity of the shrimp during its summer growth burst.

267 Despite the much lower NO_3^- -N concentration in the sediment porewater than the
268 water column (Figure 2), sediment porewater N_2O concentration was comparable to or
269 more than twice that in the overlying water (Figure 3). These observations suggest that
270 N_2O was produced mainly via denitrification and accumulated in the anoxic sediment
271 (Blackburn and Blackburn, 1992), whereas N_2O production via denitrification (using
272 NO_3^- -N) or nitrification (using NH_4^+ -N) was rather limited in the water column.

273 Movement of the shrimp in the pond bottom could disturb the sediment and
274 accelerate chemical exchanges between the sediment and the overlying water. Previous
275 experimental studies have suggested that bioturbation by *L. vannamei* released nutrients
276 from the sediment and increased oxygen consumption in the overlying water (Yang et al.,
277 2017; Zhang et al., 2015); however, those measurements were made in artificial

278 enclosures where the shrimp may not behave normally. In this study, we observed that
279 sediment porewater N₂O concentration was consistently higher than that in the bottom
280 water (~5 cm above sediment) on all but two occasions (Figure 3); therefore, there was
281 no evidence of bioturbation by the shrimp in the pond that would have destroyed the N₂O
282 gradient across the water-sediment interface.

283 Based on the N₂O distributions, we calculated the N₂O fluxes across the SWI and
284 the WAI (Figure 4). The N₂O fluxes across SWI were highest in the summer months
285 (August-October), coinciding the higher porewater NO₃⁻-N concentrations and sediment
286 temperatures (Figure 2a and 2b), both of which would have increased microbial
287 denitrification activity in the sediment (Hu et al., 2012; Murray et al., 2015; Reisinger et
288 al., 2016). The N₂O fluxes across WAI increased earlier, reaching a maximum in July,
289 which was also consistent with the earlier rise in water temperature and water column
290 NO₃⁻-N concentration (Figure 2). These explanations were further supported by
291 Spearman correlation and RDA analyses (Figures 5 and 6).

292 Averaging across the study period, N₂O flux across WAI was an order of magnitude
293 higher than N₂O flux across SWI, implying that the water column played a larger role in
294 emitting N₂O from the aquaculture pond. Overall, the N₂O emissive fluxes from the
295 shrimp ponds were comparable to other aquaculture ponds, static waters (reservoirs and
296 lakes) and running waters (rivers and estuaries), but considerably less than the highly
297 eutrophic waters (Table 1). Coastal shrimp aquaculture in China is dominated by small-
298 hold earthen ponds that cover a total area of ~2400 km² (BFMA, 2019), the majority of

299 which are poorly monitored for their greenhouse gas emissions. Using data from this
300 study, we estimate that coastal shrimp ponds in China would emit 0.2 Gg N₂O yr⁻¹ or
301 4.5×10^{-3} Gg N yr⁻¹. A recent study estimated that N₂O emission from marine and
302 freshwater aquacultures in China amounts to ~ 16.7 Gg N y⁻¹ (Zhou et al., 2021).
303 Notwithstanding the uncertainties associated with the feed conversion rates and emission
304 factors used in Zhou et al.'s study, our calculations suggest that coastal shrimp ponds
305 contribute <1 % of the aquaculture N₂O emission in China.

306 Assuming the aquaculture pond ecosystem was in steady state, we may estimate the
307 N₂O production rates in the sediment and the water column compartments as follows:
308 We considered a sediment surface area of 1 m². Assuming that deposition of nitrogenous
309 substrates for N₂O production was limited to the top 15 cm sediment and given the
310 average N₂O flux of 16 nmol m⁻² h⁻¹ across SWI and an average sediment porewater N₂O
311 concentration of 18.94 nmol L⁻¹, the N₂O turnover time for a 1 m² × 15 cm sediment
312 compartment would be 17.75 hours and the equivalent N₂O net production rate would be
313 1.07 nmol L⁻¹ h⁻¹ under a steady-state condition. For a water column of 1 m² surface area
314 × 1.5 m depth (average depth of the aquaculture pond), the N₂O concentrations were
315 similar between surface water (9.29 nmol L⁻¹) and bottom water (9.88 nmol L⁻¹);
316 therefore, we used the average concentration of 9.59 nmol L⁻¹. The net loss of N₂O from
317 the water column compartment through emission to air would be the difference between
318 the average WAI flux (216.85 nmol m⁻² h⁻¹) and SWI flux (16.00 nmol m⁻² h⁻¹), i.e.,
319 200.85 nmol m⁻² h⁻¹. The water-column N₂O turnover time would be $9.59 \times 1500 \div$

320 200.85 = 71.58 hours, and the equivalent net N₂O production rate in the water column
321 would be 0.13 nmol L⁻¹ h⁻¹. There is a lack of empirical data on ambient N₂O production
322 rates in aquaculture ponds; nevertheless, our estimates fall within the range observed in
323 eutrophic coastal waters (De Wilde and De Bie, 2000; Punshon and Moore, 2004).

324 **5. Conclusions**

325 Overall, our results show that the sediment compartment and the water column
326 compartment played opposite roles in N₂O dynamics within the aquaculture ponds: The
327 sediment had a much higher N₂O production rate than the water column, whereas the
328 water column contributed a much higher overall emission to air. Environmental
329 temperatures and nitrogenous substrates were the main controlling factors in both
330 compartments. Although our data showed that coastal aquaculture shrimp ponds were a
331 net source of N₂O emission to the atmosphere, their overall contributions were relative
332 minor compared to other aquaculture operations in China; therefore, shrimp aquaculture
333 remains a preferred solution to the growing demand for animal proteins without
334 exacerbating the climate impact.

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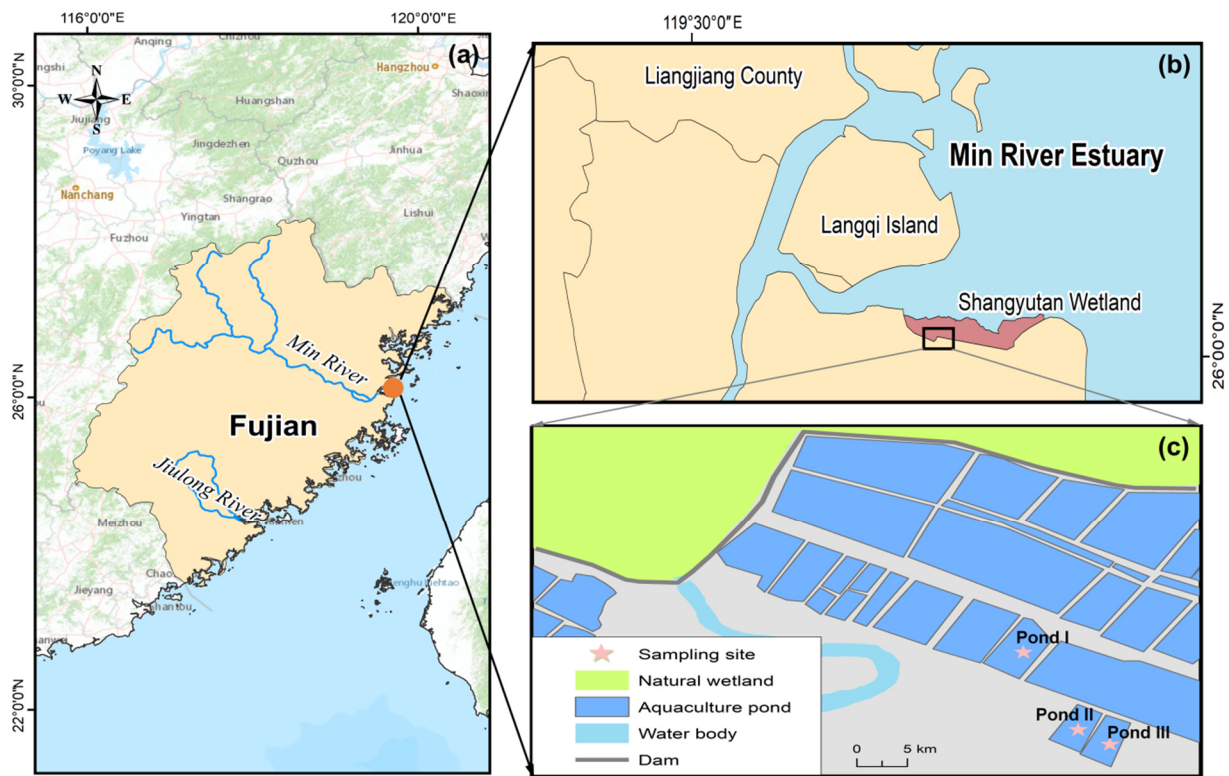
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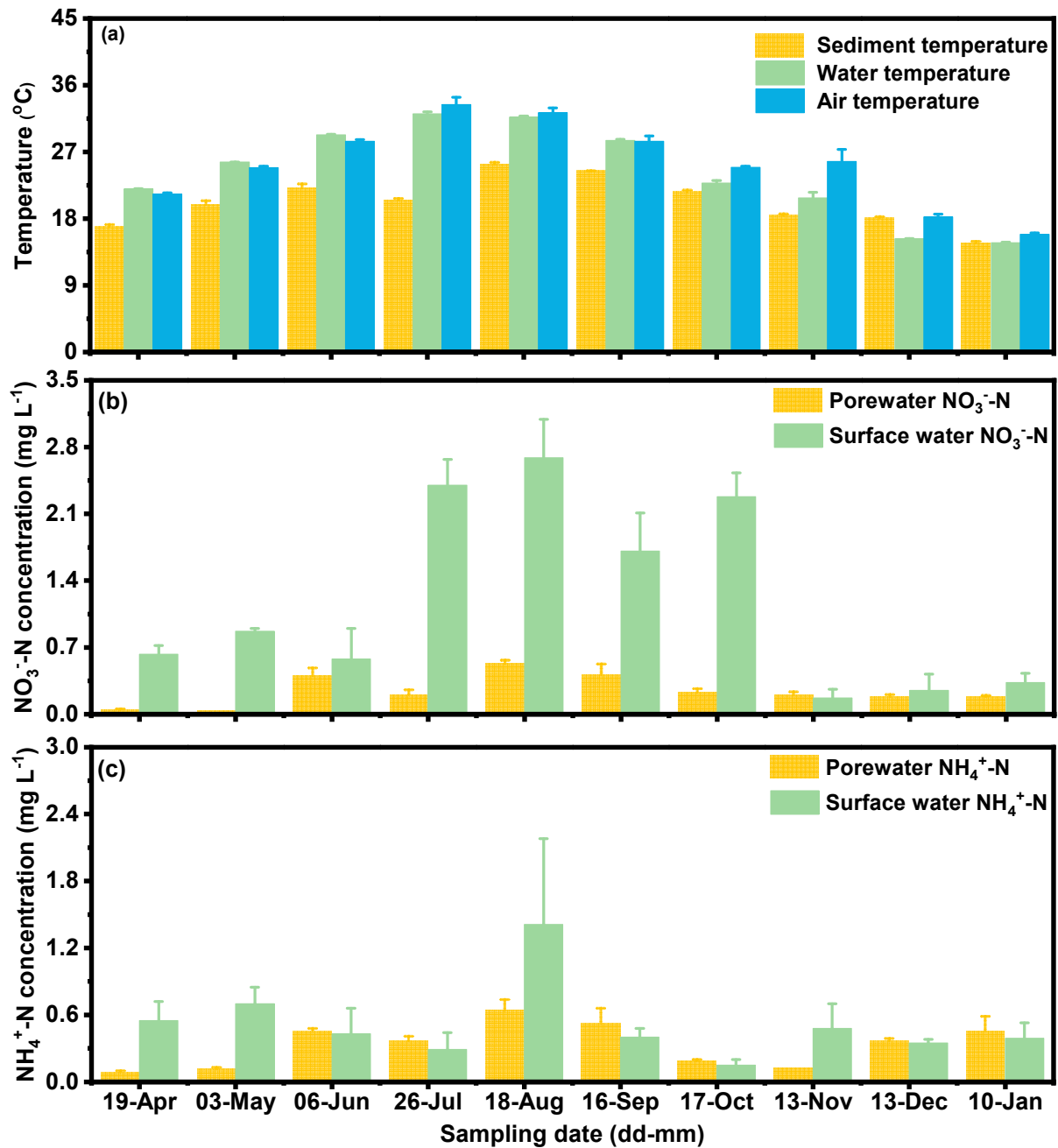
Table 1. Comparison of N₂O fluxes across the water-atmosphere interface in different aquatic ecosystems. Numbers in brackets are averages. “nd” indicates no data.

Aquatic ecosystems	Study area	N ₂ O flux (nmol m ⁻² h ⁻¹)	Reference
Aquaculture ponds	Min River Estuary, China	25.1–507.9 (216.9)	This study
	Jiulong River estuary, China	600.0–6909.1 (750.1)	Jin, 2017
	Xinghua, Jiangsu province, China	475.1–647.3 (561.6)	Fang et al., 2022
	Jurong Reservoir watershed, Eastern China	240.9–495.5 (334.1)	Xiao et al., 2019b
	Tianjin Binhai New Area, Tianjin, China	155.5–365.6 (256.6)	Hu et al., 2020
	Xinghua, Jiangsu Province, China	nd (109.2)	Liu et al., 2016
	Jiulong River Estuary, China	47.3–85.2 (66.3)	Ye et al., 2022
	Tai Lake basin, Suzhou, China	33.4–85.9 (56.5)	Yuan et al., 2021
	Wenwusha Reservoir, Min River Estuary, China	-246.1–16170.5 (2590.9)	Yang et al., 2021b
	Southeast Queensland, Australia	-3.4–80.3 (9.2)	Musenze et al., 2014
	Eguzon Reservoir, France	112.5–3229.2 (716.7)	Descoux et al., 2017
	Petit Saut Reservoir, France, and Fortuna Reservoir, Panama	5625.0–9127.3 (3800.0)	Guérin et al., 2008
Reservoirs / Lakes	Lokaa Reservoir, Finland	-79.5–259.1 (nd)	Huttunen et al., 2002
	Jurong Reservoir, Eastern China	115.9–513.6 (234.1)	Xiao et al., 2019b
	Hongjiadu Reservoir and Wujiangdu Reservoir, China	79.5–1759.1 (545.5)	Liu et al., 2011a
	Dongfeng Reservoir, China	9.1–609.1 (190.9)	Liu et al., 2017
	Lake Taihu, China	45.5–229.5 (145.5)	Xiao et al., 2019a
	Lake Mochou, Lake Tuanjie and Lake Daming, Antarctica	-809.1–1556.8 (303.0)	Liu et al., 2011b
	Poyang Lake, China	-221.1–2886.4 (nd)	Wang et al., 2017
	Lake Nakaumi, Japan	709.1–1790.9 (1070.5)	Hirota et al., 2007
	Lake Kevätön, Finland	90.0–500.1 (nd)	Huttunen et al., 2003
	China's reservoirs	nd (962.1)	Li et al., 2018
	China's reservoirs	nd (277.4)	Li et al., 2018
	Rivers / Estuaries	River Avon, River Eden and River Wensum, UK	127.3–2197.7 (912.9)

River Wensum, Norfolk, UK	nd (406.8)	Hama-Aziz et al., 2017
Chongqing metropolitan river network, China	188.6–65284.1 (10875.1)	He et al., 2017
Xin'an Tang river, China	1470.1–3130.1 (2000.1)	Xia et al., 2013
Shanghai river network, China	1040.1–2570.1 (nd)	Yu et al., 2013
River Neuse, USA	-600.0–4600.1 (nd)	Stow et al., 2005
River Millstone, USA	10.0–2140.0 (250.1)	Laursen and Seitzinger, 2004
Adyar River, India	9.1–5100.0 (nd)	Rajkumar et al., 2008
Sado estuary, Portugal	-90.9–156.8 (125.0)	Gonçaves et al., 2015
Guadalete River estuary, Spain	20.5–1200.0 (470.5)	Burgos et al., 2017
Guadalquivir Estuary, Spain	-290.9–1459.1 (354.5)	Huertas et al., 2018
Yangtze River Estuary, China	213.6–1184.1 (554.5)	Zhang et al., 2010

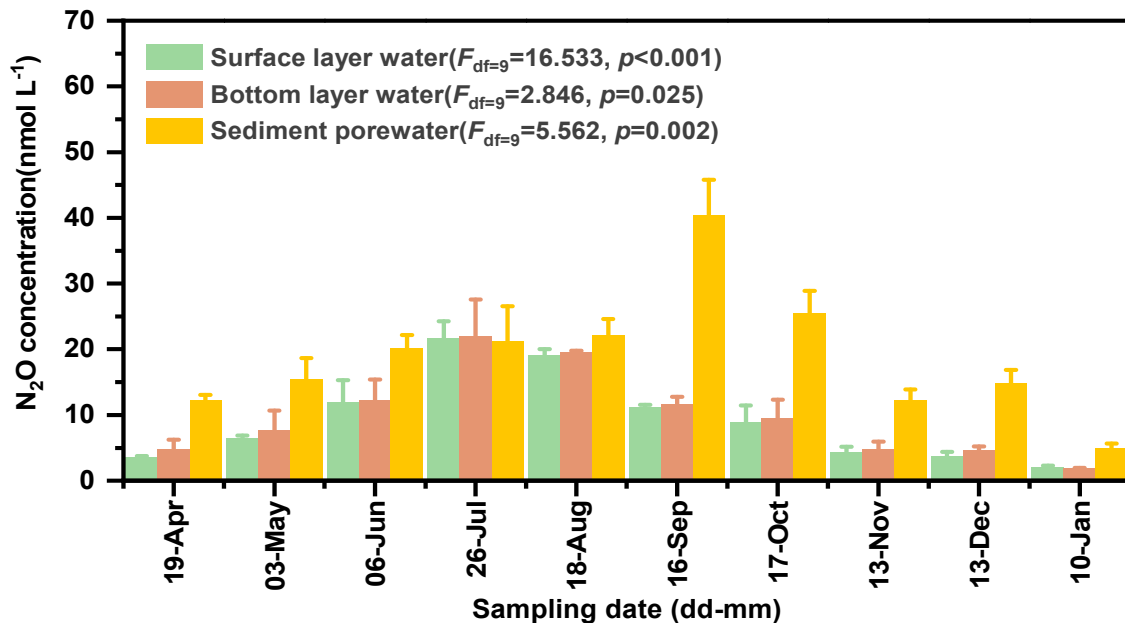


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 2 **Figure 1.** Map of the Shanyutan wetland within the Min River estuary showing the sampling
 3 sites in aquaculture ponds. The geographical coordinates of Pond I, Pond II and Pond III being
 4 26°01'43"N and 119°38'30"E, 26°01'37"N and 119°38'35"E, and 26°01'36"N and 119°38'38"E,
 5 respectively.



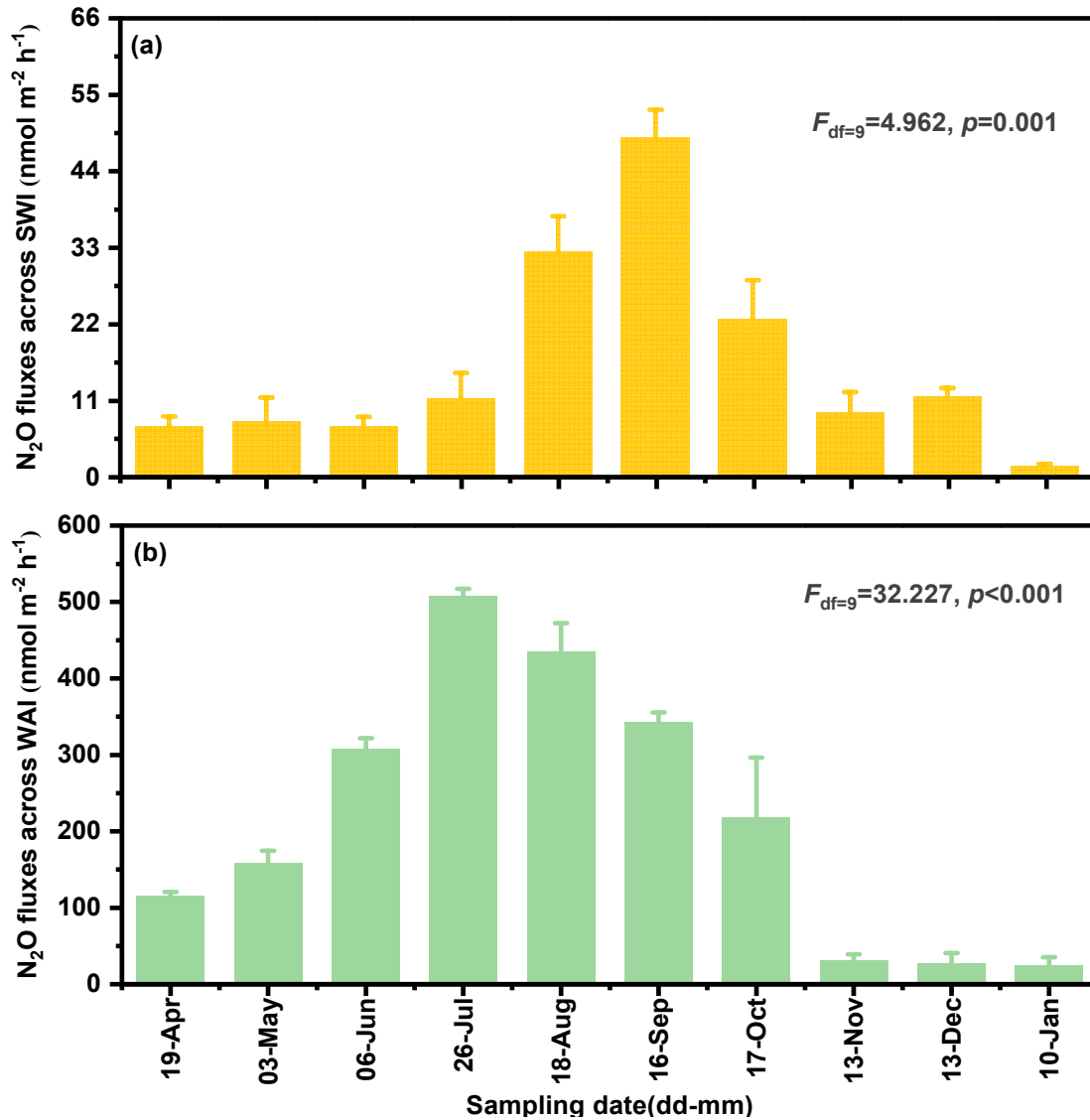
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7 **Figure 2.** Monthly temperatures, NO₃⁻-N and NH₄⁺-N concentrations in the aquaculture ponds
 8 during the farming period. The bars represent the means + 1 standard error (*n* = 3).



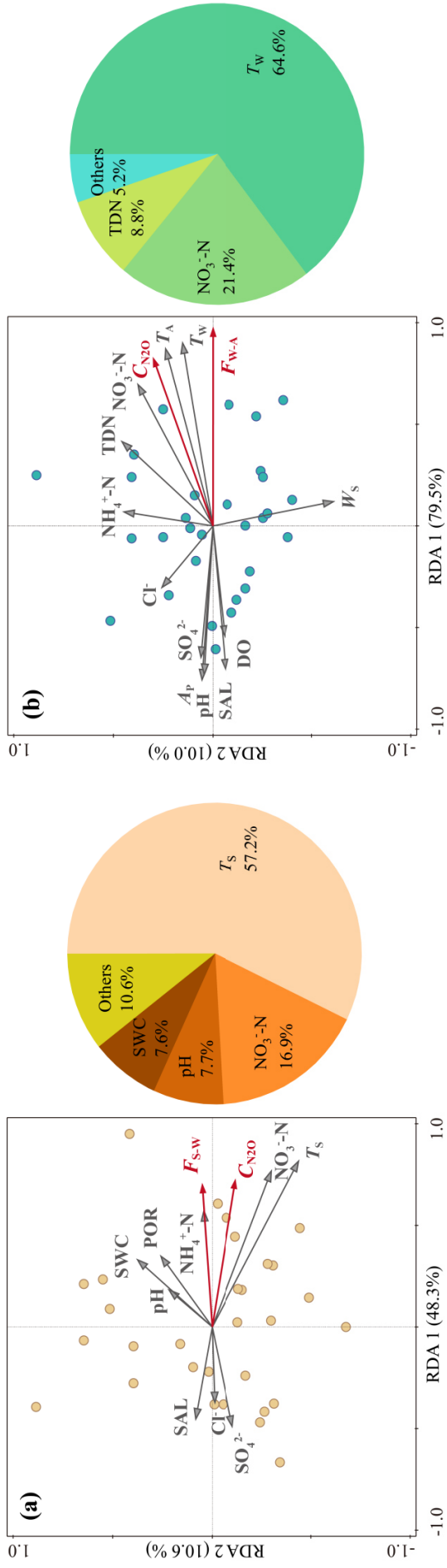
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10 **Figure 3.** Monthly dissolved N₂O concentrations in the surface layer water, bottom layer
 11 water and sediment porewater of the aquaculture ponds during the farming period. The
 12 bars represent the means + 1 standard error ($n = 3$).



13

14 **Figure 4.** Monthly diffusive N₂O fluxes across (a) sediment-water interface (SWI) and
 15 (b) water-atmosphere interface (WAI) of the aquaculture ponds during the farming
 16 period. The bars represent the means + 1 standard error ($n = 3$).



22

23 **Figure 6.** Results of redundancy analysis (RDA) of (a) N₂O fluxes across the sediment-water interface [or porewater N₂O concentration], and (b)

24 N₂O fluxes across the water-atmosphere interface [or surface water N₂O concentration], showing the loadings of the different environmental

25 variables. The pie charts show the percentages of the variance of N₂O fluxes explained by the different variables. See main text for explanation of

26 the abbreviations.