

1 **The response of steroid estrogens bioavailability to various**
2 **sorption mechanisms by soil organic matter extracted with**
3 **sequential alkaline-extraction method from an agriculture soil**

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16 **ABSTRACT:** The long-term groundwater contamination risks posed by steroidal
17 estrogens (SEs) in animal-manured agricultural soils are closely associated with the soil
18 organic matter (SOM) content and composition. In this study, the bioavailability of
19 estrone (E1) and 17 β -estradiol (17 β -E2) under different sorption mechanism in humic
20 acids (HA1 and HA2) and humin (HM) extracted with sequential alkaline-extraction
21 technique (SAET) were examined. These SOMs extracted by SAET showed various
22 properties and sorption characteristics for SEs. The alkyl carbon and condensed SOM
23 increased during SAET, but aromatic carbon decreased and the same trend for polarity.
24 Quick sorption was the major SEs sorption mechanism on HA1 and HA2, which
25 contributed more than 69%; whilst slow sorption rate was about 50% in soil and HM.
26 The $\log K_{oc}$ values were proportional to the TOC of SOM according to Freundlich fitting,
27 and the sorption capacity of sorbent for E1 and 17 β -E2 was related to the $\log K_{ow}$ values,
28 indicating that the main mechanism controlling the SEs sorption was hydrophobic
29 interaction. The larger micropore volume of HM and soil was more conducive to the
30 micropore filling of SEs. Meanwhile, the specific sorption of SEs on condensed domain
31 of SOM was the main reason for the strong desorption hysteresis and slow sorption in
32 HM and soil. The SEs degradation rate was positively correlated with the contribution
33 rate of quick adsorption and negatively correlated with the contribution rate of slow
34 adsorption, indicating that the bioavailability of SEs sorbed by hydrophobic interaction
35 was higher than that of micropore filling or specific sorption, which was also the reason
36 for the low bioavailability of SEs in HM and soil. This work confirms the regulation of
37 on-site SOM compositions and their properties on SEs sorption and bioavailability.
38 Characterization of these details is crucial for the improved prediction of long-term
39 risks to groundwater.

40 **Keywords:** Steroid Estrogens, Sorption, Bioavailability, Soil Organic Matter, Risk

41

42 **1. Introduction**

43 Steroidal estrogens (SEs) such as E1 (estrone) and 17 β -E2 (estradiol) are
44 commonly found in agricultural soils and groundwater after fertilization with livestock
45 and poultry manure (Gall et al., 2015; Fang et al., 2016; Zhao et al., 2019). Their high
46 bioactivity may cause endocrine disruption effects on the agroecological system, even
47 following exposure to low SEs concentrations (Adeel et al., 2017; Zhao et al., 2019;
48 Wei et al., 2020). Although gains have been made through organic green agriculture in
49 reducing chemical fertilizer application risks, these can be offset by greater
50 environmental risks from SEs posed by increased manure application. As a result of the
51 incomplete sorption and biodegradation in the vadose zone (Song et al., 2018; Zhao et
52 al., 2019), underlying groundwater resources are at risk of SEs contamination through
53 leaching from soils (Citulski et al., 2010; Yang et al., 2021). Improved understanding
54 of sorption and bioavailability controls on the accumulation and decline of legacy SEs
55 sources in agricultural soils and their regulation of groundwater risks is crucial.

56 The sorption mechanism of SEs in soil is complex due to the heterogeneity of soil
57 components. Especially the origin, type, and the distribution and structure of the
58 functional groups of soil organic matter (SOM), including humic acid (HA) and humin
59 (HM), play a central role in the sorption of SEs, which have a significant impact on the
60 bioavailability of SEs (Sangster et al; 2015; Alizadeh et al., 2018; Yu et al., 2020; Sun
61 et al., 2010). A few studies have confirmed that SEs sorption affinity has a positive
62 correlation with SOM content and the octanol-water partition coefficient (K_{ow}) of SEs,
63 supporting the dominance of hydrophobic interaction mechanisms
64 (Karnjanapiboonwong et al., 2010; Mashtare et al., 2011). However, further sorption
65 mechanisms must exist because of the presence of polar functional groups both in SEs
66 and humic substances (Sun et al., 2012; Gall et al., 2016). For instance, the dominant
67 role of aromatic carbon versus aliphatic carbon in SOM controlling the sorption of SEs
68 remains unresolved (Sun et al., 2008; Jiang et al., 2017). The sorption of SEs may be
69 significantly influenced by their polar hydroxyl function on the benzene ring of SEs
70 molecules or their π - π binding (Takigami et al., 2011; Bedard et al., 2014; Zhang et al.,

71 2018; Zhao et al., 2020). The feature of SEs sorption on HA or HM is often nonlinear,
72 which is attributed for the specific polar interaction (H-bonding) or pore-filling (Lima
73 et al., 2012; Zhang et al., 2017). Soil minerals and its heterogeneity also affect sorption
74 (Tong et al., 2019; Xu et al., 2020). The sorption capacity of SEs on iron oxides and
75 clay minerals is governed by the content of their ion-exchange materials
76 (Thanhmingliana et al., 2016; Shi et al., 2017), whereas the sorption on
77 montmorillonites displayed interlayer spacing decrease, thereby supporting
78 incorporation of SEs into the clay interlayer space and producing slow sorption (Sun et
79 al., 2017). Therefore, further investigation of the sorption behavior of SEs in a real
80 agricultural soil (in preference to artificial, commercial products) is needed to evaluate
81 SEs bioavailability and transport in the agricultural system. To our knowledge, related
82 research remains surprisingly limited.

83 The bioavailability of SEs in agricultural soil that are closely related to
84 biodegradation is another important factor affecting their environmental risk because
85 of their nonpersistent properties (Stumpe et al., 2010; Zheng et al., 2012). The
86 bioavailability of contaminants refers to the accessibility of microorganisms or
87 extracellular enzymes in soil and groundwater, which is controlled by the concentration
88 of directly available contaminants and the rate of conversion from potentially available
89 contaminants (Wang et al., 2021). As a result of the heterogeneity of soil components
90 and the complexity of the hydrophobic organic contaminant (HOCs) sorption
91 mechanism on SOM, there may be considerable differences in related bioavailability,
92 which leads to the different opportunities for the utilization of sorbed and free HOCs
93 by microorganisms (Zhang et al., 2018; Gámiz et al., 2018). It is generally believed that
94 only free SEs can be effectively degraded by soil microorganisms, while sorbed SEs
95 need to be desorbed from the solid phase before utilization (Fan et al., 2007). In
96 particular, soil with high SOM content may reduce the bioavailability and
97 biodegradation rates due to irreversible sorption (Lee et al., 2011). However,
98 Combalbert et al. (2010) suggested that soils with different sorption capacities have
99 little influence on the biodegradation of SEs, and the sorbed SEs could be freely
100 desorbed to the liquid phase and degraded. In addition, the degree of isolation of HOCs

101 increases significantly with the time spent in contact with the soil, a process known as
102 aging, which may increase desorption resistance and reduce the bioavailability of HOCs
103 (Xu et al., 2018). The response of SEs bioavailability to sorption mechanism and to
104 sorption aging with different SOM remains unclear, and is thus the key scientific
105 problem to be solved in the current study.

106 Thus, we hypothesized that the differences in composition and structure of SOM
107 fractions can result in different controls on the SEs sorption, thereby different
108 bioavailability of sorbed SEs. The main objectives of the present research were to (1)
109 investigate controls on the sorption and desorption of E1 and 17 β -E2 in a real
110 agricultural soil and in its component SOM fractions, (2) characterize the conformation
111 and functional groups of SOMs to gain insights into the mechanism of nonlinear and
112 slow sorption of SEs, and (3) explore the effect of sorption behavior on SEs
113 bioavailability in soil and clarify the possible mechanisms related to differences in
114 SOM properties and aging processes. The findings of this work could provide
115 theoretical support for guiding the sustainable development of agriculture and
116 controlling the risk of SEs contamination of groundwater from agricultural sources.

117 **2. Materials and methods**

118 *2.1. SEs solutions, agricultural soil, and soil bacterial suspension*

119 The selected sorbates, E1 and 17 β -E2, are commonly detected in farmyard manure
120 and agricultural soil (Adeel et al., 2017), and their physicochemical properties are listed
121 in Table S1. E1 ($\geq 99.5\%$) and 17 β -E2 ($\geq 98.4\%$) were purchased from Sigma-Aldrich
122 (USA). The stock solutions of individual SEs were prepared by dissolving each
123 compound in methanol at a concentration of 1000 mg/L before being stored at $-20\text{ }^{\circ}\text{C}$
124 prior to use.

125 An agricultural silt loam (9.71% clay, 51.91% silt, and 38.38% sand) was collected
126 from a rural farm in Shenyang, Northeastern China, where maize, vegetables, and fruits
127 are predominantly grown. In our previous investigation (Song et al., 2018), the total
128 concentrations of both SEs in the farm soil were 39.3–112.7 ng/g, compared with

129 manure at 85.6–357.3 ng/g. For this study, the soil samples without SEs used for SAET
130 of SOM and batch experiments were collected from 0–10 cm soil depth around the
131 edges of the farm where manure had not been applied; however, the soil microbial
132 suspension was prepared using the SEs-containing soil for bioavailability experiment.
133 The fresh soil was air-dried, passed through a 2-mm sieve, and stored at 4 °C before the
134 sorption and bioavailability experiments. The soil samples were treated by sterilization
135 and non-sterilization. For the sterile treatment, soils were autoclaved at 120 °C for 30
136 min three successive times.

137 The soil microbial suspension was prepared as follows: 10.0 g non-sterile soil and
138 20 mL sterile water were mixed well in a 200 mL amber reagent bottle and sealed,
139 followed by incubating at 30 °C in the dark for 3 days. After 80 mL of sterile water was
140 added, the slurry was stirred at 30 °C, 180 rpm for 1 h. Finally, the slurry was
141 centrifuged at 1000 rpm for 10 min to collect the supernatant as a soil microbial
142 suspension.

143 2.2. *Sequential alkaline-extraction technique (SAET)*

144 The SAET method, a modification of the standard method of the International
145 Humic Substances Society, was conducted for the isolation of HA and HM fractions
146 from the agricultural soil (Swift et al., 1996); in brief, (1) soil equilibration was set to
147 pH 1–2 and adjusted with 0.1 M HCl to a ratio of 10 mL liquid to 1 g dry sample, then
148 the suspension was shaken for 1 h at 70 °C; (2) residue was separated by centrifugation
149 and neutralized to pH 7 under a N₂ atmosphere; (3) the alkaline suspension was shaken
150 at 70 °C for 24 h and each supernatant was collected by centrifugation. The above steps
151 were repeated until the supernatant became light yellow or colorless, allowing the HM
152 to be obtained as a precipitate. The supernatant was acidified with 6 M HCl to pH 1–
153 1.5 and left to stand for 12 h, followed by centrifugation to separate the HA (precipitate).
154 The HA was suspended in 0.1 M HCl/0.3 M HF solution and shaken overnight. The
155 precipitate was collected by centrifugation and the HA was transferred to a dialysis tube
156 until the dialysis water gave a negative Cl⁻ test with AgNO₃, and then the HA was
157 freeze-dried. Next, 46% HF/2M HCl (1:1) was added to the HM, and the suspension

158 was shaken at 70–80 °C for 2–3 min. The precipitate was centrifuged for collection,
159 and the HM residue was washed with deionized water and then freeze-dried to obtain
160 the HM. The SAET for HA was conducted five times in this study. As the amount of
161 HA decreased with the sequential extractions, HAs obtained during the first and second
162 extraction were combined to give the sorbent HA1, and the HA fractions obtained
163 during the third to fifth extractions were combined to give the sorbent HA2. This was
164 required to fulfil the requirements for sorbent dosage in the experiments investigating
165 SEs sorption to different HA fractions.

166 *2.3. SEs detection and sorbent characterization*

167 The solid phase extraction (SPE) and gas chromatography-mass spectrometry
168 (GC-MS) approach previously developed by our team (Song et al., 2018) was applied
169 for quantification of the E1 and 17 β -E2 concentrations, and details on SEs detection
170 were provided in the Supplementary Material (S1). The C, H, O, N, and S contents of
171 the sorbent were analyzed using an Elementar Vario ELIII elemental analyzer
172 (Germany). The solid-state cross-polarization magic angle-spinning ¹³C NMR spectra
173 of all sorbents were obtained with a Bruker Avance III 600 MHz NMR Spectrometer
174 (Germany) operated at a ¹³C frequency of 100 MHz. Surface elemental composition
175 and carbon-based functionalities of the sorbent were determined using XPS (Thermo
176 ESCALAB 250Xi, USA) with a monochromatic Al KR radiation source operated at
177 225 W, 15 mA, and 15kV. Spectrum processing was accomplished with the Xpspeak41
178 software to identify the functional group corresponding to the binding energy (Nguyen
179 et al., 2009; Doskočil et al., 2015; Mylotte et al., 2015). FTIR (Perkin-Elmer 1725X)
180 was used to analyze the functional groups in soil and its organic fractions, with a
181 resolution of 4 cm⁻¹ and measurement range of 4000–400 cm⁻¹. The surface area,
182 micropore area, and micropore volume were used to examine the level of pore filling
183 and calculated from the adsorption–desorption isotherm of N₂ at 77 K by the multipoint
184 Brunauer–Emment–Teller (BET) method using a surface area and porosimetry analyzer
185 (Micromeritics ASAP 2460, USA) (Xin et al., 2012).

186

187 2.4. Batch sorption experiments

188 The batch sorption experiments of E1 and 17 β -E2 were conducted on soil
189 component fractions in screw-cap vials with aluminum foil and Teflon liners.
190 Background solutions consisted of 0.01 M CaCl₂ in deionized water to maintain a
191 constant ionic strength. The optimal liquid/solid phase ratio (25:1 for soil and 40:1 for
192 HA/HM) and initial SEs concentration range were determined through preliminary
193 experiments to obtain 20%–80% uptake of sorbate at equilibrium. For kinetic
194 experiments, the vials were placed in a thermostatic oscillator (150 rpm) for shaking
195 over 7 d at 25 \pm 1 $^{\circ}$ C with sampling throughout. Preliminary tests indicated that
196 apparent sorption equilibrium was attained within 3 d. The equilibrium sorption tests
197 were hence equilibrated for 3 d with initial concentration of SEs within the range of
198 100–3000 μ g/L. The supernatant was collected for SEs analysis after standing for 2 d.

199 Following the sorption experiments, the vials with initial concentrations of 500,
200 1000, and 2000 μ g/L were used to determine the SEs desorption. Background solution
201 was added to each vial to begin the desorption study, and these desorption steps were
202 repeated three times. All samples, including blanks, were run in duplicate. Because of
203 the negligible mass loss of sorbates confirmed in the blanks (<2%), SEs sorption uptake
204 by all samples was calculated based on mass difference.

205 2.5. Bioavailability experiment

206 To explore the bioavailability of SEs under different sorption mechanisms and
207 aging conditions simultaneously, soil, HA1, and HM were pre-equilibrated with SEs
208 solutions (E1 or 17 β -E2 1000 μ g/L, 0.01 M CaCl₂) at 25 \pm 1 $^{\circ}$ C for 3, 15, and 30 d in
209 dark, respectively. The soil, SOM and background solution were all autoclaved at
210 120 $^{\circ}$ C for 30 min three successive times in the aging processes to prevent SEs
211 biodegradation. Before the bioavailability experiment, the suspension was centrifuged
212 at 5000 rpm for 5 min, and then the supernatant was collected and an equal amount of
213 soil microbial suspension was added to the residues. The SEs content in supernatant
214 was detected to calculate the concentration of SEs sorbed on soil or SOM, which was
215 used as the initial concentration of bioavailability experiment. All treatments were

216 incubated in dark at 25 ± 1 °C, 150 rpm for 14 d. Samples containing both sorbent and
217 aqueous phases were collected at different periods and analyzed using GC-MS to
218 determine the biodegradation of SEs. Each treatment included two replicates.

219 2.6. Modelling of sorption kinetics and isotherms

220 The sorption kinetic data were fitted with a two-compartment first-order kinetic
221 model (Johnson et al., 2001):

$$222 \quad q_t = q_e [f_1(1 - e^{-k_1 t}) + f_2(1 - e^{-k_2 t})], \quad (1)$$

223 where t [h] is the sorption reaction time; q_t [$\mu\text{g/g}$] is the SEs concentration in the sorbent
224 at time t ; q_e [$\mu\text{g/g}$] is the equilibrium concentration of SEs in the sorbent; k_1 and k_2 [1/h]
225 are apparent first-order rate constants for the quick and slow sorption fractions,
226 respectively; and f_1 and f_2 are the contribution rate of quick and slow sorption,
227 respectively, where $f_1 + f_2 = 1$. The sorption equilibrium data were fitted with the
228 logarithmic form of the Freundlich model:

$$229 \quad \log q_e = \log K_f + n \log C_e, \quad (2)$$

230 where C_e [$\mu\text{g/L}$] is the equilibrium aqueous concentration; K_f [$(\mu\text{g/g})/(\mu\text{g/L})^n$] is the
231 sorption coefficient; and n is a site energy heterogeneity factor, typically used as an
232 isotherm nonlinearity indicator. The concentration dependence of the affinity of the SEs
233 to the sorbent was analyzed by calculating the solid/liquid distribution coefficients (K_D ,
234 L/g) at concentrations (C_e) equal to the aqueous solubility (S_w) of solute (at $0.001S_w$,
235 $0.01S_w$, and $0.1S_w$). The organic carbon content-normalized distribution coefficient
236 (K_{oc}) was calculated as follows:

$$237 \quad K_D = q_e / C_e, \quad (3)$$

$$238 \quad K_{oc} = (K_D / \% \text{ organic carbon}) \times 100. \quad (4)$$

239 The desorption hysteresis index (HI) was obtained by calculating the ratio of linear
240 index of the sorption–desorption isotherm, which was fitted with the Freundlich model:

$$241 \quad HI_i = N_{Di} / N_S, \quad (5)$$

242 where N_S and N_{Di} are the linear indexes of the sorption isotherm and desorption

243 isotherm, respectively.

244 The degradation of SEs was fitted to a first-order kinetic equation:

$$245 \quad \ln C/C_0 = -kt, \quad (6)$$

246 where C_0 and C [$\mu\text{g/g}$] are concentration of SEs at the initial time and time t ,
247 respectively, and k is the degradation rate constant. The half-life ($t_{1/2}$) of SEs was
248 computed from the rate constant using the following equation: $t_{1/2} = \ln 2/k$.

249 *2.7 Statistical analysis*

250 A one-way analysis of variation (ANOVA) and a least significant difference (LSD)
251 multiple comparisons ($p < 0.05$) were used to assess the significant difference among
252 different treatments (sorbents, SEs and so on). The Pearson correlation analysis was
253 performed to measure the pairwise relationship between different variables (i.e., TOC,
254 quick sorption, slow sorption, micropore volume and others).

255 **3. Results and discussion**

256 *3.1. Characteristics of SOM fractions*

257 The FTIR spectra of soil and its sequential alkaline-extraction fractions showed
258 some differences (Fig. 1). Observed functionality included (Pérez et al., 2004; Smidt et
259 al., 2007) the broad band at 3620 cm^{-1} ascribed to O-H stretching vibration; the peaks
260 at 2950 and 2860 cm^{-1} corresponding to aliphatic C-H stretching; the absorption peak
261 at 1710 cm^{-1} attributed to C=O stretching of COOH and ketones; and the peaks at 1450 ,
262 1510 , and 1645 cm^{-1} attributed to structural vibrations of C=C within the aromatic ring
263 skeleton. In addition, the strong peak at 1030 cm^{-1} was associated with stretching
264 vibration of C-O in aromatic ethers, carbohydrates, or polysaccharides, and that at 1110
265 cm^{-1} was attributed to the stretching vibrations of C-O in aliphatic ethers. Thus, soil
266 and SOM fractions all contained prominent aliphatic carbon ($3000\text{--}2800 \text{ cm}^{-1}$)
267 alongside aromatic carbon (1640 cm^{-1}).

268 The aliphatic carbon and aromatic carbon accounted for 31.2%–43.9% of the total
269 organic carbon (TOC) in bulk SOM measured by the ^{13}C NMR (Table S2; Fig. 1), which

270 was low than their surface content of 54.8%–62.1% determined using XPS (Table S2;
271 Fig. S1). These results indicated that more other O-containing polar functionalities were
272 distributed within the SOM interior, which was demonstrated by the higher bulk
273 polarity (O+N/C) compared with the surface polarity (Table S2). The total aliphatic
274 carbon content of bulk (0–109 ppm) and surface both increased with the sequential
275 extraction of HA (Table S2; Fig. 1), but aromatic carbon and O-containing polar
276 functionalities contents declined. Together these results indicate that aliphatic carbon
277 is more strongly bound to soil minerals than aromatic carbon (Zhang et al., 2014; Yang
278 et al., 2012; Wang et al. 2011). The increasing C/N values (taken to be indicative of
279 higher stability of SOM) indicated that stability increased with HA isolation (HA1 <
280 HA2 < HM; Table 1) (Zhu et al., 2005). Our work is consistent with previous studies
281 that showed that stability, compacted structure, condensation degree, and
282 intermolecular intersections of SOM increase in the following order: HA < HM <
283 kerogen < black carbon (Yang et al., 2004; Cornelissen et al., 2005).

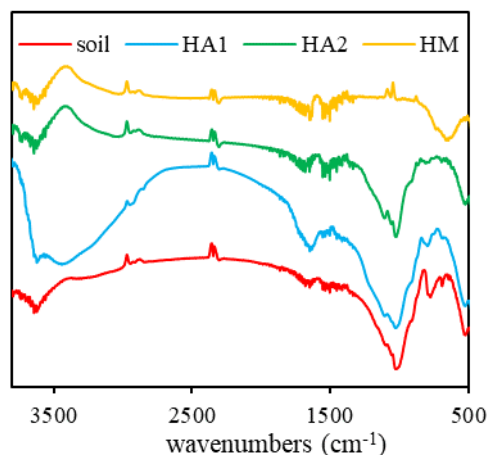


Fig. 1. FTIR spectra of soil, HAs (HA1 and HA2), and HM.

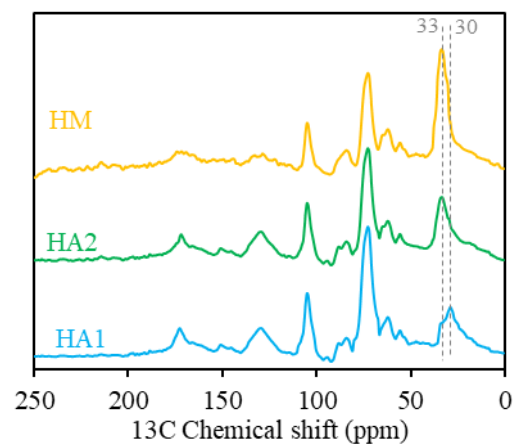


Fig. 2. Solid-state ¹³C NMR spectra of HAs (HA1, HA2) and HM.

284 3.2. Slow sorption, sorption nonlinearity, and desorption hysteresis

285 Rates of E1 and 17β-E2 sorption on soil components were relatively high, with
286 80% of the equilibrium sorption capacity reached within 24 h and the apparent sorption
287 equilibrium attained in 72 h (Table 1, Fig. 3). A typical two-stage process of quick
288 sorption was evident, followed by a slow sorption process (Fig. 3). Quick sorption
289 accounted for over 70% of sorbed mass occurring within 12 h (except for HM).

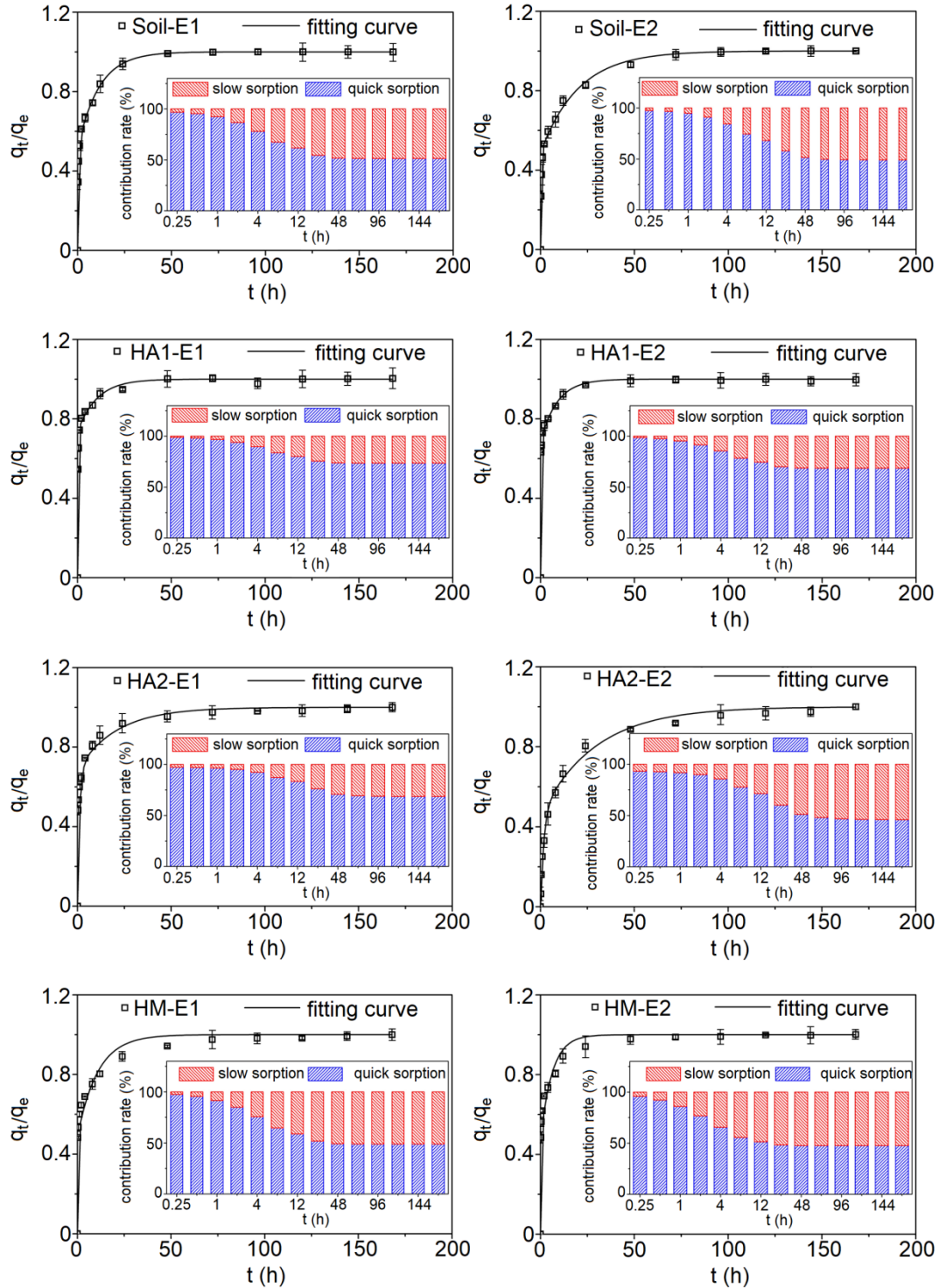
290 Thereafter, the contribution rate of slow sorption gradually increased until the sorption
 291 equilibrium was achieved. The quick sorption rate constant k_1 exceeded the slow
 292 sorption rate constant k_2 by 1–2 orders of magnitude (Table 1). For the same sorbent,
 293 the contribution rate of slow sorption (f_2) of 17 β -E2 was slightly higher than that of E1.
 294 Comparison of the f_2 value of SEs in different sorbents showed that the f_2 of HM and
 295 soil were highest, but quick sorption dominated in HA1 (Table 1). This also supports
 296 the influence of progressive SOM fractionation on SEs sorption; with the sequential
 297 extraction of SOM from HA1, HA2, and then HM, SEs sorption shifted from quick
 298 sorption (f_1) dominance to increasing slow sorption (f_2) contributions reaching similar
 299 levels.

300 **Table 1**

301 Two-compartment first-order kinetics model parameters for E1 and 17 β -E2 sorption in
 302 soil and SOM.

Sorbent	E1					17 β -E2				
	k_1	k_2	f_1	f_2	R^2	k_1	k_2	f_1	f_2	R^2
Soil	3.93	0.09	0.51	0.49	0.999	2.81	0.05	0.49	0.51	0.998
HA1	4.92	0.10	0.74	0.26	0.996	9.30	0.11	0.69	0.31	0.998
HA2	0.88	0.05	0.69	0.31	0.986	0.60	0.03	0.46	0.54	0.995
HM	13.61	0.09	0.49	0.51	0.980	15.35	0.16	0.48	0.52	0.982

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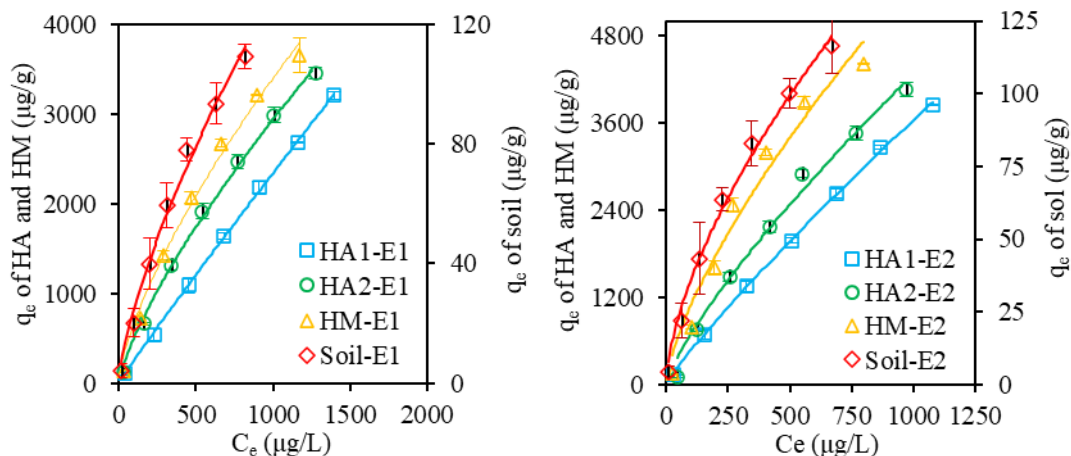
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308 **Fig. 3.** The batch data and simulated two-compartment first-order kinetic model fittings
 309 of E1 and 17β -E2 sorbed on soil and SOM, with inset graphs showing the simulated
 310 quick and slow sorption rate contributions.

311 The sorption and desorption processes fitted well to the Freundlich equation (all
 312 $R^2 > 0.96$) (Fig. 4 and S2; Tables S3 and S4), which is consistent with previous SEs

313 sorption studies. Nonlinear sorption isotherms with $n < 1$ were obtained for both SEs,
 314 especially in soil and HM. The n value of E1 was greater than that of 17 β -E2 (except
 315 for HA2) with the same sorbent (Table S3), and hence nonlinear sorption of 17 β -E2
 316 was more evident. The sorption nonlinearity of humic substances also increased with
 317 the sequential extraction of HA, consistent with the increased contribution rate of slow
 318 sorption (f_2). It is reasonable to suggest that slow sorption was primarily responsible
 319 for the observed nonlinear sorption. The sorption capacity also increased with
 320 sequential HA extraction (Table S3). The $\log K_{oc}$ values of 17 β -E2 were higher than
 321 those for E1, consistent with its greater hydrophobicity. The desorption hysteresis index
 322 (HI) was positively correlated with initial SEs concentration (Table S4). A consistently
 323 slightly lower HI of 17 β -E2, when compared to E1 for the same sorbent, indicates its
 324 greater desorption hysteresis, and hence its more difficult desorption with a higher
 325 proportion of irreversible sorption sites. The HI values of sorbents increased with
 326 sequential HA extraction, indicating that desorption became increasingly difficult and
 327 the proportion of irreversible sorption sites also increased.



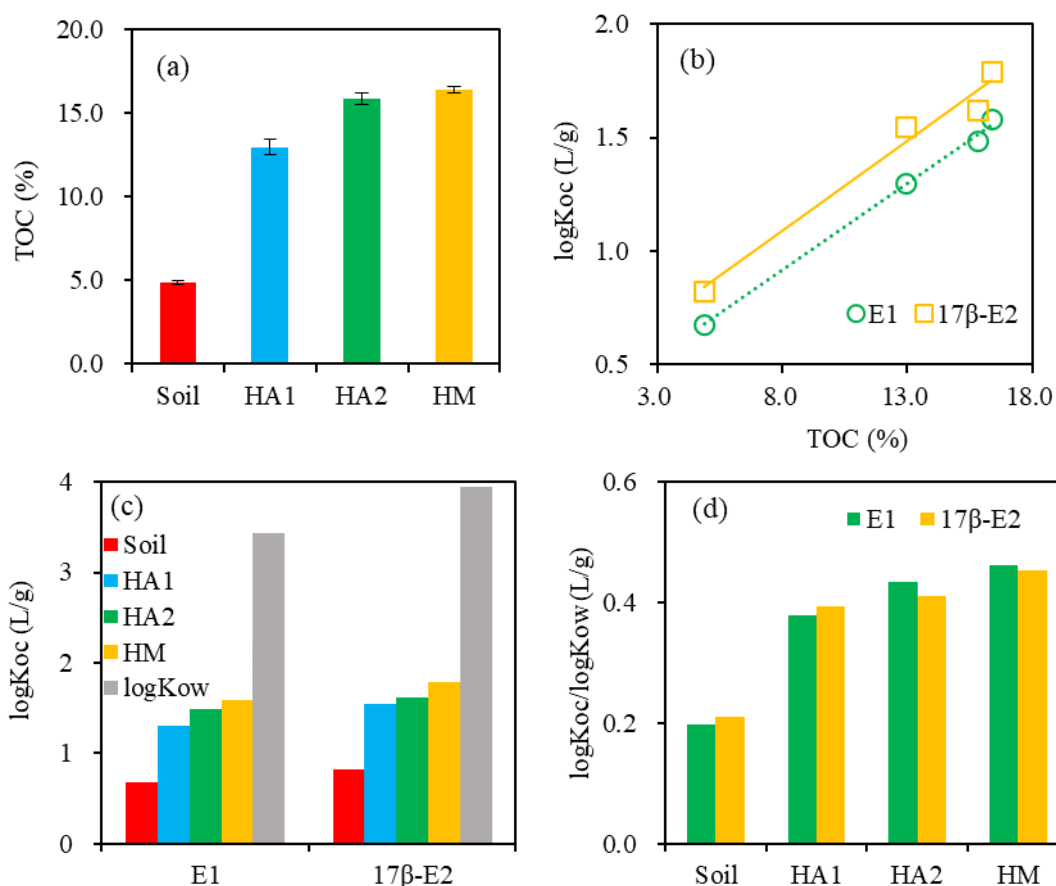
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 329 **Fig. 4.** Freundlich sorption isotherms of E1 and 17 β -E2 by soil, HAs and HM, respectively.

330 *3.3. Impact of sorbent elemental composition, polarity, functionality, and spatial*
 331 *arrangement*

332 The $\log K_{oc}$ of E1 and 17 β -E2 by soil and humic substances was positively
 333 correlated with TOC ($p < 0.05$) (Fig. 5a and 5b), and the $\log K_{oc}$ of 17 β -E2 was higher
 334 than that of E1, which was consistent with a higher $\log K_{ow}$ value of 17 β -E2 compared

335 with E1 (Fig. 5c). The hydrophobic interaction of SEs in SOM hence played a critical
336 role in their sorption. With the sequential extraction of SOM, the $\log K_{oc}/\log K_{ow}$ value
337 of 17 β -E2 was slightly higher than that of E1 in soil and HA1, but with the opposing
338 trend found in HA2 and HM (Fig. 5d). This was probably related to the functionality
339 composition of hydrophobic organic compounds generally (including SEs) and their
340 interaction with the specific sorption sites (such as hydrogen bonds and π - π bonds) of
341 condensed organic matter in soil (Zhu et al., 2004, 2005; Sun et al., 2010; Takigami et
342 al., 2011).

343 Linear and quick sorption of HOCs is generally expected in the SOM amorphous
344 organic matter domain, and nonlinear, slow sorption to the crystalline SOM (Huang et
345 al., 1997; Johnson et al., 2001). The ^{13}C NMR spectra of humic substances showed that
346 the crystalline carbon (33 ppm) became more prominent compared with amorphous
347 carbon (30 ppm) with the increased sequential extraction of humic substances (Fig. 2)
348 (Chen et al., 2017). The nonlinear sorption can hence be ascribed to the increase in
349 condensed organic matter content and the influence of these sorption sites. However,
350 the nonlinearity (n) of SEs in the bulk soil was also relatively high, which can be
351 reasonably ascribed to the complicated composition and structure and increased
352 heterogeneity. The chemical π - π interactions and hydrogen bonding also contributed to
353 the nonlinear sorption due to the aromatic ring, hydroxyl, and ketone groups for both
354 E1 and 17 β -E2 molecules (Kim et al., 2016). However, hydroxyl groups may form
355 stronger hydrogen bonds than their ketone counterparts, and the two hydroxyls within
356 17 β -E2, versus E1's single hydroxyl and single ketone group, may cause increased
357 hydrogen bonding of the former to the SOM and explain the greater nonlinearity of
358 17 β -E2 sorption observed. Therefore, the contents of SOM and its hydrophobic
359 distribution are the decisive factors controlling SEs adsorption capacity; however, the
360 difference in molecular structure between E1 and 17 β -E2 and their specific sorption on
361 SOM via the hydrogen bond and π - π bond resulted in the different sorption nonlinearity.



362

363

364 **Fig. 5.** Characteristics of humic substances and their sorption: (a) TOC variation; (b)
 365 correlations between TOC and sorption capacity of SEs ($\log K_{oc}$); (c) relationship
 366 between their octanol-water partition coefficient ($\log K_{ow}$) and $\log K_{oc}$ values of SEs;
 367 and (d) $\log K_{oc}/\log K_{ow}$ values of E1 and 17β-E2 in soil and humic substances.

368

369 Micropore area and volume of sorbents can be used to examine the sorption
 370 mechanism of pore filling, which is the main process responsible for slow sorption and
 371 subsequent desorption hysteresis due to accessibility constraints (Pan et al., 2006). The
 372 contribution rate of the slow sorption of SEs in soil and SOM were positively related to
 373 the micropore volumes of sorbents ($R^2 > 0.80, p < 0.05$), but poorly correlated with
 374 TOC content ($R^2 < 0.1, p > 0.05$) (Fig. S3a–d). It can be concluded that the slow sorption
 375 was mainly induced by pore filling. The micropore volumes of sorbents was negatively
 376 correlated with HI ($p < 0.05$) (Fig. S3d). It is speculated that more SEs can enter the
 377 sorption sites in micropores with larger micropore volumes, which cannot desorb,
 leading to enhanced desorption hysteresis (Dai et al., 2022). The sorption capacity

378 ($\log K_{oc}$) of SEs was negatively correlated with the surface and bulk polarity of sorbents
379 ($p < 0.05$) (Fig. S3e, f). Sorbents of higher polarity generally contain abundant
380 hydrophilic oxygen-containing moieties, which can provide sites for water cluster
381 formation at their surface through H-bonding (Wang et al., 2011). Water clusters may
382 further reduce the surface hydrophobicity of sorbents and the accessibility of those
383 HOCs to sorption sites, as well as compete with them; thus, the sorption of both E1 and
384 17β -E2 may be blocked by the high polar components on the sorbent surface. The
385 $\log K_{oc}$ of E1 and 17β -E2 was positively correlated with bulk alkyl carbon in SOM ($p <$
386 0.05) (Fig. S3g), which demonstrated that the aliphatic carbon content of SOM was a
387 key factor regulating the sorption of SEs. Although the soil had the highest bulk H/C,
388 its TOC content was the lowest of the studied sorbents, resulting in the lowest $\log K_{oc}$
389 value. This further indicated that SOM was the main factor controlling SEs sorption in
390 soil. The ratio of surface aliphatic carbon was lower than the bulk, indicating that the
391 aliphatic carbon was mainly distributed in the interior parts of sorbent rather than at the
392 surface. The sorption capacity ($\log K_{oc}$) of E1 and 17β -E2 was also positively correlated
393 with the content of aliphatic carbon of the sorbent surface and negatively correlated
394 with the content of surface aromatic carbon (Fig. S3h), which further demonstrates that
395 the content of aliphatic carbon in SOM and the conformation of SOM were the key
396 factors governing SEs sorption. The overall impact of sorbent elemental composition,
397 polarity, functionalities, and their spatial arrangement on SEs sorption was significant
398 and complex. Especially, the spatial arrangement characteristics of functionalities by
399 the increase of surface hydrophobic alkyl carbon, and the decrease of surface polarity
400 with sequential alkaline-extraction of SOM contributed to more SEs sorption.

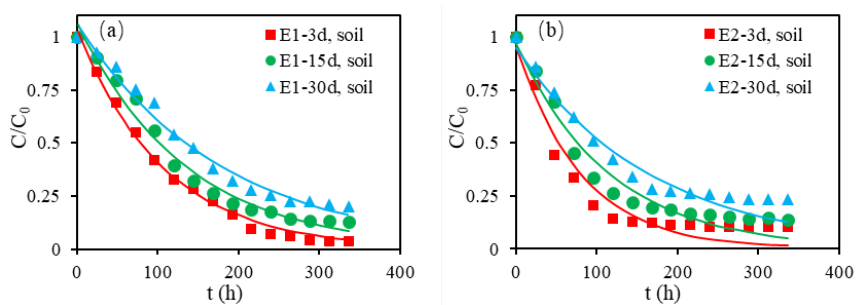
401 *3.4. Response of SEs bioavailability to sorption and aging processes with SOM*

402 The biodegradation of E1 and 17β -E2 in soil, HA1, and HM were well fitted with
403 the first-order degradation kinetics model ($R^2 > 0.85$) (Fig. 6; Table S5). The
404 degradation rate constant (k) and degradation ratio of SEs showed no significant
405 correlation with the TOC of the sorbent ($p > 0.05$) (Fig. S4), indicating that the
406 difference in SE biodegradation in soil and SOM could not be attributed to the TOC

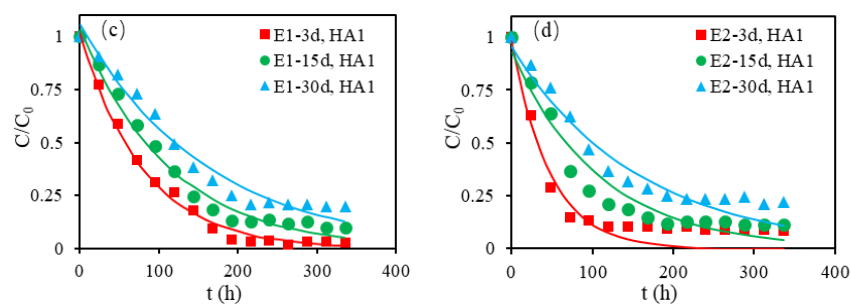
407 content. The HOCs are first sorbed on the soil surface, into the large pores, or through
408 hydrophobic interaction into amorphous organic matter of SOM. This is followed by
409 the slow and gradual adsorption of SEs into micropores or condensed organic matter of
410 SOM that cannot be entered by microorganisms. Therefore, it could be inferred that the
411 bioavailability of SEs by quick sorption is higher than that of slow sorption. The content
412 of condensed organic matter increases with sequential alkaline-extraction, and the k
413 value and degradation ratio of E1 and 17 β -E2 were both positively correlated with the
414 contribution rate of fast sorption, but negatively correlated with the contribution rate of
415 slow adsorption (Table 2). This indicates that the rapid degradation in the early stage is
416 associated with the SEs in the dissolved phase or in fast sorption, as both gradually
417 reduce as the degradation rate decreases. Subsequently, the SEs of slow sorption are
418 used by microorganisms until desorption stops, giving a stable degradation ratio. This
419 further indicates that the irreversible sorption and strong desorption hysteresis of
420 estrogen on HM with the maximum content of condensed matter organic carbon and
421 micropore volume contribute to the low bioavailability, which is the key controlling
422 factor led to the differences in SEs bioavailability among these sorbents. Overall, these
423 results suggested that the hydrophobic interaction was the key mechanism of SEs
424 sorption and was characterized with quick sorption and high bioavailability. However,
425 micropore filling and its specific sorption in condensed SOM were likely the main cause
426 of the slow or irreversible sorption and desorption hysteresis, which also contributed to
427 the low bioavailability.

428

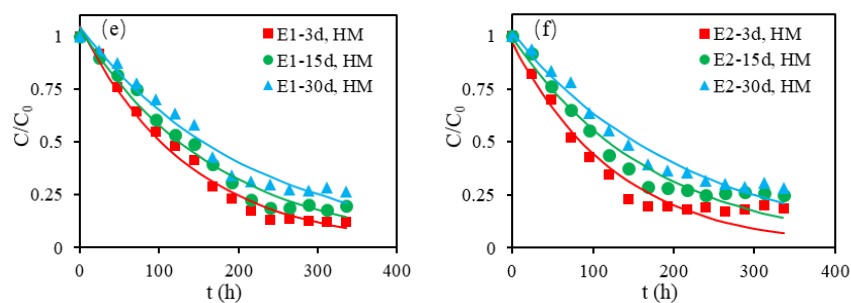
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431



432 **Fig. 6.** Effect of the aging time (3 d, 15 d, and 30 d) on E1 and 17 β -E2 degradation in
 433 soil and SOM fractions (HA1 and HM).

434 With the increase of aging time, the bioavailability and biodegradation rate
 435 constant decreased, and the degradation half-life increased (Table S5). In the aging
 436 process, the bioavailability of SEs was greatly reduced because the SEs have more time
 437 to diffuse into soil micropores and condensed organic matter, leading to an increase in
 438 the proportion of irreversible sorption and the blocking of the exposure pathway of SEs
 439 to microorganisms. This is confirmed by the well-developed micropores and high
 440 content of condensed organic matter in HM. For the same sorbent, the degradation rate
 441 of 17 β -E2 was lower than that of E1 at different aging times. This may be related to the
 442 strong hydrophobicity of 17 β -E2 and its strong affinity with soil and SOM. In addition,
 443 the 17 β -E2 molecule contains two hydroxyl groups and has a stronger specific binding

444 ability (such as hydrogen bonding) to sorbents than E1, which may be the reason for
445 the lower bioavailability of 17 β -E2.

446 **Table 2**

447 Relationship between sorption kinetic parameters and degradation kinetic parameters
448 after aging 3 days.

SEs	sorbents	k (h ⁻¹)	degradation ratio (%)	contribution rate of quick sorption	contribution rate of slow sorption
E1	HM	0.007	88.0	0.49	0.51
	Soil	0.009	96.3	0.51	0.49
	HA1	0.012	97.2	0.74	0.26
17 β -E2	HM	0.008	81.5	0.48	0.52
	Soil	0.012	89.8	0.49	0.51
	HA1	0.021	91.7	0.69	0.31

449 **4. Conclusions**

450 Application of soil sequential extraction techniques provide an important means
451 of assessing the contributions of the various SOM components to the sorption,
452 desorption, and bioavailability of SEs in real agricultural soils. These findings provide
453 an improved understanding of the risks posed to agricultural environment by SEs. The
454 sorption of E1 and 17 β -E2 was dominated by hydrophobic partitioning and quick
455 sorption. However, this does not negate the importance of other sorption mechanisms.
456 Specific sorption (hydrogen bond and π - π bond) interactions of SEs with the SOM and
457 micropore filling were responsible for nonlinear and slow sorption, as well as
458 desorption hysteresis. There was a significant effect of sorption on the bioavailability
459 of SEs. In particular, the irreversible sorption in microporous and condensed organic
460 matter improved the environmental persistence of SEs. This effect was significantly
461 enhanced with an increase in aging time.

462 The gradual accumulation and distribution of SEs in agricultural soils receiving
463 manures, and the potentially greatly delayed release of SEs, are each influenced by the
464 balance between sorption on the SOM composition and the degradation of
465 microorganisms. Even with assumptions of simple linear sorption control, lag times of

466 decades have been predicted by Gall et al. (2016) for the release of SEs from
467 agricultural soils, with attendant risks posed over these timeframes to groundwater. The
468 results of the current study suggest that these timeframes will be further extended
469 because of the need to additionally consider the hysteretic, slow, non-linear desorption
470 of SEs from the soil matrix. In particular, a fraction of SEs is subjected to pore filling
471 sorption within any crystalline carbon present, which was characterized by high
472 persistence due to the low bioavailability. Under these conditions, establishing the SEs
473 sorption on different SOM fractions associated with the response of SEs bioavailability
474 is necessary to predict the long-term risks of SEs in agricultural environment.

475 **CRedit authorship contribution statement**

476 **Xiaoming Song:** Experiment, Data curation, Writing - original draft, Writing -
477 review & editing. **Zhipeng Zhang:** Experiment, Data curation, Writing - review &
478 editing. **Yujuan Wen:** Experiment & SEs detection. **Wei Zhang:** Writing - review &
479 editing. **Yi Xie:** Modeling. **Nan Cao:** Data curation, Formal analysis. **Dong Sun:** Data
480 curation. **Yuesuo Yang:** Writing - review & editing, Supervision.

481 **Declaration of competing interest**

482 We declare that the research was conducted in the absence of any commercial or
483 financial relationships that could be constructed as a potential conflict of interest.

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489 **Appendix A. Supplementary data**

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