

Modelling of Bimodal Wave Overtopping of Coastal Defences



Stephen Tunde Orimoloye

Submitted to Swansea University
in fulfilment of the requirements for the
Degree of Doctor of Philosophy

2020

Declaration

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Signed: *stephen Orimoloye* (candidate)

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Statement 1

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I would like to dedicate this thesis to my loving parents

Acknowledgements

I will like to principally acknowledge and appreciate my supervisory Team, Professor Dominic E. Reeve and Professor Harshinie Karunaratna for their unending support during this research period. I thank them for their guidance and patience. I appreciate the assistance of Dr. Jose Horrillo-Caraballo during both my numerical and experimental works. I place much respect on my colleagues in the Energy and Environmental research group for their help and guidance during these research periods. I would also like to acknowledge the contribution of professor Thomas Lykke Andersen and Professor Peter Frigaad from the Aalborg University, Denmark for the generously-articulated workshop on the generation and analysis of waves in physical models conducted in their university during the course of my study. I sincerely appreciate Professor Javier Lara from the Institute of Hydraulics, Cantabria for providing the codes for the computational model. The PhD project was funded by the Petroleum Technology Development Fund (PTDF) with Award Ref Number: PTDF/ED/OSS/PHD/TOS/842/16.

Abstract

Wave overtopping from a few hours of violent storms can cause damage to infrastructure. Extensive knowledge is available on the prediction of wave overtopping of smooth, sloping and vertical impermeable seawalls. However, existing prediction methods provide less certain predictions when the sea state has a bimodal character, such as when there are strong components of both wind and swell. This thesis improves wave overtopping study of seawalls under storm-driven bimodal seas. Firstly, we developed an energy-conserved bimodal spectrum which contains the same energy as the unimodal spectrum derived from a sea state for easy comparison. The peak periods of the bimodal spectrum is occurring at four different spectra peak periods (11, 15, 20 and 25 secs, respectively), swell magnitude is defined with percentages (0, 25, 50, and 75 percentages). Both the unimodal and the bimodal spectra were used to perform wave overtopping experiments in a laboratory tank and numerical flume. Different bimodal spectra were created from the same sea state but the peak periods of the swell in the resulting bimodal spectrum were varied while retaining the same energy content. Bimodal wave sequences are derived from these spectra and used to investigate the overtopping of sloping seawalls. The reflection performances of the impermeable slopes were analysed in present study. The dimensionless overtopping rate for a range of seawalls was measured for bimodal storm conditions of a minimum of 1000 waves. Results for unimodal conditions agree closely with predictions using the EurOtop formulae. The current standard suggested in EurOtop (2018) underpredicts the observed wave overtopping rates for bimodal conditions even when using the recommended parameters for such cases. Under these conditions, it was found that there is a dependence between mean overtopping rate and swell occurrences. New sets of empirical influencing factors to describe these dependencies have been determined in terms of seawall designs at locations that are exposed to local storm waves and open Atlantic oceans.

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List of Symbols

Roman Symbols

α	Slope angle
α_f	Irregularity factor of the wave
β_e	beta due to swell and wind
δ_{ij}	Kronecker coefficient
ε	Dissipation rate of the turbulent energy
γ_j	spectrum peak enhancement factor for swell and wind waves
$\gamma_{n,ij}$	The damping coefficient
κ	Correlation coefficient based on first-order Markov's chain
λ	Groupiness parameter of the wave
μ	Mean of the distribution
μ_v	Molecular viscosity
μ_{ab}	Artificial viscosity
μ_{eff}	efficient dynamic viscosity
ω	Angular frequency of the wave
$\bar{\tau}_{ij}$	Mean viscous stress tensor
\bar{E}	Mean experimental dimensionless overtopping
\bar{N}	Mean numerical dimensionless overtopping
\bar{u}_1	Mean velocity

Π_{ab}	Artificial viscosity
ρ	Density of sea water
ρ_f	Fluid density
σ	Standard deviation
σ'	Coefficient of variation
σ_s	coefficient of surface tension
τ	Mean time differences between two successive crests and troughs
τ_{ij}	Total stress in the system
θ_r & θ_l	Additional model openness functions
ε	Broadness factor of the wave
ξ	The breaker parameter based on S
$\xi_{m-1,0}$	The breaker parameter based on $S_{m-1,0}$
ξ_{op}	The breaker parameter based on S_{op}
ξ_o	The breaker parameter based on S_o
$A\&B$	EurOtop's fitted coefficients for different structure slopes
A_s	Modified A-coefficient due to swell percentage
B_s	Modified B-coefficient due to swell percentage
BE	Bias error
CD	Central difference
CI	Confidence interval
D_{50}	Nominal Diameter of rock
$E(f)$	Total spectral energy in the spectrum
E_i	Experimental dimensionless overtopping
F	Frequency

f^n	Frequency in line with order of moment
F_d	Damping force
F_r	Repulsive Force
$f_{i,j}$	Peak frequencies of the swell and wind waves
$F_{n,ij}$	total normal force
FFT	Fast Fourier Transform
g	acceleration due to gravity
g_i	acceleration due to gravity due to the i^{th} component
$h(t)$	Flow thickness of overtopping wave
H_i	wave height
H_{m0S}	Spectral wave height due to swell sea
H_{m0W}	Spectral wave height due to wind sea
H_{m0}	Spectral wave height of the wave
H_s	Significant wave height of the entire wave
h_{toe}	Water depth at toe of the structure
i,j,k	dimensions of motion
$IFFT$	Inverse Fast Fourier Transform
k	The turbulent kinetic energy
K_r	Reflection coefficient
K_u	Kurtosis of the wave
$k_{n,ij}$	The stiffness parameter
L	Wave length of propagated wave
L_o	Deep water wave length
$L_{m-1,0}$	Spectral wave length in deep water

L_{om}	Mean wave length in deep water
L_{op}	Peak wave length in deep water
$m_1 \& m_2$	Slopes of the gradient of lines of empirically fitted coefficients
m_n	Spectral moments determined at different order $n = -1, 0, 2, \& 4$
mod_{fac}	Model factor of calibration
Mod_{value}	Elevation recorded at model scale
N	number of individual waves
n	Order of moment defined as -1, 0, 2, & 4
N_i	Numerical dimensionless overtopping
NSE	Nash–Sutcliffe efficiency
p	Pressure field
p'_i	Pressure fluctuations
P^*	Pseudo dynamic pressure
PPE	Poisson pressure equation
q	Average overtopping discharge per meter width of a structure
$q(t)$	Instantaneous wave overtopping
q^*	Non-dimensional overtopping rate
$q^*_{measured}$	Measured overtopping rate
q^*_{pred}	Predicted overtopping rate
Q_p	peakedness parameter
R^*	Non-dimensional crest freeboard
R^2	Coefficient of determination
R_c	Crest freeboard
R_c/H_{m0}	Non-dimensional crest freeboard $\frac{R_c}{H_{m0Toe}}$

R_s	Quality parameter of the wave
$RMSE$	The root mean square errors
RSD	Relative standard deviation
S	Wave steepness H/L
$S(f)$	Total energy in the spectrum
S_0	Swell percentage
S_k	Skewness of the wave
S_{0m}	Wave steepness based on L_o and $T_m = \frac{H_m 0}{L_{0m}}$
S_{0p}	Wave steepness based on L_o and $T_p = \frac{H_m 0}{L_{0p}}$
S_{bim}	Bimodal spectrum
$S_{m-1,0}$	Wave steepness based on L_o and $T_{m-1,0} = \frac{H_m 0}{L_{m-1,0}}$
S_{ss}	Spectrum due to swell
S_{ss}	Spectrum due to wind
σ	Standard deviation
T	Wave period
t	time steps in secs
T_i	period of wave
$T_{m-1,0deep}$	$T_{m-1,0}$ determined at deep water
$T_{m-1,0toe}$	$T_{m-1,0}$ determined at the structure's toe
$T_{m-1,0}$	Spectral wave period defined by $\frac{m-1}{m_0}$
$T_{m0,1}$	Mean wave period defined by $\frac{m_0}{m_1}$
$T_{m0,2}$	Mean wave period defined by $\sqrt{\frac{m_0}{m_1}}$
T_m	Mean wave period
$T_{Sm-1,0}$	Spectral swell-wave peak period

$T_{Wm-1,0}$	Spectral wind-wave peak period
u'_i	Turbulent velocity
u_i	velocity in x-direction
u_j	velocity in z-direction
ν	Narrowness parameter of the wave
$\nu(t)$	Depth averaged velocity
V_a	Volume of air
V_f	Volume of the fluid
ν_i	Eddy viscosity
VOF	Volume of fluid function
x_i	x coordinate measured along x-axis in the domain
D	Discretised Domain
F	VOF Function
K	Kernel Function
P	Pressure
W	Modified Kernel Function

Chapter 1

Introduction

1.1 Coastal Flooding from Wave Overtopping

Coastal flooding may occur whenever there is extreme wave overtopping of coastal defences during storms. A large proportion of the world's coastlines are exposed to local storm waves and open oceans (Figure 1.1). Long period swell waves can be present, leading to bimodal wave conditions.

Extreme wave overtopping events often accompany storms. The occurrence of wave overtopping generally threatens human lives, causes damage to infrastructures and properties. With the ever-increasing threat of climate change, coastal areas are more vulnerable to storm-driven overtopping events. Attentions have drawn towards the accuracy of prediction of wave overtopping in coastal structure and seawall designs. There is no doubt that modelling has improved the coastal seawall design, predictions and have saved construction costs.

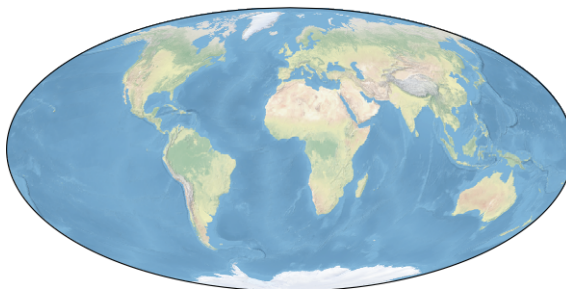


Figure 1.1 *A perspective of world's vulnerable coastlines under the influence of swell waves (Scitools, 2019)*

Sea defences are customarily being placed at shores to dissipate wave energy and prevent wave overtopping (see, van der Meer (2002), and van der Meer et al. (2016)). Most existing coastal defences have performed significantly reliably under usual wind-driven sea conditions in the absence of storms. Most recorded failures occur during

stormy sea conditions. Several processes occur during wave propagation. Obstacles in the path of waves create diffraction, including attenuation dispersion shoaling refraction refractions.

During storms, locally generated wind waves would naturally be combined with long-period ocean swells to produce bimodal waves. It has been noted that a bimodal sea state could be the worst-case sea conditions that sea defences may encounter. Currently, a great deal of research guidance is available on wind-driven wave overtopping caused by unimodal sea states; there is fewer available data focusing on bimodal seas. Thus, good knowledge of physics behind bimodal wave overtopping is vital to improve the design standards and reliability of coastal structures.

1.2 Waves Characterisation

The air-sea interaction concept has been a rapidly growing concept since the 1970s. Waves are generated in deep water under the influence of the wind. The generated waves are then propagated from the offshore to the shoreline while transferring energy along the way. As described in Goda (2010), swell waves are formed from wind-wave attenuation as propagates distances away from generating region.

The travelling waves consists of infinite number of wavelets (see, Goda et al. (1975) which are distributed across different frequencies and directions. Actions of storms would generally increases the sea surface fluctuations which have been described as sea randomness (e.g Ochi (2005)). Also Reeve et al. (2015) suggests that a random sea can be attributed to the transfer of energy from the wind to the sea. It therefore suggests that energy transfer and wave transformations would occur while creating more fluctuations as waves approach the shoreline.

Times series representation of random waves has been observed to be inefficient because wave-by-wave analysis of random waves appears impossible and unrealistic (Goda, 2010). The stochastic approach using the energy flux model is often preferred because the potential and the kinematics of random waves can be better represented by using the wave spectral density function (Ochi, 2005). Moreover, waves can be significantly represented in the energy spectrum phase. In this case, we can define a sea in only a form of spectral characteristics of the spectral shape and the corresponding moments computed from the wave energy spectrum. As highlighted by Ochi (2005), spectral shapes could vary concerning the severities of fetches, wind velocity, and duration of the wind blowing in the propagation area. The stochastic prediction of a linear and non-linear system is important to determine the information on the intrinsic

spectral analysis and the resonant condition of coastal seawalls with waves (Goda, 2010; Ochi, 2005).

At present, the most widely used techniques of describing wave spectrum depends largely on H_{m0} and $T_{m-1,0}$ or $T_{m0,2}$. However, there is more analysis of the spectrum and calculations that determine the type of sea-state. Spectrum representations of random seas are required for engineering applications and analysis. With spectrum analysis of waves, mixtures of swell and the wind can be easily identified from the energy peaks in the spectrum. Spectra representation of wind waves possess simple configuration and are easier to predict because they are characterised by one spectral peak (Unimodal spectrum) with one significant wave height and one peak period Battjes and Groenendijk (2000); Biesel F (1951); Burcharth (1978); Didier et al. (2014). A bimodal (Double-peaked) spectrum is usually formed during the transition from wind sea to swell because the wind speed would rise and start generating another wind sea after exceeding the separation frequency.

1.3 Modelling of Wave-Structures Interaction

The computational modelling techniques involve the use of computer models to investigate complex phenomena by incorporating various physics involved in the process. Conversely, a physical model replicates the actual flow parameters physically by applying scaled-up version of key parameters in the process. While a larger number of different model configurations can be easily initiated within a computational model, physical models are generally limited in terms of cost and time. However, accurate simulations in the numerical models can be computationally expensive.

Modelling techniques (computational and physical) have been applied extensively to investigate wave-structure interactions and wave overtopping of coastal structures under unimodal sea conditions, (e.g. Briganti et al., 2011; Franco et al., 1994; Goda et al., 1975; Owen, 1982; Tuan and Oumeraci, 2010). These studies provide guidance and data for wave overtopping estimates of coastal defences. (De Rouck et al., 2009a) provide a summary of numerous waves overtopping datasets compiled on different physical experiments covering various seawall geometries. More details of these studies can be found in (De Rouck et al., 2009a; Franco et al., 1994; Troch et al., 2004). These databases have been valuable and have been presented by (van der Meer et al., 2005, 2009) or more comprehensively in Pullen et al. (2007).

Recently, a series of wave overtopping tests have been performed at Ghent University (Belgium). These are presented in Gallach Sanchez et al. (2016) and Victor and

Troch (2012) whose focus is primarily on steep low-crested coastal seawalls in deep and shallow water conditions. Based on these, and other studies, some significant modifications to wave overtopping formulations have been made (van der Meer et al., 2018; van der Meer and Bruce, 2013). Concerning wave overtopping estimates for bimodal seas, EurOtop (2018) suggests using $T_{m-1,0}$ to account for bimodality as this weights the wave period towards higher values. While previous work on this topic is scarce, notable contributions include Hedges and Shareef (2003); Kashima et al. (2010); P. J. Hawkes and T. Coates and R. J. Jones (1998) who presented physical model tests, and Thompson et al. (2017) who presented a numerical assessment of overtopping performance of a simple coastal seawall under bimodal seas. The influence of wave directionality on wave overtopping oblique wind and swell waves was recently investigated by van der Werf and van Gent (2018). They found that the presence of swells with low steepness contributes more to wave overtopping in obliquely intersecting wind and swell seas.

1.4 Wave Overtopping Predictions

Prediction of wave overtopping is a crucial component of coastal structure and seawall design. The EurOtop tests (and most others) have been performed with unimodal conditions. Wave overtopping of sea defences occur when the local still water level *SWL* exceeds the crest freeboard R_c of the structure (Figure 1.3).

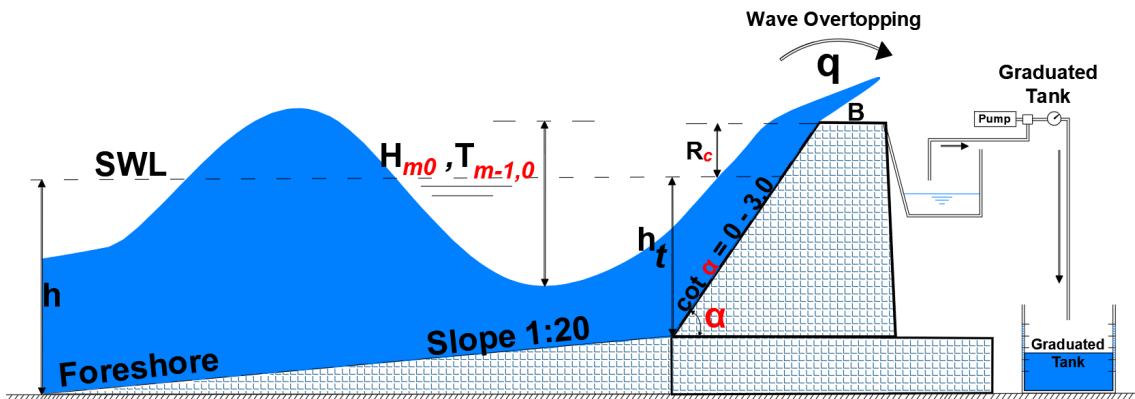


Figure 1.2 A schematic representation of wave overtopping phenomenon

Wave overtopping Q can be influenced by increasing random fluctuations of the sea surface. The Eurotop manual (van der Meer et al., 2018) acknowledges that the presence of swell can increase overtopping and recommends the use of $(T_{m-1,0})$ to account for

this. Also, the primary purpose of coastal defences is to provide efficient protection to the coast. It could be in the form of embankments, concrete, and composite structures.



Figure 1.3 Selected photographs of Aberystwyth Promenade, Wales, before and during the overtopping events on Dec. 2014

According to HR Wallingford report in the UK coastlines alone, coastal defences are protecting up to 23 per cent of the English coast, approximately 860 kilometres. Coastal defences provide efficient protection against erosion, inundation of the hinterland, and wave overtopping. Figure 1.3 demonstrates the relationship between a safe coast, wave overtopping as related with the 3-hour storm occurrence in Aberystwyth on December 2014. There is a significant failure of coastal defences during these 3 hours against wave

overtopping due to the extreme swell storms. The wave action causes lots of damages of coastal defences (Figure 1.4). The failure occurrence in England and Wales costs the UK government up to 1500 million pounds during this time (Thompson et al., 2016).

Sea states are irregular, being composed of a combination of harmonics with varying frequencies and directions, especially for shorelines exposed to very large fetches. These locations are influenced predominantly by wind and swell waves, and the energy characteristics are bimodal (Brodtkorb et al., 2000; Rychlik et al., 1997b). In these locations, direct application of overtopping formulations should include the presence of swell waves to prevent under-design. As observed by Goda (2010), random states are mostly created by the influence of the wind (from the local wind sources) and grow systematically until they decay.



Figure 1.4 An example of damage by the 3-hour storm occurrence in Aberystwyth on Dec. 2014

The decay processes are gradual and accompanied mainly by the transfer of energy from higher to lower frequencies (in the form of swell). Also described in P. J. Hawkes and T. Coates and R. J. Jones (1998) and Reeve et al. (2015), combinations of swell and wind waves would often result in more randomness of the sea. The few experiments of wave overtopping with bimodal conditions (e.g. Hedges and Shareef, 2003; P. J. Hawkes and T. Coates and R. J. Jones, 1998) suggest that overtopping can be enhanced relative to what would be expected from formulations based on unimodal wave conditions. Present design formulae are based predominantly on unimodal experiments because it is harder to include longer period waves. This is the main motivation of present work. In this research work, we present the first detailed modelling study of overtopping under bimodal conditions using carefully constructed spectra that relate the overtopping to the energy of the entire sea state.

1.5 Aims and Objectives

The overall aim of this research is the modelling of wave overtopping induced by bimodal waves. Creation of bimodal sea conditions containing swell has been constructed from the superposition of two JONSWAP spectra; one representing the wind sea and one the swell. With this construction, there will be a degree of overlap between the spectra and a sharp break between the swell and wind sea as a separation frequency is not present. It contrasts with the practice-based approach in which the swell and wind-sea components have no overlap (Bradbury et al., 2007). Here, we use the former method to create a set of bimodal sea states that have varying proportions of swell, while at the same time containing a fixed amount of energy. It refers to a set of sea states that have a fixed amount of energy but varying proportions of swell and wind sea as "energy conserved" bimodal waves. The advantage of this method is that it allows us to relate overtopping to the energy of the complete sea state. The extremes of these conditions are "pure wind sea" at one end and "pure swell" at the other; with both cases resulting in a unimodal spectrum. It was specifically targeted at whether changes in how the distribution of wave energy across swell and wind wave components quantifiably affects the properties of bimodal waves.

1.5.1 Specific Objectives

The following objectives will be considered in this thesis:

1. To develop a bimodal spectrum that will contain the same energy as the unimodal spectrum for easy adaptation of previously published formulations on wave overtopping.
2. Adapt the numerical model IH2VOF to simulate the formulated bimodal waves accurately.
3. Conduct experiments to increase data availability on predictions of wave overtopping under bimodal conditions.
4. Examine the significant changes caused by swell and different spectra parameters on the wave overtopping behaviour by the bimodal spectrum.
5. Modify existing formulations on the reflective behaviour of coastal seawalls to impinging wave bimodality
6. Examine a more practical numerical method of reducing wave overtopping by using storm walls or parapets.

1.6 Thesis Outline

This thesis comprises of ten chapters which will examine the aforementioned objectives. The main body of the work is covered in Chapters 2 to 9, and the conclusions of every chapter are presented in Chapter 10. A brief description of the contents of these chapters below:

Chapter 2 reviews commonly used wave overtopping numerical and experimental models. The aim is to apply both numerical and experimental approaches to model wave overtopping under the influence of bimodal waves. A typical bimodal wave is composed of the combined wind and swell waves. Understanding the previously applied modelling approach would enhance the selection of a suitable model that could accurately describe the physics of wave hydrodynamics of bimodal seas. Chapter 3 is devoted to the development of the model input. The chapter firstly introduce how a bimodal spectrum can be formulated by applying a well-established principle of double Jonswap spectra. A brief description of how bimodal energy can be combined from the wind and the swell waves was given. Description of conserved modes of the bimodal spectrum which bears the same energy as the unimodal sea states are presented. The presence of swell waves in the sea state is varied in terms of percentages and peak periods. The spectra were converted into random waves in the time domain. This

chapter finishes with a comparison of the distinguishing characteristics of the energy conserved-bimodal waves with those in unimodal waves.

Chapter 4 introduces the numerical model applied for the numerical simulation. A RANS-Based wave hydrodynamics model, validated for several overtopping studies under unimodal wave conditions was selected for this study. The aim is to examine how well the model performs under bimodal wave scenarios. The governing equations with the model, initial conditions, and boundary conditions are fully described. Chapter 4 also describes different methods of wave generation, absorption, and free surface tracking techniques in the RANS model. The chapter was concluded by propagating simple waves in the model and inspecting the performance of the wave under these conditions. Chapter 5 explores the capability of the RANS model to accurately propagate both pure unimodal waves and bimodal waves of the same energy component that was developed in Chapter 3. A smooth, impermeable seawall was installed in the model domain, and surface elevations were recorded using numerically-installed wave gauges. Influences of swell on the wave steepness, skewness, kurtosis, autocovariance function of numerically generated bimodal waves are presented. Finally, the effect of swell waves on wave height distribution near the wavemaker and the structures was investigated.

In Chapter 6, an experimental programme was conducted to validate the numerical model developed in Chapter 4, and to increase the research guidance on wave overtopping of coastal defences under swell-driven bimodal seas. A great deal of research guidance is available on wind-driven unimodal sea states; fewer data with a focus on bimodal seas is available. Three smooth, impermeable, sloping seawalls were constructed, where different bimodal spectra are tested. This chapter describes the experiment and bimodal wave interactions with coastal seawalls. Measurements of wave overtopping of smooth, impermeable seawalls recorded under different unimodal and bimodal wave conditions are outlined and discussed. Moreover, an overview of the successful results obtained from more than 800 overtopping tests was presented. The experimental datasets obtained from chapter 6 was analysed in Chapter 7 using the presence of swell waves and spectra characteristics of the bimodal spectra involved. Special attention is placed on the analyses of the direct effect of the existence of swell in the sea state. Moreover, the effects of spectra shapes imposed by different spectra types were examined using the spectra characteristics of narrowness parameters, broadness parameters and peakedness parameters of each spectra types. The chapter was completed by presenting a new empirical influencing factor to describe the effects of swell percentages on wave overtopping discharges.

Chapter 8 presents the results of the reflection properties of smooth, impermeable structures. It was demonstrated that previously published formulations are only based

on unimodal sea conditions and cannot accurately describe reflection coefficients under bimodal waves. Simplified empirical equations that better represents the reflection coefficient of coastal seawalls under bimodal wave conditions than in previous literature are presented. Chapter 9 explores the validation of the numerical model IH2VOF in simulating wave overtopping processes. The extended capability of the numerical model to simulate wave overtopping behaviours of modified structures than simple ones were tested. Wave overtopping due to simple structures was validated, and then simulated for complex structures using different storm walls to reduce excess wave overtopping induced by bimodal waves.

Chapter 10 summarises all the results of the preceding chapters, and outlines possible research perspectives and applications. The chapter discussed the key contribution to research, and presented some directions into future work.

Chapter 2

Review of Wave Overtopping Modelling

This chapter will summarise the state of the art of experimental and numerical modelling of wave overtopping of coastal seawalls. First, an introduction of a comprehensive history of research into an experimental study of wave overtopping models of smooth, impermeable seawalls. In choosing an appropriate wave overtopping model suitable for bimodal seas, it is crucial to assess the performances of some novel wave overtopping models. Both numerical and experimental views are considered here. The numerical part is focus primarily on the physical theories involved, the computational wave generation pattern, model boundary conditions and proven limitations.

2.1 Experimental Modelling of Wave Overtopping

The experimental study of wave overtopping of impermeable coastal seawalls has a long history. A comprehensive histories of the research be found in Hedges and Reis (1998), and Burcharth and Hughes (2002). According to several previous studies (e.g. Burcharth and Andersen, 2007; Franco et al., 2009; van der Meer et al., 2018) and as earlier illustrated in Fig. (2.1), random waves can be expressed spectrally in terms of the spectral wave height H_{m0} and the spectral wave period $T_{m-1,0}$). The random waves would generally overtop a sloping impermeable structure whenever the wave run-up exceeds the crest freeboard R_c .

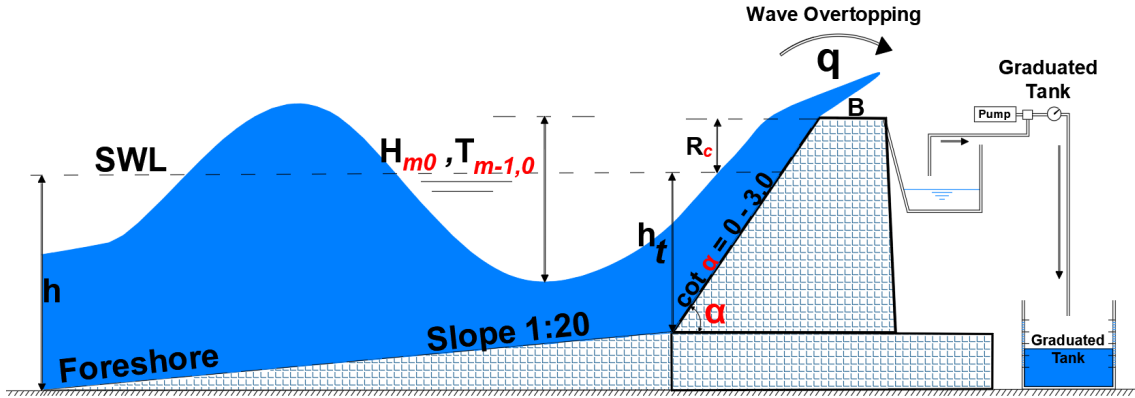


Figure 2.1 A schematic representation of wave overtopping phenomenon

As the waves are random the overtopping will vary from wave to wave. The mean overtopping rate, averaged over many waves, is denoted by Q_m . Firstly, Owen (1980) proposed a decreasing exponential relationship between the wave overtopping Q by a coastal seawall with respect to the crest freeboard R_c . Equation (2.1) shows the mathematical representation as:

$$Q_m = Q^* T_m g H_s \quad (2.1)$$

where $Q^* = A \exp^{-\frac{BR^*}{r}}$, $R^* = \frac{R_c}{T_m \sqrt{g H_s}}$, and $Q_m = g H_s T_m A \exp^{-\frac{BR_c}{r T_m \sqrt{g H_s}}}$

In this equation, H_s represents the significant wave height at the toe of the structure, T_m denotes the mean of the random wave period at the toe, A and B are empirically fitted coefficients, and r is a roughness factor. Since Owen's first formulation, several definitions have been given to coefficients A and B as additional experimental results became available. The relationship between wave overtopping and the crest freeboard has been refined by different authors. A summary of wave overtopping formulae have been presented by (Burcharth and Hughes, 2002; Hedges and Reis, 1998) and these are used as the basis of the extended tabulation in Table 2.1. One of the most significant definitions of wave overtopping formulae, was summarised by Pullen et al. (2007).

Table 2.1 Design formulations on dimensionless discharge, dimensionless freeboard and overtopping model (modified from Burcharth and Hughes, 2002; Hedges and Reis, 1998)

Reference	Dimensionless Discharge Q^*	Dimensionless Freeboard R^*	Overtopping Model	Applicability
Owen (1980)	$\frac{q}{gH_s T_m}$	$\frac{R_c}{T_m \sqrt{gH_s}}$	$Q^* = A \exp(-BR^*)$	Impermeable smooth, rough, straight and bermed slopes
Bradbury et al. (1988)	$\frac{q}{gH_s T_m}$	$\frac{R_c^2}{T_m \sqrt{gH_s^3}}$	$Q^* = A (R^*)^{-B}$	Rock armoured impermeable slopes with crown walls
Aminti and Franco (1988)	$\frac{q}{gH_s T_m}$	$\frac{R_c}{H_s}$	$Q^* = A (R^*)^{-B}$	Rock, cube, tetrapod, double layered armoured impermeable slopes with crown walls
Ahrens and Heimbaugh (1988)	$\frac{q}{\sqrt{gH_s^3}}$	$\frac{R_c}{H_s L_{0p}^{1/3}}$	$Q^* = A \exp(-BR^*)$	Different seawall designs
Pedersen and Burcharth (1993)	$\frac{q T_{0m}}{L_{0m}^2}$	$\frac{R_c}{H_s}$	$Q^* = AR^*$	Rock armoured impermeable slopes with crown walls
De Waal and van der Meer (1993)	$\frac{q}{\sqrt{gH_s^3}}$	$\frac{R_c - R_{2\%}}{H_s}$	$Q^* = A \exp(-BR^*)$	Impermeable smooth, rough, straight and bermed slopes
van der Meer (1993)	$\frac{q}{gH_s T_m}$	$\frac{R_c}{T_m \sqrt{gH_s}}$	$Q^* = A \exp(-BR^*)$	Impermeable smooth, rough, straight and bermed slopes
Franco et al. (1994)	$\frac{q}{\sqrt{gH_s^3}}$	$\frac{R_c}{gH_s} \frac{1}{\gamma}$	$Q^* = A \exp(-BR^*)$	Vertical wall breakwaters with/without perforated front wall
Smith et al. (1995); van der Meer (1993, 1995b)	$\frac{q}{\sqrt{gH_s^3}} \frac{\tan \alpha}{\xi_{0p}} \text{ for } \xi_{0p} < 2;$ $\frac{q}{\sqrt{gH_s^3}} \text{ for } \xi_{0p} > 2$	$\frac{R_c}{H_s} \frac{\tan \alpha}{\xi_{0p}} \cdot \frac{1}{\gamma} \xi_{0p} < 2;$ $\frac{R_c}{H_s} \cdot \frac{1}{\gamma} \text{ for } \xi_{0p} > 2$	$Q^* = A \exp(-BR^*)$	Impermeable smooth, rough, straight and bermed slopes
Hedges and Reis (1998)	$\frac{q}{gR_{c,3} v_{max}}$	$\frac{R_c}{R_{d,max}}$	$Q^* = A(1 - R_c)^B$ for $0 \leq R_c \leq 1$; $Q = 0$ for $R \geq 1$	Impermeable smooth, rough, straight and bermed slopes
Van Gent and Smith (1999)	$\frac{q}{\sqrt{gH_{m0}^3}}$	$\frac{R_c}{H_{m0} \xi_{m-1,0}}$	$Q^* = A \exp(-BR^*)$	Shallow and very shallow foreshore slopes
Pullen et al. (2007)	$\frac{q}{\sqrt{gH_{m0}^3}}$	$\frac{R_c}{H_{m0}}$	$Q^* = A \exp(-BR^*)$	Different seawalls designs including armoured rubble mounds and composite structures
Altomare et al. (2016)	$\frac{q}{\sqrt{gH_{m0}^3}}$	$\frac{R_c}{H_{m0} (0.33 + 0.022 \xi_{m-1,0})}$	$Q^* = A \exp(-BR^*)$	Shallow and very shallow foreshore slopes
van der Meer et al. (2018)	$\frac{q}{\sqrt{gH_{m0}^3}}$	$\frac{R_c}{H_{m0} \xi_{m-1,0} \gamma_c} \text{ for breaking waves;}$ $\frac{R_c}{H_{m0}} \text{ for non - breaking waves}$	$Q^* = A \exp(-BR^*)$	Different seawalls and slope designs including armoured rubble mounds and composite structures
van der Werf and van Gent (2018)	$\frac{q}{\sqrt{gH_{m0}^3}}$	$\frac{R_c - 0.5 H_{m0,well}}{H_{m0} \xi_{m-1,0} \gamma_c} \text{ for breaking waves;}$ $\frac{R_c - 0.5 H_{m0,well}}{H_{m0}} \text{ under non - breaking waves}$	$Q^* = A \exp(-BR^*)$	Smooth impermeable sloping dike under oblique wind and swell waves

The more recent work of Victor and Troch (2012) and van der Meer and Bruce (2013) has enhanced the wave overtopping formulation of impermeable structures to low crest freeboards. Modifications based on varieties of smooth slopes are presented for breaking waves in Equation (2.2) and for non-breaking waves in Equation (2.3) (see van der Meer et al., 2018):

$$\frac{q}{\sqrt{gH_{m0}^3}} = \frac{0.023}{\tan\alpha} \gamma_b \cdot \xi_{m-1,0} \cdot \exp \left[- \left(2.7 \frac{R_c}{\xi_{m-1,0} \cdot H_{m0} \cdot \gamma_b \cdot \gamma_f \cdot \gamma_\beta \cdot \gamma_v} \right)^{1.3} \right] \quad (2.2)$$

Also, maximum estimates of wave overtopping values are generally represented by the non-breaking wave cases as:

$$\frac{q}{\sqrt{gH_{m0}^3}} = 0.09 * \exp \left[- \left(1.5 \frac{R_c}{\xi_{m-1,0} \cdot H_{m0} \cdot \gamma_f \cdot \gamma_\beta \cdot \gamma_v} \right)^{1.3} \right] \quad (2.3)$$

Refinements have generally been made to previous formulations to accommodate several other influence factors such as the presence of berms γ_b , permeability and roughness γ_f , for oblique wave attack, γ_β , influence factor by a vertical wall γ_v and for different geometries γ_* as shown in the equations. $\xi_{m-1,0}$ represents the surf similarity parameters of the wave. Equation (2.2) and (2.3) are the same general form as Equation (2.4) but take the form of a Weibull-shaped function as the power of the argument to the exponential is 1.3 instead of 1.0 as in overtopping Equations of sloping seawalls previously presented in Pullen et al. (2007). The study by Pullen et al. (2007) provides detailed guidelines for the estimation of dimensionless overtopping using different coefficients contained in Equation (2.2) and (2.3). The latest general form of these equations can then be summarised as in Equation (2.4) below:

$$\frac{q}{\sqrt{gH_{m0}^3}} = A \cdot \exp \left(-B \frac{R_c}{H_{m0}} \right)^C \quad (2.4)$$

Coefficients A and B for non-breaking waves across different slopes are defined deterministically by Equation (2.5) and Equation (2.6):

$$A = \begin{cases} 0.09 - 0.01(2 - \cot\alpha)^{2.1} & \text{for } \cot\alpha \leq 2 \\ 0.09 & \text{for } \cot\alpha \geq 2 \end{cases} \quad (2.5)$$

Also, for coefficient B :

$$B = \begin{cases} 1.5 + 0.42(2 - \cot \alpha)^{1.5} \text{ with a maximum of } 2.35 & \cot \alpha \leq 2 \\ 1.5 & \text{for } \cot \alpha \geq 2 \end{cases} \quad (2.6)$$

2.2 Numerical Wave Overtopping Models

The numerical modelling of the wave overtopping is of interest to many applications of engineering such as coastal structures designs, ship decks stability and Floating Production Storage and Offloading FPSO units because they provide a more complex solution to the flow equations in a numerical domain than the empirical and semi-empirical methods (Zanuttigh et al. (2005), Zanuttigh et al. (2014), Battjes and Groenendijk (2000) and van der Meer et al. (2016)). Intensive research has been performed in the last four decades using different models to estimate wave overtopping. Some of these efforts are statistical in nature (Aziztayfun, 2009; Longuet-Higgins, 1952; Shi-Igai and Rong-Chung, 1977). These have been suitably applied to overtopping and flooding of berm breakwaters and longitudinal revetments (Coelho et al., 2016; Cox and Scott, 2001; van der Meer and Veldman, 1992). Some other modelling used different theories of physics in coupled Solutions of Non-Linear Shallow Water Equations NLSW and Navier-Stokes NS equations. Examples of these are the RBREAK, ODIFLOCS, AMAZON, the ANEMONE (Tuan and Oumeraci (2010) and van der Meer et al. (2016)). Other examples combined Eulerian field and depth averaging methods as proposed by Greenwood and Osborne (1990). The NLSW solver has been introduced in wave overtopping computations (Briganti et al., 2011).

Some other numerical models combined higher-order Boussinesq, Stoke and the non-linear dispersion theories. For example, Park and Cox (2016), Ohyama and Nadaoka (1994), Ohyama et al. (1995), and Tsung et al. (2012) to model wave-structure interaction and determine possibilities of wave overtopping. Mizanur and Hanif (1995) identified the method of St. Venant one-dimensional numerical model for wave-structure interaction and Hu et al. (2000) assessed the application of NLSW Equation in modelling wave overtopping AMAZON Model. Hubbard and Dodd (2002) and Brocchini et al. (2001) have applied an old and a newer 2-dimensional OTT-2D NLSW model towards computing wave run-up and overtopping on a sloppy seabed and an arbitrary beach respectively. Wavelet transform analysis has also been used to examine aerial landslide generated waves. Panizzo et al. (2002), and Lara et al. (2008) applied two-dimensional Volume-Averaged Reynolds Average Navier-Stokes (VARANS) equation to investigate

waves interaction with low-mound breakwaters in an experimental flume. Andersen and Burcharth (2009) conducted a three-dimensional Overtopping performance of rubble mound structures placed in the direction of incident oblique waves. Dong and Zhan (2009) proposed a wave run-up and evolution in shallow water using the Navier-Stokes equations and the Volume of Fluid (VOF) functions. Guanche et al. (2009) proposed a two-dimensional numerical approach to evaluating how wave loads affect the stability of a coastal structure. The COBRAS-UC model was adopted to simulate wave overtopping behaviour of rubble mound breakwaters.

Zheng et al. (2009) and Gao et al. (2012) established a numerical Smoothed Particle Hydrodynamics (SPH) model to depict the turbulence characteristics that occur at the time of wave breaking. Du et al. (2010) applied a depth-averaged COAST2D model to examine the effects of overtopping on the hydrodynamics and morphodynamics within the spatial shore-parallel breakwaters. Bakhtiary et al. (2010); Hsiao and Lin (2010); Peng and Zou (2011) used RANS to investigate wave overtopping caused by impinging and rough waves like the tsunami in front of a breakwater. Gao et al. (2012) developed a numerical model of the behaviour of regular waves slamming on an open pile structure. This model was based on and improved SPH model that incorporates CSPM and Riemann solution discretised into the Navier-Stokes equations. Garrido and Medina (2012) presented a new semi-empirical formula to depict the reflection patterns of waves on Jalan-type breakwaters using Neural Network model. Sun et al. (2012) improved the Boussinesq Equations and the Finite Element Method (FEM) to investigate wave run-up on a vertical surface cylinder. Tsung et al. (2012) combined some sets of laboratory datasets with the higher-order Boussinesq-type equation to model solitary wave run-up and to overtop on a plane sloppy experimental terrain.

In most of these approaches, applying coupled solutions of the RANS, VOF, non-linear shallow water equation and the higher-order Boussinesq algorithm to estimate wave overtopping appeared to be familiar to most of the modelling efforts McCabe et al. (2013), Casella et al. (2014), and Gui et al. (2015). Sometimes the turbulence field may be included or ignored to reduce the modelling complexities, but these usually affect the expected overtopping estimates (van der Meer et al., 2016). This report reviews the performance of four wave overtopping numerical models, namely NEWRANS model, IH2VOF, DSPH, and OpenFoam. Table 2.2 provides a detailed summary of this review.

2.3 New Reynolds-Averaged Navier-Stokes (NEWRANS)

The NEWRANS model is a mesh-based model that was proposed by Batchelor (1974). This model provides solutions to the Reynolds-averaged Navier-Stokes and the modified version of the turbulence equations. Initially, the model uses the boundary element method (BEM), which limits the design to the wave breaking point. However, the Boussinesq Depth average method and length turbulence have been incorporated into the model and have made it possible for the model to survive beyond the breaking points.

2.3.1 Theory and Assumptions in NEWRANS Model

NEWRANS model applies the RANS equation in a turbulence field. The fluid is assumed to be incompressible in a domain ω with motion in three dimensions $i, j = 1, 2$ respectively. These have been represented as a numerical solution of equations (2.7) to (2.16) Batchelor (1974) that are expressed as follows:

$$\frac{\partial u_i}{\partial x_i} = 0 \quad (2.7)$$

$$\frac{\partial u_i}{\partial t} + u_j \frac{\partial u_i}{\partial x_j} = -\frac{1}{\rho} \frac{\partial p}{\partial x_i} + g_i + \frac{1}{\rho} \frac{\partial \tau_{ij}}{\partial x_j} \quad (2.8)$$

The density change ρ_i that is being used to track the free surface can also be expressed as:

$$\frac{\partial \rho}{\partial t} + u_i \frac{\partial \rho}{\partial x_i} = 0 \quad (2.9)$$

and the turbulence model $k - \varepsilon$ in all the three directions is given by:

$$\frac{\partial k}{\partial t} + u_j \frac{\partial k}{\partial x_j} = \frac{\partial}{\partial x_j} \left[\left(\frac{v_t}{\sigma_k} + v \right) \frac{\partial k}{\partial x_j} \right] + 2v_t \sigma_{ij} \frac{\partial u_i}{\partial x_j} - \varepsilon \quad (2.10)$$

$$\frac{\partial \varepsilon}{\partial t} + u_j \frac{\partial \varepsilon}{\partial x_j} = \frac{\partial}{\partial x_j} \left[\left(\frac{v_t}{\sigma_\varepsilon} + v \right) \frac{\partial \varepsilon}{\partial x_j} \right] + C_{1\varepsilon} \frac{\varepsilon}{k} 2v_t \sigma_{ij} \frac{\partial u_i}{\partial x_j} - C_{2\varepsilon} \frac{\varepsilon^2}{k} \quad (2.11)$$

The total stress τ_{ij} in equation (2.2) can be linked to $k - \varepsilon$ and σ_{ij} as:

$$\tau_{ij} = 2(v + C_d \frac{k^2}{\varepsilon}) - \frac{2}{3} k \delta_{ij} \quad (2.12)$$

These equations are solved using the two-step projection method proposed by Chorin (1968) and Chorin (1969):

$$\frac{\tilde{u}_i^{n+1} - u_i^n}{\Delta t} = u_j^n \frac{\partial u_j^n}{\partial x_j} + g_i + u_j^n \frac{\partial \tau u_{ij}^n}{\partial x_j} \quad (2.13)$$

The velocity is also projected to the divergence free plane as:

$$\frac{\tilde{u}_i^{n+1} - u_i^n}{\Delta t} = -\frac{1}{\rho^n} \frac{\partial p^{n+1}}{\partial x_i} \quad (2.14)$$

$$\frac{\partial u_i^{n+1} - u_i^n}{\partial x_i} = 0 \quad (2.15)$$

which can be combined with the Poisson Pressure Equation PPE:

$$\frac{\partial}{\partial x_i} \left(\frac{1}{\rho^n} \frac{\partial p^{n+1}}{\partial x_i} \right) = \frac{1}{\Delta t} \frac{\tilde{u}_i^{n+1}}{\partial x_i} \quad (2.16)$$

Table 2.2 A review of numerical wave overtopping models

Disparity Criteria	NEWRAN Model	DualPhysics Model	IHFOAM Model	IH2VOF Model
1. Model structure & possible Dimension	Mesh based & 2-D	Meshless langragian approach with 2D & 3D functionalities	Meshed and poly meshed 2-D & 3D possibilities	Mesh-based with 2-D & 3D
2. Possible Wave Generation Methods	Wave paddle & internal point sources	Wave paddle & internal point sources	Wave paddle & internal point sources to generate both regular and irregular waves.	Piston wave type generator. Wave Paddle generation and internal point wave sources
3. Equation Solved	Reynolds-averaged Navier-Stokes and the turbulence ($k-\epsilon$) equations	Navier-Stokes equation in a langragian meshless mode. Fluid is treated as particles, artificial viscosity and turbulence	RANS equation containing the momentum and the continuity equations. Turbulence has recently been incorporated.	2DV Reynolds Averaged Navier-Stokes (RANS) equations and VARANS in porous media. Turbulence Kinetic κ and dissipation ϵ . VOF for tracking free surface movement
4. Model Inputs & Initial Conditions	Takes in a RIPPLE fortran based file containing a well structured input wave file, directional data, input parameters related with acceleration due to gravity, density, viscosity, initial velocity and wave generation and behaviour parameters	The model takes a .xml file that contains the input parameters, VOF, the wave maker information, piston movement info, and other configuration parameters. The VTK file and the object files are also accepted as input.	Initial velocity, VOF, and bouyancy pressures are applied to generate pressures at the wave generation phase.	Takes in a RIPPLE fortran based file containing a well structured input wave file, directional data, input parameters related with acceleration due to gravity g , density, viscosity, initial velocity and wave generation and behaviour parameters
5. Boundary Conditions	Open, moving and closed boundary conditions are possible. The Active Wave Absorption System, AWAS is adequately implemented in the model.	Absorption boundary condition, wave generation BC are specified at each faces which are owned by a unique cell. The gravity field is assumed to be negative direction of Z axis.	Dirichlet boundary conditions is applied. AWAS has just been recently been incorporated but no known publication on this yet.	Dirichlet boundary condition, moving boundary and the internal wave maker. The AWAS is applied to avert wave reflection at the boundary
6. Turbulence and AWAS Consideration	Turbulence, AWAS and bottom friction were adequately treated	No turbulence and AWAS in the model	Turbulence, AWAS and bottom friction were adequately treated	Turbulence, AWAS and bottom friction were adequately treated
7. Model Output	VOF Function, horizontal and vertical velocity field, pressure field, turbulent component, free surface elevation, free gauges, spectral outputs, wave run-up, mean overtopping discharge, instantaneous velocity and volume discharge, wave breaking results and gravity	Output files in types of Binary, csv, ascii are generated based on specified position, particle type, velocity limit, fixed or moving boundary, and variables like velocity, pressure, mass, volume, acceleration, and vorticity. Processes relying on these quantities can be simulated.	Volume averaged velocity and position, mass density, heat capacity, volume, surface, and emissivity of the thermocouple bead, pressure, mass, volume, acceleration, and vorticity can also be obtained for any time step.	VOF Function, horizontal and vertical velocity field, pressure field, turbulent component, free surface elevation, free gauges, spectral outputs, wave run-up, mean overtopping discharge, instantaneous velocity and volume discharge, wave breaking results and gravity
8. Maintaining the geometric and kinematic similarities for experimental flume Replication (Rank 1 to 5, 5 is most complex)	3	4	2	1
9. Model Limitation	It is computationally expensive. The absence of open sourced referenced materials makes it more difficult using the models for model implementation.	Absence of turbulence and wave absorption consideration	No proven updated publication has been made on the limitation of this version of Openfoam because it has just been recently released. The previous version of the model deviates from the 2D wave absorption theory. Turbulence was not considered also.	It provides an easier mode of replicating a wave overtopping experimental flume numerically. Most practitioners currently using Mike 21 BW to the toe of the structure and then use Eurotop to compute overtopping which might not work well in this case

2.3.2 Wave Generation in the NEWRANS Model

NEWRANS model is capable of generating waves using both internal point sources and by using wave paddles at one end. No dynamic boundary condition is applied to the free surface, but tangential stress, normal stress and inflow boundary condition are applied to one cell down from the surface. New values of quantities are obtained by the Finite difference method. The domain are generally discretised into a m by n domain. The centre of each cell is identified and defined with scalar quantities (p , k , e , and F). Under the assumption that the density ρ_f of the fluid remain unchanged but being averaged into the individual cell. Each cell is assumed to contain some volume of the fluid V_f and some V_a volume of air which would share up the density ρ_f on the basis that:

$$\rho = \frac{\rho_f V_f}{V_f + V_a} \quad (2.17)$$

The normalised average density for the centre of each cell can be described as the VOF function F expressed as:

$$F = \frac{\rho}{\rho_f} \quad (2.18)$$

Cell identification and classification becomes possible with the aid of the VOF function F . It is feasible to determine the exact density status of each cell to determine whether or not VOF F is full at $F = 0$ or $F > 0$. Final boundary conditions will then be applied to take care of the newly filled cells that were empty during the previous run.

2.3.3 The Limitation of the NEWRANS Model

NEWRANS model lacks useful open source materials. The model is also computationally expensive to run. At moment, the model cannot be cheaply implemented on GPU-based parallel computers. More coding on parallelization is required.

2.4 The IH2VOF Model

IH2VOF model provide a modified approach to the aforementioned NEWRANS model. The solutions to the 2DV RANS and the Volume-Averaged Reynolds Averaged Navier-Stokes (VARANS) equations in a porous media are provided using a meshed model set up. The model include a CORAL mesh generator which made it easily possible to generate a two-dimensional designed flume and structures. The instantaneous velocities and pressure field are decomposed to establish their mean and equivalent turbulent components. The VOF method has been applied to track the free surface movement

of particles used. Because this model has been selected for this research, elaborated details about the model is presented in Chapter 4.

The IH2VOF was chosen over the NEWRANS model because it runs on the Linux operating system. In this way, it leveraged on the linux operating speeds and functionalities. The compiled executables in the IH2VOF are reusable with different input files. In contrast with the window-based NEWRANS model, each simulation is accompanied with new sets of compilation making it difficult for batch simulation. Several simulations can be run simultaneously in the IH2VOF model while saving computational cost in time. However, new sets of post-processing scripts are developed to analyse model results and produce visualisations in IH2VOF model.

2.5 DualSPHysics (DSPH) Wave Overtopping Model

DSPH is an upgrade of the Smoothed Particle Hydrodynamics Model Sphysics. It is based on a set of C++, CUDA and Java codes that solved free surface flow phenomena such as waves and breakwaters. The model operates in a fully Lagrangian approach with the fluid being discretised into a set of particles in a meshless mode. Computations of physical quantities of each particles (at nodes) are based on the aggregate interpolation of the nodal values of nearest neighbours. These are computed by using a kernel function (W) (Monaghan, 1992) and a well determined smooth length (h) generated from a weighted distance from the influence of nearest inter-particles on the targeted node. The aggregate contribution of the nearest neighbour can then be seen in an integral form as F expressed as:

$$F(r) = \int (F(r') \cdot W(r - r') \cdot h) dr'. \quad (2.19)$$

A modified version of the Kernel function K has been presented as $W(q)$ by Wendland (1995). Inter-neighbour contribution to particle values at each nodes were then set not to exceed $2h$ in this case as $0 \leq q \leq 2$. In a general form:

$$W(q) = \alpha_D \left(1 - \frac{q}{2}\right)^4 (2q + 1) \quad (2.20)$$

Wendland (1995) then modified $q = r/h$ and α_D as $7/14\pi h^2$ for a two dimensional approximation.

In a discretised domain D for a point a , this would be aggregated from nearest neighbour b and Equation (2.21) (for a volume computation obtained from b using

m_b/ρ :

$$F(r_b) \approx \sum_b (F(r_b) \cdot W(r_a - r_b) \cdot h) \frac{m_b}{\rho_b} \quad (2.21)$$

2.5.1 DSPH Physical Theory and Assumptions

This was based on the solution to the Navier-Stokes equation. Water waves are treated like particles that are weakly compressible and the mass is kept constant in order to determine the density changes. The velocities of particle a at node a is computed by applying the kernel function to solve the law of conservation of momentum obtained from the nearest neighbours:

$$\frac{dv_a}{dt} = - \sum_b m_b \left(\frac{P_b^2}{\rho_b} + \frac{P_a^2}{\rho_a} + \Pi_{ab} \right) \nabla_a W_{ab} + g \quad (2.22)$$

The artificial viscosity Π_{ab} is also given by:

$$\Pi_{ab} = \begin{cases} 0 & \text{if } \overline{C_{ab}} \mu_{ab} \leq 0 \\ \frac{-\alpha \overline{C_{ab}} \mu_{ab}}{\rho_{ab}} & \text{otherwise} \end{cases} \quad (2.23)$$

$V_{ab} \cdot r_{ab} < 0$; $V_{ab} \cdot r_{ab} > 0$ with the values of μ_{ab} computed by using

$$\mu_{ab} = \frac{h v_{ab} \cdot r_{ab}}{r_{ab}^2 + \eta^2} \quad (2.24)$$

while the

$$\overline{C_{ab}} = 0.5(C_a + C_b); \quad \overline{\rho_{ab}} = \rho_a^2 + \rho_b^2 \quad (2.25)$$

Pressure changes are also usually computed from the Taint's equation presented by Batchelor (1974):

$$P_a = B \left[\left(\frac{\rho_a}{\rho_0} \right)^y \right] - 1 \quad (2.26)$$

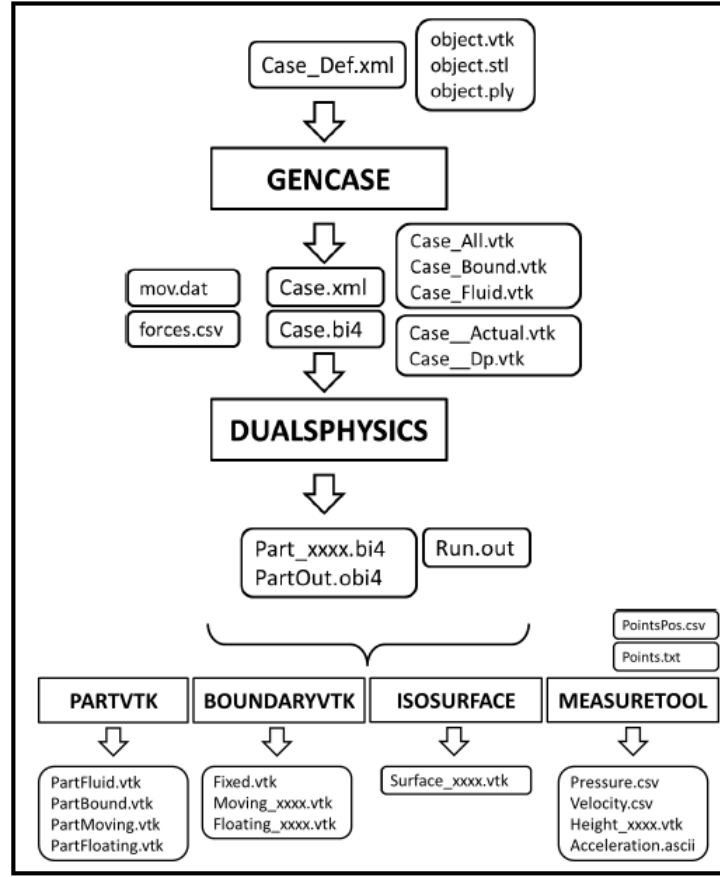


Figure 2.2 General Workflow of DSPH (Canelas et al. (2016))

2.5.2 Wave Generation and Particle Coupling in DSPH

The general workflow of DSPH is shown in Figure 2.2. The model is capable of generating a real wave flume for both long crested regular and random waves with both a point source and piston-type wavemaker. The first order wave equation proposed by Biesel F (1951) and second order wave equations by Madsen (1971) are possible in the model. Irregular waves based on the Jonswap spectrum have also been introduced (Liu and Frigaard, 2001).

The particles are coupled using the Discrete Element Method (DEM). The interaction between the particles are resolved numerically using the integration of SPH with DEM approach. These are decomposed into both damping force F_d and repulsive force F_r that are determined from the total normal force $F_{n,ij}$ respectively:

$$F_{n,ij} = F_n^r + F_n^d = k_{n,ij} \delta_{ij}^{\frac{3}{2}} e_{ij}^n - \gamma_{n,ij} \delta_{ij}^{\frac{1}{2}} \delta_{ij} e_{ij}^n \quad (2.27)$$

The damping coefficient can be computed with

$$\gamma_{n,ij} = -\frac{\log e_{ij}}{\sqrt{\pi^2 + \log e_{ij}^2}}$$

and the stiffness parameter is determined as:

$$k_{n,ij} = \frac{4}{3}E^*\sqrt{R^*}$$

2.5.3 DSPH Integration and Boundary Conditions

The model makes use of a second order Verlet algorithm-based time stepping that adopts Courant Friedrich Lewy condition Verlet (1967) and Monaghan and Kos (1999). The fusion and diffusion terms are considered during computations in variable time steps and are expressed in the form:

$$V_a^{n+1} = V_a^n + 2\Delta t F_a^n; \quad r_a^{n+1} = r_a^n + \Delta t V_a^n + 0.5\Delta t^2 F_a^n; \quad \rho_a^n + 2\Delta t D_a^n \quad (2.28)$$

The dynamic boundary conditions for the DSPH model was introduced by Crespo et al. (2007). This considered that positions of the particles before waves are generated. The continuity and state equations were obeyed in the process. The particles moves according to the exerted force. This come in as a point floating object, or as gates of wave makers. The particles would then move outward from the position of displacement. The pressure of the system increases with time as the particles approach a boundary because of increased density as the separation distance between individual particles decreases than the smoothing length (h). After some DSPH model runs, the pressure builds up as described before. This created pressure usually poses a reflection effect along the boundary because there is no provision for active wave absorption (AWAS) in the system.

2.5.4 Limitations of DualSPHysics Model

The model composition described shows that the model is capable of modelling wave-structure interaction. The model still lacks in providing the wave absorption at the boundary after some sets of runs (Altomare et al., 2015). Reflection tends to occur which could impair the overtopping estimate that is expected to be captured. Small series of time run can be simulated which is not also good if a matured sea is to be modelled mostly for bimodal overtopping assessment. Turbulence created by the

particles after some model runs has not been considered. The disturbances could pose false observation of the expected overtopping result.

The model is also computationally expensive when compared to other mesh-based CFD methods, but this has been partially alleviated by introducing the GPU systems that can run on cluster machines.

2.6 OpenFOAM Model for Wave Overtopping

The Open-source Field Operation And Manipulation (OpenFOAM) has been described as one of the leading free and open source software for computational fluid dynamics (CFD). DSPH was introduced by Henry Weller (1989) based on C++ libraries and codes to solve turbulence and other fluid flow complex problems by applying the finite volume discretization method del Jesus et al. (2012). The model can run both on single system CPU and parallel cluster modes.

Some first sets of practical attempt to introduce OpenFOAM to simulate coastal engineering processes was made by (Higuera et al., 2013b). More studies on wave-structure interaction have been carried out using the same model (Higuera et al., 2013a, 2014, 2015; Maza et al., 2015). The latest form of the OpenFOAM for coastal modelling is the OpenFOAM v1612+/IHFoam that was released by IH Cantabria in 2016.

2.6.1 IHFoam Theory and Assumptions

Just like in DSPH, the model solves the RANS equation containing the momentum and the mass conservation equations. The continuity equation can be expressed as:

$$\nabla \cdot u = 0 \quad (2.29)$$

$$\frac{\delta \rho u}{\delta t} + \nabla \cdot (\rho u u) - \nabla \cdot (\mu_{eff} \nabla u) = -\nabla p^* - g \cdot X \nabla \rho + \nabla u \cdot \nabla \mu_{eff} + \sigma_s \kappa \nabla \alpha \quad (2.30)$$

The pseudo dynamic pressure p^* and the vector velocity u are derivable from the two equations: σ represents the coefficient of surface tension, while μ_{eff} is the efficient dynamic viscosity which is obtained by adding the molecular dynamic viscosity to the turbulent kinetic viscosity. In IHFoam, The coefficient matrices are assembled for variables on the left hand side of the equality sign while independent equation terms are used to solve the variables on the right side of the equation to be solved by explicit

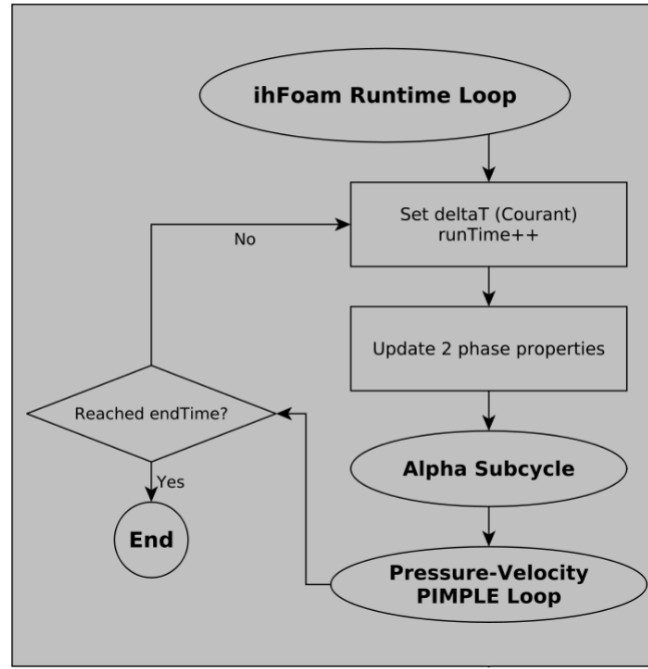


Figure 2.3 A Simple PIMPLE Computation Loop (Adapted from Higuera et al. (2013b))

calculations. The density of the cell is given by:

$$\rho = \alpha \rho_{water} + (1 - \alpha) \rho_{air} \quad (2.31)$$

The advection effect for fluid movement tracking is given by

$$\frac{\delta \alpha}{\delta t} + \nabla \cdot u \alpha + \nabla \cdot u_c \alpha (1 - \alpha) = 0 \quad (2.32)$$

These equations are bonded together by applying the Multidimensional Universal Limiter for Explicit Solution (MULES) proposed by Rusche (2002).

2.6.2 Integration of IHFoam and Boundary Conditions

IHFoam makes use of a new algorithm PIMPLE (Figure 2.3). A PIMPLE is a combination of the OpenFOAM-generated Semi-Implicit Method for pressure linked equations (SIMPLE) with Pressure Implicit Splitting Operators (PISO). PIMPLE was proposed by Kissling et al. (2010) and make provision for easy under relaxation and convergence condition for the variables that are being solved in the equations. The relationship between A PIMPLE and other model parameter runs is provided in Figure 2.3. The boundary conditions in IHFoam is done in modules for easy modification. These modules allow for possible automation of different wave theories in the model set up.

Boundary conditions are based on the faces because of its limited volume nature. It is feasible to access the centre of each face because different cells have been defined for them. Waves could be generated based on this and absorption can be controlled.

2.6.3 Wave Generation in IHFoam

Simple wave theories based on Stoke's I and II can be implemented in the usual OpenFOAM (freely available online). It has been observed that this model is too simple because it makes use of the VOF method to account for wet/dry cells but could not account for turbulence at an earlier stage of time runs and wave absorption at the boundary. Accurate wave generating and absorption boundary conditions were well represented in the model. Extensive coding will be required to accurately describe the boundary conditions that can generate most waves hydrodynamics. A modified approach was introduced by Higuera et al. (2013a). It has now been possible to replicate different wave theories and forms ranging from Stokes I and II, regular, solitary and irregular random waves. The piston wavemaker can also be replicated in this model. Dirichlet boundary conditions are specified based on the free surface and velocity functions only while pressures are solved based on special boundary conditions available in OpenFOAM. First order irregular waves are assumed to be obtained from the aggregate sum of N numbers of individual waves with each having their period T_i , wave height H_i and phases ω_i respectively:

$$\eta = \sum_{i=1}^n \frac{H_i}{2} \cos(k_{xi}x + k_{yi}y - \omega_i t + \psi_i) \quad (2.33)$$

$$\begin{aligned} u &= \sum_{i=1}^n \frac{H_i}{2} \omega_i \cos^2(\Delta\beta_i) \frac{\cosh[k_i(h + z^*)]}{\sinh(k_i h)} \cos(\theta_i) \cos(\beta_i) \\ v &= \sum_{i=1}^n \frac{H_i}{2} \omega_i \cos^2(\Delta\beta_i) \frac{\cosh[k_i(h + z^*)]}{\sinh(k_i h)} \cos(\theta_i) \sin(\beta_i) \\ w &= \sum_{i=1}^n \frac{H_i}{2} \omega_i \cos^2(\Delta\beta_i) \frac{\cosh[k_i(h + z^*)]}{\sinh(k_i h)} \sin(\theta_i) \end{aligned}$$

Active Wave absorption criteria proposed by Schäffer and Klopman (2000) has been implemented for both 2-D and quasi 3-D modes.

2.6.4 Proven Limitations of IHFoam

The previous version of the model deviates from the 2-D wave absorption theory. The older model also radiates waves at the boundary instead of absorbing them. The new 3-D absorption criteria would show a great performance for incident waves at the perpendicular directions (close to 0 or 90 degrees) because it might be difficult to anticipate the actual direction of the incident waves. Even though the quasi3-D takes care of this abnormality, it might be very difficult using this for a two dimensional overtopping estimates that is needed for this research.

2.7 Neural Network Overtopping Models

The numerical and empirical models described above can efficiently predict overtopping rates based on specific structure geometries and wave conditions. Obviously, greater errors and uncertainties emerges at applying such models to other structure types or wave conditions that are different from the original model conditions. In this way, overtopping model formulations can reflect only most dominant parameters. On this premise, the prediction of overtopping was extended to the use of data-driven neural network (NN) models. Unlike the empirical models, the NN models can easily adapt existing training features to further interpolate and extrapolate to further predict overtopping for other structure geometries and test conditions. However, accuracies of predictions reduce with further departure from the training datasets. NN-based overtopping models are developed using training datasets from the CLASH database (De Rouck et al., 2009b). The single-phased NN was developed by the Delft Hydraulics (DNN) and is fully described in van Gent et al. (2007) and Coeveld et al. (2005). Verhaeghe et al. (2008) presented the double-phase overtopping prediction mode. The double-phase model consists of a classifier (to classify whether overtopping is significant or not), while a quantifier is provided to compute the overtopping. Once overtopping is significant in the double-phase mode, the two NN models are similar.

2.8 Review of Reflection Characteristics of Smooth Impermeable Slopes

Wave incidents on coastal seawalls will be partially reflected unless fully absorbed by the structure. The reflected wave component will interact with the incoming wave

creating interference. This can lead to wave amplification, wave breaking, and standing waves Andersen et al. (2014). In the case of vertical walls, standing waves can be pronounced (Zanuttigh and Martinelli, 2008). Standing waves lead to an amplification of wave-induced velocities which can lead to exacerbated scouring of sediments near the toe of the structure, and eventually to its failure and collapse. At locations exposed to local storm waves and open oceans, long period swell waves can be present leading to bimodal wave conditions. Existing literature provides little guidance on reflection characteristics in this situation, which P. J. Hawkes and T. Coates and R. J. Jones (1998) considers might constitute the worst-case in terms of wave conditions. Recent studies by Thompson et al. (2017) and Polidoro et al. (2018) provide evidence that bimodal wave conditions constitute worse storm conditions than pure wind wave conditions of similar total energy content. There remains a gap in our understanding of seawall performance under bimodal wave conditions. This thesis presents the results of laboratory experiments of bimodal waves impinging on impermeable seawalls.

Previous studies including Miche (1951), Ursell et al. (1960), and Battjes (1974) propose the reflection coefficient of monochromatic waves on a plane beach. Miche's hypothesis was reformulated by Battjes (1974) in terms of waves breaker parameter index after an earlier study conducted by Iribarren and Nogales (1949). Zanuttigh and Martinelli (2008) pointed out that both studies from Ursell et al. (1960) and Battjes (1974) indicated overestimation of reflection coefficient criterion compared by Miche (1951) formulations. One of the main advantages of Zanuttigh and Martinelli's results is the fact that large reflection datasets covering several structures, have new formulations. The prediction formulae are applicable to both breaking and non-breaking waves. The formulation is also valid for breaker parameters between 1 and 4.1 and for wall slopes of $\cot \alpha$ between 1.5 and 4.0. Studies by Battjes (1974) and Seelig and Ahrens (1981) are limited to breaker parameters less than 2.3 due to the limitations of their laboratory studies.

Some other notable studies were described in Numata (1976) and Losada and Gimenez-Curto (1981). Numata (1976) presented reflection and transmission performances of artificial blocks in a dimensionless form by comparing the ratio of breakwater width to the diameter of the armour. In Losada and Gimenez-Curto (1981), flow behaviours created by reflection and transmission were represented in an exponential form. Well-defined interaction curves obtained from wave heights and periods of regular waves were applied to obtain probabilistic standards which serve as input to predict flow patterns of equivalent irregular wave actions. The exponential probability model proposed by Losada and Gimenez-Curto (1981) were validated using experimental datasets. Other notable studies include, Postma (1989), Davidson et al. (1996), and

Allsop et al. (2005) in which reflection performances of different rock slopes and rock armours were presented. Postma (1989) investigated rock slopes under irregular wave attack which showed a greater dependence of the reflection coefficient K_r on the breaker index ξ . However, a weaker correlation was obtained in the relationship between K_r with spectral characteristics and depth at the structure-toe.

Relationships between K_r and ξ were later improved by van der Meer (1995a) by applying a multiple regression analysis combining influences of the characteristics of waves (in terms of heights and periods), and the structure (slope and permeability). Similarly, a modified version of the study by Seelig and Ahrens (1981) was presented by Allsop and Hettiarachchi (1989). Here values of wave steepness from 0.043 to 0.042 were investigated, which corresponds to ideal wind-sea states. Lower wave steepness which correspond to swell-driven sea conditions under bimodal wave conditions have not been considered. Newer coefficient values for predicting reflection performances of random waves were derived. A more recent study was presented by Neelamani and Sandhya (2003). Predictive equations were proposed based on a series of experimental tests derived from wave reflection measurements of several wave heights and wave periods. Different seawall types including plane, dentated and serrated and one water depth were used in all measurements obtained.

As observed in some previous studies including Thompson et al. (2017), Polidoro et al. (2018), and Orimoloye et al. (2019b) bimodality in sea waves generally increase the wavelength of the wave train. It would, in turn, reduced the wave steepness and could usually alter the breaker parameter indices. These occurrences have not yet been fully investigated for bimodal sea cases. In this thesis, the physical model tests were performed on three different seawalls to examine reflection performances under bimodal wave scenarios. Unimodal and bimodal cases were compared with previously formulated formulas. A new prediction formula which considers reflection coefficients under wave bimodality is proposed.

2.9 Input to Wave Overtopping Studies

In all the numerical and experimental modelling described above, accurate wave generation techniques are essential to wave overtopping predictions. In most cases, the JONSWAP spectrum is applied to produce irregularly randomised surface elevations of waves. A minimum of 1000 irregular waves are usually applied in most cases for wave overtopping studies (e.g. Burcharth and Andersen, 2007; Franco et al., 2009; van der Meer et al., 2018). As described in chapter 1 EurOtop (2018) suggests the use

of $T_{m-1,0}$ to account for bimodality in wave overtopping studies. It has been found that more weights can be assigned to wave period of higher values. There is limited prior research focusing on wave overtopping study that considers wave bimodality. Notable examples are Hedges and Shareef (2003); Kashima et al. (2010); P. J. Hawkes and T. Coates and R. J. Jones (1998) who presented physical model tests, and a numerical wave overtopping assessment by Thompson et al. (2017). The influence of wave directionality on wave overtopping oblique wind and swell waves was recently investigated by van der Werf and van Gent (2018). They found that the presence of swells with low steepness contributes more to wave overtopping in obliquely intersecting wind and swell seas. A great deal of research guidance is available on wind-driven unimodal sea states but on bimodal seas, less data can be found. It is imperative to consider the knowledge gap on bimodal wave overtopping of smooth impermeable seawalls.

2.10 Conclusion

This chapter presented a comprehensive review of wave overtopping models. A great deal of research guidance is available on wind-driven unimodal sea states, but on bimodal seas, less data can be found. Large portions of coasts across the globe are exposed to more complicated sea states under very large fetches that are bimodal in nature. The knowledge gap on bimodal wave overtopping of smooth, impermeable seawalls will be considered in subsequent chapters in order to improve the wave overtopping formulations representing the bimodal wave cases. For easy comparison of the bimodal sea states with the unimodal sea states already presented in previous literatures, a bimodal spectrum that will contain the same energy as the unimodal spectrum will be developed in the next chapter.

Chapter 3

Development of Energy-Conserved Bimodal Spectrum

This chapter presents the analytical development of a bimodal spectrum that contains the same energy as the unimodal wind-sea states. In bimodal seas, different percentages of swell can be present and the swell can appear at different frequencies while the energy content of the sea state remains unaltered. Commonly, bimodal sea conditions containing swell have been constructed from the superposition of two JONSWAP spectra; one representing the wind sea and one the swell. With this construction, there will be a degree of overlap between the spectra and a sharp break between the swell and wind sea as a separation frequency (usually between 0.09 and 0.1Hz) is not present. It contrasts with the practice-based approach in which the swell and wind-sea components have no overlap (Bradbury et al., 2007). Here, we use the former method to create a set of bimodal sea states that have varying proportions of swell, while at the same time containing a fixed amount of energy. We refer to a set of sea states that have a fixed amount of energy but varying proportions of swell and wind sea as "energy conserved" bimodal waves. To achieve this purpose, three different techniques were applied until the energy in the sea state became conserved. The formulated energy-conserved bimodal spectrum was created from unimodal sea states and converted into random waves time series using the Inverse Fast Fourier Transform (IFFT).

3.1 Development of a Bimodal Spectrum

In order to create an energy-conserved bimodal spectrum for this study, the four-parameter analytical approach proposed by Guedes Soares (1984) was adapted for this

purpose. Figure 3.1 contains the flow chart of the MATLAB algorithm for generating the bimodal spectrum. Firstly, a bimodal spectrum was created from the arithmetical combination of double modified JONSWAP spectrum, (due to Hasselmann et al. (1973). This is described in Equations (3.1) and (3.2) which detail the superposition of modified Jonswap spectra for swell and wind sea components.

By superposition, the bimodal spectrum, $S_{bim} = S_{ss} + S_{ww}$ (3.1)

$$S(f)_{ij} = S(f)_i + S(f)_j$$

$$S(f)_{ij} = \beta_e H_s^2 T_{ij}^{-4} f_{ij}^{-5} \exp[-1.25(T_{ij} f_{ij})^{-4}] \gamma_{ij} \exp\left[-\frac{(T_{ij} f_{ij} - 1)^2}{2\sigma_{ij}^2}\right] \quad (3.2)$$

$$\beta_e = \frac{0.0624}{0.230 + 0.0336\gamma - 0.185(1.094 - 0.01915\ln\gamma_{ij})} \left[1.094 - 0.01915\ln(\gamma_{ij})\right]$$

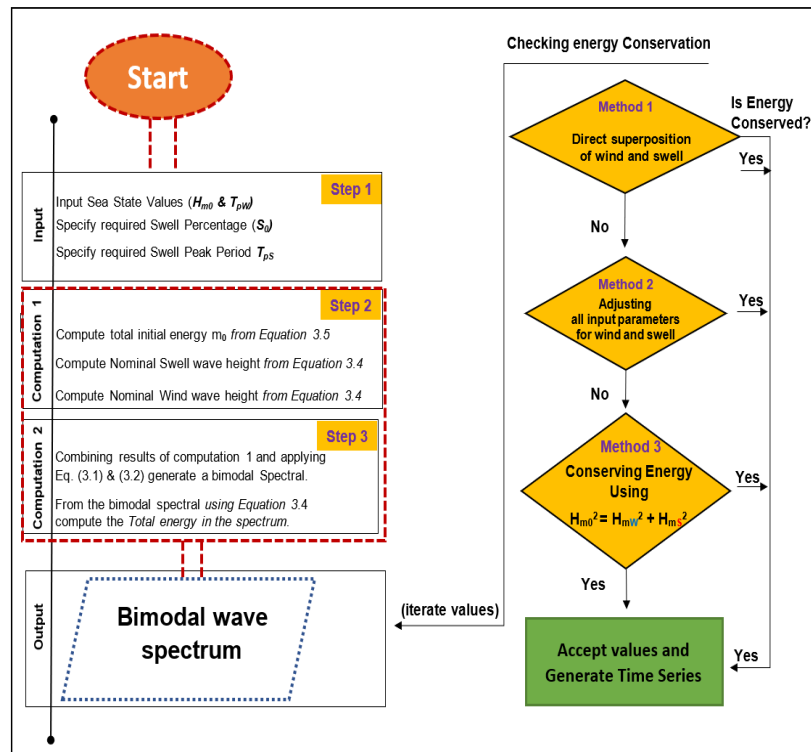
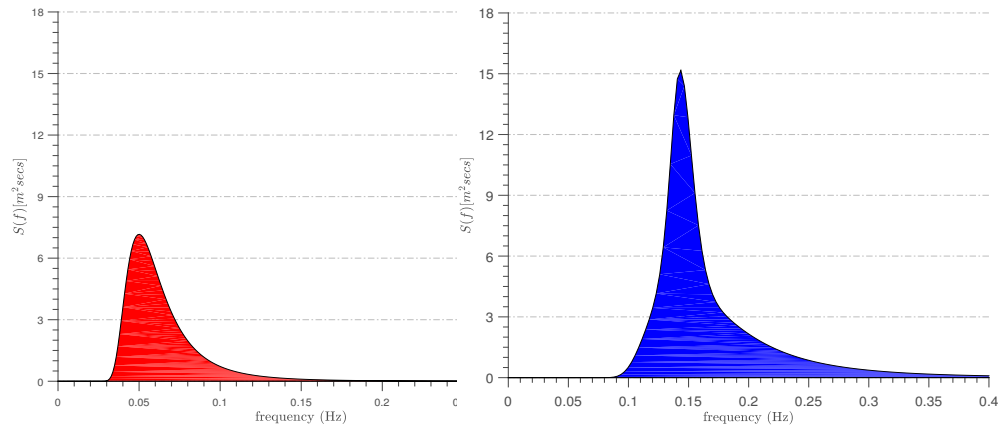
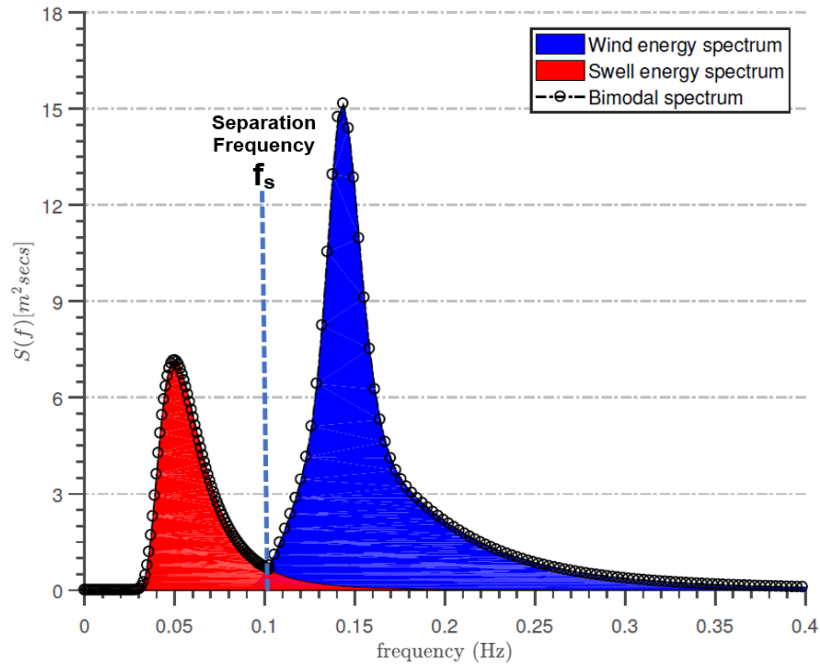


Figure 3.1 Flow chart on the generation of an energy conserved bimodal spectrum

In Equations (3.1) and (3.2) above, H_s represents the significant wave height, T_{ij} the peak period, and γ_{ij} is the peak enhancement factor of the spectrum with i , and j representing the equivalent swell and wind sea components. σ is sigma defining the frequency ranges of the spectra widths σ_a to the left and σ_b to the right of the peak frequencies f_{ij} .



(a) Swell wave spectrum from Equation (3.2) for j (b) Wind wave spectrum from Equation (3.2) for i



(c) Bimodal wave spectrum from wind and swell

Figure 3.2 Building up of energy-conserved bimodal spectrum from separate swell and wind components

The spectra widths σ can be expressed as:

$$\sigma = \begin{cases} \sigma_a : 0.07 \text{ for } f \leq f_{ij} \\ \sigma_b : 0.09 \text{ for } f \geq f_{ij} \end{cases} \quad (3.3)$$

To apply these sets of equations in this study, we first compute the equivalent wavelength for both swell and the wind and then combine the two wave systems as shown in Figure 3.2. Further details can be found in Goda (2010).

$$\begin{aligned} \text{For wind sea, } H_{m0W} &= 4 * \sqrt{1 - 0.25m_0} \\ \text{and for swell, } H_{m0S} &= 4 * \sqrt{0.25m_0} \end{aligned} \quad (3.4)$$

The H_{m0S} is the spectral wave height due to swell sea, H_{m0W} is the spectral wave height due to wind sea, and m_0 is the total energy in the overall bimodal spectrum.

It can be shown in Figure 3.2b that the wind wave spectrum and the swell can be separated by the separation frequency f_s . Goda (2010) suggested the frequency of occurrence to be approximately 90 percent of the spectral peak frequency of the resulting bimodal spectrum. This is equivalent to $0.124 H_z$ in Figure 3.2b.

3.1.1 Spectra Shape Parameter

Spectral shape changes according to the severities of the sea state. Ochi (2005) suggested that spectral shapes are dependent on the wind speed, the severity of wind velocity, the duration of wind blowing, and the fetch length. Values of spectral shapes (represented by γ) generally vary from 1 to 7. With each shape value representing the age of the sea. A fully developed sea-state is assumed to have a γ value of 1. Different studies have examined these fundamentals (for example see Goda, 2010; Hasselmann et al., 1973). However, it is challenging to determine whether or not sea-states are fully developed in practice because as a sea-state grows towards maturity, another sequence of swell waves is forming to disrupt the process. In this regard, a γ value of 3.3 was chosen for the wind sea according to the recommendations specified in Hasselmann et al. (1973) and Goda (2010). Similarly, a shape parameter value 2.5 was selected for the swell waves based on the recommendations in Bradbury et al. (2007). Figure 3.3 presents different assessment of γ values.

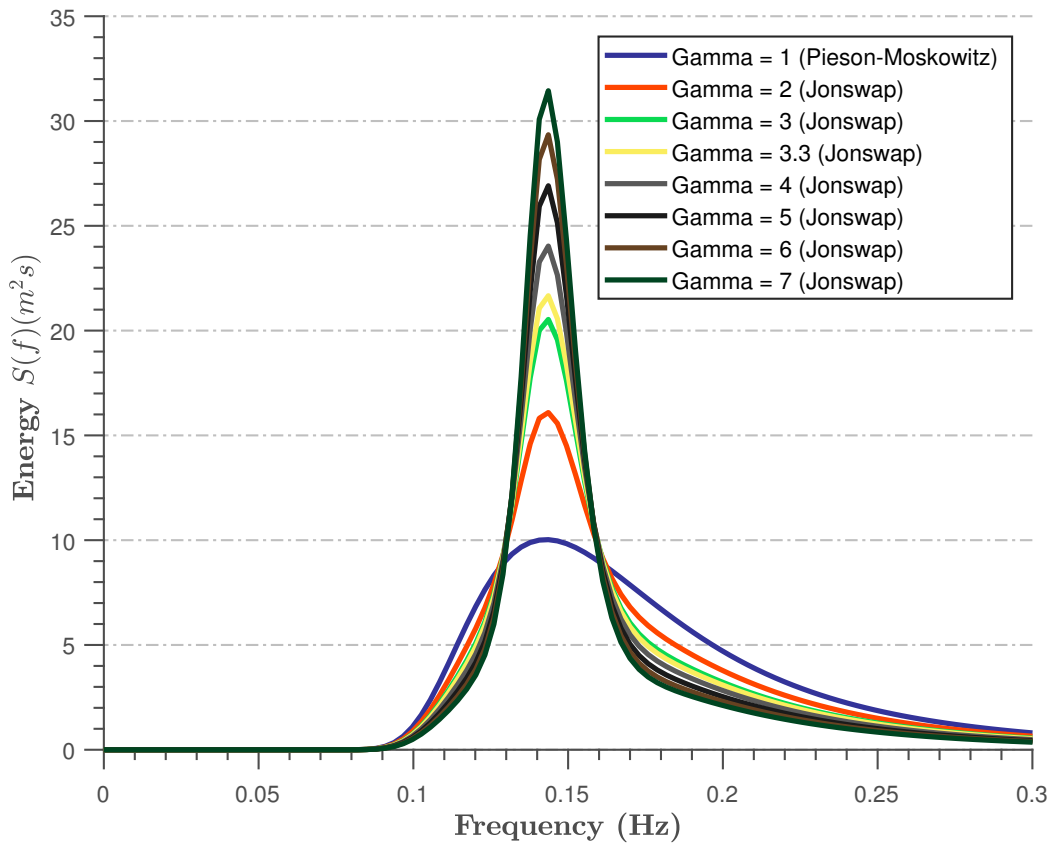


Figure 3.3 Different peaks enhancements definitions for a sea state.

3.2 Estimation of Total Energy in the Bimodal Spectrum

The total energy m_0 in the bimodal spectrum system was computed using Equation (3.4). As shown in step 1 of Figure 3.4, a direct combination of the energy of the superposition of swell waves on the local wind waves were estimated. The input parameters due to swell and wind waves were varied within some specified tolerance. The tolerance was specified in the beginning of the computation to define the degree of closeness of the computed parameters with the input parameters.

Inspecting the total energy as,

$$m_0 = \int f^n E(f) df \quad (3.5)$$

In Equation (3.4), n represents the n^{th} order of the spectral moment of the combined spectrum for total energy, n takes value of zero.

There will be a degree of overlap between the spectra and a sharp break between swell and wind sea as a separation frequency is not present. This contrasts with the practice-based approach, (see, Reeve et al., 2015), in which the swell and wind-sea components have no overlap, thus ensuring that the energy in the sea states is completely consistent with what it will be computed directly from the significant waveheight H_{m0} .

3.2.1 Choices of Bimodal Spectra

Different options can be applied to developing a bimodal sea state. The most conventional approach is based on a direct superposition of the swell wave on the wind-wave using Equation (3.1). An example of this is presented in Figure 3.4. Only the swell component of the bimodal spectrum is influenced. The growth of swell in the bimodal spectrum indicates an overall general increase in the total energy in the sea.

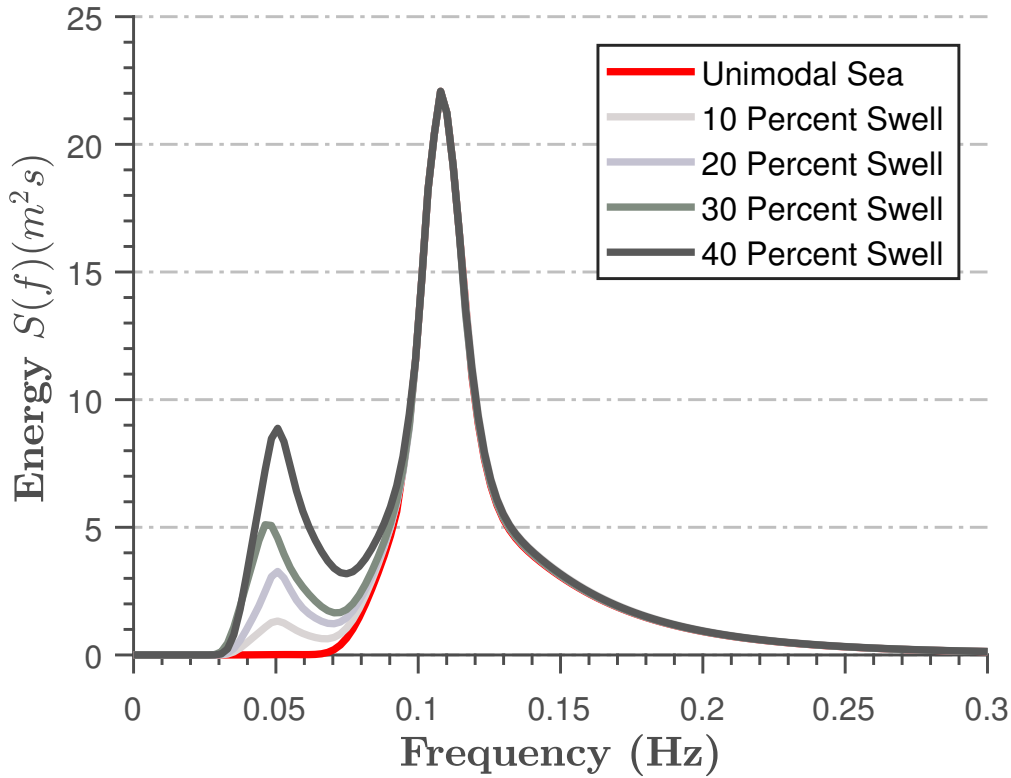


Figure 3.4 An example of spectrum development using direct combination of different swell contents to the wind-sea states without any energy modification

To avoid the geometric energy growth, the wave parameters were internally adjusted (gradual by increasing or decreasing the spectral wind wave height H_{mW} , the spectral wind wave period T_{pW} , the spectral swell wave height H_{mS} , and the spectral wind wave period T_{pS}) from the input bimodal wave components. It produced a more refined energy balance but not enough to conserve the sea state (Figure 3.5). Using this technique, the energy performance of the bimodal spectral was initially erratic but tended to balance to the percentages within the tolerance value. These were not enough to conserve the energy because the wave parameters in the bimodal sea do not consistently converge. The wave parameters are dynamically adjusted using the matlab iteration process described in the Figure 3.1. In this way, the intersection of the swell energy and the wind wave energy are required in order to conserve the total energy m_0 in the new bimodal spectrum.

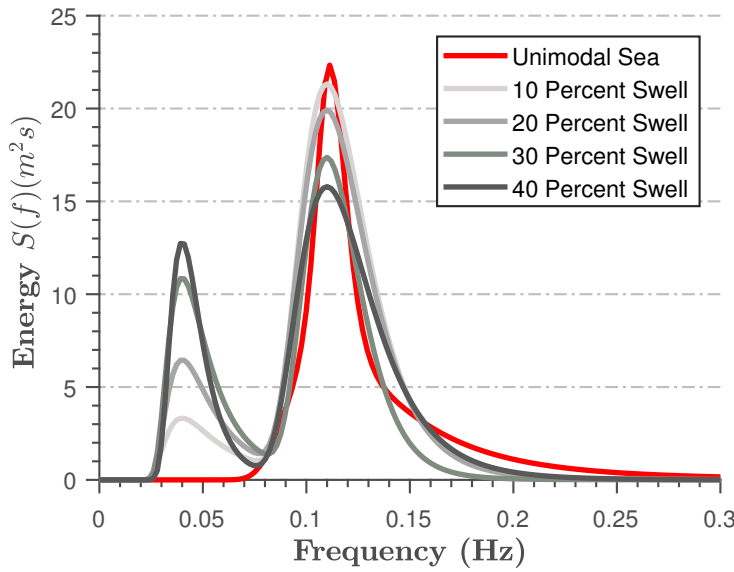


Figure 3.5 An example of spectrum development by modifying parameter inputs while still combining different proportions of swell in the bimodal spectrum

3.3 Ensuring Energy conservation in the Bimodal Spectrum

As proposed in the previous section, the influence of swell presence in a sea state condition can be best examined with bimodal sea states when the overall energy in the sea can be conserved by both components of the wind and the swell present in the sea. This was achieved in this case by solving Equations (3.1) – (3.4) iteratively as illustrated in Figure 3.1. The resulting bimodal spectrum yields the same total energy m_0 given by Equation (3.4). As shown in Figure 3.7, the newly computed significant wave height H_{m0} of the bimodal spectrum was ensured to be consistent with any swell percentage and swell frequency introduced.

To achieve this condition, the entire Equations (3.1) – (3.4) were solved iteratively until the energy m_0 obtained for the significant wave height H_{m0} in Equation (3.5) is fulfilled in step 3 of Figure 3.1. The resulting bimodal spectrum was converted into time series using the Inverse Fast Fourier Transform IFFT technique.

$$H_{m0} = \sqrt{H_{m0S}^2 + H_{m0W}^2} \quad (3.6)$$

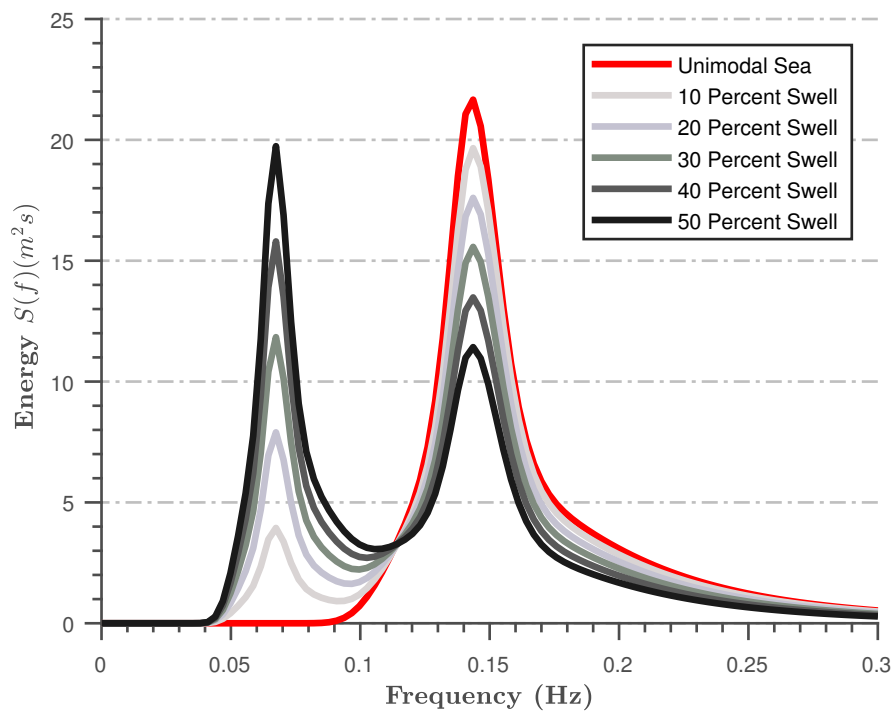


Figure 3.6 An example of an energy-conserved bimodal spectrum

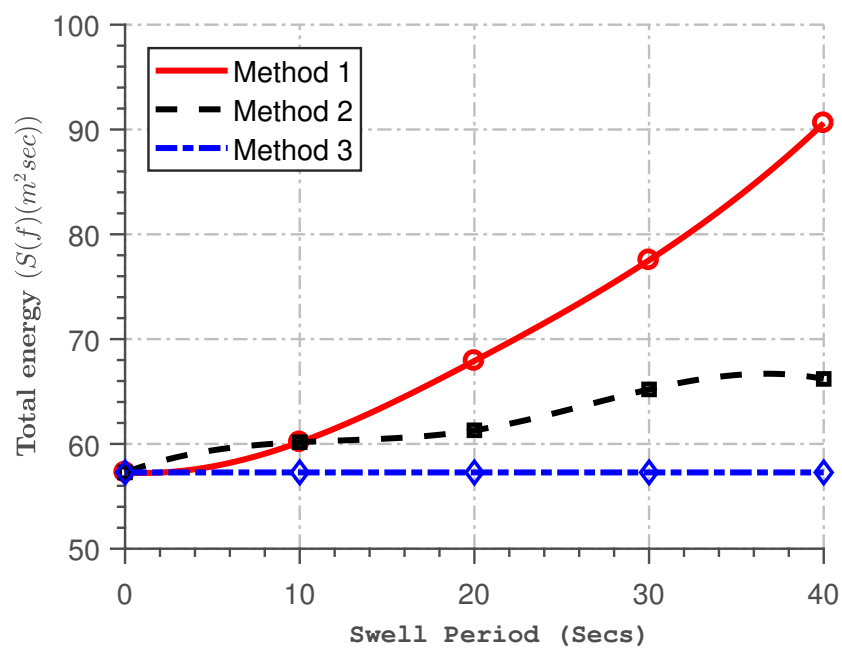


Figure 3.7 Comparison of total energy obtained for different methods of combining swell and wind.

3.4 Variations in Energy-Conserved Bimodal Spectrum

Figure 3.8 presents an example of a sea state with H_{m0} of 4 m and $T_{pW} = 7$ s. The swell component was fixed at a period varied from 11 s to 25 s while maintaining the same H_{m0} . Assuming an introduction of a 25 percent swell component into the sea, the remaining 75 % of the energy will be allocated to the wind-sea using the SSER approach proposed by Guedes Soares Guedes Soares (1984) as stated in Equation (3.3). The exception to the rule here is that the analytical relationship was solved iteratively until the total energy in the system was conserved.

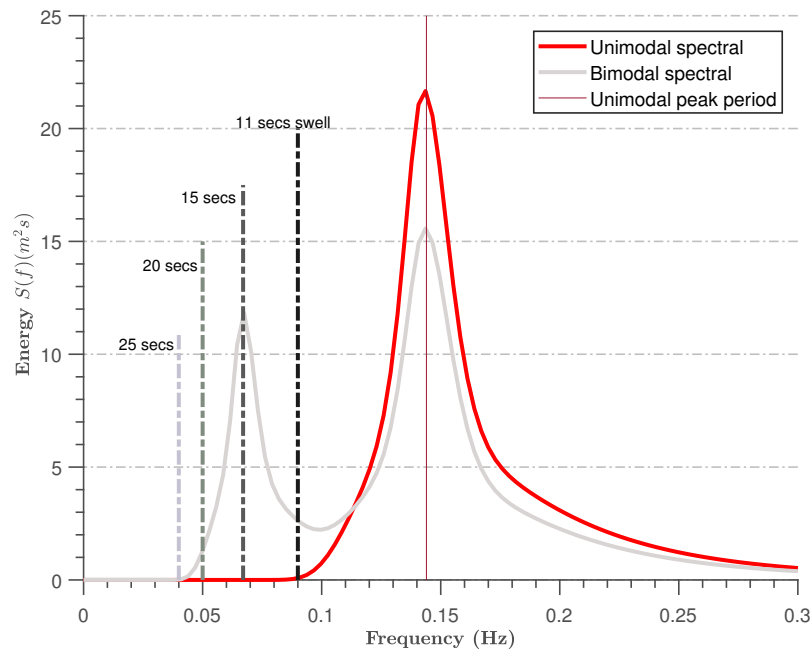


Figure 3.8 An example of the development of energy-conserved showing different frequencies of swell

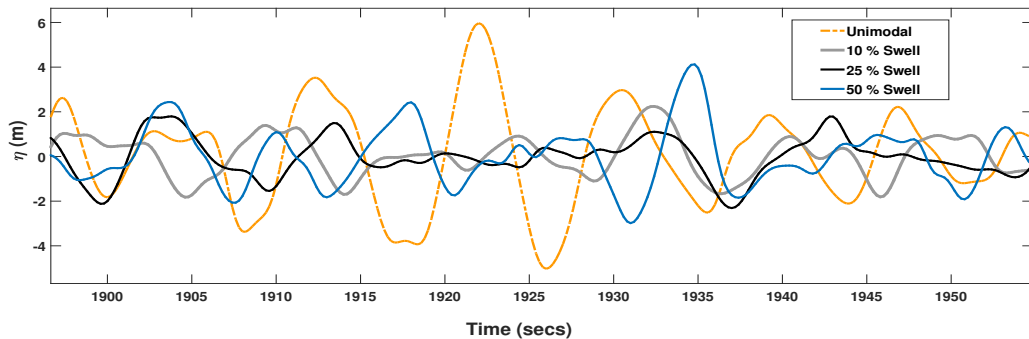


Figure 3.9 An example of the development of energy-conserved bimodal spectrum showing surface elevation obtained by Inverse Fast Fourier Transform (IFFT).

This method gives a different definition of swell compared to that based on separation of frequency because the overlap between the swell and the wind waves was preserved in this method. As earlier mentioned, different bimodal spectra were converted into time series using the Inverse Fast Fourier Transform IFFT technique. Figure 3.9 gives an example of the converted surface elevations for different swell content of a single bimodal spectrum containing the same energy. In this study, considering a sea state with H_{m0} of 4 metres and $T_{pW} = 7$ seconds, Table 3.1 gives a full representation of the number of wave conditions derived from this single state imposed by different swell components at differently occurring swell peak periods and percentages.

Table 3.1 Generation of 13 different wave conditions from the same-energy sea state with H_{m0} of 4 metres and $T_{pW} = 7$ secs.

ID	Test-No	T_{pW} (secs)	T_{pS} (secs)	Swell (%)	Mode
					(Unimodal = 1 Bimodal = 2)
1	T001-1	7	11	25	2
2	T001-2	7	15	25	2
3	T001-3	7	20	25	2
4	T001-4	7	25	25	2
5	T001-5	7	11	50	2
6	T001-6	7	15	50	2
7	T001-7	7	20	50	2
8	T001-8	7	25	50	2
9	T001-9	7	11	75	2
10	T001-10	7	15	75	2
11	T001-11	7	20	75	2
12	T001-12	7	25	75	2
13	T001-13	7	N/A	0	1

As clearly presented in this table, the presence of swell in a single sea state can be investigated using different bimodal spectra while the sea state still retains the same energy content. Generally in present study, more attention will be devoted to four main T_{pS} (11, 15, 20, & 25) and three swell percentages (25, 50, & 75) all matched up with the unimodal spectrum as presented in Table (3.1).

3.5 Conclusion

In this chapter, the formation of different bimodal spectra with constant energy was the unimodal sea states was proposed. These spectra are based on varying swell peak periods and percentages while preserving the overall energy content. In this way, the resulting spectra will be comparable to the original unimodal sea state rather than applying a sea state with continuously increasing energy due to swell. In contrast to the conventional methods, there will be a degree of overlap between the spectra and a sharp break between the swell and wind sea as a separation frequency is not present. It contrasts with the practice-based approach, in which the swell and wind-sea components have no overlap, thus ensuring that the energy in the sea states is entirely consistent with what it will be computed directly from the significant wave height H_{m0} . Specified bimodal spectra were converted into time series using inverse FFT. The energy-conserved bimodal spectrum will be applied as input into both the numerical and the experimental model to simulate the effect of swell waves in such simulations that will be carried out in the rest of the chapters. Moreover, how the present study conforms with previous formulations due to the presence of swell waves will be thoroughly investigated in this study.

Chapter 4

Numerical Implementation

This chapter examines the description of the IH2VOF computational model used in this study. This includes: the mathematical formulation of the model, boundary and initial conditions, methods of wave generation, the numerical domain, surface tracking techniques and the numerical resolution schemes. Basically, pre-computed bimodal wave conditions in the form of time-varying surface elevations, horizontal and vertical velocities were generated and simulated within the numerical model. As initial conditions, the model considers mean flow in the whole domain with still water level without waves under zero velocities and hydrostatic pressures. The model is also represented by the turbulence kinetic energy that produces a numerical perturbation. These conditions are specified in the input file of bimodal waves that are passed into the model boundary.

4.1 Governing Equations of IH2VOF

The IH2VOF model has been proposed by Lara et al. (2011) and Torres-Freyermuth et al. (2014). The numerical model is based on the Volume of Fluid (VOF) technique to create a modified version of the first generation RANS model introduced by Lin and Liu (1998). Fundamentally in turbulent flow, the instantaneous velocity field u_i and pressure field p can be expressed in terms of the mean velocity and pressure components (\bar{u}_i and \bar{p}), and complemented by the turbulent velocity u'_i and pressure fluctuations p' . All these can be expressed mathematically as Equations (4.1)a and (4.1)b. For instantaneous (mean) velocity,

$$u_i = \bar{u}_i + u'_i \quad (4.1a)$$

and for pressure components,

$$p = \bar{p} + p'_i \quad (4.1b)$$

Noting that from Equations 4.1(a and b), $i = 1, 2$ for two dimensional flow. By assuming fluid incompressibility, the Navier-Stokes equations can be further decomposed into Equation 4.2:

$$\frac{\partial \bar{u}_i}{\partial x_i} = 0 \quad (4.2a)$$

$$\frac{\partial \bar{u}_i}{\partial t} + \bar{u}_i \frac{\partial \bar{u}_i}{\partial x_j} = -\frac{1}{\rho_f} \frac{\partial \bar{p}}{\partial x_i} + g_i + \frac{1}{\rho} \frac{\partial \bar{\tau}_{ij}}{\partial x_j} - \frac{\partial \overline{u'_i u'_j}}{\partial x_j} \quad (4.2b)$$

where ρ_f is the density of the fluid, g_i is the i^{th} component of gravitational acceleration, with μ describing the molecular viscosity and $\bar{\tau}_{ij}$ is mean viscous stress tensor. It can be represented mathematically as:

$$\bar{\tau}_{ij} = 2\mu \bar{\sigma}_{ij} \quad (4.3)$$

Equation 4.4 presents the mathematical aggregation of the rate of strain tensor described by the mean flow $\bar{\sigma}_{ij}$:

$$\bar{\sigma}_{ij} = \frac{1}{2} \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \quad (4.4)$$

Also, the pattern of the Reynolds stress tensor R_{ij} of the momentum equation (already presented in Equation (4.2)b) can be defined in terms of the product of the density of the fluid ρ and combined turbulent velocities $\overline{u'_i u'_j}$.

$$R_{ij} = \rho \overline{u'_i u'_j} \quad (4.5)$$

IH2VOF performs calculation of complicated flows under breaking waves. The Reynolds stress term considered in Equation (4.5) is closed with a second order model described in Jaw and Chen (1998). The stress tensor is considered to be related to the strain rate of the mean flow through the algebraic non-linear $\kappa - \varepsilon$ model described in

Lin and Liu (1998):

$$\begin{aligned} \rho \overline{u'_i u'_j} = & \frac{2}{3} \rho k \delta_{ij} - C_d \rho \frac{\kappa^2}{\varepsilon} \left(\frac{\partial \bar{u}_i}{\partial x_i} + \frac{\partial \bar{u}_j}{\partial x_j} \right) \\ & - \rho \frac{\kappa^3}{\varepsilon^2} \left[C_1 \left(\frac{\partial \bar{u}_i}{\partial x_l} \frac{\partial \bar{u}_i}{\partial x_i} + \frac{\partial \bar{u}_j}{\partial x_l} \frac{\partial \bar{u}_j}{\partial x_i} - \frac{2}{3} \frac{\partial \bar{u}_l}{\partial x_k} \frac{\partial \bar{u}_k}{\partial x_l} \delta_{ij} + \right) \right. \\ & \left. + C_2 \left(\frac{\partial \bar{u}_i}{\partial x_k} \frac{\partial \bar{u}_j}{\partial x_k} - \frac{1}{3} \frac{\partial \bar{u}_l}{\partial x_k} \frac{\partial \bar{u}_l}{\partial x_l} \delta_{ij} \right) + C_3 \left(\frac{\partial \bar{u}_k}{\partial x_i} \frac{\partial \bar{u}_k}{\partial x_k} - \frac{1}{3} \frac{\partial \bar{u}_l}{\partial x_k} \frac{\partial \bar{u}_l}{\partial x_k} \delta_{ij} \right) \right] \quad (4.6) \end{aligned}$$

In Equation (4.6), C_d , C_1 , C_2 , and C_3 are empirical coefficients, while δ_{ij} is the Kronecker coefficient. The turbulent kinetic energy k can be expressed as:

$$k = \frac{1}{2} \overline{u'_i u'_i} \quad (4.7)$$

and the dissipation rate of the turbulent kinetic energy ε is given by:

$$\varepsilon = \nu \overline{\left(\partial u'_i / \partial x'_j \right)^2} \quad (4.8)$$

The molecular kinematic viscosity can be expressed in terms of the ratio of the molecular viscosity μ_v and the density ρ of the fluid is given by (μ_v / ρ) .

Equating empirical coefficients $C_1 = C_2 = C_3 = 0$ in Equation (4.6), results in the conventional linear (isotropic) eddy viscosity model for the Reynolds closure expressed in Equation (4.9):

$$\overline{u'_i u'_j} = -2\nu_i \overline{\sigma_{ij}} + \frac{2}{3} k \delta_{ij} \quad (4.9)$$

Similarly, the eddy viscosity ν of the model can be mathematically represented as:

$$\nu_i = C_d \left(\frac{k^2}{\varepsilon} \right) \quad (4.10)$$

Champagne et al., (1970) determined the values of the coefficients C_2 and C_3 from results of experiments conducted on turbulent shear stress flow. Shih et al., (1996) assumed that $C_1 = 2C_3$. Rodi (1980) proposed the value of C_d . Altogether, the following values of coefficients have been proposed so far:

$$\begin{aligned}
C_d &= 0.09 \\
C_1 &= 0.0054 \\
C_2 &= -0.0171 \\
C_3 &= 0.0027
\end{aligned}
\tag{4.11}$$

However, it was observed in Lara et al. (2003) that these coefficients could lead to numerical inconsistencies when implemented in Equation (4.2)b. Therefore modified expressions were implemented in the IH2VOF as follows:

$$\begin{aligned}
C_d &= \frac{2}{3} \left(\frac{1}{7.4 + S_{max}} \right) \\
C_1 &= \left(\frac{1}{185.2 + D_{max}^2} \right) \\
C_2 &= \left(\frac{1}{58.5 + D_{max}^2} \right) \\
C_3 &= \left(\frac{1}{370.4 + D_{max}^2} \right)
\end{aligned}
\tag{4.11}$$

In these equations, S_{max} is given by:

$$S_{max} = \frac{k}{\varepsilon} \max \left[\left| \frac{\partial \bar{u}_i}{\partial x_i} \right| \right]
\tag{4.12}$$

and,

$$D_{max} = \frac{k}{\varepsilon} \max \left[\left| \frac{\partial \bar{u}_i}{\partial x_j} \right| \right]
\tag{4.13}$$

According to Rodi (1980), and also reiterated by Lin and Liu (1998), newer values are proposed for k and ε as shown in Equation (4.14) – (4.15):

$$\frac{\partial k}{\partial t} + \bar{u}_j \frac{\partial k}{\partial x_j} = \frac{\partial}{\partial x_j} \left[\left(\frac{v_i}{\sigma_k} + \nu \right) \frac{\partial k}{\partial x_j} \right] - \overline{u'_i u'_j} \left(\frac{\partial \bar{u}_i}{\partial x_j} \right) - \varepsilon \quad (4.14)$$

$$\frac{\partial \varepsilon}{\partial t} + \bar{u}_j \frac{\partial \varepsilon}{\partial x_j} = \frac{\partial}{\partial x_j} \left[\left(\frac{v_i}{\sigma_\varepsilon} + \nu \right) \frac{\partial \varepsilon}{\partial x_j} \right] + C_{1\varepsilon} \frac{\varepsilon}{k} \nu_i \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \frac{\partial \bar{u}_i}{\partial x_j} - C_{2\varepsilon} \frac{\varepsilon^2}{k} \quad (4.15)$$

In Equations (4.14) & (4.15), σ_k , σ_ε and $C_{2\varepsilon}$ are empirical coefficients. The first and second terms represent the convection and advection terms respectively. Similarly, the third and fourth terms refer to the production and the dissipation of kinetic turbulent energy.

The IH2VOF model considers a better simulation of the porous nature of materials than most previously represented RANS models. It is because flows in porous media were described by using the volumes of the medium as described by the VARANS equation instead of the ordinary rigid boundary of obstacles (for example, Lin and Liu (1998)). The final form of the VARANS equation can be presented as Equation (4.16):

$$\begin{aligned} \frac{1 + C_A}{n} \frac{\partial(\tilde{u}_i)}{\partial t} + \frac{\tilde{u}_j}{n^2} \frac{\partial \tilde{u}_i}{\partial x_j} = & -\frac{1}{\rho} \frac{\partial \tilde{p}}{\partial x_i} + \rho g_i + \frac{\nu}{n} \frac{\partial^2 \tilde{u}_i}{\partial x_j \partial x_j} \\ & - \frac{1}{n^2} \frac{\partial \tilde{u}_i \tilde{u}_j}{\partial x_j} - \frac{\alpha \nu (1 - n^2)}{n^3 D_{n50}^2} \tilde{u}_i - \frac{\beta \nu (1 - n)}{n^3 D_{50}} \tilde{u}_i |\vec{\tilde{u}}| \end{aligned} \quad (4.16)$$

The porosity n , rock mean size D_{n50} , and the three coefficient of frictions (C_A , α and β parameters are used to establish the effects of porous media on the entire model behaviour.

The turbulence k - ε and the total stress τ_{ij} in all the directions in the model are given by:

$$\begin{aligned}
 \frac{\partial k}{\partial t} + u_j \frac{\partial k}{\partial x_j} &= \frac{\partial}{\partial x_j} \left[\left(\frac{\nu_t}{\sigma_k} + \nu \right) \frac{\partial k}{\partial x_j} \right] + 2\nu_t \sigma_{ij} \frac{\partial u_i}{\partial x_j} - \varepsilon \\
 \frac{\partial \varepsilon}{\partial t} + u_j \frac{\partial \varepsilon}{\partial x_j} &= \frac{\partial}{\partial x_j} \left[\left(\frac{\nu_t}{\sigma_\varepsilon} + \nu \right) \frac{\partial \varepsilon}{\partial x_j} \right] + C_{1\varepsilon} \frac{\varepsilon}{k} 2\nu_t \sigma_{ij} \frac{\partial u_i}{\partial x_j} \\
 &\quad - C_{2\varepsilon} \frac{\varepsilon^2}{k} \\
 \tau_{ij} &= 2\left(\nu + C_d \frac{k^2}{\varepsilon}\right) - \frac{2}{3} k \delta_{ij}
 \end{aligned}
 \tag{4.17}$$

The model has the capability of replicating the scalability and geometrical compatibility linking a numerical simulation and experimental studies. It can also replicate nearshore processes with minimal assumptions and complications. The initiation control of wave breaking in the design is accomplished by applying an advanced turbulence kinetic energy k and the equivalent dissipation ε as expressed in Equation (4.17).

4.2 Initial Conditions and Boundary Conditions

This model implements three different boundary conditions for wave generation and absorption. In IH2VOF, different boundary conditions are defined. Basically as in other 2 dimensional models, it consists of solid boundaries, free surface boundary, and the open boundary conditions. These will be considered in details here:

4.2.1 Initial Conditions

In the initial state, the model considers mean flow in the whole domain as still water level with no wave or current motion. This refers to a state of zero velocities and hydrostatic pressure. The model can also consider a pre-determined form for the free surface displacement or mean velocity field. In this mode, variations in the free surface and velocity conditions can be fully described. To prevent singularity errors due to turbulence during simulation in the model (defined in Equation (4.14)), the turbulence was initialised with a "seeding" value that produces a numerical perturbation as:

$$k = \frac{1}{2} u_i^2 \tag{4.18}$$

where $u_i = \delta c_i$, c_i the wave celerity in the wave generation zone, and δ is a numerical constant. The value of δ is defined by Lin and Liu, (1998) as equivalent to 0.0025.

For turbulent dissipation rate ε , IH2VOF initial value can be defined by this expression:

$$\varepsilon = C_d \frac{k^2}{v_i} \quad (4.19)$$

where $v_i = \xi v$ and ξ defined by Lin & Liu, (1998) as 0.1. The variations of δ and ξ have been determined to be insignificant on the final result of the computational model. Lin and Liu (1998) concluded that δ could induce slight delays in the initiation of wave breaking performances of flow conditions.

4.2.2 Solid Boundary Condition

Two types of conditions for mean flow are considered. These are:

- Free slip: $\overline{u_n} = 0$, $\frac{\partial \overline{u_\tau}}{\partial n} = 0$
- No slip: $\overline{u_n} = 0$, $\overline{u_\tau} = 0$

where n and τ are directions normal and parallel to the boundaries respectively. Also for turbulent flows, a log-law distribution for the mean tangential velocity is implemented in the model. This can be expressed as:

$$\frac{\partial \overline{u}}{y} = \frac{u_n}{\kappa y} \quad (4.20)$$

Where κ represents the von Karman constant ($\kappa = 0.41$), y is the distance from the solid domain, u is the friction velocity. By integration of Equation (4.20), we can obtain individual values of κ and ε as:

$$\kappa = \frac{u_n^2}{\sqrt{C_d}} \quad (4.21)$$

and,

$$\varepsilon = \frac{u_n^2}{\kappa y} \quad (4.22)$$

The values of u_n can also be determined from the values of mean flow as:

$$u_n^2 = v \frac{\partial \overline{u}}{\partial y} \Big|_{y=0} \quad (4.23)$$

4.2.3 Free Surface Boundaries

The IH2VOF computational model ignores the mean density fluctuations near the free surface. During laminar flow, the zero stress and zero pressure conditions are imposed at the free surface as:

$$\bar{p} = 0 \quad (4.24a)$$

$$\frac{\partial \bar{u}_\tau}{\partial n} = 0 \quad (4.24b)$$

for the turbulent field, zero-gradient boundary condition was applied to both k :

$$\frac{\partial k}{\partial n} = 0 \quad (4.25a)$$

and to ε :

$$\frac{\partial \varepsilon}{\partial n} = 0 \quad (4.25b)$$

where n is a unit normal to the free surface and its assumed that there is no turbulence exchange between water and air.

4.2.4 Open Boundaries

The open boundary condition implemented in IH2VOF can be expressed as:

$$\frac{\partial \phi}{\partial t} + c_0 \frac{\partial \phi}{\partial x} = 0 \quad (4.26)$$

where ϕ represents the variables to be evaluated during simulation such as \bar{u} , \bar{v} , k , and ε . The wave celerity of log waves at any position can be defined as:

$$c_0 = \sqrt{g(d+a)} \quad (4.27)$$

and for short waves:

$$c_0 = \sqrt{\frac{d\lambda}{2\pi} \tanh\left(\frac{2\pi}{\lambda}(d+a)\right)} \quad (4.28)$$

where a is the wave amplitude, d is the water depth at the location, and λ the wave length.

4.3 Methods of Wave Generation

Three different wave generation methods are possible in the IH2VOF computational model. These are internal wave maker, static paddle using Dirichlet boundary condition, and the virtual force method employing dynamic paddle techniques. A more complete description has been described in Lara et al (2006).

4.3.1 Dirichlet Boundary Condition

The simplest wave generation method in IH2VOF is the Dirichlet boundary condition. This is implemented in most wave generation models because this method provides analytical expressions for free surface and velocity distribution throughout the water column. Here waves are generated on the one side of the domain and it is actively absorbed at the same side of the numerical flume. The VOF and the horizontal and vertical velocity components are two variables needed at each time step. The velocities are generated initially and are then interpolated in a linear fashion into the entire model system. The Dirichlet method is suitable to replicate the behaviour of any laboratory wave paddle at a cell size resolution such as the piston-type wave generators. When replicating a piston-type wave maker, evanescent modes are shown in the wave generation area as will be expected in a physical flume.

4.3.2 Internal Wave Maker

IH2VOF computational model is capable of generating waves at any specific region using the mass function defined within this region. The mass source method is developed to avoid re-reflection of waves in the boundary where they are generated. The introduction of a mass source function at any point in the fluid domain creates a physical effect. Alternatively, fluid is introduced or sucked into this region to generate waves in terms of crests and troughs. Full details of this wave generation are described in Lin and Liu (1998). The target wavelength is always larger than the width of the mass source function, and the waves are described to have started right above the centre of the region. The relationship between the source function $s(x, z, t)$ and the expected free surface displacement $\eta(t)$ can be expressed as:

$$\int_0^t \int_{\Omega} s(x, z, t) d\Omega dt = 2 \sum \int_0^t c_i \eta_i(t) dt \quad (4.29)$$

In above equation, c is the target wave phase velocity and the factor of two shows that waves are propagating in both directions of the domain. Only one component $n = 1$ is described for regular (monochromatic) waves. In randomly irregular waves with n components, Equation (4.29) consists of sum of n linear waves, with each of them having a wave height, wave period and phase.

4.3.3 Dynamic Boundary Condition

The dynamic boundary conditions can replicate piston type wave generators. A simulated solid object within the mesh is usually used to push the fluid in motion. An openness coefficients would then be used to absorb the movement of the solid object to the wave paddle position $X(t)$ and velocity $U(t)$ at each time step. The virtual body force is applied to bodies of partial cells ($0 < \theta < 1$). This method differs from the Dirichlet method because only the horizontal velocity component and free surface locations are specified within the mesh according to the wave maker movement. The mesh should be large enough to accommodate the movement of the wave maker.

4.4 Wave Absorption Methods

The wave absorption method implemented in the IH2VOF is efficient to cater for reflection of any physical structure present in the model and also to ensure that the simulation could run for longer duration. Both active and passive absorptions are implemented in the model. The active absorption is efficient for the Dirichlet and moving boundary conditions. On the other hand, the passive absorption helps to efficiently absorb reflected waves propagating backwards towards the wave generating end.

4.4.1 Active Wave Absorption

The details of the active absorption implemented in IH2VOF can be found in Schäffer and Klopman (2000). This method involves a clear identification of incident waves approaching the domain and generation of equivalent reverse waves to cancel out the incoming waves. In this way, it possible to implement both the moving boundary and the Dirichlet conditions simultaneously. Cancellation of incident waves by the generated reverse reflected waves in IH2VOF was first implemented in Torres-Freyermuth et al. (2010). The difference between the targeted theoretical free surface η_t and the one

measured in the model η_m is computed. Then the wave maker velocity is modified to reproduce the target wave in agreement with the waves to be absorbed. The correction to the target wave velocity U_r can be expressed mathematically as:

$$U_r = \sqrt{\frac{g}{h}} \eta \quad (4.30)$$

where $\eta_r = \eta_t - \eta_m$. In the case of the moving paddle method, it is necessary to adjust the position of the wave maker by integrating the corrected velocity $U_c = U_t + U_r$

$$X(t) = X(t - dt) + \int_{t-dt}^t (U_t + U_r) dt \quad (4.31)$$

One of the simplest active absorption cases occurs with using the open boundary condition. In this case the target setting of the IH2VOF will be set to zero to absorb waves that approached the boundary.

4.4.2 Passive (Sponge Layer) Wave Absorption

The passive absorption implemented in IH2VOF model follows after Israeli and Orszag (1981). It made it possible to absorb waves that are generated above the source region when using the source generation method. In this case, a dissipation method is applied which can be described mathematically as:

$$\gamma_d(x) = 1 - \alpha_d \beta_d^3 - (1 - \alpha_d) \beta_d^6 \quad (4.32)$$

where $\alpha_d = 0.4$ and

$$\beta_d(x) = \frac{x_0 - x}{\lambda} \quad (4.33a)$$

Also, x_0 is the coordinate of the origin of the source region seaward boundary and λ is the wavelength at the same point. The introduction of sponge layer for wave absorption increases the length of computational domain. This makes the point source method unsuitable for long storm simulation as desired for this study.

4.5 Free Surface Tracking

In the IH2VOF model, free surfaces are tracked by using the Volume of Fluid (VOF) method first suggested by Hirt and Nichols (1981). In this technique, the exact location of the free surface cannot be accurately traced by the model. Free surface is identified by tracking the change in density of each cell. Different cell types are identified within the model and values of VOF are assigned to them: interior (I), surface (S) or empty (E). It is illustrated in Figure 4.1:

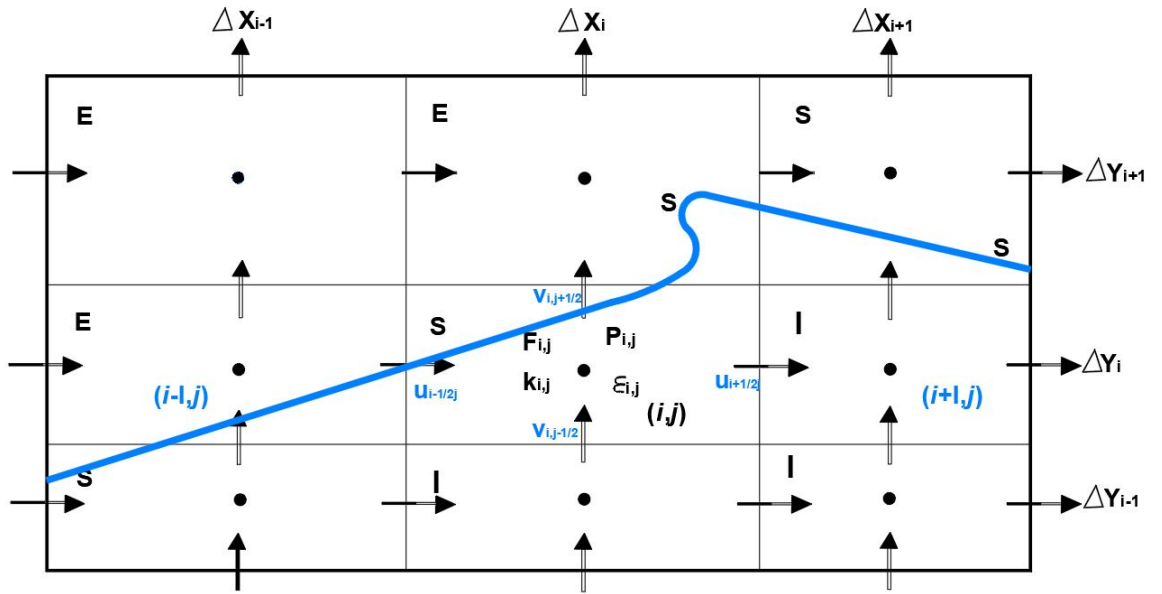


Figure 4.1 The computational domain and the free surface tracking in the model

The density ratio can be stated as,

$$F = \frac{\rho}{\rho_f} \quad (4.34)$$

where,

$$\rho = \frac{\rho_f V_f}{V_f + V_a} \quad (4.35)$$

ρ_f , represents the fluid density, V_f , the volume of the fluid in the cell and V_a being the volume of air within the cell. $F = 1$ represents the interior cell, $F > 0$ for surface

cells, and $F = 0$ for empty cells. As a result of this modification, the mass conservation equation can be transformed into the transport equation in the form of:

$$\rho(x, y, t) = F(x, y, t)\rho_f \quad (4.36)$$

or can be expressed as:

$$\frac{\partial F}{\partial t} + \frac{\partial}{\partial x}(\bar{v}F) = 0 \quad (4.37)$$

Hirt and Nichols (1981) presented an algorithm to eliminate error in the convection of F by evaluating F gradients in two directions to identify the free surface location. VOF algorithm was modified by Lin and Liu (1998). For full details of VOF implementation, readers should consult Hirt and Nichols (1981) and Lin and Liu (1998).

4.6 Numerical Resolution of RANS Equation

Within the numerical domain of the IH2VOF, RANS equations are solved by applying the Chorin (1968, 1969) finite difference two-step projection approach (Maza et al., 2017). Firstly, the following intermediate velocity was introduced into the momentum equation as:

$$\frac{\tilde{u}_i^{n+1} - u_i^n}{\Delta_t} = -u_j^n \frac{\partial u_j^n}{\partial x_j} + g_i + \frac{\partial \tau_{ij}^n}{\partial x_j} \quad (4.38)$$

This equation is a forward time difference equation where \tilde{u}_i^n denotes the intermediate velocity, Δ_t is the time step that correspond to $(n + 1)^{th}$ time level. Equation (4.38) does not include the pressure gradient term and the intermediate velocity does not satisfy the continuity equation.

Using Equation (4.39), the intermediate velocity can be well represented on a divergence-free plane to obtain the final velocity:

$$\frac{u_i^{n+1} - \tilde{u}_i^n}{\Delta_t} = -\frac{1}{\rho^n} \frac{\partial p^{n+1}}{\partial x_j} \quad (4.39)$$

$$\frac{\partial u_i^{n+1}}{\partial x_i} = 0, \quad (4.40)$$

Moreover, the pressure gradient at the next time step $((n+1)^{th})$ can be estimated in the form of the Poisson Pressure Equation (PPE):

$$\frac{\partial}{\partial x_i} \left(\frac{1}{\rho^n} \frac{\partial p^{n+1}}{\partial x_i} \right) = \frac{1}{\Delta t} \frac{\partial \tilde{u}_i^{n+1}}{\partial x_i} \quad (4.41)$$

By combining Equations (4.39)–(4.41), we can derive a complete momentum conservation equation as:

$$\frac{\tilde{u}_i^{n+1} - u_i^n}{\Delta t} + u_j^n \frac{\partial u_j^n}{\partial x_j} = \frac{1}{\rho^n} \frac{\partial p^{n+1}}{\partial x_i} + g_i + \frac{\partial \tau_{ij}^n}{\partial x_j} \quad (4.42)$$

With the application of the two projection method, velocity components and pressure fields can be spatially derived in finite difference forms. To obtain the convection terms, both the central difference (CD) and the upwind method (UM) are applied. The advantage of the combined methods is to complement the individual drawbacks of implementing each method separately. In this way, the numerical instability and tendencies for wave damping are reduced. Weighing factors are used in the numerical spatial scheme to regulate the influence of each of the two methods (Maza et al., 2017).

For example, for pressure gradients terms, and the stress gradient terms only the central difference method was applied. Detailed implementation of the numerical model can be found in Lin & Liu (1998).

4.7 The Computational Domain

The numerical domain is usually discretised into rectangular cells as shown in Figure 4.1. Sub-meshes can be defined within the domain to represent areas of special interest. Within every cell in the domain, all scalar quantities are defined (including the VOF function (F), pressure (P), turbulent kinetic energy (k), and the rate of dissipation (ϵ). Vector quantities including, the mean velocity components \bar{v} , and the additional openness function (θ_r & θ_t) are also defined.

To test capability of the model at propagating waves, a simple numerical two-dimensional wave flume was set up. As shown in Figures 4.2 & 4.3, the domains each 25 metres long and 1 metres high. The computational domain discretised into 0.004 m along the x-axes (34,394 nodes in green) and 0.01 m along y-axes mesh-sizes (81 meshes in yellow colour) were used to run the simulation with a time step of 0.004 s to maintain a stable simulation over a storm duration. As shown in Figure 4.2, wave gauges spaced out in the numerical flume (W1, W2, W3, and W4) were used to obtain the time series of the surface elevation.

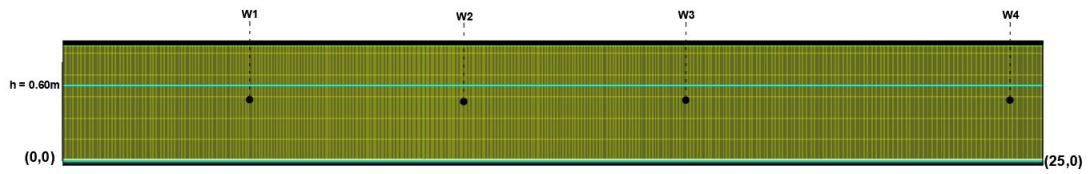


Figure 4.2 An example of simple IH2VOF model domain without any structure

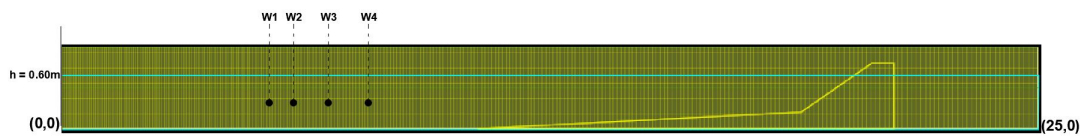


Figure 4.3 An example of simple IH2VOF model domain with a sloping simple seawall

Also, the capability of the model at reproducing a simple sloping seawall was tested. A simple sloping seawall with $\cot \alpha = 1.5$ was positioned in the numerical domain. This configuration will be further described in Chapter 6.

4.7.1 Mesh Convergence Settings

For optimal numerical simulation, it is essential to apply suitable mesh resolutions ($\Delta x: \Delta y$) that can balance the simulation time, and still able to accurately model the nu-

merically determined wave parameters. For this purpose, six different mesh resolutions are applied to determine the numerical domain that will optimally reproduce the input wave conditions. An H_{m0} of 0.125 meters was used as a base for this purpose. As a rule of thumb, $\Delta x < \Delta y$ has been recommended to avoid premature wave breaking and damping (Maza et al., 2017). Also, mesh sizes along the vertical axis must accommodate at least one-tenth of the wave height to be propagated. Along the x-axis, the domain must be stretched to evolve a minimum of 4 wavelengths before intersecting any obstacles.

Figure 4.4 shows an example of the H_{m0} obtained for tests conducted using six different mesh sizes. There is more accurate convergence of H_{m0} between mesh sizes of 7.5 mm to 12.5 mm in every tested cases against the benchmark. The convergence rates are slower for larger number of mesh sizes. Consequently, vertical mesh sizes less than one tenth of the H_{m0} are often selected along the vertical axis as previously suggested in (Maza et al., 2017).

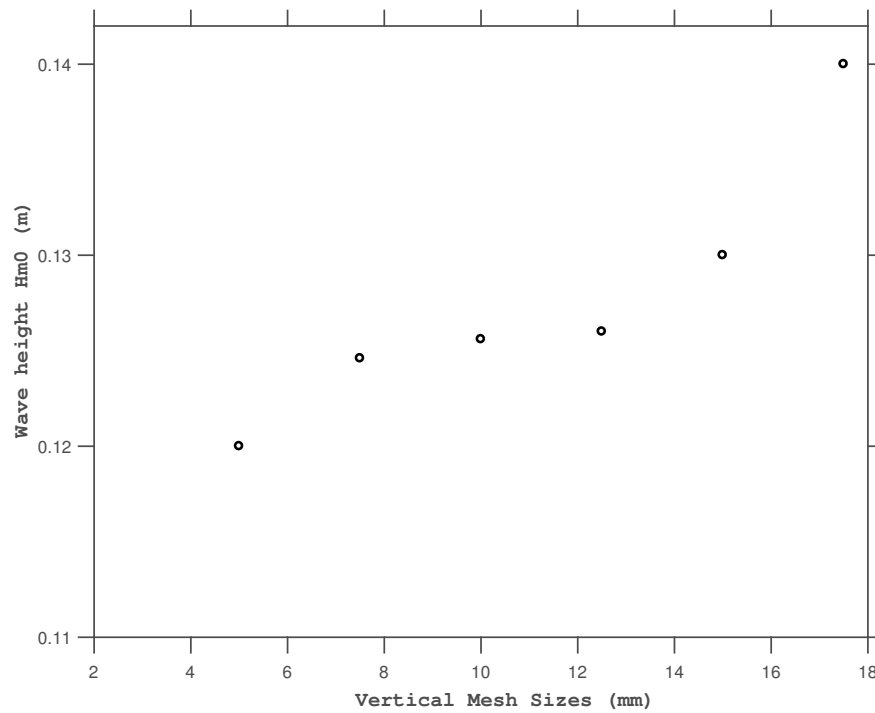


Figure 4.4 An example of the RMSE of the mesh resolution settings for optimal mesh sizing applied

For the regular wave cases, only the empty computational domain was applied. Figure 4.5 presents a representation propagation of a wave with significant waveheight $H_s = 0.08$ metres and peak period $T_p = 1.27$ secs in a water depth $h = 0.6$ metres. As clearly shown, the model is capable of reproducing this wave. The performances were good as the wave arrives each of the wave gauges and the energy build up appear.

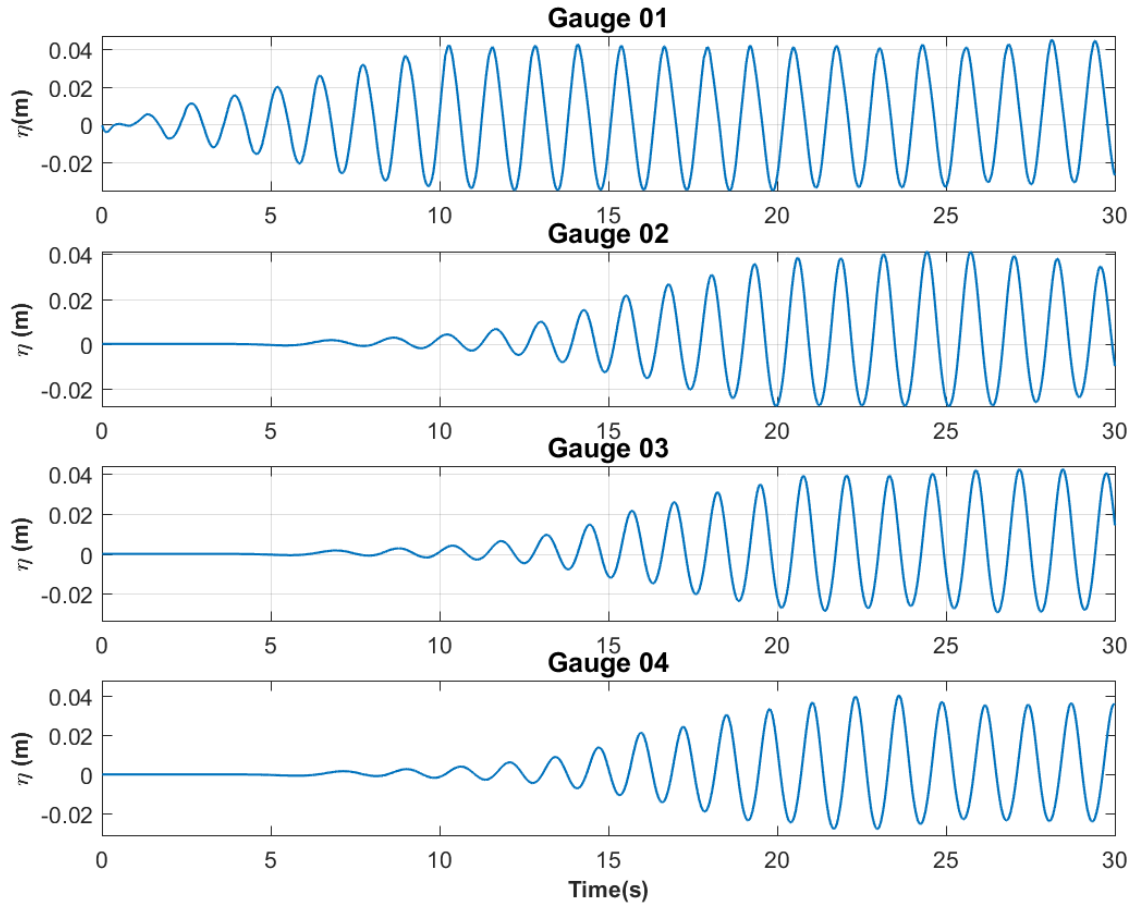


Figure 4.5 An example of regular wave propagation in the computational model for $H_s = 0.08$ metres and $T_p = 1.27$ secs in a water depth $h = 0.6$ metres recorded with wave gauges placed at W1, W2, W3 and W4.

For all the cases considered, 7 wave gauges are used. The numbers and positions of the gauges are listed in Table 4.1. In every cases in this study, waves are generated from the position of $Y = 0$ (Figure 4.3), and are propagated progressively towards the positive X-axis as presented. From the wave elevations across the wave gauges in Figure 4.5, earlier arrival of waves can be seen from earlier wave gauges (Gauge 01) than later ones (Gauges 02 to 04). During this simulation, active absorption of waves is defined at the other side of the domain to avoid excessive inherent reflections.

Table 4.1 Location of wave gauges

Gauge Number (X)	Location (m)
1	10
2	10.3
3	11.1
4	11.6
5	18.1
6	23.6

4.7.2 Generation of Solitary Waves

An example of surface elevations representing the capability of the computational model to accurately generate solitary waves is presented in Figure 4.6. The model propagates the inputted waveheight and wave periods through selected wave gauges. A 10 seconds solitary wave was simulated within 85.71 seconds on an Intel Core *i7* – 6700 CPU processor with 3.40 GH_z \times 8 cores and 15.5 GiB RAM.

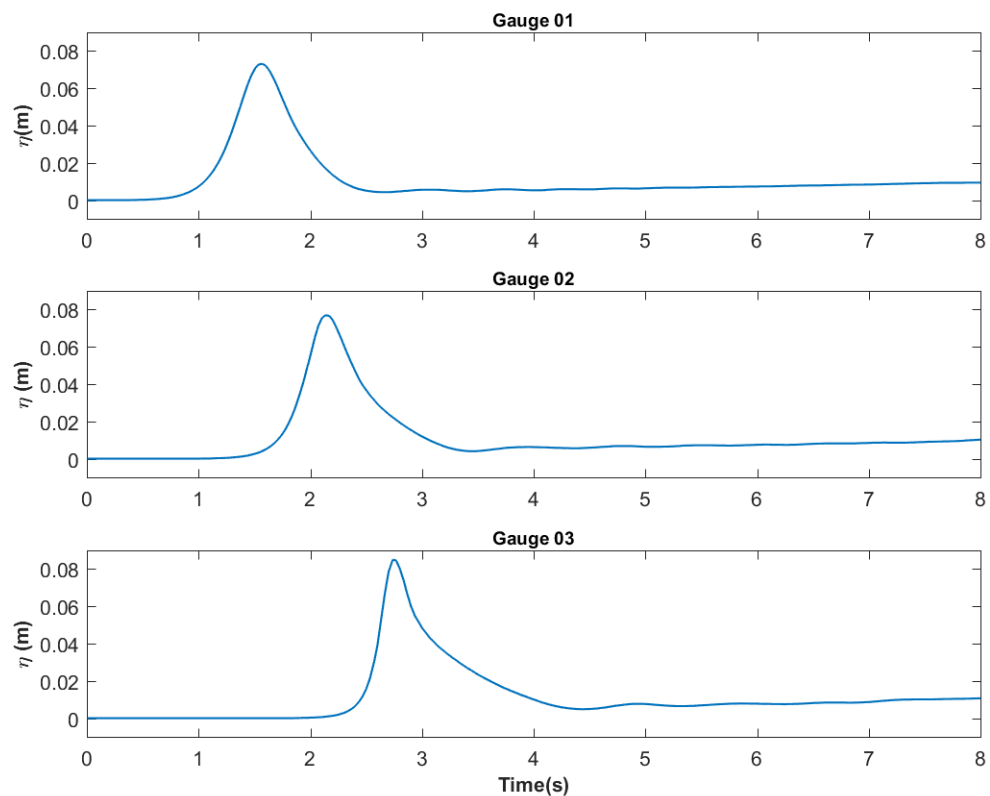


Figure 4.6 Solitary wave propagation in the numerical flume for $H_s = 0.1$ metres and $T_s = 2.01$ secs in a water depth $h = 0.25$ metres recorded at wave gauges placed at W1, W2 and W3.

4.7.3 Generation of Irregular waves

An example of an accurately generated irregular wave by the model is presented in Figure 4.7. Individual irregular wave components are well defined.

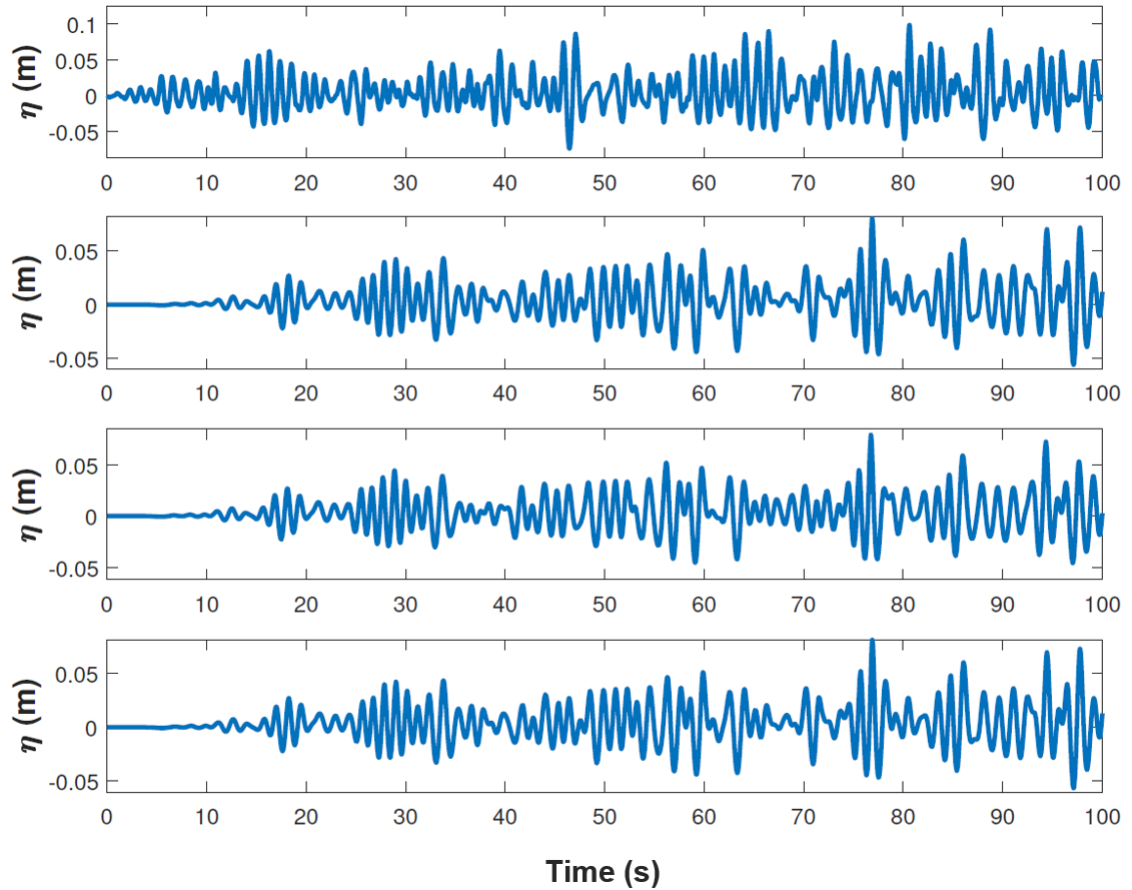


Figure 4.7 Propagation of irregular wave using the dynamic paddles in the numerical flume for $H_s = 0.08$ metres and $T_p = 1.27$ secs in a water depth $h = 0.6$ metres recorded at wave gauges placed at W1, W2, W3 and W4.

4.8 Evaluating Model Performance

Performance of an hydrodynamic model is essential to determine the reliability of the model at propagating and absorbing waves. To achieve this aim, the mass conservation of the model is evaluated in terms of the energy balance equation (Equation (4.43)) in both active and inactive absorption mode. Firstly by operating a closed boundary domain and also by applying the open ones against wave absorption requirements. To assess the performance of this model, the porous switch was set to zero. A wave height

of 0.1 metres was propagated through an inflow and outflow water depth each of 0.6 metres.

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho u) = \odot \quad (4.43)$$

By total integration of Equation (4.43) across the entire domain, and then applying the divergence theorem we have:

$$\int_{\omega} \frac{\partial \rho}{\partial t} + \nabla \cdot (\rho u) dV = \int_{\omega} \frac{\partial \rho}{\partial t} dV + \int_{\partial \Omega} (\rho u) \cdot n dS \quad (4.44)$$

The turbulence and roughness functionalities were initiated. The active absorption were set active at both ends of the computational flume to avoid poor judgement due to reflection effects. From Figure 4.8, both the energy build up with time and dissipation are presented in the Figure.

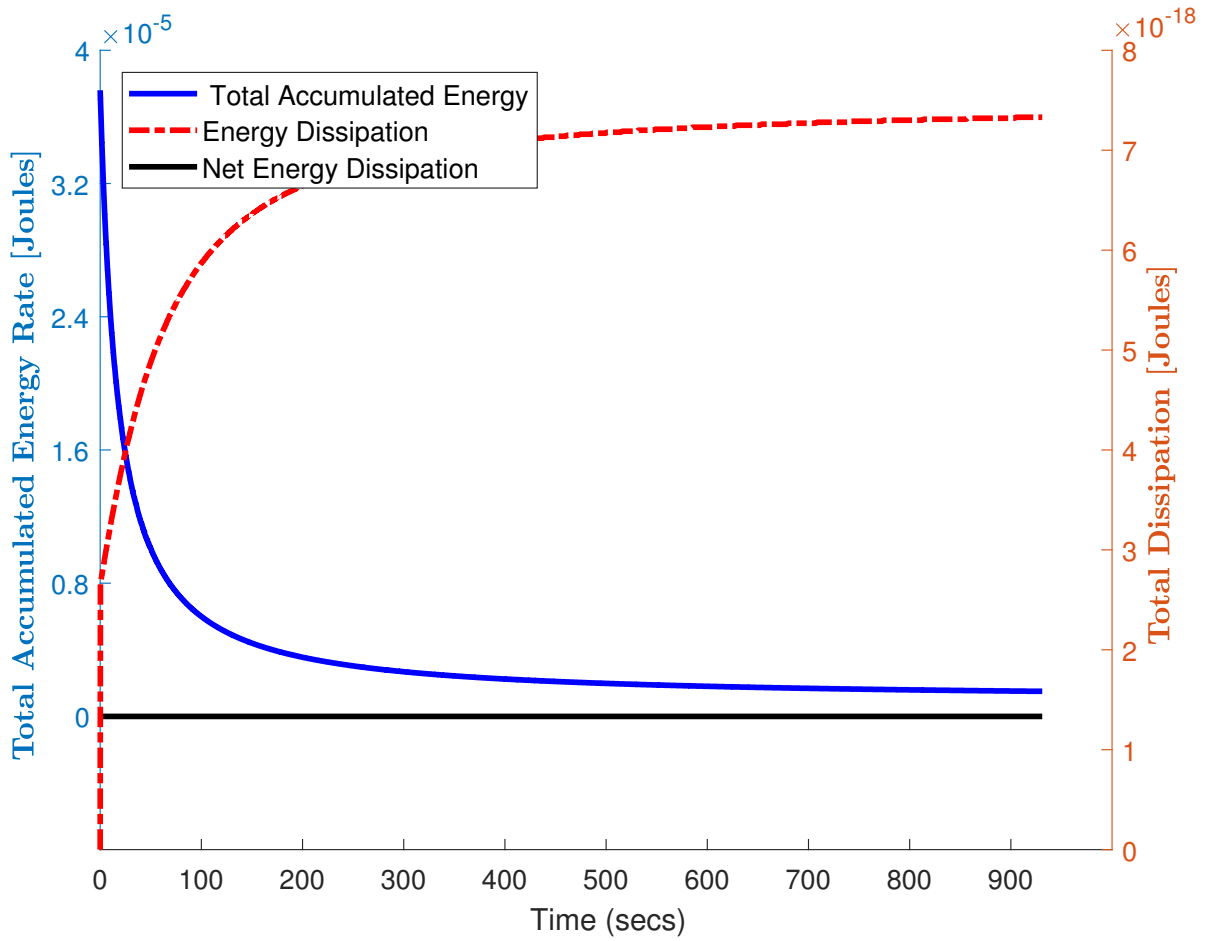


Figure 4.8 An example of the energy rate in IH2VOF with respect to time.

It can be seen in Figure 4.8 that the total accumulated energy rate vanishes as soon as the steady state is reached. There is a progressive increase in the energy dissipation until it reaches a steady state. The black line is the energy balance absolute error. There is a good agreement in the net energy balance which should be zero in this case.

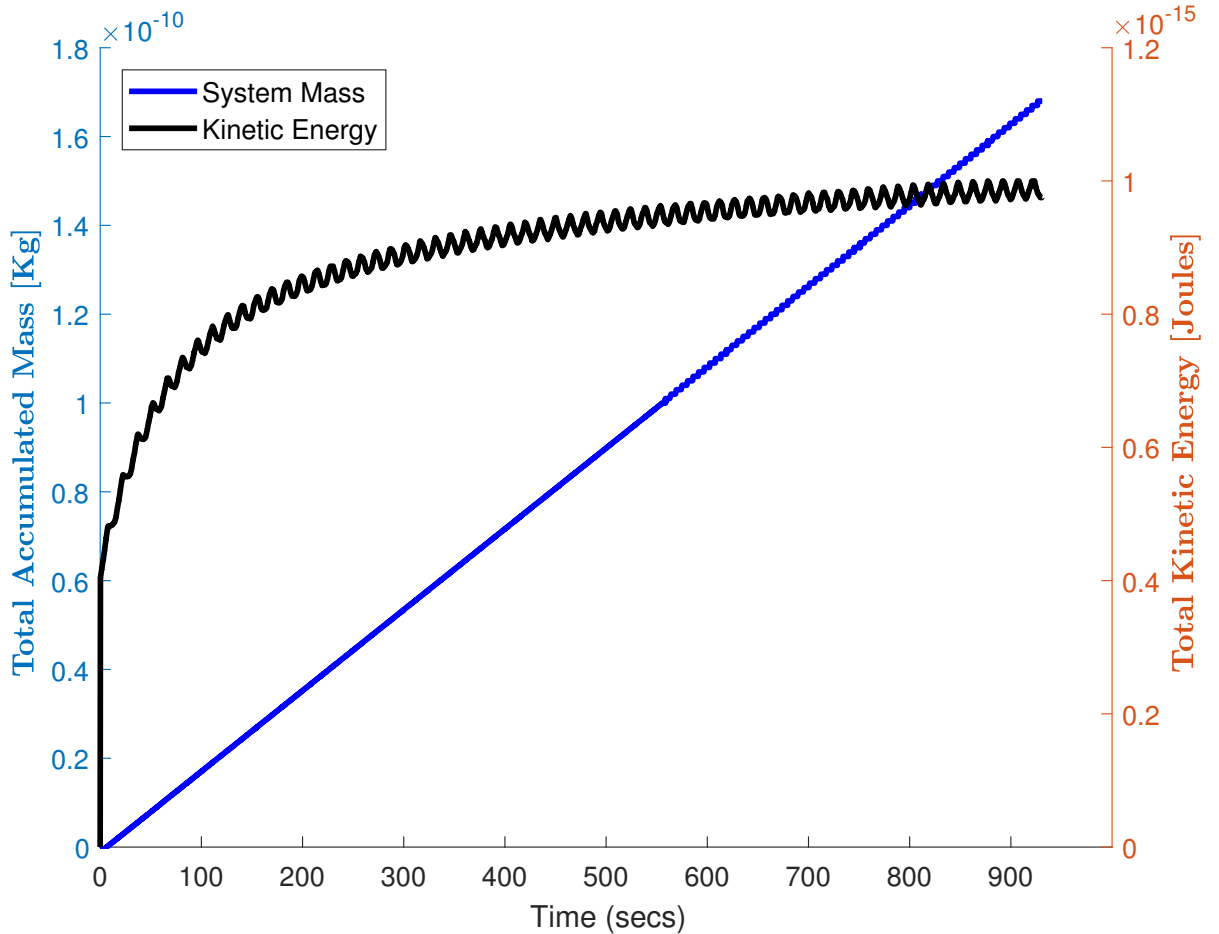


Figure 4.9 Variations of total mass (Kg) and kinetic energy (Joules) with time

The mass growth in the computational model is presented in Figure 4.9. There is a progressive growth in the mass component of the system with time. In the closed domain, there is mass accumulation with time. This evidence is confirmed with the variations of the kinetic energy KE in the system. The (K E) also grows systematically until the model reaches a stable state. It can be evidenced that the model performed accurately in conserving both mass and energy.

4.9 Conclusion

The numerical implementation of the IH2VOF computational model has been examined in this chapter. The functionalities of the model at generating different wave types, including regular, irregular waves, and solitary waves. The generation of bimodal waves will be discussed in Chapter 5. Both active and inactive absorption boundary conditions can be implemented. Waves can be both pre-determined at the boundaries using appropriately designed still water levels, or generated using dynamic piston-like paddle motions as in physical flumes. Because the model is very stable after a long period of the simulation period, it makes the model suitable for simulating a long period of bimodal storm conditions. Moreover, several wave gauges can be introduced into the domain to capture several occurrences of instantaneous velocities and water levels; it is possible to determine wave overtopping on top of the crests. More details of the computational performances of the model are presented in Chapters 5 and 9.

Chapter 5

Numerical study of Swell Effects in Bimodal Seas

An understanding of the wave height distribution of a sea state is essential in forecasting extreme wave height and lifetime fatigue predictions of marine structures. As earlier described in Chapter 3, swell can be present at different proportions and at different frequencies, while the energy content of the sea state remains unaltered. This chapter applies the numerical model (already described in Chapter 4) to investigate how the wave height distribution is affected by different swell percentages and long swell periods in an energy-conserved bimodal sea both near a wave maker and in shallow water. More specifically, wave elevations, steepness, auto-covariances, skewness, kurtosis obtained from the time variance analysis of bimodal waves are shown and discussed. Effects of swell waves on wave height distribution near the wavemaker and the structures are presented. This chapter is important because the interactions of these parameters with wave overtopping estimates will be discussed in later chapters.

5.1 Introduction

The IH2VOF computational model is applied to assess the distribution of wave heights in energy-conserved bimodal sea states described in chapter 3. There have been studies on wave height distribution in deep water and shallow water. Longuet-Higgins (1952) proposed the Rayleigh distribution of wave heights and several modifications have been made to low-wave-height-exceedance distributions (Boccotti, 2000; Forristall, 1978; Guedes Soares, 1984; Naess, 1985; Tayfun, 1981; Vinje, 1989). Specifically, a depth-modified version of the Rayleigh distribution was proposed in Battjes and

Groenendijk (2000) which is applied only to unimodal waves. Similarly, Rodriguez et al. (2002) studied the wave height probability distributions using extracted Gaussian bimodal waves from numerical simulations.

Petrova and Guedes Soares (2011) applied a linear quasi-deterministic theory to compare energies from the wind and swell seas using a simplified Sea Swell Energy Ratio (SSER) on the assumption of wave nonlinearity. Nørgaard and Andersen (2016) developed a slope dependent version of the Rayleigh distribution based on an Ursell number criterion. Nørgaard and Andersen (2016) only consider swell wave height distribution of highly non-linear swell waves due to shoaling on gentle slopes ranging from 1 in 30 to 1 in 100 slopes without considering the energy built-up in the overall bimodal spectrum. None of these studies have applied bimodal sea states that have varying proportions of swell, while at the same time containing a fixed amount of energy to investigate wave height distribution in shallow water close to a structure, which is what occurs very often in practice.

None of these studies has examined steep sloping structures in shallow water conditions that are close to a structure, a situation that is often found in practice. We focus on this critical case where the wavefield contains incoming and reflected waves which may also have been modified by breaking. The energy-conserved bimodal spectrum that was created from unimodal sea states are converted into random waves time series using the Inverse Fast Fourier Transform (IFFT).

The resulting time series was used to drive a Reynolds-Averaged Navier Stokes (RANS) computational model. Wave height values were then extracted from the model results (both away from and near to the structure) using down-crossing analysis to inspect the non-linearity imposed by wave-wave interactions and through transformations as they propagate into shallow waters near the structure.

5.2 Wave Height Extraction Using Numerical Model

A long flume is required to extract values of wave height for the assessment of the non-linearity characteristics of bimodal seas. The wave propagation model used for the present study is the validated IH2VOF model that has been fully described in Chapter 4. IH2VOF model has been validated against several hydrodynamic scenarios including plane beaches and rubble mound breakwaters (Lara et al., 2011; Ruju et al., 2014).

Surface elevations (time series) obtained from the inverse FFT were used to create sets of mesh-dependent pre-designed horizontal and vertical velocities that were used to determine the model boundary conditions. These computations were based on the

linear wave theory. The computational domain discretised into 0.005 m on x-axes (33,514 meshes) and 0.01 metres along y-axes mesh-sizes (51 meshes) were used to run the simulation with a time step of 0.004 s to maintain a stable simulation over a storm duration. As shown in Figure 5.1, wave gauges placed near both the wave-maker (W1, W2 and W3) and near the structures (W10, W11, and W12) were used to obtain the time series of the surface elevation. First three wave gauges are placed at 2 to 3 metres from the wave maker while the last three are positioned near the structure at 20 to 22 metres away from the wave maker.

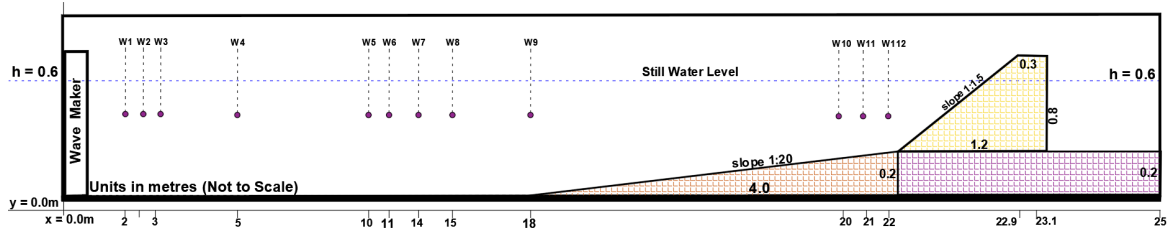


Figure 5.1 Layout of a 2-D numerical model (not to scale) for a waveheight extraction using the wave gauges near the wavemaker and the structure.

The wave conditions tested in this study are presented in Table 5.1. The same grid sizes were applied in each case. Varying percentages of swell waves were introduced into the wind sea states and at different swell periods. An equivalence of storm duration of 3 hours was propagated as represented by 1000 waves in the numerical model. The bimodal spectrum formed in this process were then converted into surface elevations which are used to drive the IH2VOF model as previously described in Chapter 4.

Table 5.1 Tested bimodal wave conditions

H_{m0} (m)	T_{pW} (secs)	T_{pS} (secs)	Swell Percentages	Gamma
1.0	7.0	15, 20, 25	25	Wind = 3.3; Swell = 2.5
1.5	8.0	15, 20, 25	50	Wind = 3.3; Swell = 2.5
2.0	9.0	15, 20, 25	75	Wind = 3.3; Swell = 2.5

5.2.1 Determination of Kurtosis and Skewness of the Bimodal Spectrum

The relationship between the kurtosis K_u and the skewness S_k of the energy conserved bimodal seas was examined using the method proposed by Marthinsen and Winterstein

(1992). A second-order approximation of the resulting bimodal spectrum S_{bim} is determined using Equation (5.1). Moreover, different orders of spectrum moments are computed from order 1 to 4 designated as m_0 , m_1 , m_2 , and m_4 using the approaches detailed in Cavannie et al. (1976).

The kurtosis, K_u , can be computed using m_3 and m_4 , and is defined as,

$$K_u = \frac{m_4}{m_3^2} \quad (5.1)$$

1st, 2nd, 3rd and 4th moments can be computed as:

$$m_n = \int_0^\infty f^n S(f) df \quad (5.2)$$

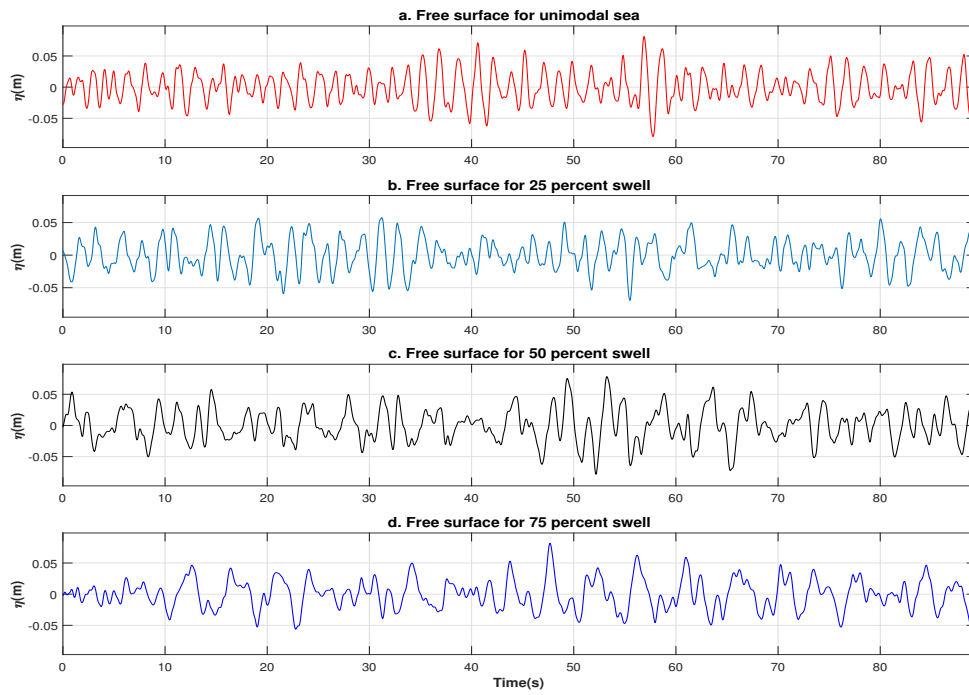
Skewness is defined according to

$$S_k = 6 \int \int f_{sum}(f_{ij}) + f_{diff}(f_{ij}) \frac{S(f_i)df_i * S(f_j)df_j}{m_0^{(3/2)}} \quad (5.3)$$

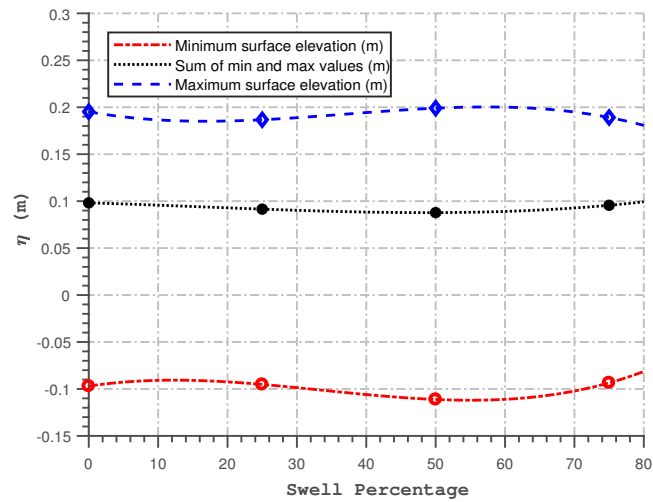
where $f_{sum}(f_{ij})$ and $f_{diff}(f_{ij})$ represent the sum and the difference of frequency effects observed from individual spectra combined.

5.2.2 Analysis of Free Surface From the Energy-Conserved Bimodal Spectrum

Examples of free surface derived from bimodal spectra are presented in Figure 5.2. The bimodal sea time series contained different swell percentages and the swell peak periods that were gradually varied as detailed in Table 5.1. As shown in Figures 5.2a, and b, there is no significant graphical noticeable differences between the surface elevation of waves time series generated by different swell percentages that was simulated for 1000 waves. This could happen because the overall energy contained in the sea states are preserved even though different random waves that was applied when obtaining the time series as presented. Energy conservation was evident in Figure 5.2b by further inspection of the minimum and maximum values of the surface elevations. The values are fluctuating around a specific mean of minimum and maximum values of -0.1 m and 0.1 m respectively for the sea state (4 m, 7 s) investigated at a scale of 1:20. These results agree with the standard weakly stationarity assumptions of a sea state with zero mean described by Rychlik et al. (1997a) because the overall wave spectra properties are preserved.



(a)



(b)

Figure 5.2 Comparison of the different free surface elevations for different swell percentages at swell peak period of 25 s. (a) Free surface elevation obtained for different swell percentages as shown; (b) Analysis of the surface elevation shown in (a).

5.3 Influence of Swell on Time Variance Analysis of Bimodal Waves

Time variance analysis of waves provides an essential background towards variability of time series. Regular and irregular characteristics of surface elevations can be identified for further analysis. Moreover, better understanding of the physical processes that is responsible for the variations can be provided. In this study, the influence of swell on the autocovariances of waves have been analysed from the extracted time series. Figures 5.3 and 5.4 presented the relationship between the auto-covariance function, (ACF), with increasing overall swell percentages and swell periods of the individual wave characteristics extracted from the computational model. The cross-correlation function in Figure 5.4 reveals that by increasing the swell percentages of a sea state, the individual waves are still largely independent of each other. There is a slight decrease in the independence as the swell percentages increase from 25 to 75 percent swell sea conditions. It can be described in terms of the variance of different swell percentages in as shown by the time lag between these waves. This effect is similar to what is observed as the swell peak period increased from 11 s to 25 secs in Figure 5.5.

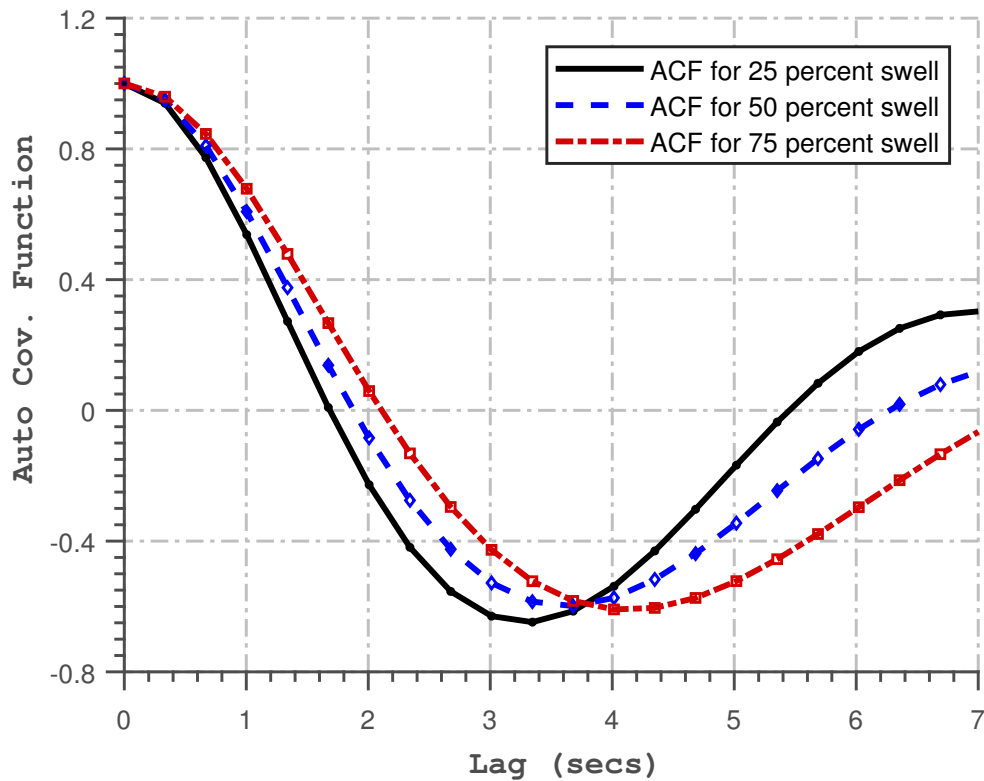


Figure 5.3 Time variance analysis of the effects swell periods and percentages in a bimodal sea with $H_{m0} = 4$ metres and $T_{pw} = 7$ secs showing the relationship between the auto-covariance function ACF with increasing overall swell percentages in a bimodal sea

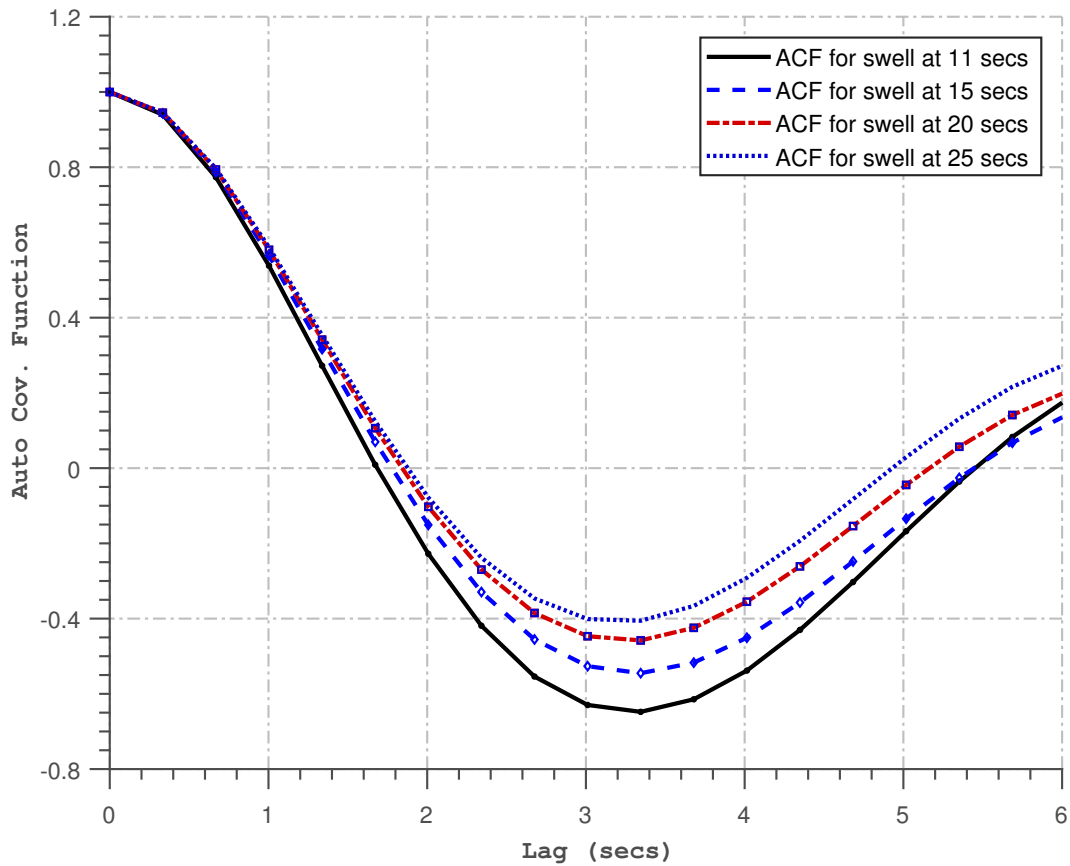


Figure 5.4 Time variance analysis of the effects swell periods and percentages in a bimodal sea with $H_{m0} = 4\text{metres}$ and $T_{pw} = 7\text{secs}$ showing the relationship between the ACF with increasing swell periods in a bimodal sea

5.4 Influence of Swell on Wave Skewness and Kurtosis

Wave skewness and kurtosis derived by applying Equations (5.1) – (5.6) presented the performance of skewness parameter with varying proportions of swell periods. The results are presented in Figures (5.5 & 5.6). As clearly shown in Figure 5.6, there is an inverse relationship between the wave kurtosis and the skewness. Skewness reduces with increasing swell peak periods. Skewness behaviour was tending towards a constant value as swell peak periods increases in the same sea state. Skewness behaviour was highest for the unimodal sea than for equivalent bimodal sea conditions; reduction of skewness was observed with increasing swell percentages.

As shown in Figure 5.5, under the conserved energy bimodal sea states, there is a dynamic reduction in the relationship between the kurtosis K_u and skewness S_k as the swell percentage and the peak period increases. This is an increasing smooth and

continuous exponential relationship that can be expressed as:

$$K_u = 2.99 \exp^{0.11S_k} \quad (5.4)$$

A sea state with 25 percent swell at a peak period of 11 s has the largest skewness-kurtosis relationship as compared with the lowest affinity relationship portrayed by the 75 percent swell at a much lower frequency of 25 s. The coefficient of kurtosis deviate significantly from linear predictions. The results obtained here near the structure are similar to those Forristall (1978) observed on a shoaling beach.

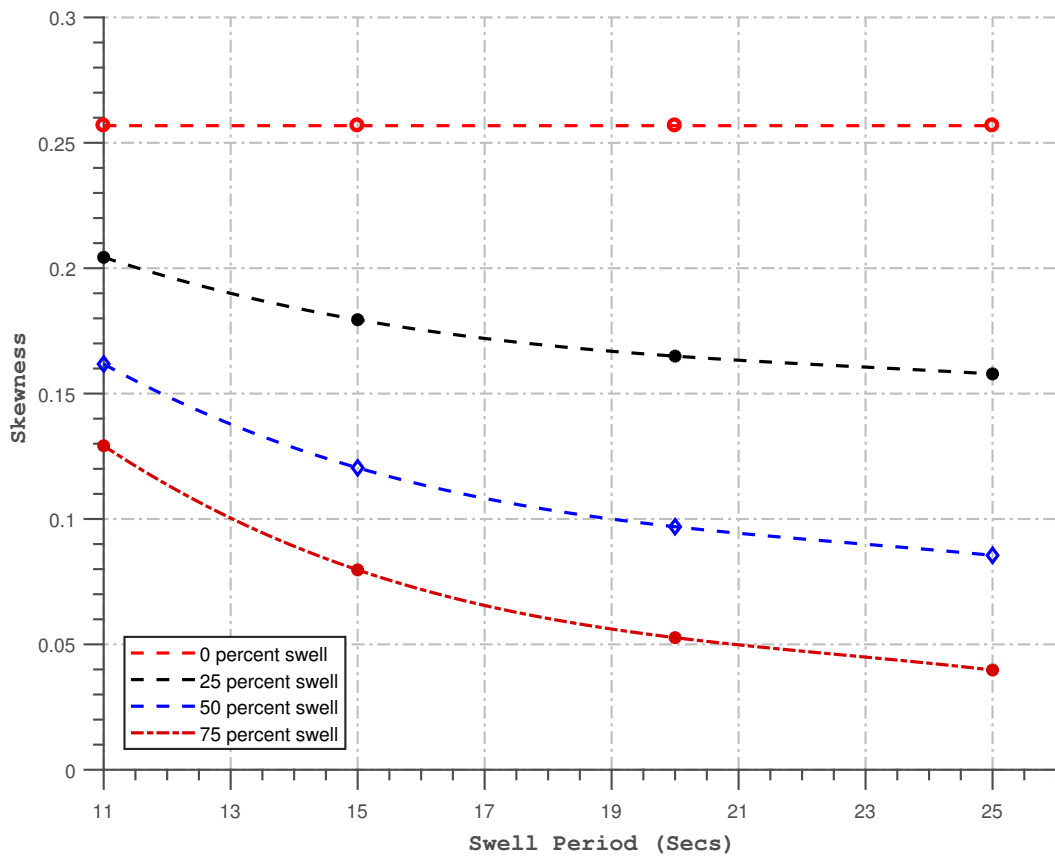


Figure 5.5 Time variance analysis of the effects swell periods and percentages in a bimodal sea with $H_{m0} = 4m$ and $T_{pW} = 7s$ showing the variation of skewness with swell periods and percentages

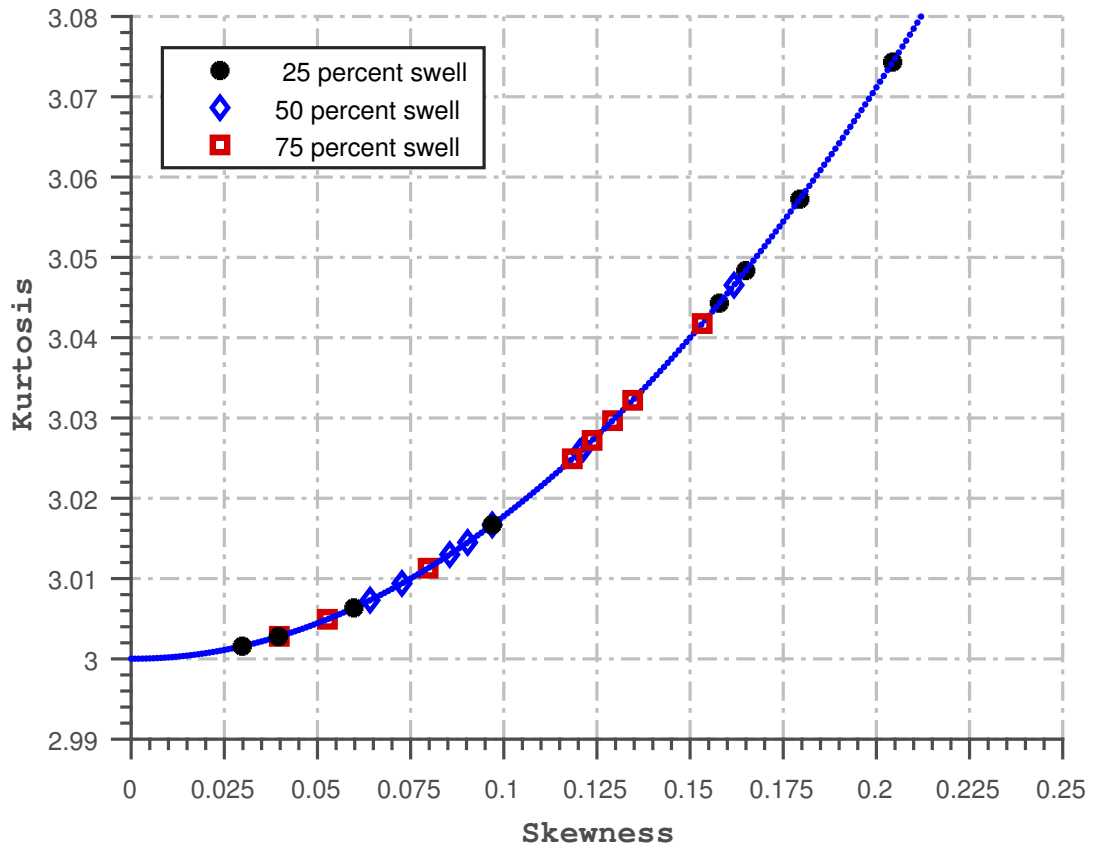


Figure 5.6 Time variance analysis of the effects of swell percentage in a bimodal sea with $H_{m0} = 4$ metres and $T_{pW} = 7$ secs showing the variations of kurtosis against skewness for different swell.

5.5 Influence of Swell on Wave Height Distribution

5.5.1 Influence of Swell Period

The probability distribution curves in Figures 5.7 and 5.8 show how wave height distribution is influenced by different frequencies of swell in the bimodal energy spectrum as observed both near (at 2 metres) from the wave maker and at farther distances away (at 20 metres) near the structure.

If the waves followed a Rayleigh distribution, all the points would fall on the full straight line shown in each figure. Wave height distribution is greatly affected by the peak period of the swell. Non-linearity patterns are indicated by deviation from the Rayleigh linear scale, and strongly related to the swell peak period. This is because the Rayleigh distribution better represents narrow-band sea states such as observed by Toffoli et al. (2006). These deviations increase with the closeness of

the two peak frequencies (for swell and wind) on the energy spectrum. Increasing the swell peak period in the energy spectrum expands the width occupied by the bimodal spectrum. This occurrence would in turn increase the overall spectra period obtainable from the resulting spectrum when transformed. In general, higher wave values deviates significantly from Rayleigh predictions. Boccotti (2000) suggests that the waves become more unstable because of the non-linearities from long crested conditions. Further observations are that the non-linearities of wave height became more pronounced near the structure because of wave-wave interactions and through shoaling as they propagate into shallow waters.

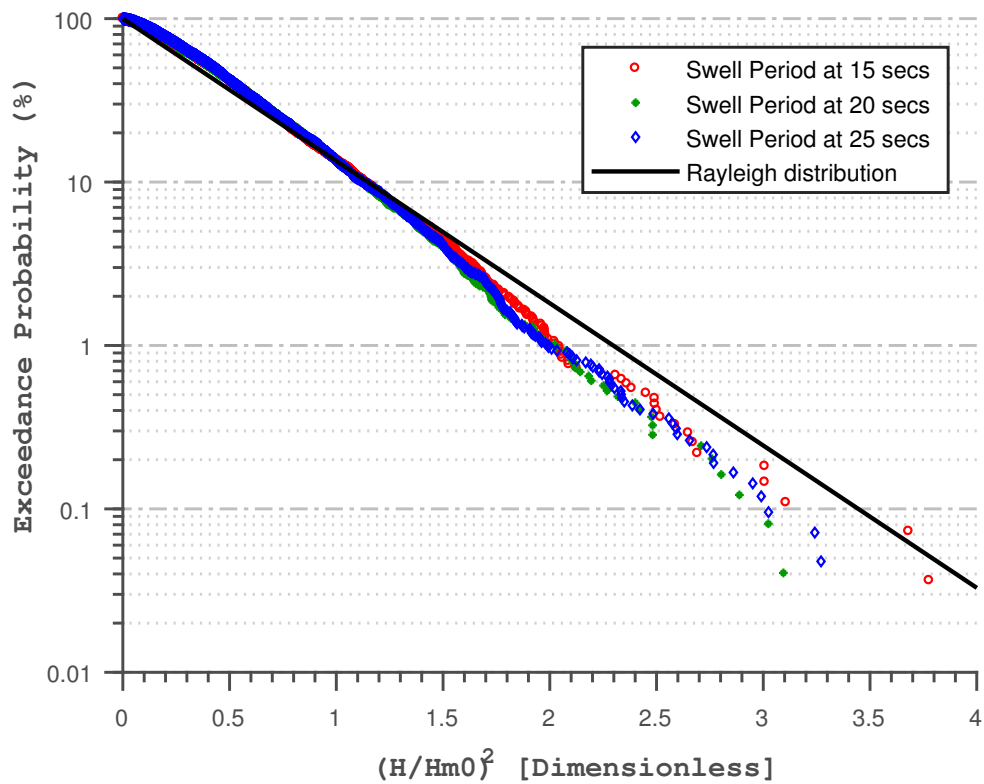


Figure 5.7 Effects of swell periods on wave height distribution near and away from the structure. Wave height distribution for different swell periods observed away from the structure

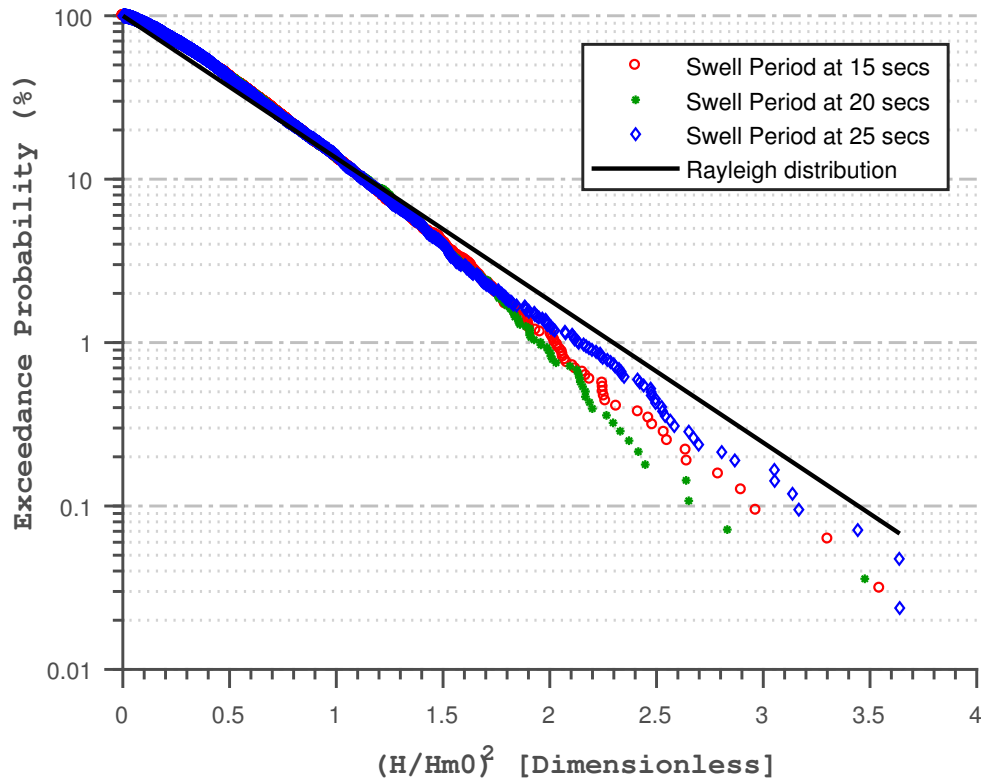


Figure 5.8 Distribution of wave height for different swell periods observed near the structure.

5.5.2 Influence of swell percentages

Following the previous section, Figure 5.9 depicts the wave height statistics within the domain for different swell percentages introduced into an energy-conserved sea state. The non-linearity patterns were extended to a unimodal sea. In all the sea cases considered (unimodal and bimodal), the non-linearity pattern was highest near the seawall than near the wave maker. This is expected because of the shoaling process and reflection created by the wave interaction with the sea-wall. As shown in Figure 5.10, non-linearity behaviour is higher in the unimodal seas state than for the bimodal seas (in this case). This trend of non-linearities reduces with increasing swell percentages. However, some deviations from this trend are observed in 50 percent swell. This may be due to the uneven distribution of energy formed by the double Jonswap spectra and may require further study.

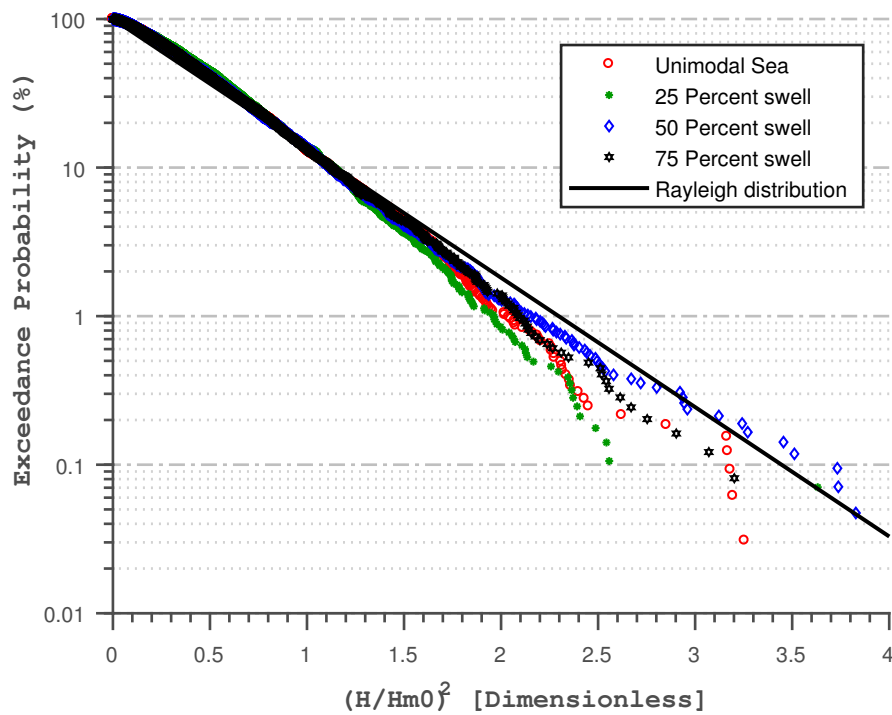


Figure 5.9 Effects of wave bi-modality caused by wave height distribution for different swell percentages observed near the structure

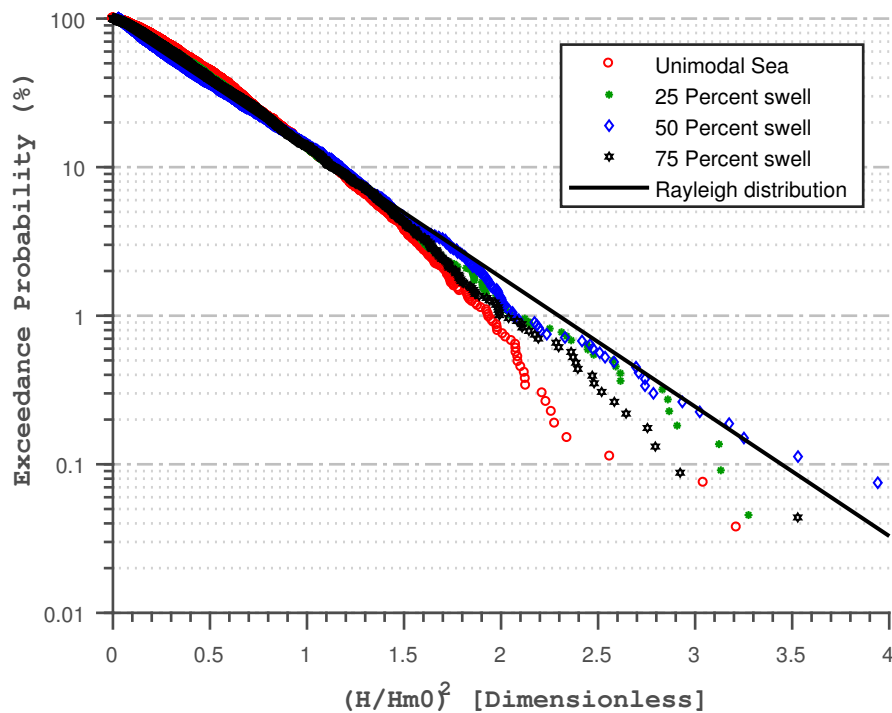


Figure 5.10 Effects of wave bi-modality caused by the distribution of wave heights for different swell percentages observed near the wave-maker.

5.5.3 Patterns of Wave Height Distribution

Figures 5.11(a–c) and 5.12(a–d) shows Weibull and the Rayleigh plots depicting the wave height distribution performance of different swell percentages in the unimodal and the bimodal sea states. In all the cases considered (both near and away the structure), wave height is better predicted by the Weibull distribution than the Rayleigh plots.

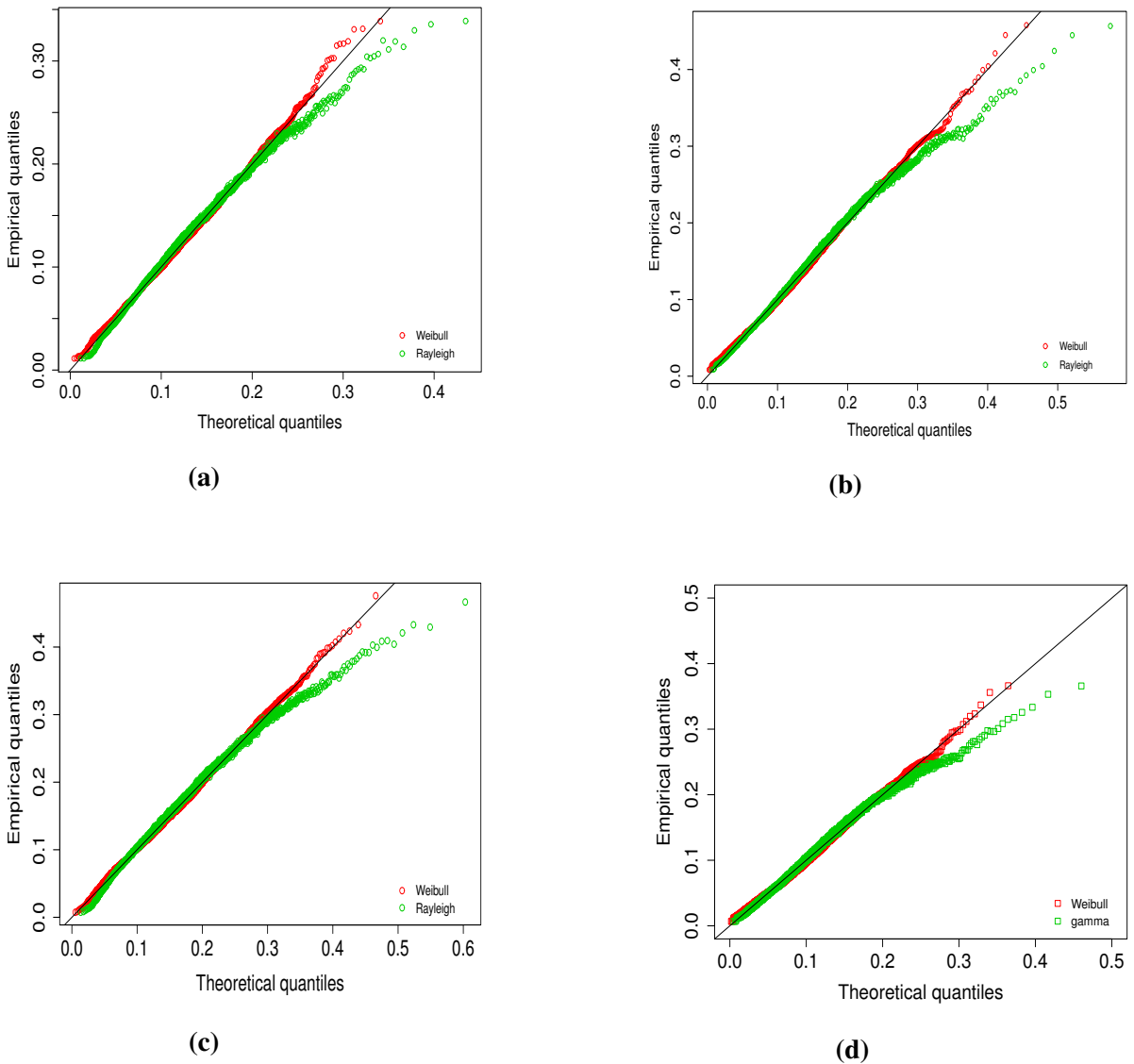


Figure 5.11 Comparisons of performance of wave height distribution using Weibull and Rayleigh plots for different swell percentages away from the structure slope for (a) 25 percent swell, (b) 50 percent swell, (c) 75 percent swell and (d) the unimodal sea

However, both distributions show great agreement at representing lower to middle part of wave height values. Some deviations from normal were evident for Rayleigh distributions in the lower quantiles.

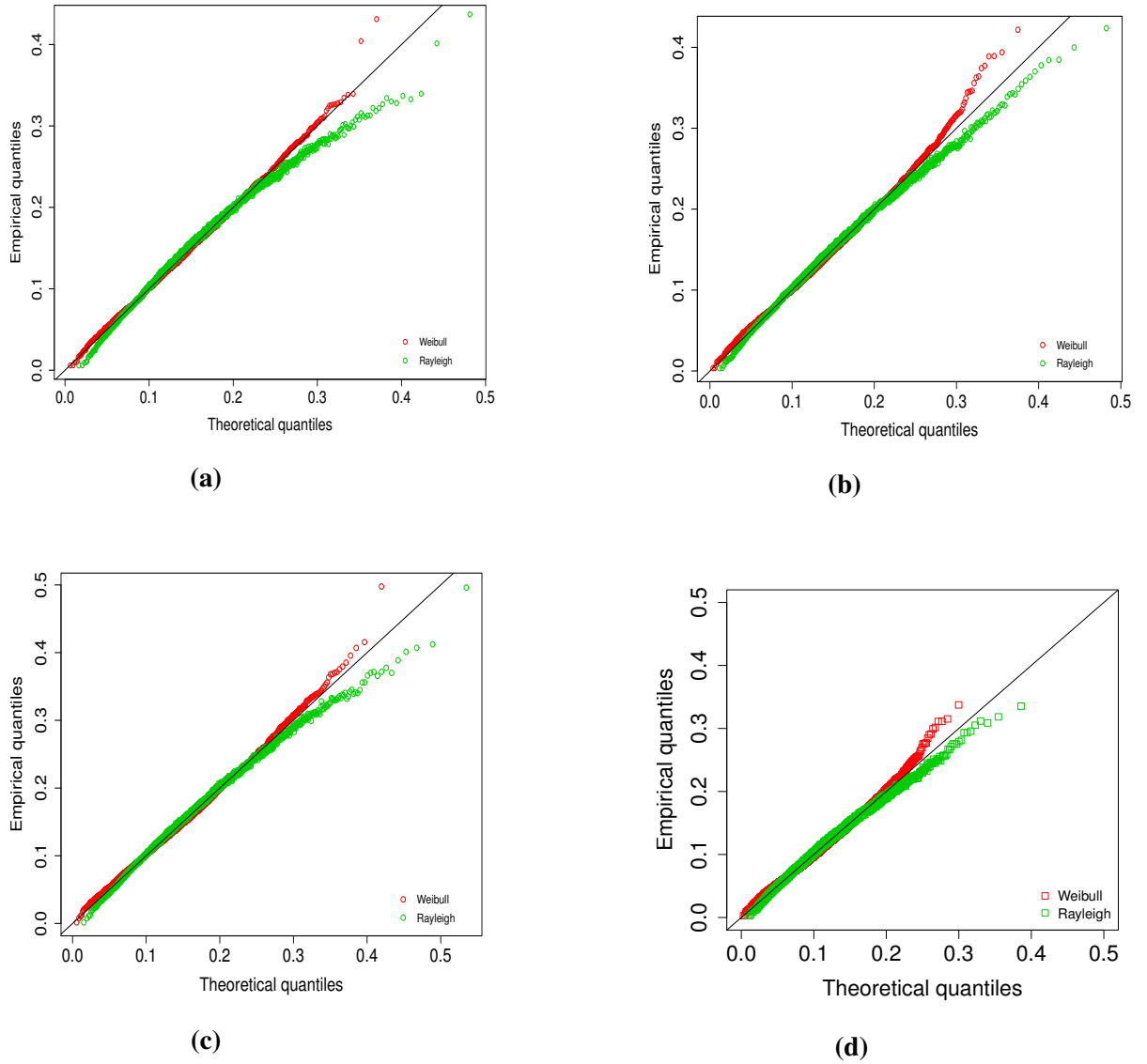


Figure 5.12 Comparisons of performance of wave height distribution using Weibull and Rayleigh plots for different swell percentages near the structure for (a) 25 percent swell, (b) 50 percent swell, (c) 75 percent swell, and (d) the unimodal sea.

The degree of fitness achieved by the weibull plots decreases slightly as the wave approaches the structure as shown in Figure 5.12(a–d) because more non-linearity patterns are possible by waves transformation caused by the wave-structure interactions. Also as expected, Rayleigh distribution underpredict higher waveheight values while the

Weibull plots are over-predicting in a way. As observed in all cases, non-linear trends are greatest in the 25 percent swell in Figure 5.12(a) and gradually reduce as the swell percentages increase towards the 75 percent swell in Figure 5.12(b–c). However as shown in Figure 5.12(d), the unimodal sea states is characterised by an earlier departure from the fit line. This is another evidence of more non-linearities in the unimodal sea states. The observed behaviour is similar to the Petrova and Guedes Soares (2014), wherein the non-linearities in the wave height distribution decreases with increasing swell percentages in the bimodal sea conditions. The energy within both initial unimodal and bimodal cases inputted into the model are identical in present study.

5.6 Conclusion

This chapter has addressed the influence of swell in a bimodal sea state with constant energy content while varying swell peak periods and percentages. Specified bimodal spectra were converted into time series using inverse FFT. The time series was used to drive the VOF-model in a numerical domain. The downcrossing analyses of numerical wave-gauge results sampled from near the wavemaker and the structure were used to assess the physical non-linearity of the bimodal sea state. There is an increasingly smooth and continuous exponential relationship between the skewness and kurtosis as the swell percentage and the peak period grows in an energy conserved bimodal sea. The extracted wave height distribution were represented by the Weibull and Rayleigh distribution to inspect the non-linearity imposed by wave-wave interactions and through dispersions as they propagate into shallow waters near the structure. The Weibull distributions produced the best fitness for wave patterns of high swell percentages considered. As the swell frequency and percentages increases in a bimodal sea, the relationship between the skewness and the kurtosis are tending towards a unimodal sea. The distribution of the waveheight on the Rayleigh model shows that the non-linearity slightly increases with the distance away from the wave maker. This could be caused by the development of other physical processes like reflection, wave breaking, and other shoaling processes arising from wave-structure interactions. In agreement with previous studies, the unimodal sea states showed the greatest non-linear behaviour however, non-linearity patterns vary as the swell grows. The Rayleigh distribution could not properly represent bimodal seas with wider spectral width and also overpredicts waveheight at low exceedances in these conditions.

Chapter 6

Experimental Test Programme

Much research guidance is available on wind-driven unimodal sea states, but fewer data can be found on bimodal seas. To bridge this research gap, the following experimental programme was conducted. Results of the 823 datasets obtained in present study will be applied to validate numerical results. This chapter is devoted to capturing new experimental datasets using incident bimodal sea. Models of three different seawalls were designed and experimented at the Swansea University Coastal Laboratory flume. The bimodal overtopping of different unimodal sea states were tested on the incorporation of swell waves at 25, 50 and 75 percentages arriving from different peak periods of 11, 15, 20, and 25 secs and propagating with 1000 random waves. Three levels of the crest freeboard R_c were tested using three different water levels of 0.60, 0.65 and 0.70 meters. Response of three different seawalls to the incoming waves were then assessed by introducing varying slopes of 1:1.5, 1:3 and vertical seawalls. In order to compare data to previous overtopping estimates, the equivalent unimodal spectra of the sea conditions were also used in each test. The quantity of water waves that over-topped the crest of the proposed breakwater were captured into a wave overtopping tank using a chute.

6.1 The Wave Flume at Swansea University

The Coastal Laboratory wave tank at Swansea University consists of an Armfield wave tank which is 30 metres in length, 0.8 metres in width and 1.2 metres in depth. Waves are generated with a HR Wallingford computer-controlled piston paddle, which has the capability to reproduce user-defined spectra of different types, includes a second-order wave correction due Schaeffer (1996), and is also equipped with an active

wave absorption system to minimise the wave reflection emerging during the wave propagation (Figure 6.1).

As a result of the configurations, the flume is particularly suitable for wave overtopping formulation and assessment studies. To represent a more realistic bathymetry for the experiment, the seabed was modified to a slope of 1:20 and the seawall was positioned to accommodate a large number of wavelengths to allow proper development of the random wave train (Figure D.1). This also gave room for a deeper water at the wavemaker than at the structure-toe. The model is constructed with stainless steel with adjustable edges to facilitate easy reformation into other geometries. A constant crest width of 0.40 metres was tested in all the experiments.

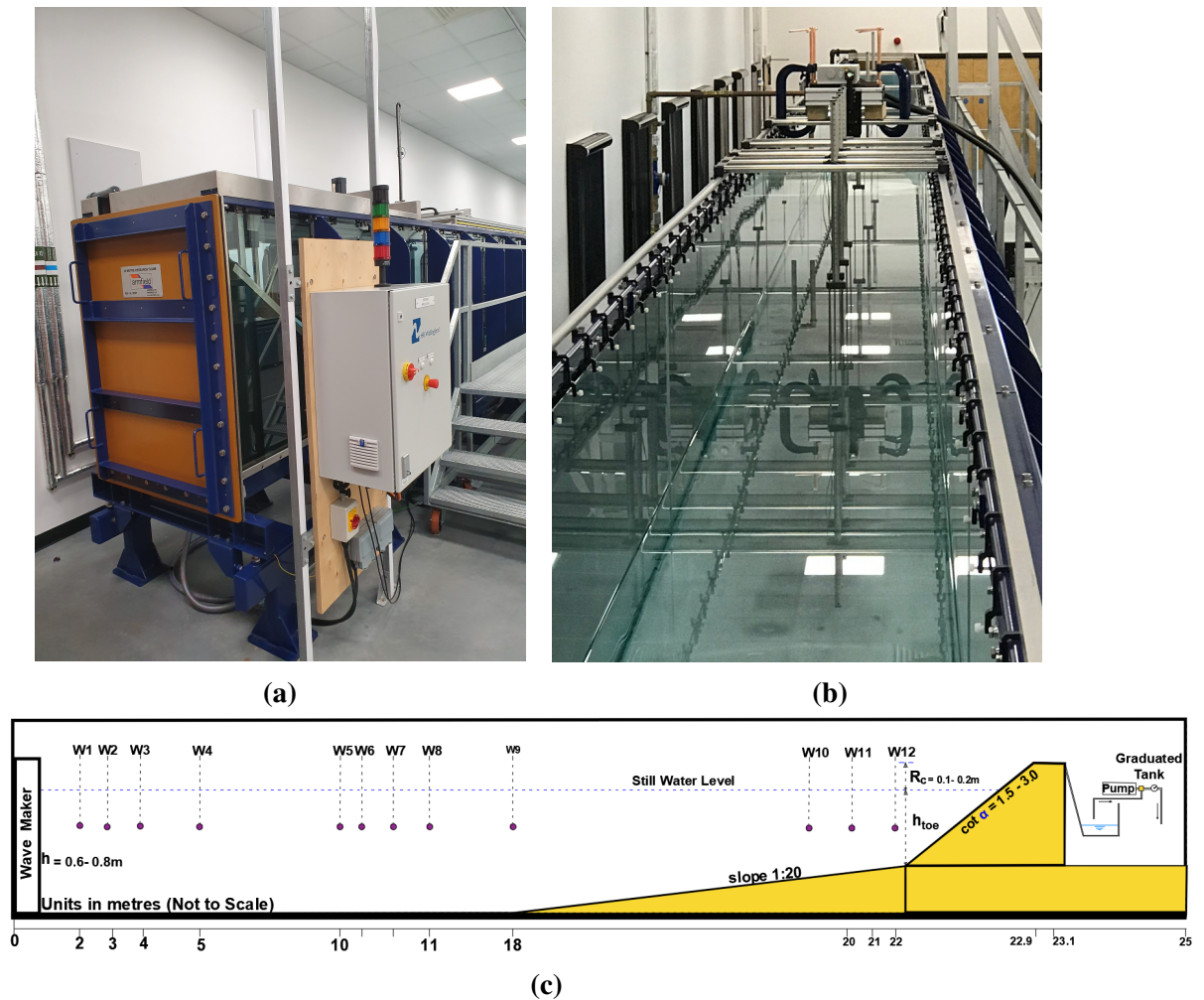


Figure 6.1 Photographs showing (a) the wave paddle system, (b) array of wave gauges, and (c) an annotated sketch of the structure set-up for the wave overtopping experiment in the Swansea University Coastal Laboratory

6.1.1 Wave Measurement Systems

The entire wave generating system is integrated into HR Daq suite Software. It comprises a range of integrated programs for use with hydraulic models. It can acquire and analyse a range of instrumentation. Accurate spectral analysis is possible within the software to generate wave spectra outputs and associated parameters. Wave probes (shown in Figures 6.3 and 6.4) are driven using the HR Daq software. Whenever the wave probes are immersed in water, flowing current between two stainless wires are converted into an output voltage to determine the surface elevation. The surface elevation is directly proportional to the output voltage. Because of variations in water conductivity due to changes in laboratory temperature conditions, periodic calibration is required for reliable results.

HR Daq systems are equipped with the calibration routines to determine accurate relationships between observed voltages and the water level. For these set of experiments, calibrations of wave probes were performed mostly twice a day or whenever there was a noticeable drop in temperatures within the laboratory.

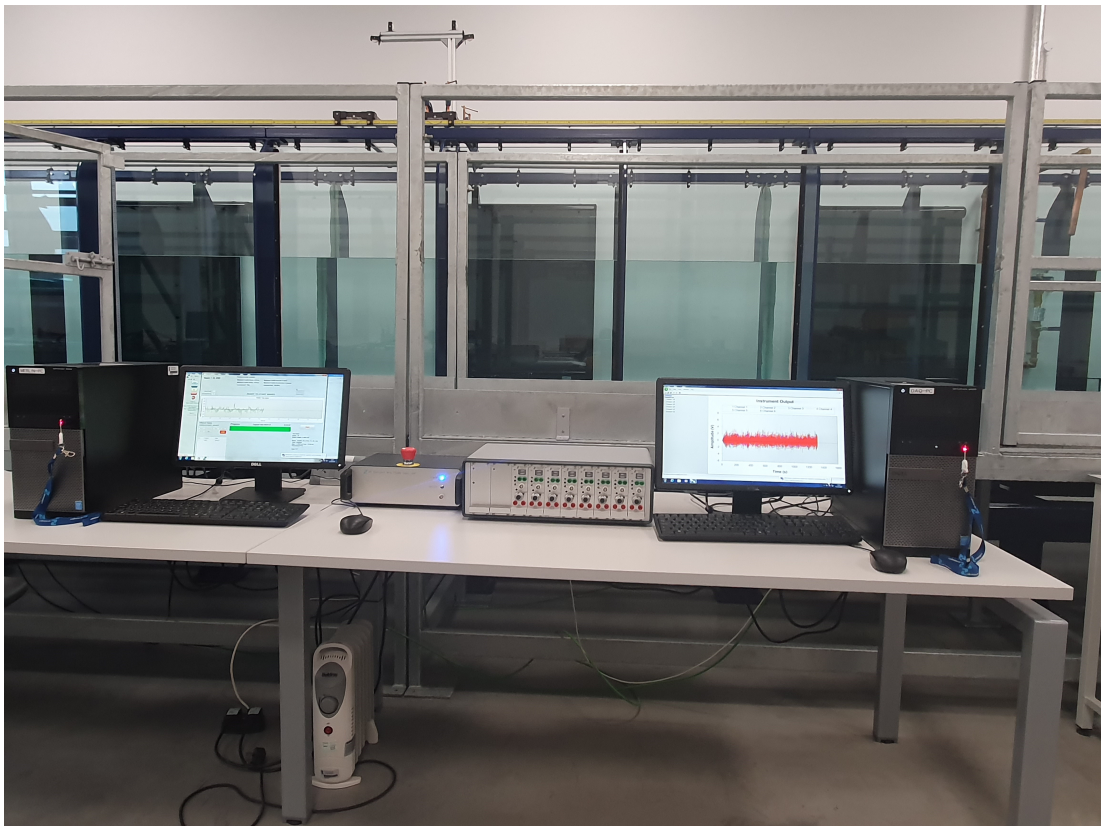


Figure 6.2 A photograph of the HR Daq system integration of various wave generation and capturing systems



Figure 6.3 Connection pattern of the HR Wallingford wave probe system

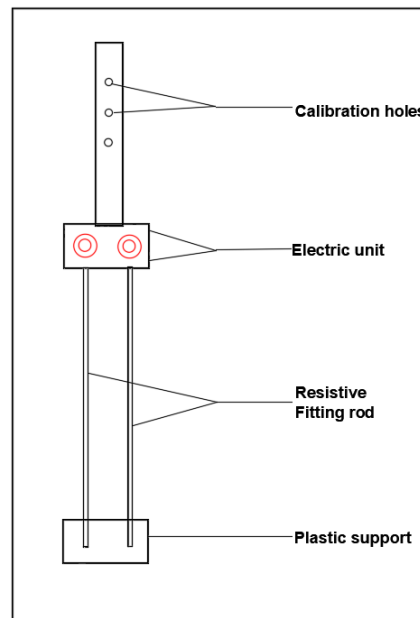


Figure 6.4 A sketch of HR Wallingford wave probe system for measuring surface elevation

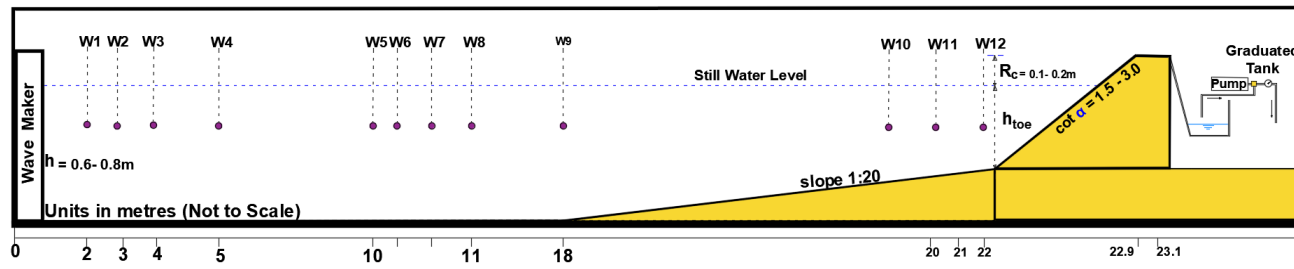
Apart from real time acquisition of wave data by the Daq system, a module is also provided for reflection analysis. In this module, a minimum of four well positioned wave gauges are required.

6.2 Experimental Procedure

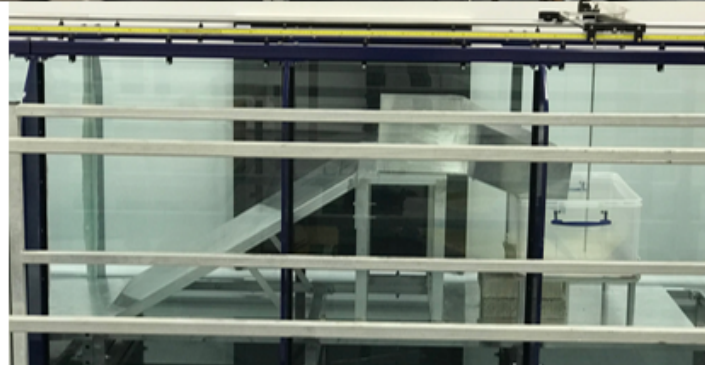
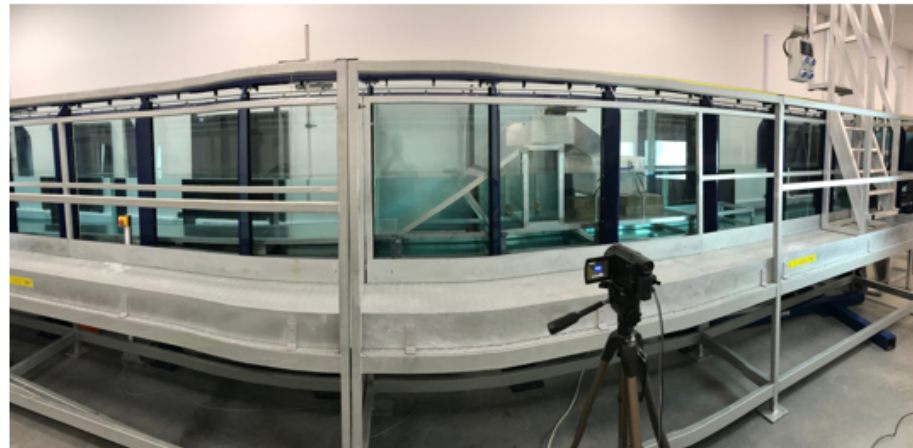
The bimodal spectra, (discussed in chapter 3), were input into the wave generator in a user-defined format to generate bimodal waves in random mode. As recommended in previous studies, a minimum of 1000 waves were tested for each wave condition. Overtopping volumes were collected using a chute fitted directly on top of the structures placed at the rear end of the crest (Figure 6.5). The chute emptied the water into the overtopping collection tank using a simplified pumping method that has been previously adopted (see, for example, Kortenhaus et al., 2005; Troch et al., 2004). The overtopping volume was obtained by measuring the water level in the calibrated tank before and after each test. A direct volumetric measurement in a calibrated tank was applied to minimise uncertainties due to multiple conversions.

In order to obtain the dimensionless wave overtopping, H_{m0toe} is required. It is determined by applying the bulk reflection coefficient obtained from the reflection analysis. The reflection analysis of acquired signals was performed using the reflection module incorporated with the HR-Daq data acquisition and processing software. This package removes reflected waves from the total signals using the method Zelt and Skjelbreia (1992) with four wave gauges. The method is an extension of the three-wave gauges least-squares solution of reflection analysis, which was first introduced by Mansard and Funke (1980). In this experiment, the reflection analysis of wave signals was analysed using four wave gauges that were placed at the flat, and deeper end of the flume. These wave gauges are WG5, WG6, WG7, and WG8 (Figure 6.5a).

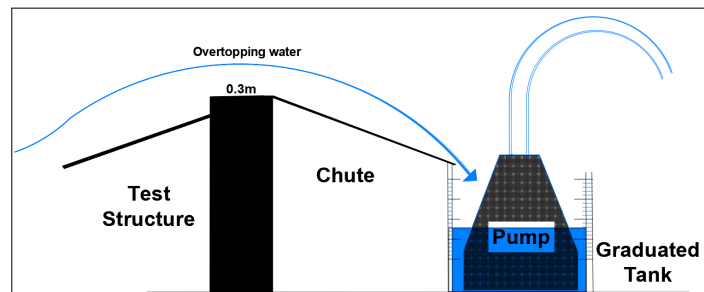
Separation distances between the gauges were computed using offshore based linear wavelength $L_{m-1,0}$ of the incident waves. This can be expressed as $X_1 = 0$, $X_{12} = L/10$, $X_{13} = L/4$, $X_{14} = L/3$ as shown in Figure 6.5a. During these tests, the wave gauges were fixed at distances 10, 10.33, 10.83 and 11.1 metres from the wave maker respectively. These distances satisfy the normalised denominator criteria of the HR-Daq reflection analysis module (Figure 6.6a). The wave signals were analysed using Fast Fourier Transform (FFT) producing amplitude components in the frequency domain. Some portions of wave elevation at earlier and later parts of each simulation were ignored to allow for consistency of wave elevations. Bandpass filtering was applied to isolate the frequency band of $0.33f_p \leq f_p \leq 3f_p$ (f_p is the peak frequency of the spectrum). Three levels of the crest freeboard were tested at 0.1, 0.15 and 0.20 metres respectively for seawalls with slopes of 1:1.5, 1:3 and vertical seawalls.



(a)

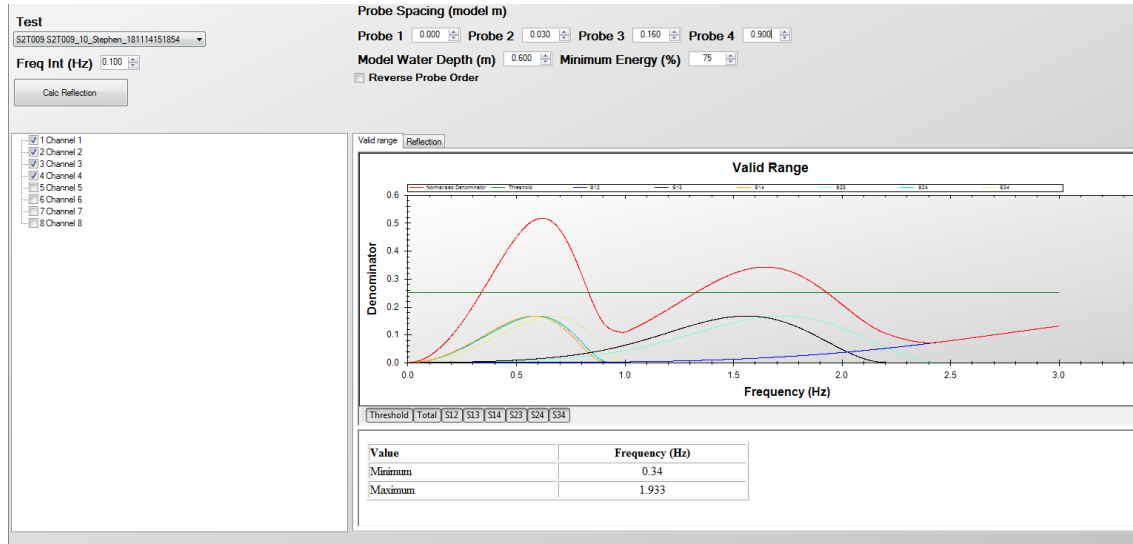


(b)

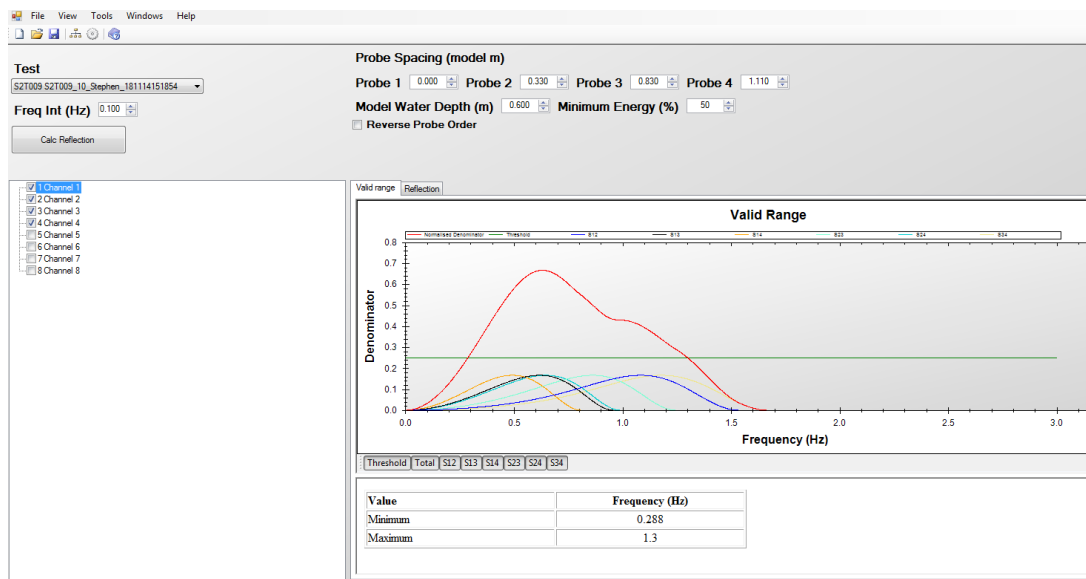


(c)

Figure 6.5 (a) Sketch of the model set-up, and (b) Photographs showing the panoramic view of the flume depicting a section the seaward slope, and (c) the schematic sketch of the overtopping tank.



(a)



(b)

Figure 6.6 A snapshot of the HR-Daq system showing the normalised denominator for reflection analysis using different wave gauge positions with (a) poor wave gauge position, and (b) suitably calibrated wave gauge position based on the frequency band of $0.33f_p \leq f_p \leq 3f_p$

6.2.1 Tested Wave Conditions

Full details of the methodology, and how the spectra were defined while conserving total wave energy were discussed in Chapter 3 of this thesis. They are also available in Orimoloye et al. (2019b).

Table 6.1 Generation of 13 different wave conditions from the same-energy sea state with H_{m0} of 4 metres and $T_{pW} = 7$ secs.

ID	Test-No	T_{pW} (secs)	T_{pS} (secs)	Swell (%)	Mode
					(Unimodal = 1 Bimodal = 2)
1	T001-1	7	11	25	2
2	T001-2	7	15	25	2
3	T001-3	7	20	25	2
4	T001-4	7	25	25	2
5	T001-5	7	11	50	2
6	T001-6	7	15	50	2
7	T001-7	7	20	50	2
8	T001-8	7	25	50	2
9	T001-9	7	11	75	2
10	T001-10	7	15	75	2
11	T001-11	7	20	75	2
12	T001-12	7	25	75	2
13	T001-13	7	N/A	0	1

Different bimodal spectra with the same energy were then developed by introducing different swell components of different percentages at different frequencies, as shown in Figures 3.6 & 3.8. Table 6.1 contains the ranges of parameters tested in this study. Different spectral components H_{m0} and ($T_{Wm-1,0}$, and $T_{Sm-1,0}$) each for wind, and, swell seas were selected within the ranges shown in the Table. The spectral wave height H_{m0} obtained from the bimodal spectrum were varied from 0.05 to 0.18 metres. In this case, the spectral peak period of the wind-wave $T_{Wm-1,0}$ (varying from 1.106 to 1.581 secs) were combined with different corresponding spectral peak period $T_{Sm-1,0}$ of the swell wave. Only three swell percentages (25, 50 and 75) were investigated to accommodate more swell condition. Also, four different swell peak periods of 1.739, 2.372, 3.162 and 3.953 secs at a model scale of 1:40, which are equivalent to 11, 15, 20 and 25 seconds at the prototype scale were considered. These were tested across three

crest freeboards as stated in Table 6.1. The crest width of the structure are fixed at 0.3 metres (Figure 6.5c). Therefore, there are 13 different spectra, (including the unimodal case). As seen from Table 6.1, a broad range of $S_{m-1,0}$ (from 0.01—0.07) were achieved from current experiments. It gives room for investigating a large number of bimodal wave cases.

Table 6.2 Ranges of Parameters tested.

Parameter (unit)	Range
Spectral wave height H_{m0} (m)	0.05, 0.075, 0.10, 0.125
Crest freeboard R_c (m)	0.1, 0.15, 0.2
Water depth at wave maker h (m)	0.60, 0.65, 0.70
Water depth at structure toe h_{toe} (m)	0.49, 0.54, 0.59
Spectral wind-wave peak period $T_{Wm-1,0}$ (s)	1.106, 1.265, 1.423, 1.581
Spectral swell-wave peak period $T_{Sm-1,0}$ (s)	1.739, 2.371, 3.1622, 3.952
Swell percentage S_w	0, 25, 50, 75
Surf similarity Parameter $\xi_{m-1,0}$	2.4 - 5.5
Non-dimensional freeboard R_c/H_{m0}	0.8 - 4
Wave steepness $S_{m-1,0}$	0.8 - 4.0
$\cot \alpha$	0, 1.5, 3.0
Slope angle	90, 33.70 and 18.43
Total successful tests	823

Implementing different percentages and periods of swell in the sea states yields different spectra and time series with the same energy content. Details of varying wind wave spectra that were combined with their equivalent swell periods and percentages are presented Table 6.2. Datasets in Table 6.2 are based on a Froude's scaling of 1:40 for all test cases. These tests are chosen based on the bimodal characteristics of UK coastline.

Table 6.3 Combinations of wave conditions tested in the present study.

Test No	H_{m0} (m)	T_{pW} (secs)	T_{pS} (secs)	$h(m)$	$cot\alpha$
T001	0.125	1.107	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T002	0.125	1.264	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T003	0.125	1.423	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T004	0.125	1.581	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T005	0.1	1.107	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T006	0.125	1.264	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T007	0.1	1.423	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T008	0.125	1.581	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T009	0.1	1.107	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T010	0.1	1.264	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T011	0.125	1.423	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T012	0.125	1.581	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T013	0.075	1.107	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T014	0.075	1.264	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T015	0.075	1.423	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T016	0.075	1.581	1.739, 2.3712, 3.162, 3.953	0.65	0, 1.5, 3.0
T017	0.1	1.107	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T018	0.1	1.264	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T019	0.1	1.423	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T020	0.1	1.581	1.739, 2.3712, 3.162, 3.953	0.7	0, 1.5, 3.0
T021	0.075	1.107	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T022	0.075	1.264	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T023	0.075	1.423	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0
T024	0.075	1.581	1.739, 2.3712, 3.162, 3.953	0.6	0, 1.5, 3.0

6.3 Overview of Experimental Tests

To ensure that appropriate formulations are applied in analysing observed results, it is important to examine the data range of the captured datasets perceptually. Figure 6.7 presents a general representation of the relative wave height H_{m0}/h and the non-dimensional freeboard R_c/H_{m0} of the successful tests across the three slopes $cot\alpha = 0.0$, $cot\alpha = 1.5$, and $cot\alpha = 3.0$. H_{m0} was determined from the spectral conversion of the wave surface elevation at the toe of the structure. Figure 6.7, it can be seen at a glance

that the majority of the observations fall within the deep-water range $H_{m0}/h \leq 0.2$ (as suggested in the EurOtop (2018) manual). It is this ratio that describes the shallow water effects expected in the results afterwards. It also affirms that wave breaking effects would not be considered in later analyses of these test cases.

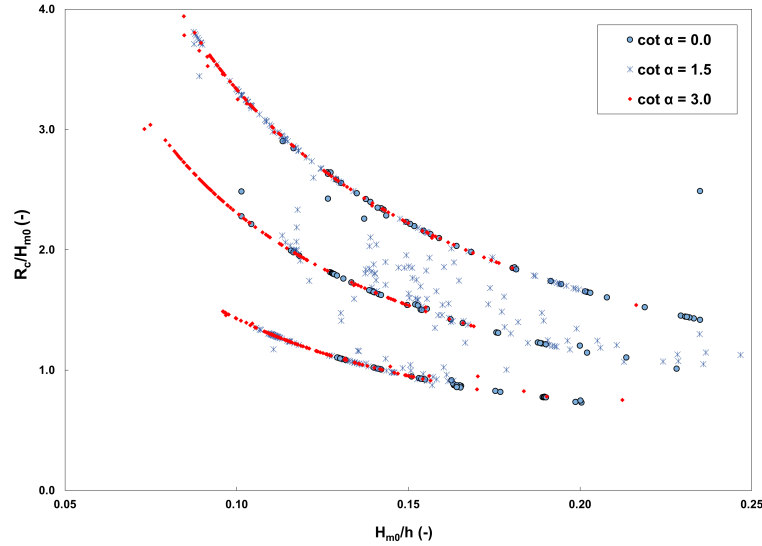


Figure 6.7 An overview of comparison of the relative wave height H_{m0}/h with the non-dimensional freeboard R_c/H_{m0} of successful tests across the three slopes $\cot \alpha = 0.0$, $\cot \alpha = 1.5$, and $\cot \alpha = 3.0$

6.3.1 Spectra Validation

It is crucial to examine how well the wave maker can reproduce the input bimodal spectrum. The theoretical spectra were validated in this section by comparing the results of the target wave conditions with the recorded output by wave gauges. Figure 6.8 shows example of comparisons between the target (red line) and the reproduced (black lines with markers) incident bimodal spectra under laboratory conditions. The result is presented for all the gauges before the structure toe after reflection analysis has been carried out. As shown in Figure 6.8, a good overall agreement is obtained for different swell content. The frequencies of the modal values are well captured, but the spectral values are under-predicted, an outcome of the limitations imposed by scaling. It is challenging to generate long period bimodal waves in the laboratory. Consequently, the wave periods have to be scaled down significantly for the wave paddle motion.

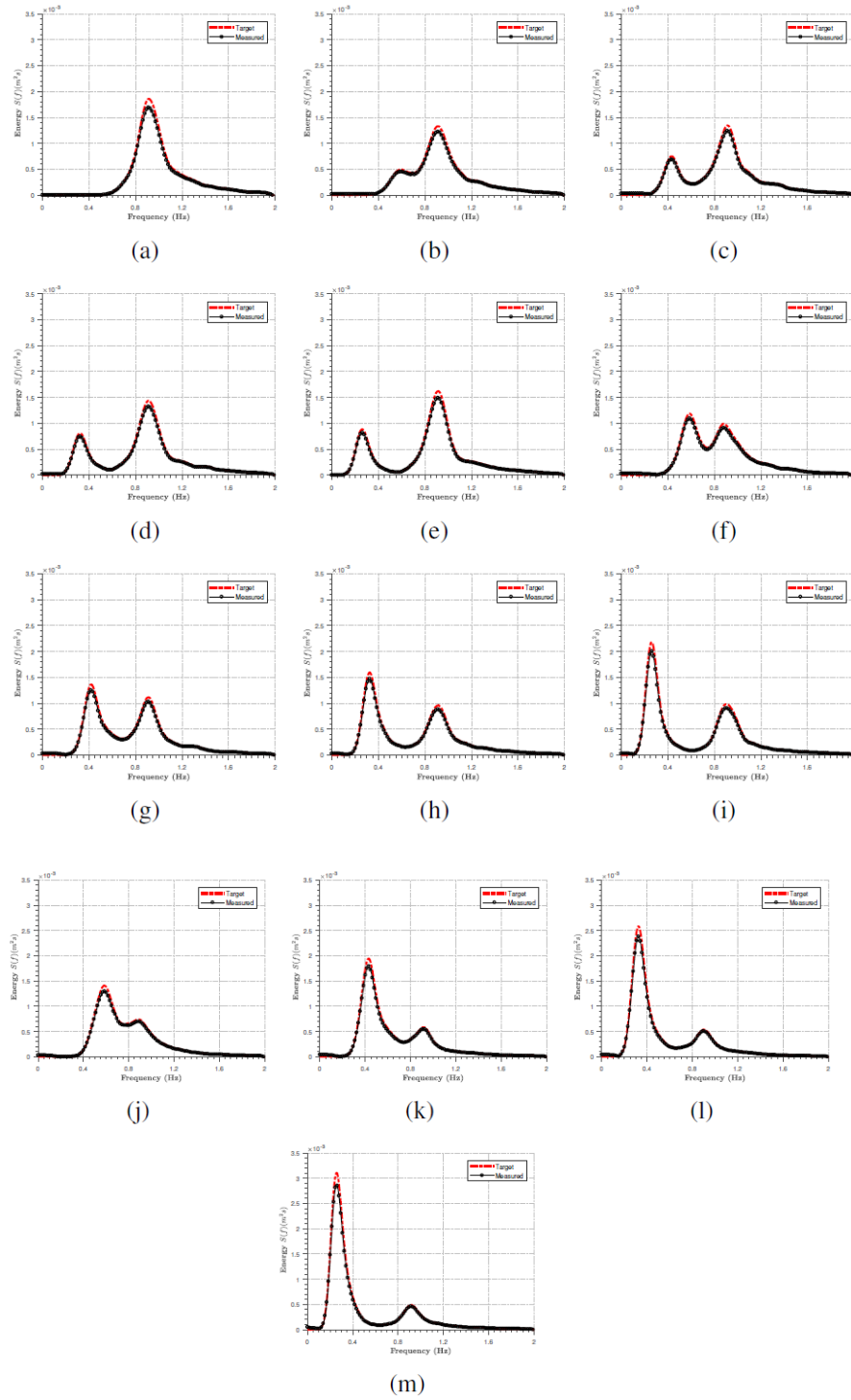


Figure 6.8 Comparison between the target bimodal spectrum (in dashed lines) and measured (with markers) for (a) unimodal wind-sea with 0 percent swell, (b – e) 25 percent swell, (f – i) 50 percent swell, (j – m) 75 percent swell obtained for a sea state with wave height of 4 metres and wind wave period of 7 secs with equal energy but different swell percentages and swell peak periods

6.3.2 Repeatability of Experiments

Repeatability of measurements is essential in experimental studies to ensure the acceptability of results. In this way, uncertainties of measurements can be reduced and in turn, enhance the reliability of outputs. Kortenhaus et al. (2005) suggested the use of the coefficient of variation (*CV*) to assess the variability of wave overtopping parameters. The following parameter σ' are proposed as the *CV* with the assumption that the individually observed datasets are each normally distributed around a mean μ , and standard deviation σ :

$$\sigma' = \frac{\sigma}{\mu} \quad (6.1)$$

Equation (6.1) has also been regarded as the Relative Standard Deviation (*RSD*) of each repeated tests in terms of the nominal test. Because a minimum of 1000 random waves are generated in this experiment, it is essential to inspect the repeatability of the tests. It can be achieved by examining the wave parameters and the overall wave overtopping captured during the test. Based on FFT, from the inputted wave spectra, an infinite number of random wave elevations can be produced from these sets. Whenever waves are generated with the HR Daq software, a random seeding number is generated and assigned to the collection of observation. The number keeps track of all different wave sequences generated.

Table 6.4 Repeated observations captured at different wave gauge positions for a wave case with $H_{m0} = 0.125$ metres

Location (m)	Set 1 (m)	Set 2 (m)	Set 3 (m)
0	0.128	0.127	0.128
3.075	0.128	0.127	0.128
6.15	0.128	0.128	0.128
9.225	0.128	0.127	0.128
12.3	0.128	0.128	0.127
15.375	0.128	0.127	0.127
18.45	0.128	0.128	0.129
21.525	0.14	0.139	0.135
24.6	0	0	0

For a repeatable wave sequence, the seeding number must be applied while specifying waves. Inputting another seeding number will create a new set of wave sequences even if the spectral characteristics of the waves are the same. The same random numbers are maintained while conducting all these experiments.

During this study, using each of the slopes, three random repeated tests were carried out to ascertain the repeatability of random waves within the system across three slopes. Tables 6.3 and 6.4 shows example of the repeated wave height captured at different wave gauges across the flume. Recorded wave height values are precise and accurate.

Table 6.5 An example of the coefficient of variations obtained for overtopping and H_{m0} for the repeated observations captured for different sets of observations

Parameter Measured	Set 1 σ' (%)	Set 2 σ' (%)	Set 3 σ' (%)
$q(l/s/m)$	22.3	24.1	20.6
H_{m0}	3.28	3.22	2.14

Repeatability of wave height evolution across different wave gauges in the domain are represented. This is presented in Figure 6.9. H_{m0} of 0.125 metres were very well described across the wave gauges in the deep-water zones. Moreover as shown in the Figure, patterns of wave shoaling across the and the wave breaking are repeatedly well represented by the three sets of observation.

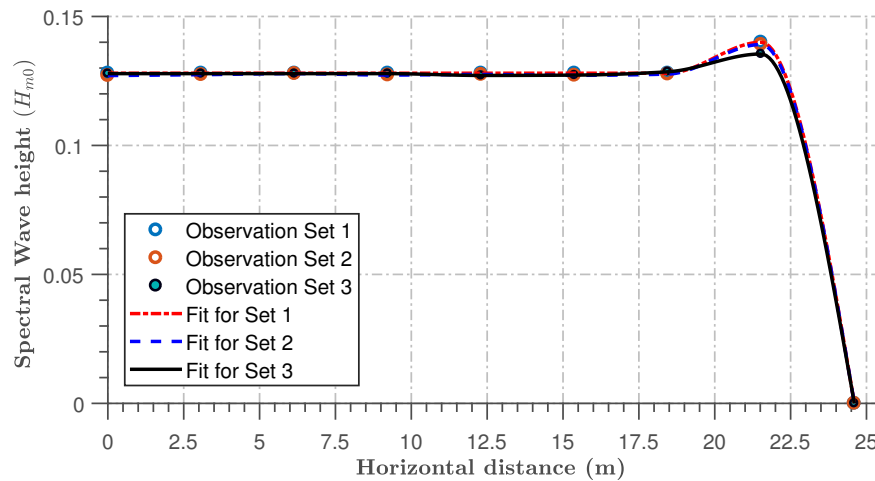


Figure 6.9 Performance of repeated wave heights captured at different wave gauge positions during the experiment for $H_{m0} = 0.125$ metres with $\cot \alpha = 3.0$, and 75 % swell

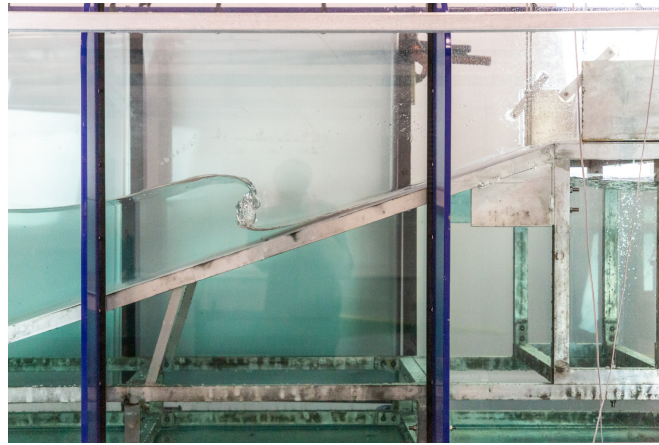
Table 6.6 Summary of the slope angle $\cot \alpha$, observation average μ , standard deviation σ and CV σ' for the same sets of observations with $H_{m0} = 0.125$ metres and 75% swell

Slope Angle	Parameter Measured	μ	σ	σ' (%)
$\cot \alpha = 3.0$	$q(l/s/m)$	0.000832	0.000041	4.93
	H_{m0}	0.121	0.0066	1.82
	$\frac{q}{\sqrt{gH_{m0}^3}}$	0.00631	0.0613	9.71
$\cot \alpha = 1.5$	$q(l/s/m)$	0.000562	0.000033	5.87
	H_{m0}	0.116	0.001	0.86
	$\frac{q}{\sqrt{gH_{m0}^3}}$	0.0045	0.0143	3.15
$\cot \alpha = 0.0$	$q(l/s/m)$	0.000172	0.000019	11.05
	H_{m0}	0.112	0.0005	0.45
	$\frac{q}{\sqrt{gH_{m0}^3}}$	0.00147	0.0096	6.55

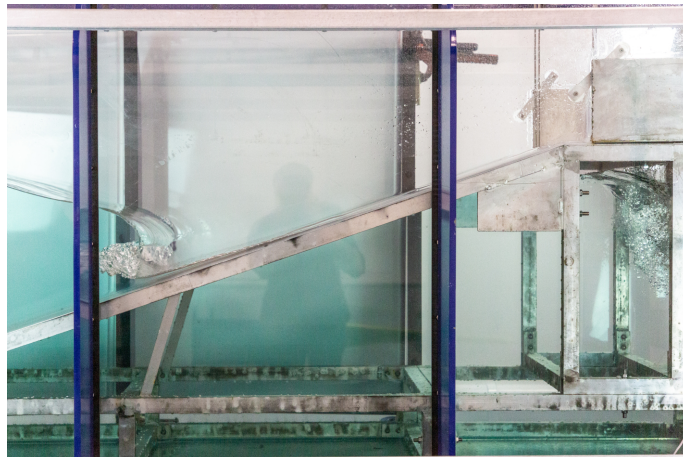
Full summaries of repeatability tests conducted in this study are shown in Table 6.6. Three sets of data were repeatedly observed across the three slopes for $H_{m0} = 0.125$ metres and 75% swell. It can be observed that there is a minimal variation observed between different sets of the CV σ' obtained across slopes. As shown in Table 6.6, it is expected that the CV σ' values measured for the overtopping should be higher than for observed H_{m0} . It is because there are more uncertainties associated with the overtopping estimation than the H_{m0} derived from the wave elevations recorded by wave gauges. Lesser values of variations are observed in the present study than ones recorded by Kortenhaus et al. (2005).

6.3.3 Wave Breaking and Overtopping Events

In this section, examples of snapshots of waves interacting with coastal structures, obtained from this overtopping experiment are presented (Figures 6.10, 6.11, & 6.12). These figures demonstrate different wave breaking patterns. In Figure 6.12(b), wave breaking accompanied a wave overtopping event. Conversely, photographs in Figures 6.10a and 6.10b, show the seawall dissipating the wave energy, and so there is not sufficient energy for wave overtopping to occur. The wave propagating pattern shown in the figure is hypothetical of a unimodal sea states with less energy. As the wave energy builds up over time, the flow becomes more turbulent. This in turn yields progressively higher wave surges, overtopping the seawall (Figure 6.12).

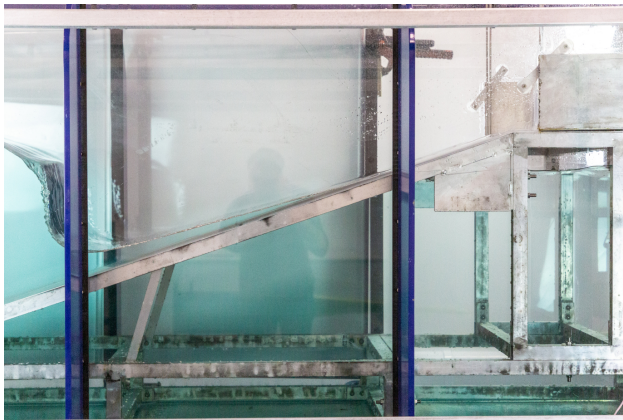


(a)

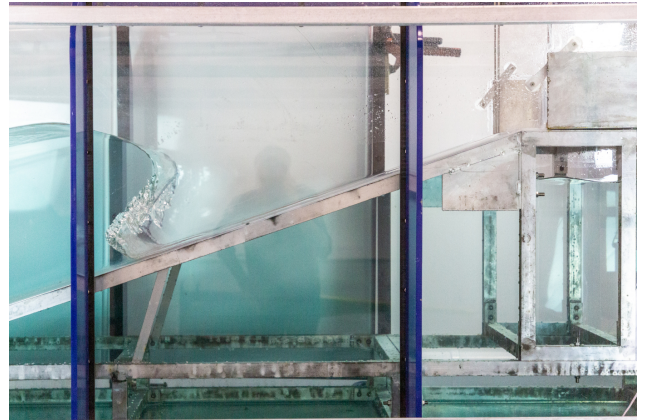


(b)

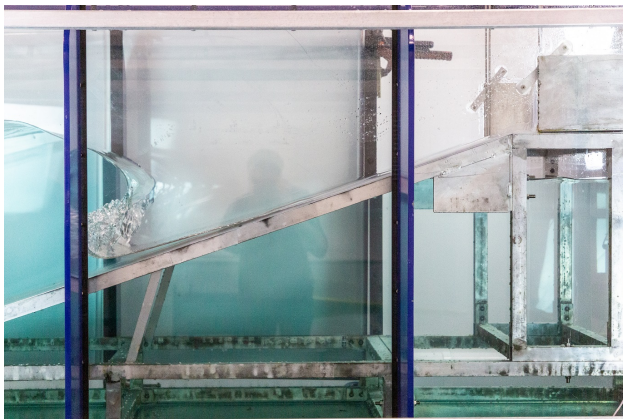
Figure 6.10 Snapshots from the experimental test showing: some wave breaking events at the seawall without wave overtopping by a wave with $H_{m0} = 0.1$ metres at a water depth 0.5 metres approaching a slope with $\cot \alpha = 3.0$



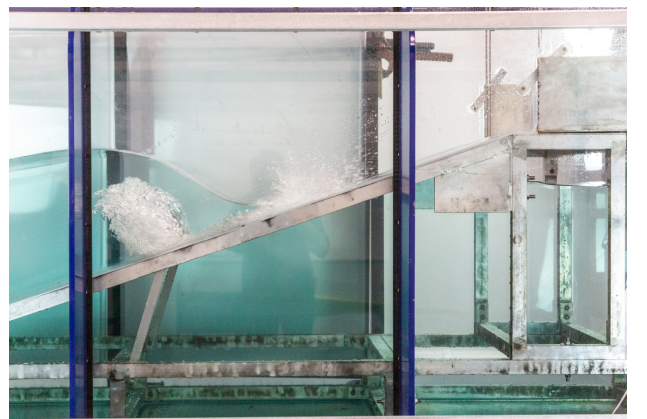
(a)



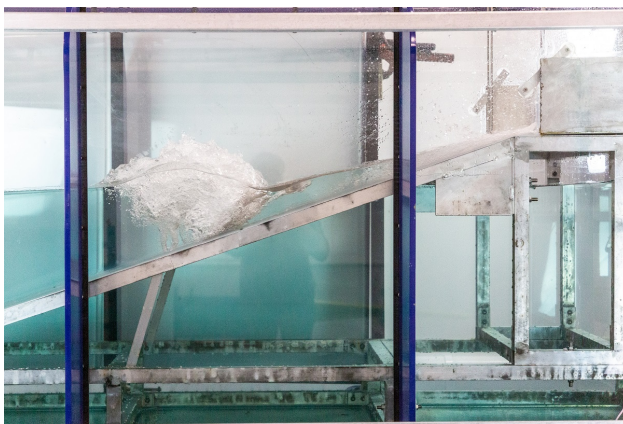
(b)



(c)



(d)



(e)

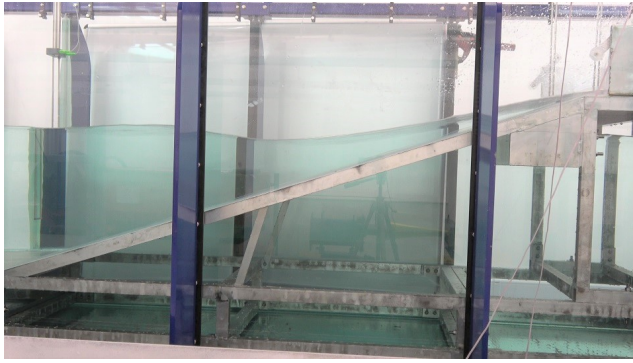


(f)

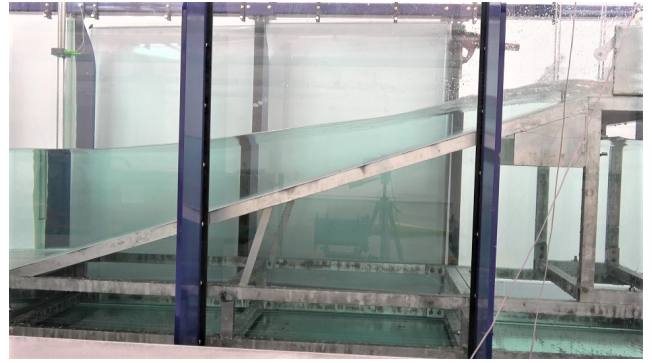
Figure 6.11 Snapshots from the experimental test showing some wave breaking events at the seawall with fewer wave overtopping events by a wave with $H_{m0} = 0.1$ metres at a water depth 0.58 metres approaching a slope with $\cot \alpha = 3.0$

As soon as more progressive waves are involved in this process, due to increasing swell, the wave overtops the crest of the structure straightforwardly, without breaking

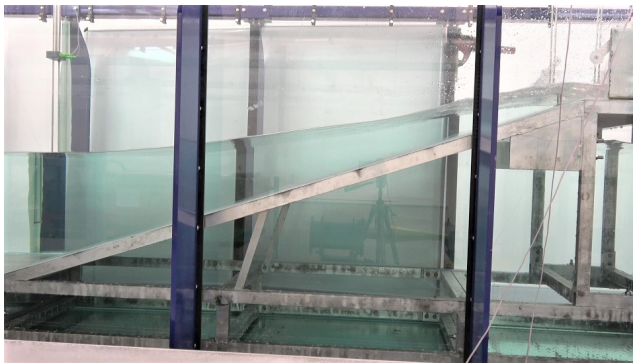
(Figure 6.12). This occurs because the forward thrust of the wave enhances more overtopping. At the time these experiments were conducted wave overtopping, caused by breaking waves, was not considered. As such formulations for non-breaking waves have been applied to each case retrospectively.



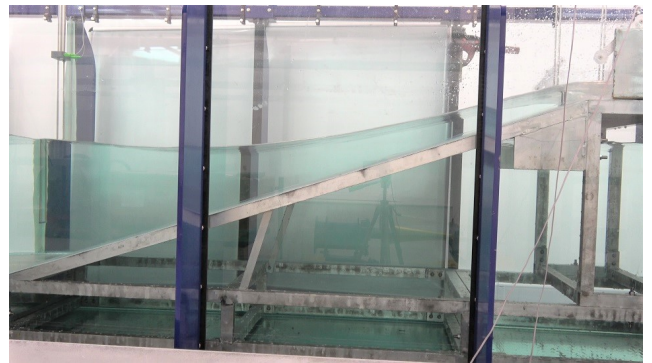
(a)



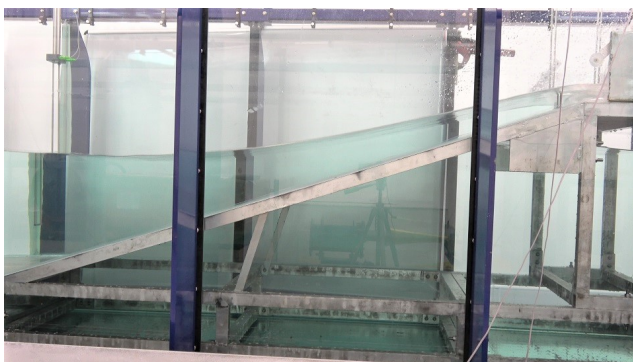
(b)



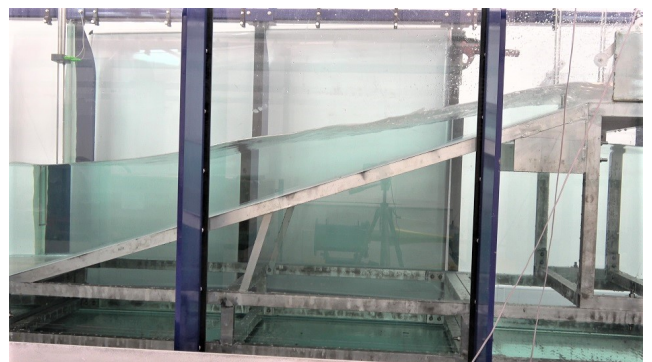
(c)



(d)



(e)



(f)

Figure 6.12 *Snapshots from the experimental test showing: wave overtopping from non-breaking waves by a wave with $H_{m0} = 0.1$ metres at a water depth 0.6 metres approaching a slope with $\cot \alpha = 3.0$*

6.4 Non-dimensional Wave Overtopping

Figure 6.13 shows the combined results of all conducted tests. As such, it depicts wave overtopping, caused by non-breaking waves, on three sloped seawalls. These estimates are derived using the EurOtop (2018) overtopping formula (Equation (2.4)). Coefficients A and B applied in Equation (2.4), also change with individual slopes using Equations (2.5) and (2.6) for the non-breaking wave cases.

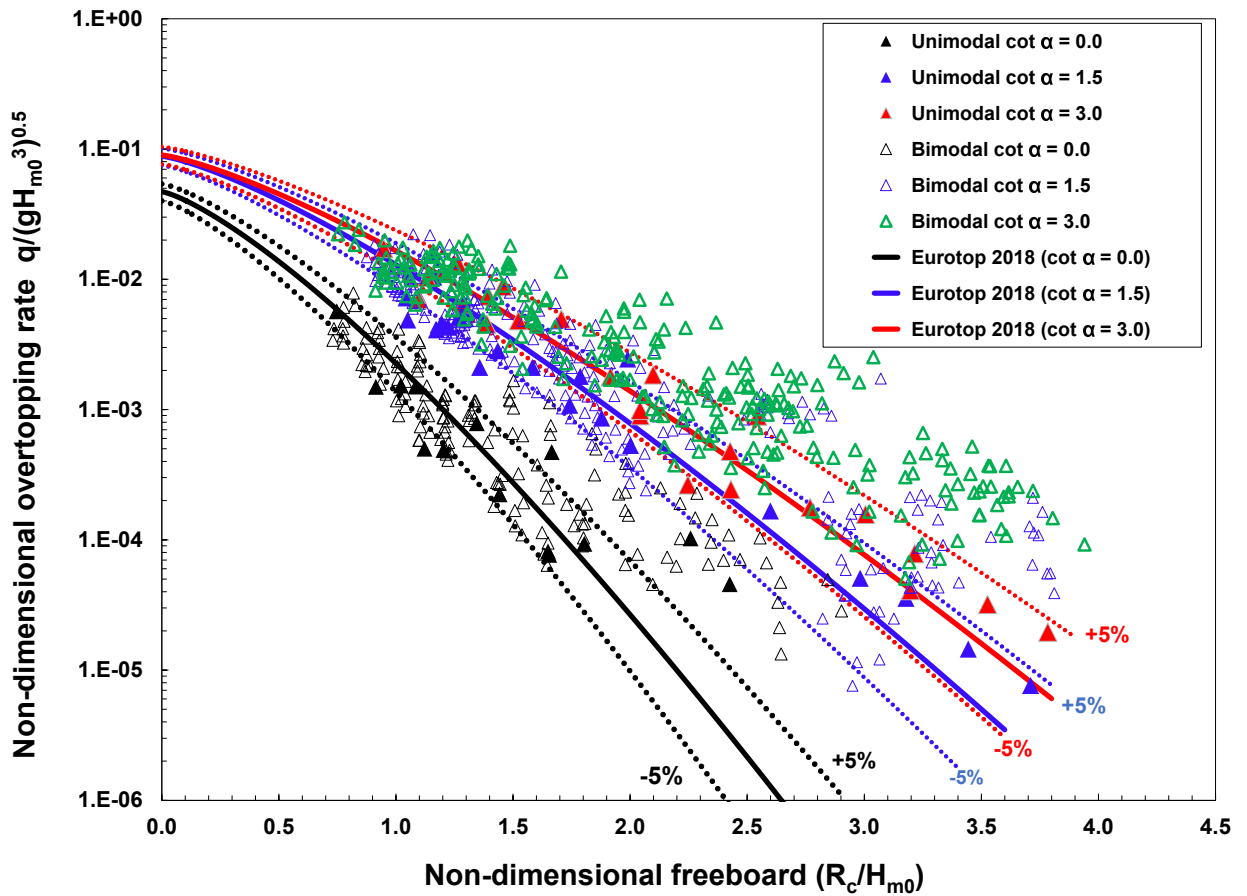


Figure 6.13 Non-dimensional average overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against relative crest freeboard $\frac{R_c}{H_{m0}}$ for all tests captured for both unimodal and bimodal sea states across slopes with $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$

6.4.1 Bimodal Wave Overtopping of Sloping Seawalls

For clarity, a replotted version of Figure 6.13 for sloping seawalls has been reproduced as Figures 6.14 and 6.15. The predicted overtopping and corresponding 90 % confidence limits of EurOtop (2018), for both unimodal and bimodal cases, are presented in each

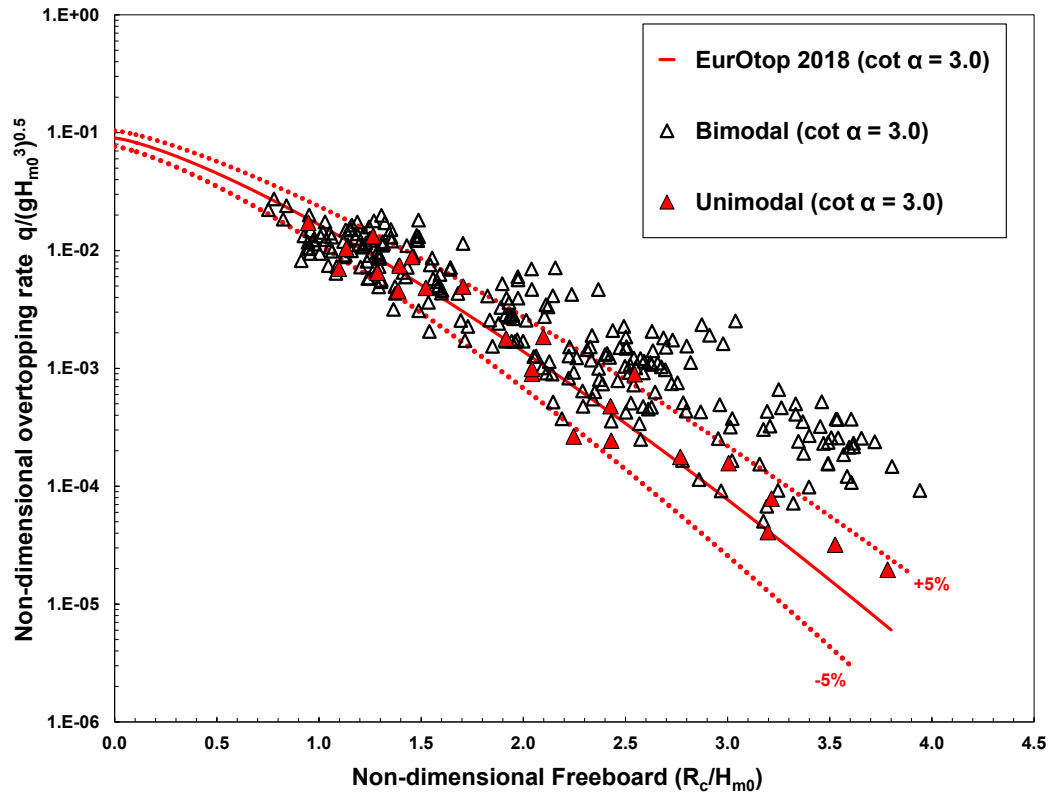


Figure 6.14 Results of non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$ plotted within the corresponding prediction band of EurOtop (2018) obtained for a sloping seawall with $\cot \alpha = 3.0$.

plot. Figure 6.14 represents a gentle seawall slope ($\cot \alpha = 3.0$), and Figure 6.15, a steeper seawall slope ($\cot \alpha = 1.5$). Different coefficients A, and B corresponding to individual slopes were obtained from Equations (2.5) and (2.6) for non-breaking overtopping cases. These were also applied to predict the confidence limits. As is visible from these results, the majority of non-dimensional wave overtopping results observed for unimodal cases fitted very well within the prediction and uncertainty intervals from the EurOtop (2018) non-breaking overtopping prediction formula that corresponds to individual slopes.

Using the conventional 90 percent prediction band in the figures, a large number of bimodal cases lie outside these intervals, especially for larger dimensionless freeboard (Figures 6.14 and 6.15), and are largely underpredicted by the EurOtop (2018) formula.

In Figures 6.14 and 6.15, the overtopping for bimodal waves spreads beyond the confidence intervals and is more pronounced for values of $R_c/H_{m0} > 2.0$. There seems to be more scatter at lower R_c/H_{m0} for bimodal sea cases than in the unimodal results.

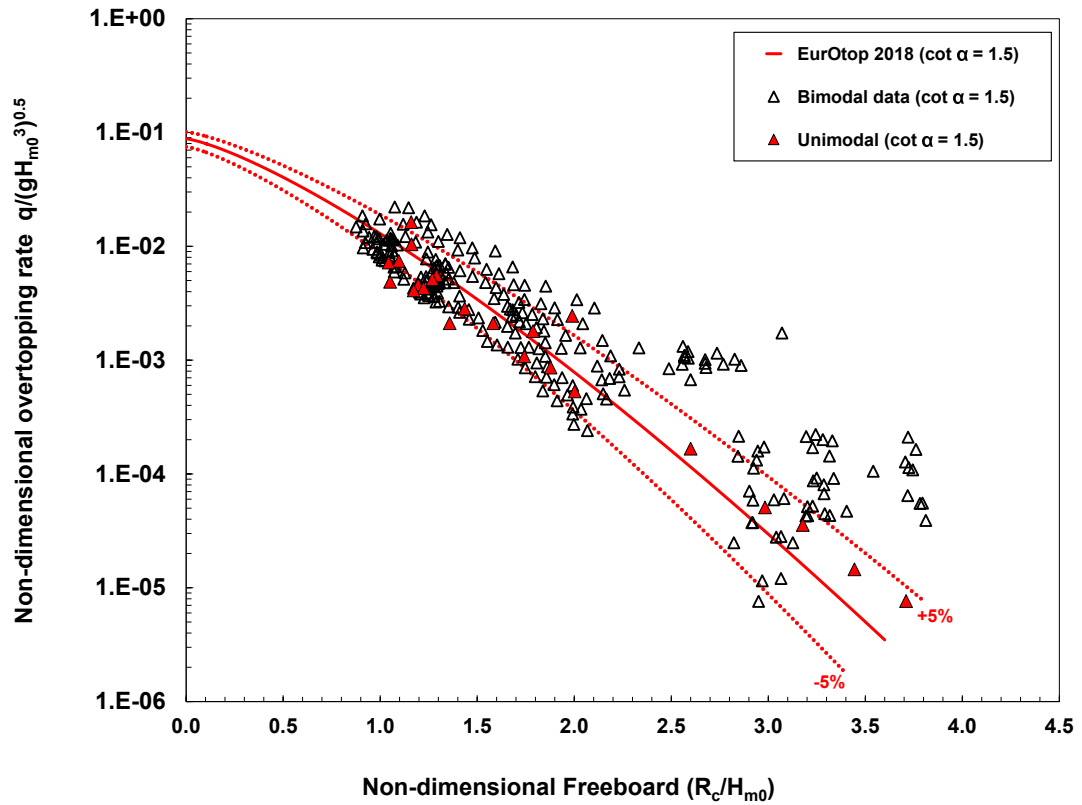


Figure 6.15 Results of non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$ plotted within the corresponding prediction band of EurOtop (2018) obtained for a sloping seawall with $\cot \alpha = 1.5$

Victor and Troch (2012) suggested that for unimodal seas, the scatter values could be due to the influence of wave periods of the tested wave conditions. However, for bimodal seas, the scatter effects due to wave periods are more significant. The effect seems to be more significant in terms of overtopping for larger crest freeboards for steeper wall slope (see Figure 6.16 in the next section) than for milder wall slopes with $\cot \alpha = 3.0$ (Figure 6.14).

6.4.2 Wave Overtopping of Vertical Seawalls

Results for the vertical seawall are presented in Figure 6.16 for both unimodal and bimodal sea conditions. In all the tests considered in this study, the wave overtopping discriminator for a vertical seawall h^* recommended in EurOtop (2018) was applied. With this assessment, all our experiments may be categorised as simple vertical walls, with influencing foreshore, under non-impulsive conditions (i.e., $h^2/H_{m0} \cdot L_{m-1,0} > 0.23$). A similar expression to Equations (2.5) and (2.6) was applied for both unimodal and

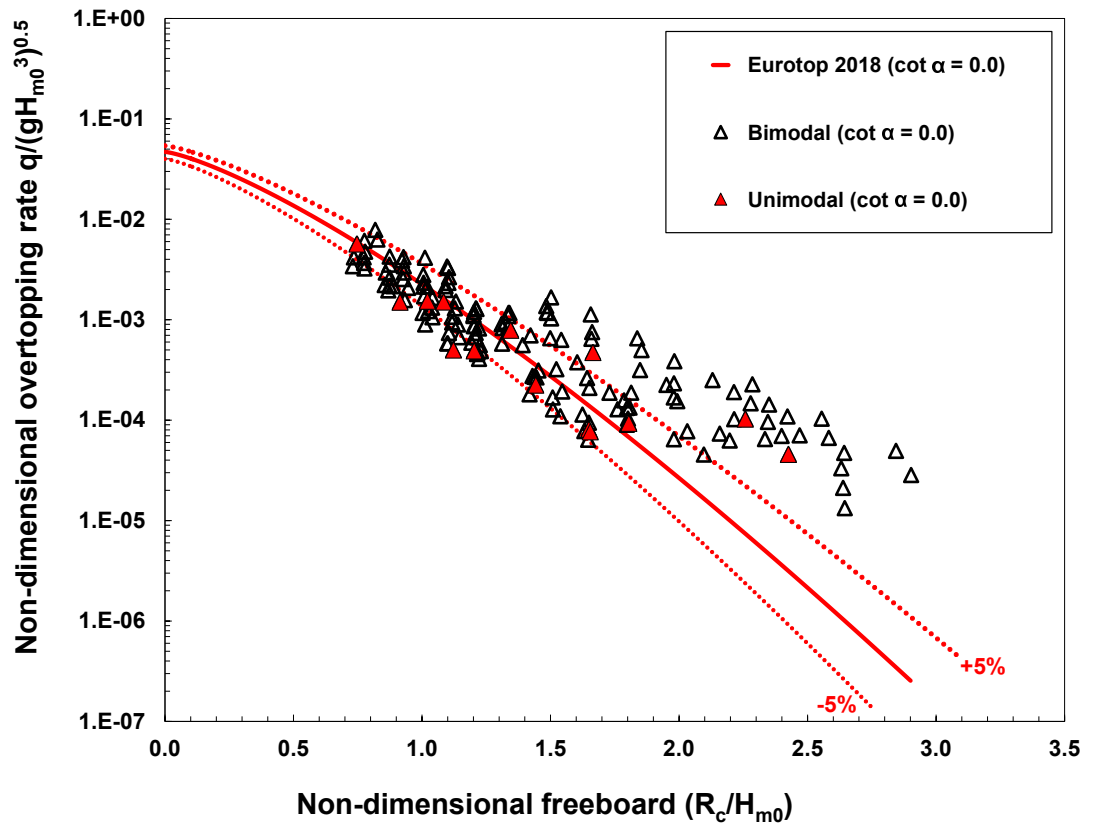


Figure 6.16 Results of non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$ plotted within the corresponding prediction band of EurOTop (2018) obtained for a vertical seawall with $\cot \alpha = 0.0$

bimodal conditions. As described in Figures 6.14 and 6.15, the 5 percent lower and upper confidence limits have been plotted, unimodal sea results are shown as red shaded triangles while bimodal seas are represented by black triangles. As can be seen from Figure 6.16, sets of predicted unimodal results are mostly within the 90 % confidence limits of the EurOTop (2018) model. As with the sloping walls most of the results for bimodal seas are outside the confidence limits. Although it is evident from Figure 6.16 that overtopping under both unimodal and bimodal waves were very well reduced by the vertical seawalls, they are less effective at dissipating wave energy than smooth, sloping seawalls (Neelamani and Sandhya, 2003; Van Gent, 1994) presented by Figure 6.14. In this case, more standing waves are created, which reflects waves in a seaward direction with less chance of overtopping the crest.

6.5 Assessment of Prediction Accuracy

It is informative to quantify the accuracy of the predictions provided by the EurOtop formulations, as presented in Figures 6.13 – 6.16. In this section, the root mean square errors RMSE and the bias of the errors BE are both applied to assess the accuracy provided by the EurOtop (2018) as presented in Equations (2.4) – (2.6). The RMSE can be generally represented as:

$$RMSE = \sqrt{\frac{1}{N_t} * \sum_{n=1}^{N_t} [q_{pred}^* - q_{measured}^*]^2} \quad (6.2)$$

where n is test identifier, N_t is the total number of tests under consideration, q_{pred}^* is the predicted dimensionless overtopping, and $q_{measured}^*$ is the measured dimensionless overtopping value. The bias of the measurement can be defined in terms of the average of the overall dispersion of individual values of q_{pred}^* and the $q_{measured}^*$:

$$BIASError = \frac{1}{N_t} * \sum_{n=1}^{N_t} [q_{pred}^* - q_{measured}^*] \quad (6.3)$$

With the bias estimates of the results, it is easier to describe both overprediction patterns (when bias is positive) and under-predictions (negative bias values).

By combining predicted values from Equations (2.4) – (2.6) with Equations (6.2) – (6.3) using observed experimental datasets across the three slopes, the accuracy assessment by RMSE and Bias are shown in Table 6.7. It is evident in this table that the unimodal datasets are very well predicted in terms of the RMSE and Bias values using the EurOtop (2018) formulations. As earlier presented in previous figures for bimodal datasets, the RMSE and bias values are not very well predicted.

Table 6.7 Values computed for *RMSE* (Equation (6.2)) and bias (Equations 6.3) by assessing the EurOtop Equation (2.4) for different ranges of relative crest freeboard $R_c = H_{m0}$ and slope angles

Formulation Equations (2.4) – (2.6)		
Tested Cases	RMSE (-)	Bias (-)
All tests (cot $\alpha = 3.0$)	0.00365	-7.46×10^{-6}
Unimodal	0.00284	8.83×10^{-4}
Bimodal	0.00463	-7.41×10^{-5}
All tests (cot $\alpha = 1.5$)	0.00336	-1.69×10^{-3}
Unimodal	0.00191	1.36×10^{-4}
Bimodal	0.00304	-5.18×10^{-4}
All tests (cot $\alpha = 0.0$)	0.00972	-3.84×10^{-4}
Unimodal	0.00103	5.80×10^{-4}
Bimodal	0.00967	-3.65×10^{-4}

The least bias values as indicated by the vertical seawall (cot $\alpha = 0.0$) in present study. It may be due to the accuracy of Equations (2.4)–(2.6) at predicting the overtopping of simple vertical walls, with influencing foreshore, under non-impulsive conditions (i.e., $h^2/H_{m0}.L_{m-1,0} > 0.23$).

6.6 Influence of Shallow Water Conditions on Non-Breaking Wave Overtopping

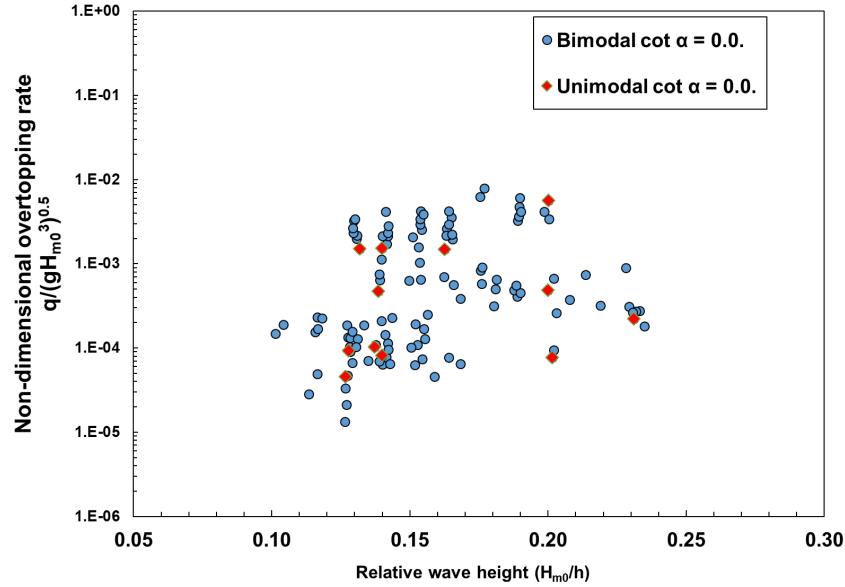


Figure 6.17 The relationship between the non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against the relative wave height H_{m0}/h for the dataset of $\cot \alpha = 0.0$

The influence of the relative wave height on the dimensionless wave overtopping across the two slopes for both unimodal and bimodal cases are each presented in Figure 6.17 (for $\cot \alpha = 0.0$), Figure 6.18 (for $\cot \alpha = 1.5$), and Figure 6.19 (for $\cot \alpha = 3.0$). The figures reveal that the majority of the observations fall within what may be categorized as deep-water conditions ($H_{m0}/h \leq 0.2$) as anticipated for non-breaking waves that are studied in this thesis. The shallow water effects can be described in terms of this ratio. This ratio is also useful in describing the water depth condition within different test groups Orimoloye et al. (2019a). As previously observed in Figures 6.14 and 6.15 that results for the steeper slope, ($\cot \alpha = 1.5$), show more scatter than for the milder slope, ($\cot \alpha = 3.0$), when the dimensionless freeboard is greater than 2.0 as shown in the figures. In deep-water conditions, no clear trend is observed in these results except for a small increase in the non-dimensional wave overtopping within the range $0.08 \leq H_{m0}/h \leq 0.12$ for the steepest slope ($\cot \alpha = 1.5$). Figure 6.18 shows that shallow water effects are less significant for the milder slope ($\cot \alpha = 3.0$). Nevertheless, there seems to be an increase in the lower values of the dimensionless wave overtopping with increasing relative wave height conditions (Orimoloye et al. (2019a)).

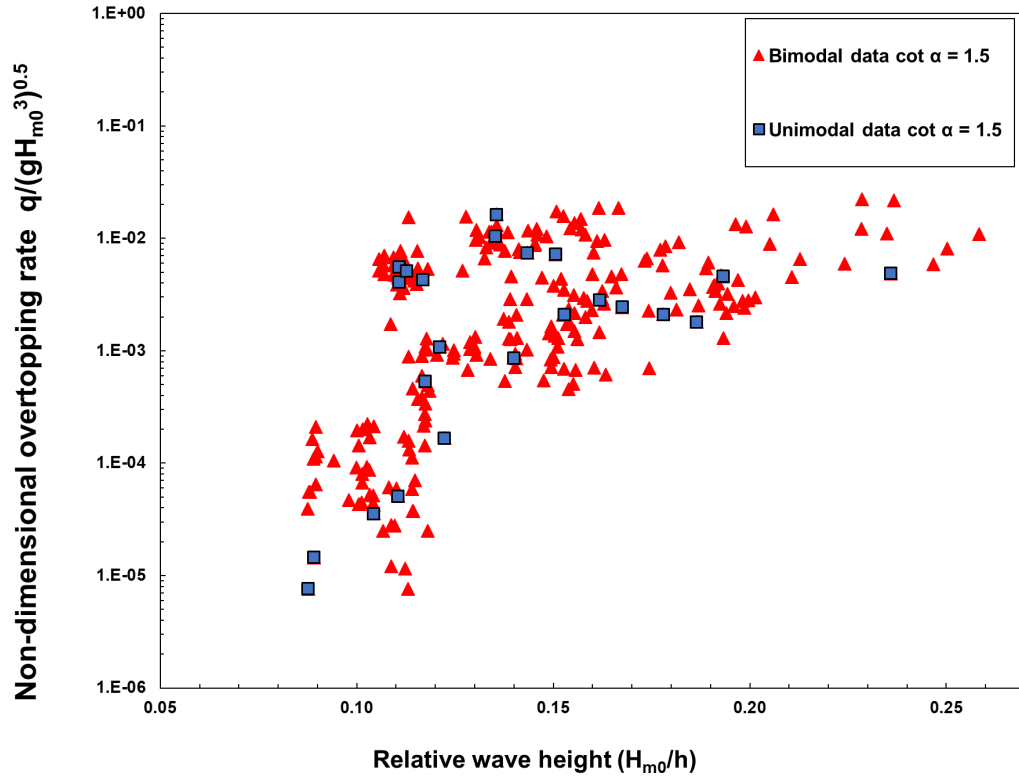


Figure 6.18 The relationship between the non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against the relative wave height H_{m0}/h for the dataset of $\cot \alpha = 1.5$

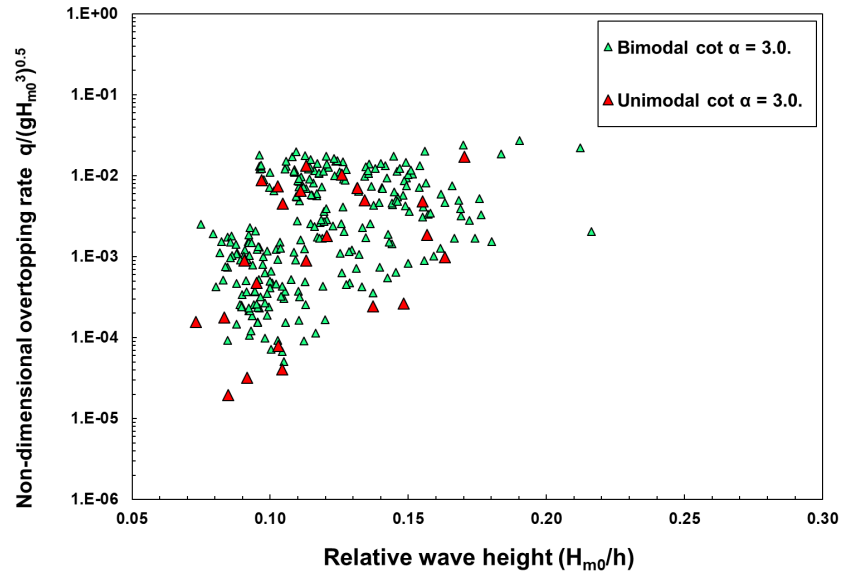


Figure 6.19 The relationship between the non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against the relative wave height H_{m0}/h for the dataset of $\cot \alpha = 3.0$

6.7 Conclusion

823 successful wave overtopping tests have been conducted to study the wave overtopping performance of sloping impermeable seawalls under the influence of bimodal sea conditions. Series of bimodal spectra that contain the same energy as the equivalent wind sea have been applied in all the tests. Different percentages of swell components were derived for each sea-state and were evaluated at different peak periods of each of the swell sea. Both unimodal and equivalent bimodal waves were tested. A minimum of 1000 waves were used for each wave condition to obtain the average wave overtopping in each case. For the fixed wave conditions, more wave overtopping was observed with increasing slope mildness, as established in previous studies. The unimodal tests agree closely with predictions using the EurOtop formulae. EurOtop formulae underpredict the observed wave overtopping rates for bimodal conditions, (even when using the $T_{m-1,0}$ period for such cases). Improvement of the EurOtop formulae reflecting effects of presence swell will be presented in Chapter 7.

Chapter 7

Effects of Swell and Spectra Parameters on Wave Overtopping

In this chapter, the experimental datasets obtained from chapter 6 are analysed using the presence of swell waves and spectra characteristics of the bimodal spectra involved. Special attention is placed on the analyses of effects of swell percentages and swell peak periods on overtopping features of coastal seawalls under bimodal sea conditions. Moreover, the effects of spectra shapes imposed by different spectra types were examined using the spectral characteristics of narrowness parameters, broadness parameters and peakedness parameters of each spectra types. Both unimodal and corresponding bimodal waves that have been simulated in Chapter 6 under different slopes and crest freeboards are applied. Until now, wave overtopping equations in the EurOtop (2018) have not specifically include the direct effect of swell percentages under each wave conditions. Because the proportions of swell in the bimodal sea have been varied in each test, it is possible to present the specific influence of swell percentages on overtopping rates. From analysis of the experimental results, a new empirical influencing factor to describe the effects of swell percentages on wave overtopping discharges are presented.

7.1 Spectral Characterisation

To investigate effects of spectral parameters on wave overtopping estimates, surface elevation η recorded by wave gauges from experiments are further inspected and analysed. According to Ochi (2005), each surface elevations can be described as an equivalence of individually statistically independent random variables X_i , which have

the same probability distribution of the form:

$$\eta = X_1 + X_2 + X_3 + \dots X_n. \quad (7.1)$$

Also, the general description of the wave spectrum $S(f)$ can be estimated. If according to Goda (2010) we assume $S(f)$ as a major input towards estimating the spectral moments m_n of n^{th} order:

$$m_n = \int_0^\infty f^n E(f) df. \quad (7.2)$$

In turn, if values of n are defined as -1, 0, 2, and 4 respectively as presented in Reeve et al. (2015), the spectral significant wave height H_{m0} and different spectra periods ($T_{m-1,0}$ & $T_{m0,2}$) can be determined from m_n as:

$$H_{m0} = 4.004\sqrt{m_0} \quad (7.3)$$

where m_0 is the zeroth moment of the spectrum and

$$T_{m-1,0} = \frac{m_{-1}}{m_0} \quad (7.4)$$

The following spectral parameters were also estimated from surface elevations recorded during each test:

1. Peakedness factor

According to Goda et al. (1975), the grouping of successive wave height can be expressed in terms of its peakedness parameter Q_p with frequency f :

$$Q_p = \frac{2}{m_0^2} \int_0^\infty E(f) df. \quad (7.5)$$

Bradbury et al. (2007) suggested that values of Q_p increase with the narrowness of the spectrum. Fully developed wind-seas were observed to have values Q_p within the range of 2 while swell seas may possess Q_p of higher values in some cases. Battjes and Groenendijk (2000) verified the correlation κ that was earlier formulated by Kimura (1980). It is the κ coefficient that was based on the first-order Markov's chain. Values of κ can be determined as follows:

$$\kappa = \frac{1}{m_0} \left| \int_0^\infty E(f) e^{i2\pi f \tau} df \right|. \quad (7.6)$$

κ also increases as the spectral bandwidth decreases, and τ denotes the mean time difference between two successive wave crests or troughs. For a well-represented wave spectrum, the peakedness of the wave spectrum would show a high correlation with the wave groupiness parameter λ . The groupiness of individual wave group with angular frequency ω can be determined using an expression suggested by Rychlik et al. (1997a):

$$\lambda = \frac{\text{abs}(\int_0^\infty S(f) * \exp^{2\pi f \tau} df)}{m_0} \quad \text{and } \tau \text{ is equivalent to } T_{m02} = \sqrt{\frac{m_0}{m_2}} \quad (7.7)$$

2. Narrowness parameter

Longuet-Higgins (1983) proposed the Gaussian process of the narrow-banded spectrum. The narrowness parameter ν of a sea-state can be defined in terms of the measurement of the distribution of the frequency component in the sea-state. It can be expressed mathematically as:

$$\nu = \sqrt{\left(\frac{m_2 m_0}{m_1^2} - 1\right)} \quad (7.8)$$

3. Spectral width parameter

The spectral width parameter or the broadness factor ε of a wave spectrum is essential to describe the frequency band where the wave energy mostly concentrates. As presented by Rychlik et al. (1997a), this could be narrow-banded with $\varepsilon \approx 0$ or broad-banded with $\varepsilon \gg 0$:

$$\varepsilon = \sqrt{\left(1 - \frac{m_2^2}{m_0^2 m_4}\right)} \quad (7.9)$$

4. Quality and irregularity parameters

In this study, the quality parameter R_s is expressed in terms of the total wave spectrum $S(f)$ while the irregularity factor α_f is extracted using some moment ratio. It is mathematically defined by Rychlik et al. (1997a) as follows:

$$R_s = \frac{S(0.092) + S(0.12) + S(0.15)}{3 * \max(S(f))} \quad \text{also, } \alpha_f = \frac{m_2}{\sqrt{m_0 * m_4}} \quad (7.10)$$

7.1.1 Wave transformation

Figure 7.1 describes the wave evolution pattern for both unimodal and bimodal spectral with the same H_{m0} of 0.125 metres propagating across the 1:20 slope interacting with a slope of 1:1.5. After removing the effect of reflection (full details are presented in Chapter 6), wave shoaling and breaking are fully described in the figure. A distance of 0 metres represents the wavemaker position while the waves propagate towards the toe of the structure. The water depth is kept at 0.6 metres at offshore while reduces by the sloping structure to 0.39 metres at the toe. Similar to Lorenzo et al. (2001), the swell percentages greatly influences shoaling and breaking over a short distance. As shown, the expected wave height of 0.125 metres was accurately captured by the wave gauges in both unimodal and bimodal sea conditions. However, effects of shoaling are more significant in the bimodal sea (to a maximum of 0.16 metres for bimodal waves) than in the unimodal sea (which shoals to a maximum of 0.14 metres) just before breaking. As previously described by van der Meer et al. (2016), it is the lower steepness that is responsible for the higher shoaling occurrences in the bimodal seas than observed in unimodal cases.

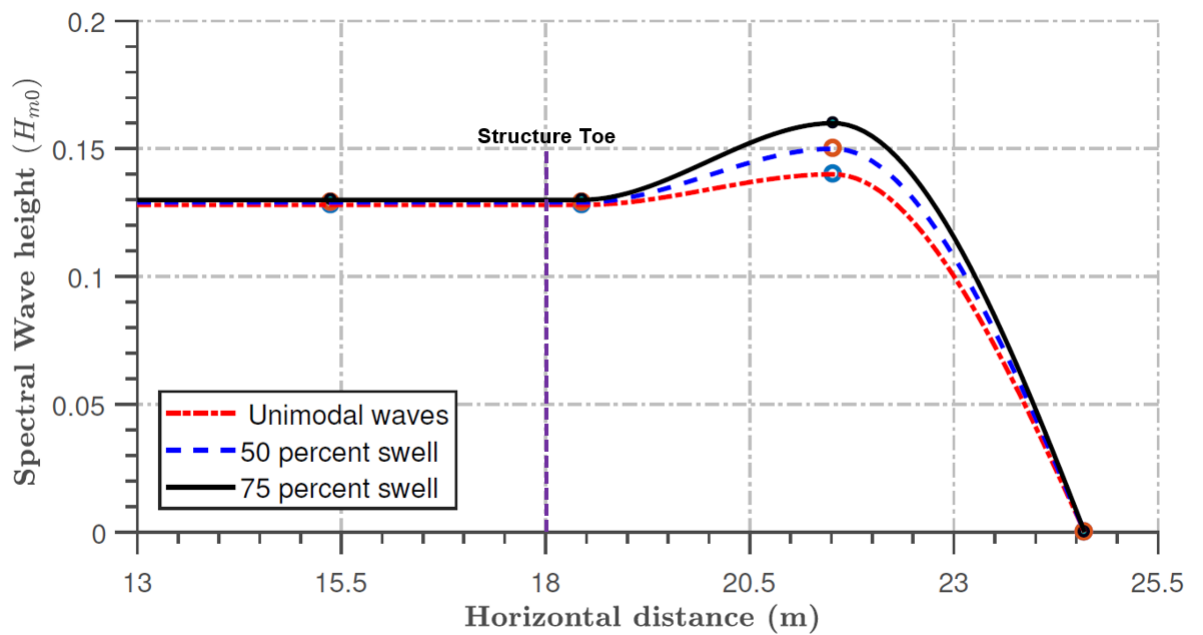
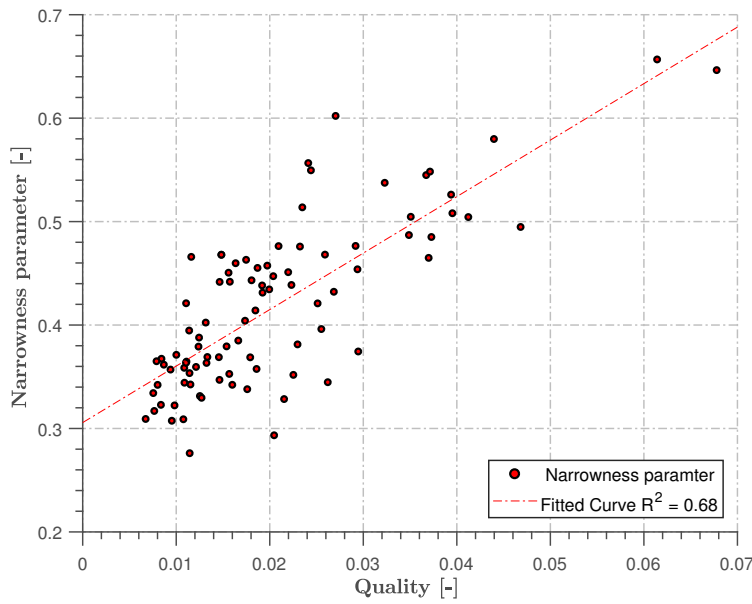


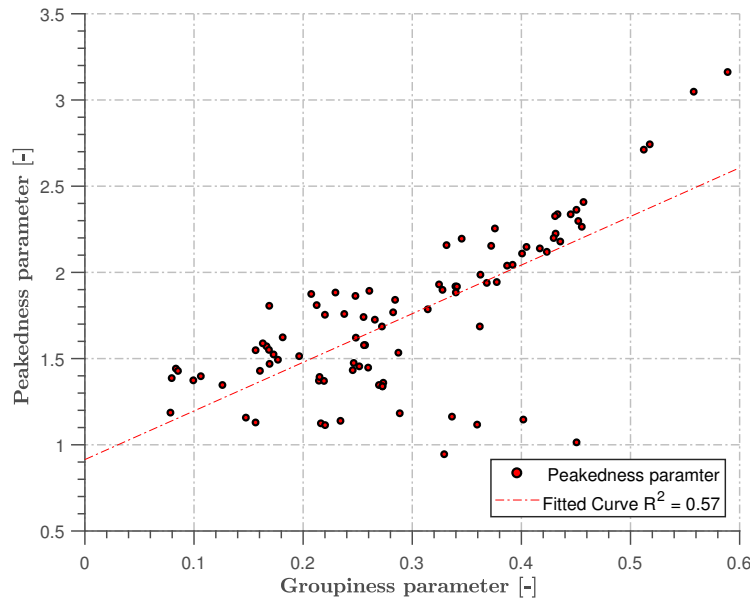
Figure 7.1 Wave height deformation observed on a foreshore slope of 1:20 at a constant water depth of 0.6 metres for unimodal and bimodal waves

7.2 Spectral effects on wave overtopping

Spectral characteristics of a sea-state is an essential requirement in describing the severity of the sea. In this section, we describe the complexities of the sea-state in terms of the swell content (percentage) of the sea-state. Firstly, in Figure 7.2, we presented how well the experimental result correlates with different spectra characteristics earlier described.



(a)



(b)

Figure 7.2 Performances of shape parameters computed for: (a) narrowness, and (b) peakedness for different bimodal tests obtained from selected datasets

Altogether, as shown in Figure 7.2, the narrowness and the peakedness parameters of the wave show a positive correlation with wave quality and groupiness parameters respectively. The wave peakedness is well associated with the groupiness parameter of the wave. There are variations in the wave spectral characteristics as shown by the spectra values on the vertical axis.

The peak period at which a swell wave occurs has the potential to alter the characteristics of the resulting bimodal wave. In Figure 7.3, a cross-section of the potential influence of different computed spectral characteristics of swell waves occurring at 11, 20, and 25 secs peak periods, as presented.

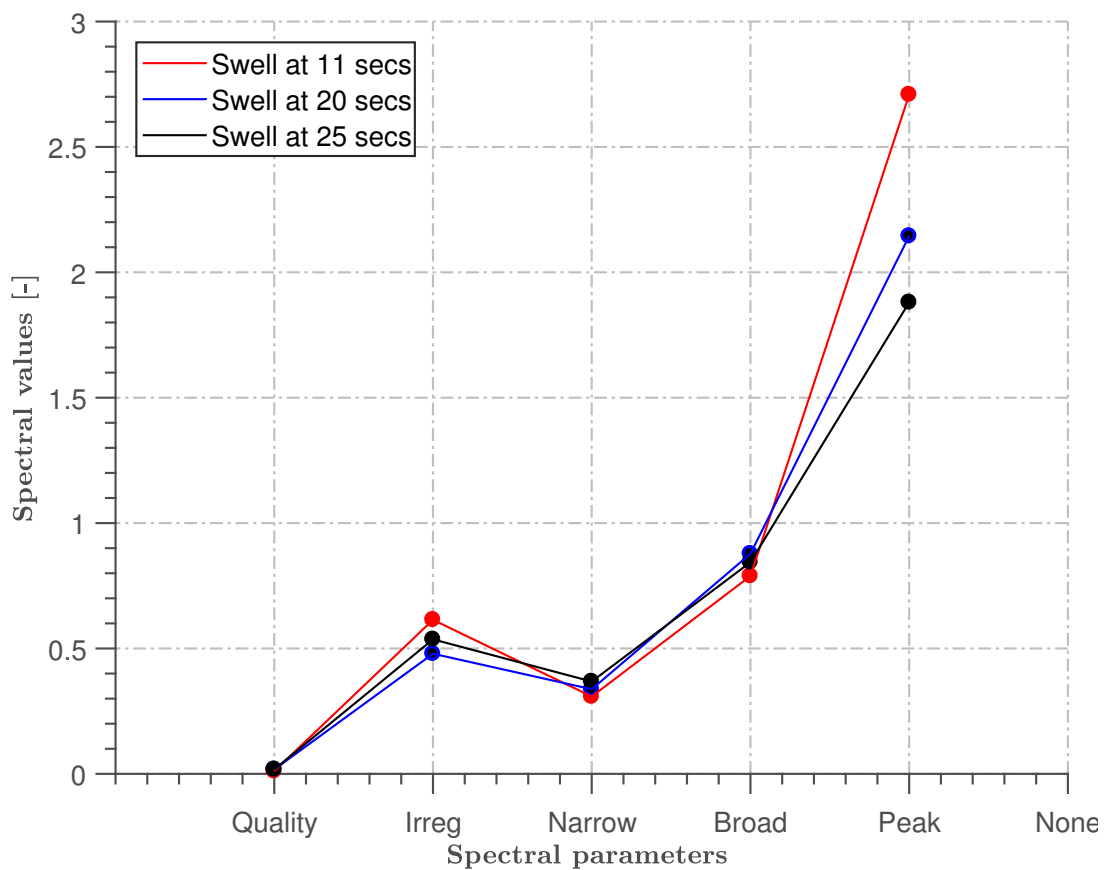


Figure 7.3 Variations of spectral parameters with swell peak periods in a bimodal sea with $H_{m0} = 0.125$ metres at a water depth of 0.60 metres.

Non-Linear behaviour exists in the spectral values portrayed by various spectra aspects. There are no significant observable differences in the broadness and narrowness parameters across swell peak periods. However, more significant variations are observed in spectral irregularities and narrowness parameters that are present in the low period swell (11 secs) than higher ones (20 & 25 secs). It could be due to continuous occurrences of bimodal waves due to a smaller value of wave periods compared to higher

ones. Also, disparities in spectral behaviour is an indication that wave characteristics would be different as the swell in the sea state changes. These observations are similar to the work of Bradbury et al. (2007) where it was suggested that a minimal difference occurs in both narrowness and broadness parameters of different spectral types.

On the other hand, relationships between the rate of wave overtopping and different spectral parameters are presented in Figures 7.4 – 7.6. Overtopping estimates are plotted on the primary axis while other spectral characteristics are plotted on the secondary axis. From the three figures, overtopping rates are increasing with swell percentages while the spectral values reduce for all the three parameters. However, in Figures 7.4 and 7.5, the trends are tending towards asymptotic values as swell percentages increase from 0 to 75 whereas, waves peakedness parameter approaches zero for higher swell percentages (Figure 7.6). These observations, however, suggested that seas that are characterised by low-peaked wave spectrum tend to impose more severe flooding of extensive coastlines than higher-peaked seas. Similarly, less flooding tendencies are expected for broad-banded seas. Overtopping is highly sensitive to the succession of wave groups and wave patterns in the random wave sequence than overtopping caused by individual peaks. The evidence of extreme wave events have been previously observed most especially in Hunt (2003) and Weston et al. (2005).

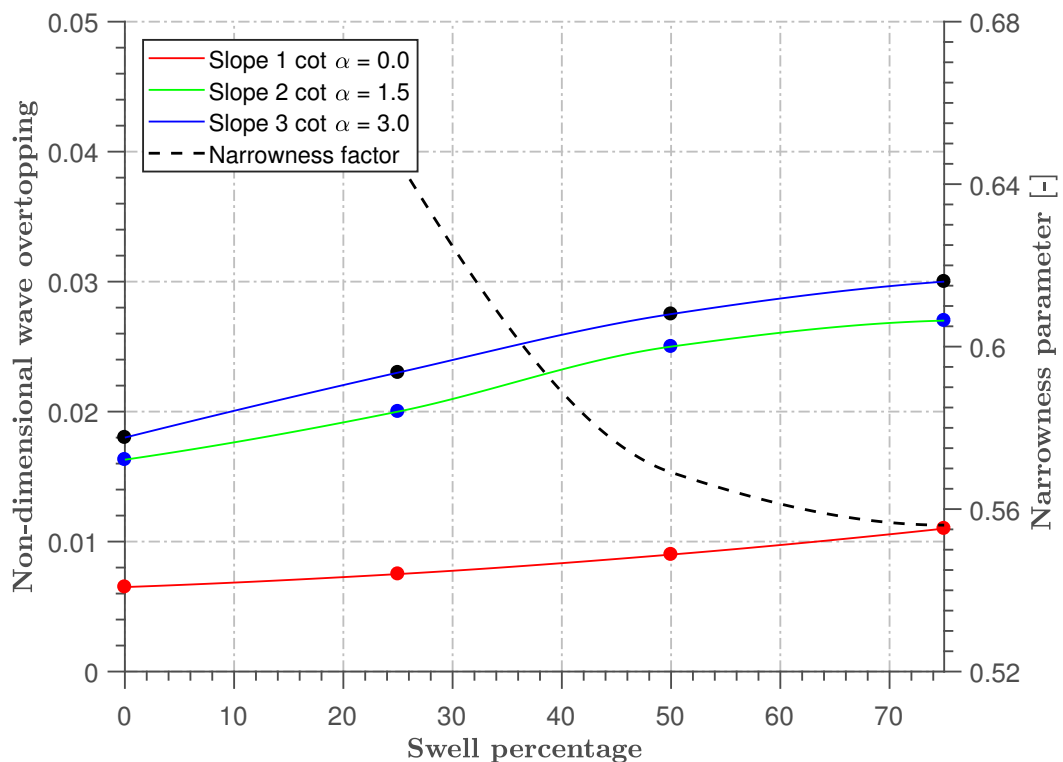


Figure 7.4 Influences of narrowness parameters on the wave overtopping of sloping seawall under different swell percentages

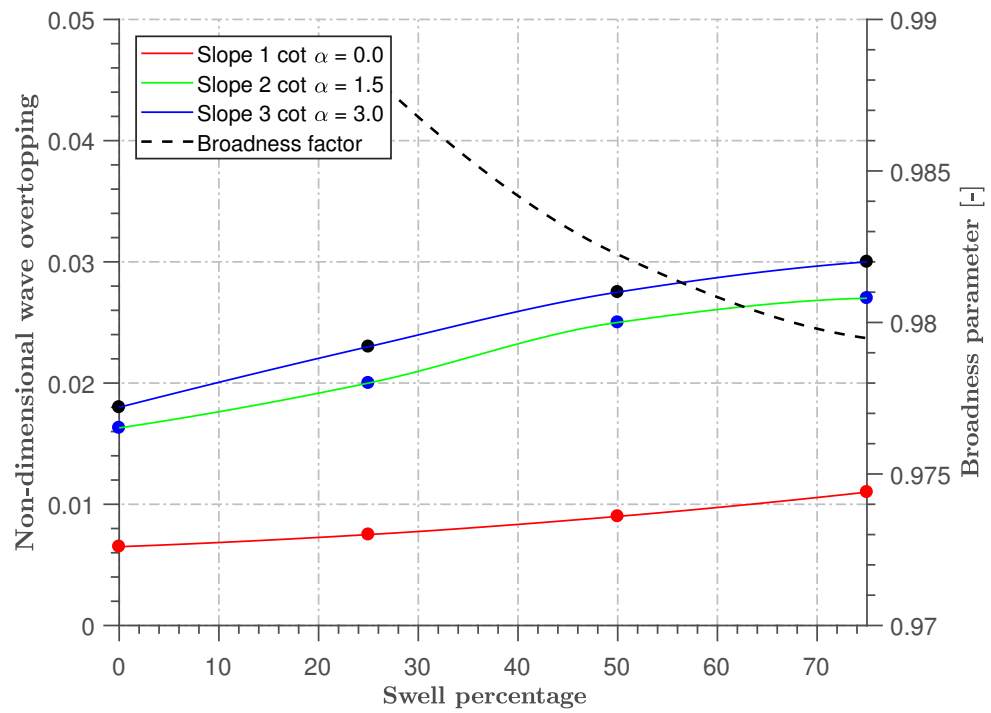


Figure 7.5 Influences of broadness parameters on the wave overtopping of sloping seawall under different swell percentages

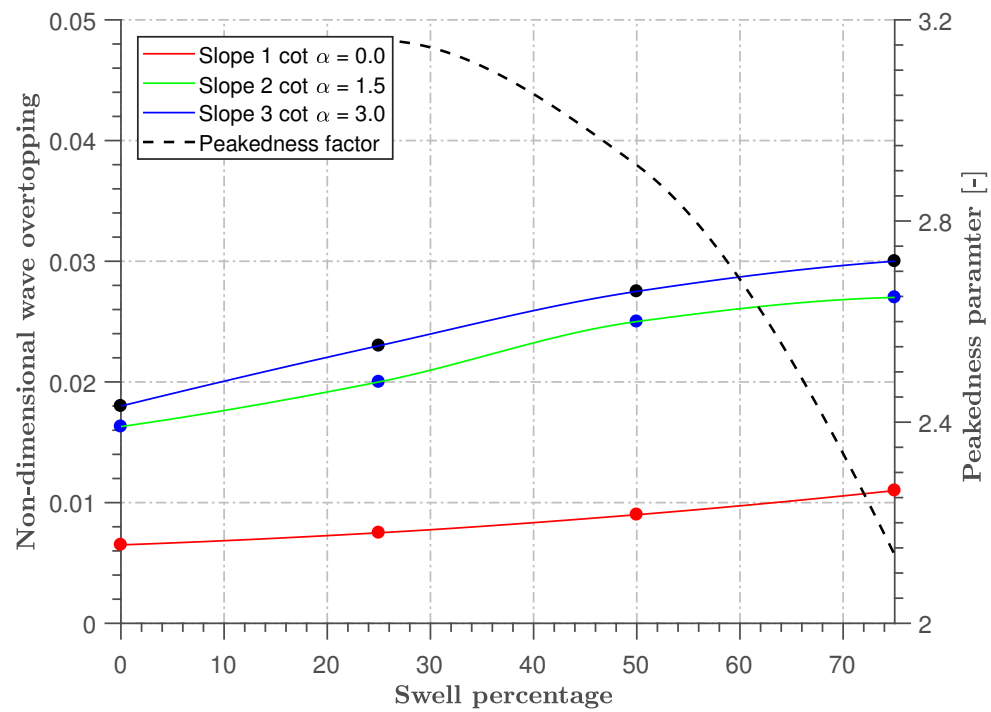


Figure 7.6 Influences of peakedness parameters on the wave overtopping of sloping seawall under different swell percentages

7.2.1 Variations of Wave Steepness with Swell Percentages

The wave steepness $S_{m-1,0}$ is an important physical parameter in describing the severity of a sea state. Most especially defining the physical indication of wave breaking for a design of coastal structures. In order to accurately describe both unimodal and bimodal sea states in this study, a wide range of $S_{m-1,0}$ between 0.01 and 0.07 has been tested as calculated from the ratio of wavelength H_{m0} and wavelength $L_{m-1,0}$.

Figures 7.7 – 7.8 show an example of the relationship between wave steepness and different swell components that were applied during this study. As suggested in the EurOtop manual (EurOtop 2018), wave steepness reduced with increasing swell percentage and R^* respectively.

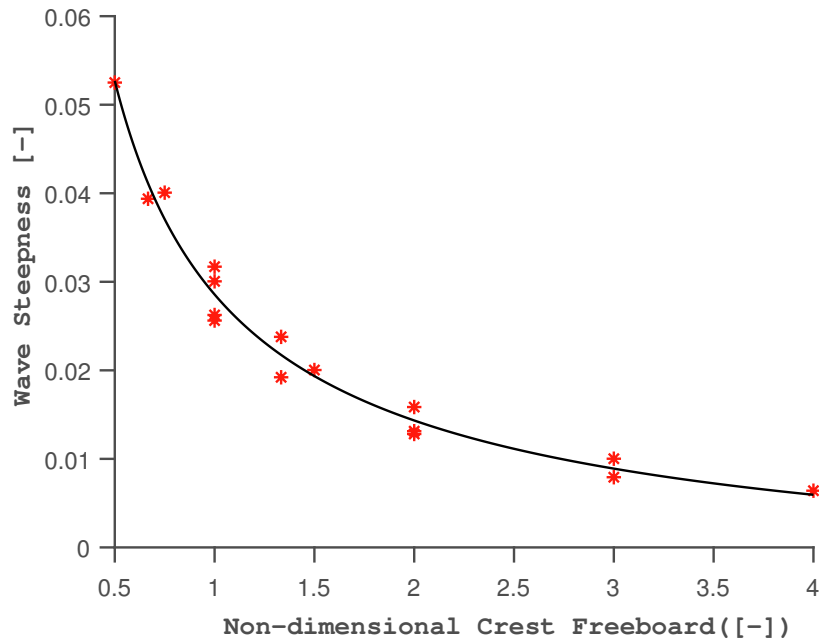
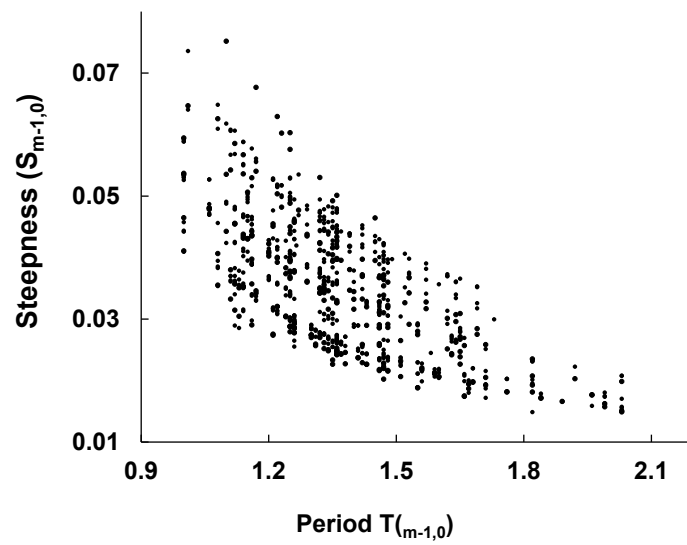
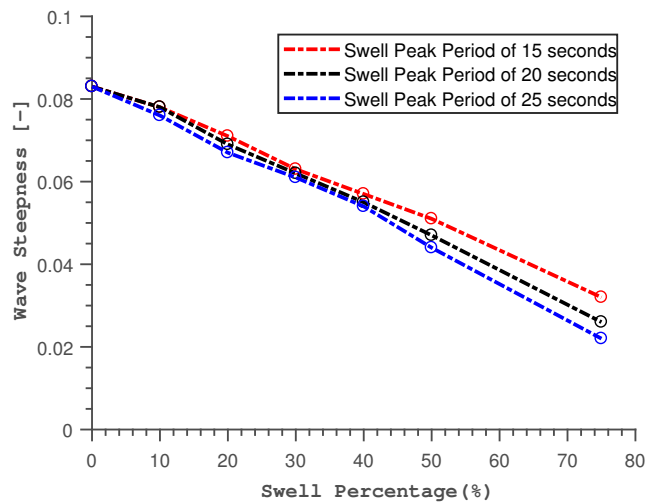


Figure 7.7 Relationship between tested wave steepness and the R^* with swell percentages in a sea-state with equal energy

Figure 7.8 shows that the highest presence of swell (at 75 %) gives the smallest wave steepness Figure 7.8b. It is the low wave steepness $S_{m-1,0}$ in the bimodal seas that accounts for the increase in the breaker parameter $\xi_{m-1,0}$ through the relationship of $\xi_{m-1,0} = \tan\alpha/S_{m-1,0}$. It is larger values $\xi_{m-1,0}$ that corresponds to surging non-breaking waves that cause more overtopping in bimodal seas.



(a)



(b)

Figure 7.8 Variations of wave steepness with (a) spectra peak periods $T_{m-1,0}$ and (b) swell percentages in a sea-state with equal energy

7.3 Influence of Swell Percentage

In each of the sea states investigated, swell percentages were varied between 0 (unimodal), 25, 50, and 75 percents. Figure 7.8 shows the influence of swell percentage on wave overtopping estimates for all the cases investigated across the three slopes. Figure 7.8 represents the box plots created based on the dimensionless overtopping obtained from the experiments. Within the boxplots are the distributional characteristics as well as the level of the dimensionless overtopping obtained across different groups of swell percentages.

The boxplots describe the minimum (lower whisker) in each group, interquartile range (depicted by the boxes), and the upper whisker of the non-dimensional overtopping groups. To show the patterns of the effects of swell, the dimensionless overtopping are sorted according to the amount of the swell content in the sea state. There is obvious differences between the dimensionless overtopping presented in Figure 7.8. Shorter box plots are obtained for unimodal cases and 25 percent swell.

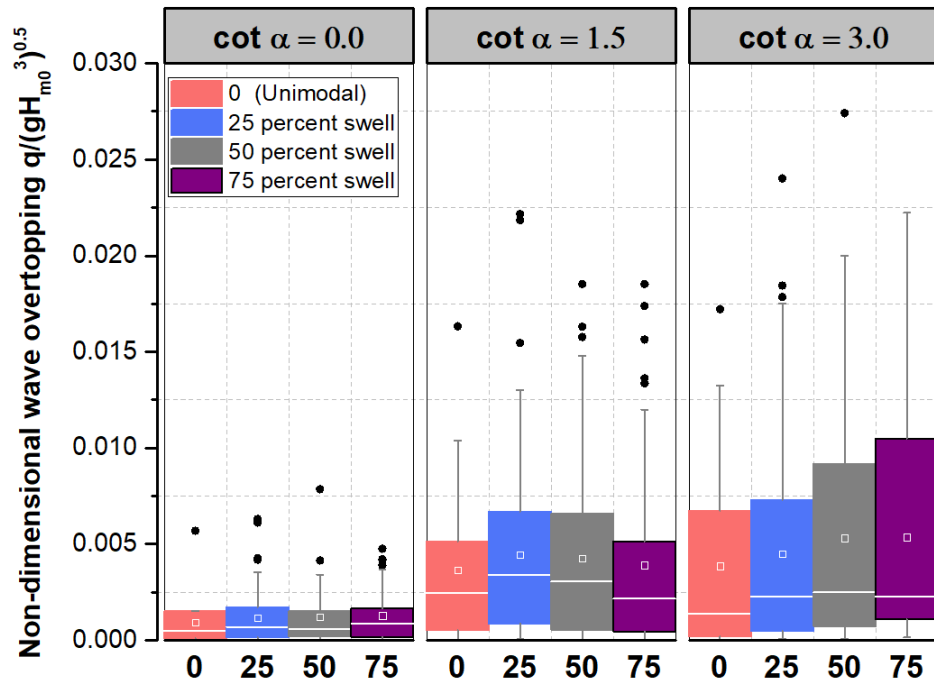


Figure 7.9 The relationship between the non-dimensional overtopping rate $\sqrt{H_{m0}^3}$ with the swell percentages for $\cot \alpha = 0.0$ and $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$

As described in Section 7.2, the overtopping is greater for the mildest slope. With further reference to Figure 7.8, as swell percentages increase in the bimodal sea cases, wave steepness reduces, overtopping increases, creating more non-breaking wave overtopping situations presented here in Figure 7.9.

The highest wave overtopping rates were observed for the sloping seawall with $\cot \alpha = 3.0$ while compared to the unimodal cases in the absence of swell. There is slight difference in the trend across the slopes. Considering the steeper slope with $\cot \alpha = 1.5$, this behaviour does not seem to follow directly as overtopping does not seem to increase monotonically as swell percentage reaches 50 percent. There seems to be a slight reduction in the wave overtopping for the extreme outliers seen in these results. In the case of the vertical seawall with $\cot \alpha = 0.0$, the overtopping rate appears fairly insensitive to the swell percentage.

For the unimodal test cases, effects of $Tm - 1, 0$ on overtopping results are not as significant for non-breaking waves compared to the effects of crest freeboards and slope angles. The observation follows previous conclusions for example, Victor and Troch, (2012), Van Doorslaer et al., (2015) and EurOtop (2018). The influence of $Tm - 1, 0$ on overtopping are negligible for non-breaking waves. This can also be seen that $Tm - 1, 0$ are not included in the expressions for the non-breaking wave cases. However, for the non-breaking bimodal waves, effects of $Tm - 1, 0$ become more significant. These were not shown in present manuscript because relationship of $Tm - 1, 0$ with overtopping seem to be affected by combined interactions of swell percentages and the appearances of swell at different peak periods. Details of effect of swell period will be presented in section 7.5.

7.3.1 Modified Overtopping Formula for Swell Percentages

The trends in overtopping with respect to swell percentage suggest the existing formula for overtopping Equations (2.4) – (2.6) might be modified to reflect this influence. Four swell percentages that are considered are 0, 25, 50, and 75 percent, (with 0 percent representing the full wind-sea state). Figures 7.10–7.13 show results obtained by non-linear least squares regression analysis by percentage of swell applied to the observations.

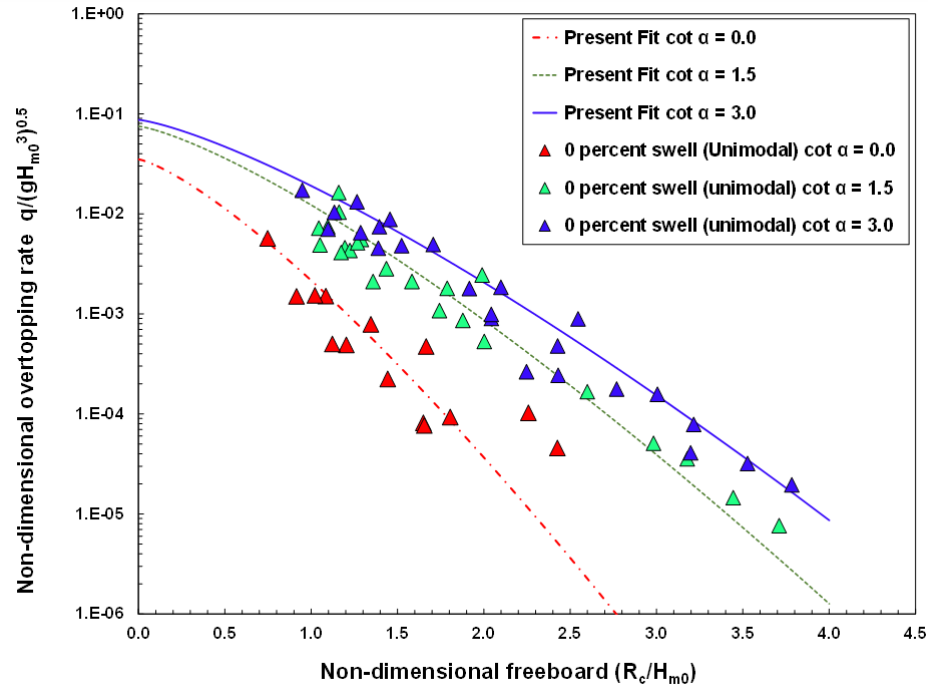


Figure 7.10 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for unimodal seas with 0 percent swell

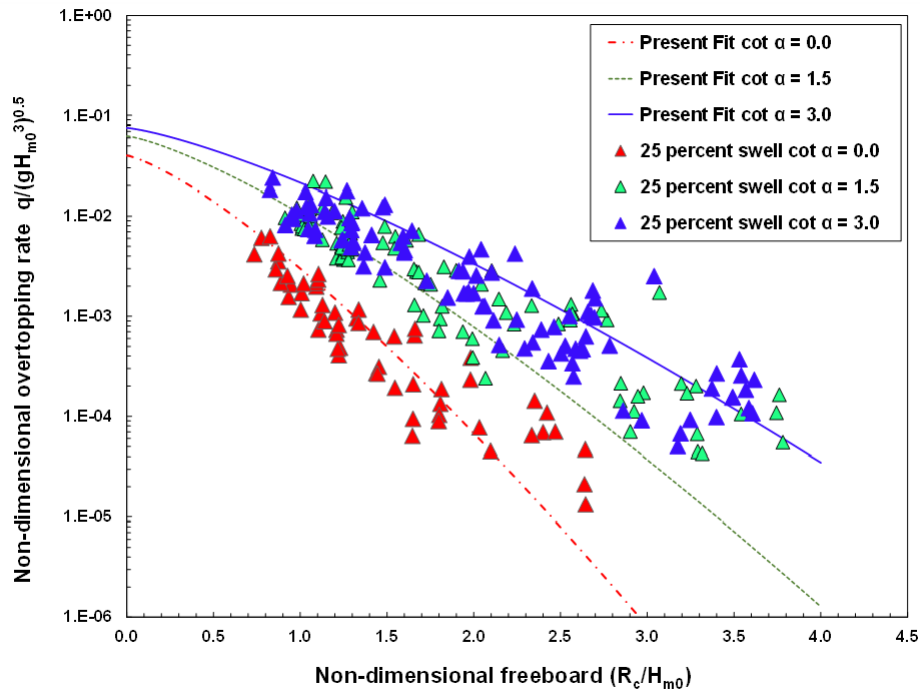


Figure 7.11 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for bimodal seas with 25 percent swell

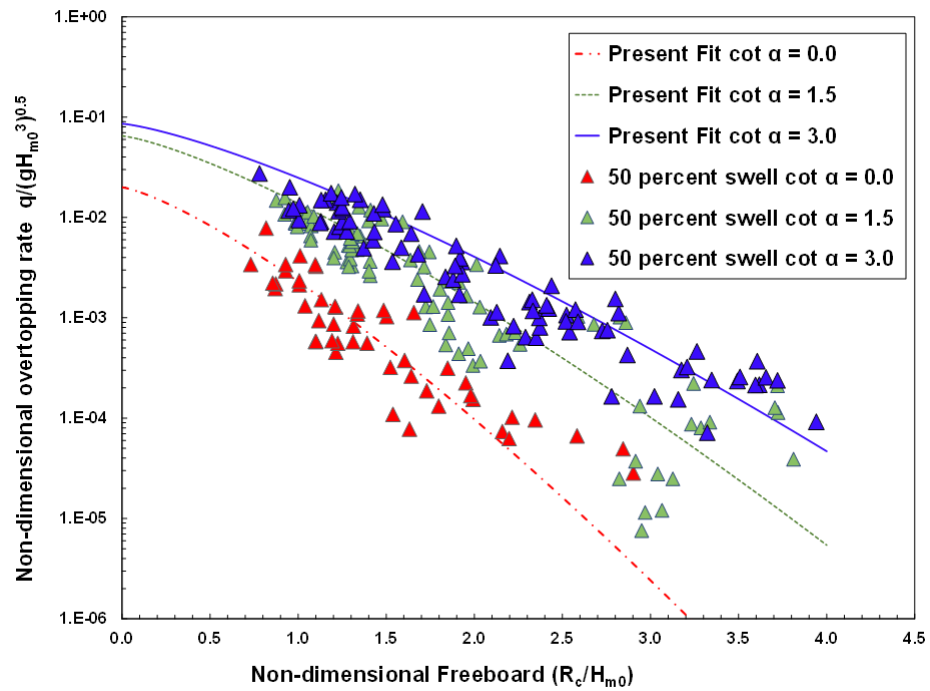


Figure 7.12 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for bimodal seas with 50 percent swell

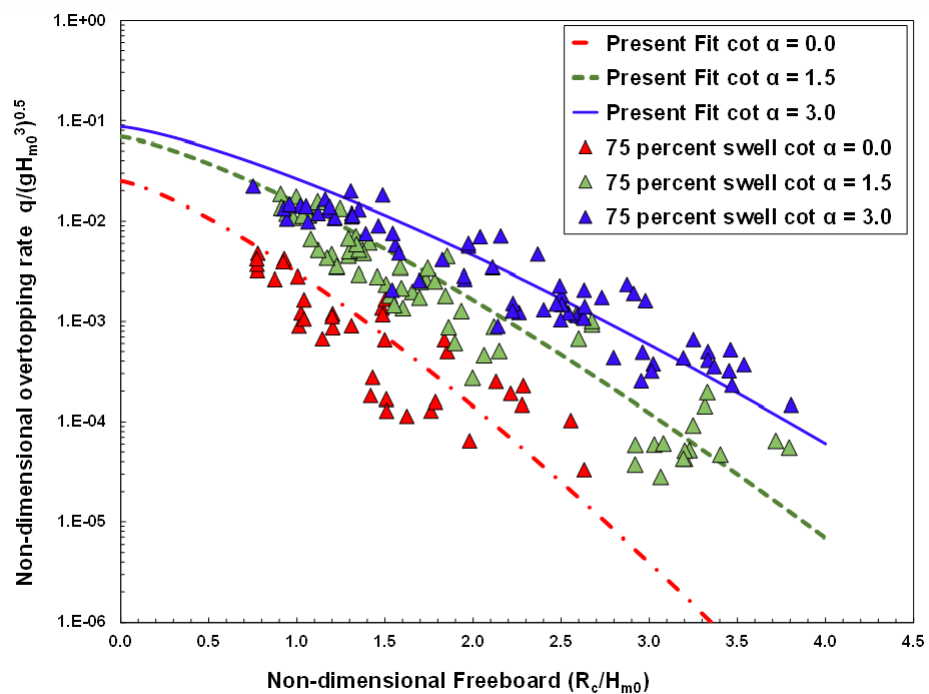
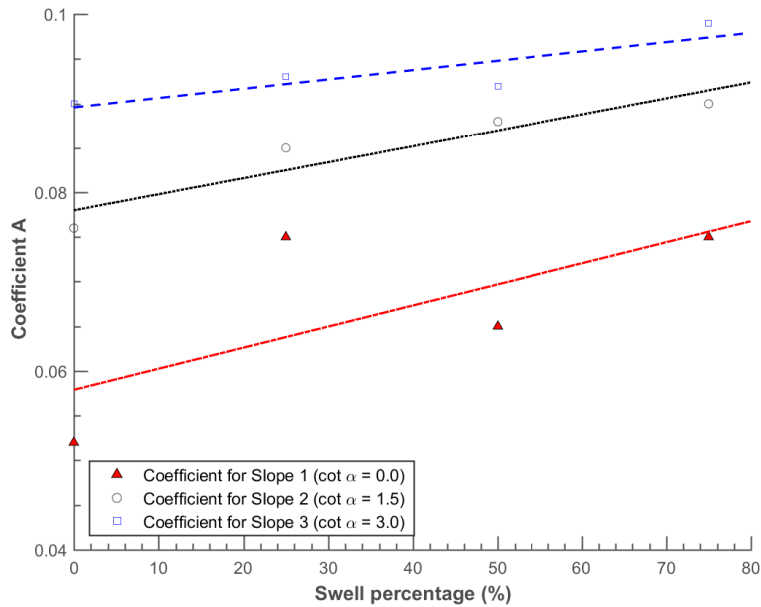
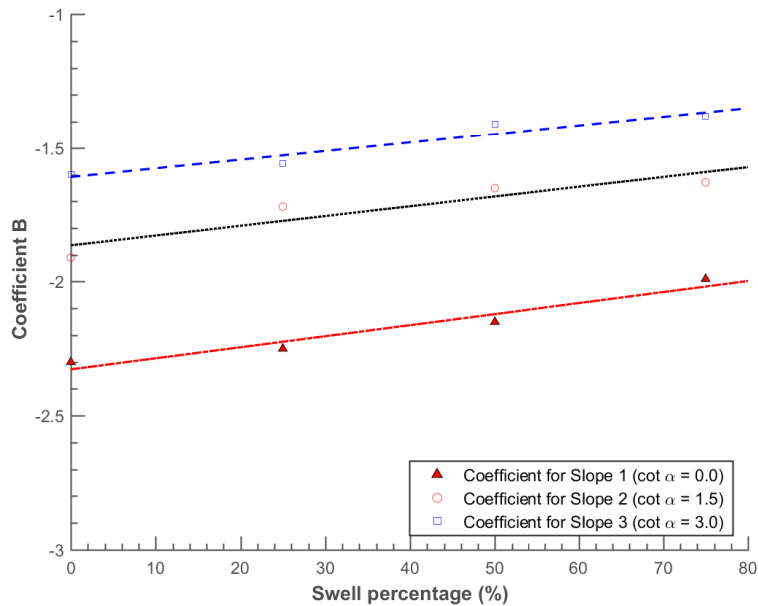


Figure 7.13 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for bimodal seas with 75 percent swell

Figure 7.14 represent plots of the coefficient values of A and B obtained as a result of the non-linear regression fitted on different percentages of swell obtained across different slopes. The fit for coefficient B has little scatter whereas the fits for coefficient A have wider scatter.



(a)



(b)

Figure 7.14 (a) Coefficient A as a function of swell percentage, varying from 0 to 75 percent, (b) Coefficient B as a function of swell percentage, varying from 0 to 75 percent

The linear variation of A and B with respect to swell percentage can be expressed as:

$$A_s = A + m_1 S_0 \quad (7.11)$$

$$B_s = B + m_2 S_0 \quad (7.12)$$

where values of A and B follow the same definition provided in EurOtop (2018) for different impermeable sloping seawalls whose values are dependent on structure slope $\cot \alpha$. m_1 and m_2 are defined by the slope of the lines in Figure 7.14. A modified version of Equation (2.4) based on the regression against our new observations, which accounts for bimodal conditions may be written as:

$$\frac{q}{\sqrt{gH_{m0}^3}} = A_s * \exp \left(-B_s \frac{R_c}{H_{m0}} \right)^{1.3} \quad (7.13)$$

where the values of A_s can be empirically redefined in terms of swell percentage S_0 as:

$$A_s = \begin{cases} A + 0.0002S_0 & \text{for } \cot \alpha \leq 1.5 \\ 0.09 + 0.0001S_0 & \text{for } \cot \alpha \geq 1.5 \end{cases} \quad (7.14)$$

and, for coefficient B_s :

$$B_s = \begin{cases} B + 0.0004S_0 & \text{for } \cot \alpha \leq 1.5 \\ -1.609 + 0.0003S_0 & \text{for } \cot \alpha \geq 1.5 \end{cases} \quad (7.15)$$

It must be noted that Equations (7.11) – (7.15) are defined to cater for the effect of swell percentage on the overtopping computed for bimodal cases only. The applicability is limited to slopes within the ranges shown ($0 \leq \cot \alpha \leq 3.0$). The usage also does not contradict the use of Equations (2.4) – (2.6) above. Coefficients for different slopes are calculated using the suggestions of Victor and Troch (2012) and van der Meer and Bruce (2013) in Equations (2.4) – (2.6), and then determine the equivalent effects due to swell percentage by applying Equations (7.11) – (7.15). It is because swell percentages follow a linear increasing relationship with the dimensionless overtopping. The trends of this slope are the same for both $\cot \alpha = 0$ and $\cot \alpha = 1.5$. However, there is a significant drift in this case for $\cot \alpha \leq 3.0$. In applying the new formulations proposed

in present study, various coefficients due to swell effects can be interpolated for other slope angles that are within with the limit discussed in this study.

7.4 Assessment of New Prediction Accuracy

Table 7.1 presents a follow up to the prediction of accuracy that was previously discussed in Section 6.5. The root mean square errors RMSE and the bias of the errors BE are both applied were determined using Equations (6.2) – (6.3) to assess the accuracy provided by the EurOtop (2018) and compared with the new swell formulations that was represented by Equations (7.13) – (7.15). The new formulation presented a better accuracy in predicting both the unimodal and bimodal dimensionless overtopping across the three slopes. Equation (7.13) – (7.15) shows an improvement in the prediction of bimodal overtopping values in terms of the RMSE and the bias values described by the new formulation.

Table 7.1 Values computed for *RMSE* in Equation (6.2) and bias Equation (6.3) by assessing the EurOtop for different ranges of relative crest freeboard $R_c = H_{m0}$ and slope angles

Formulation	Equation (2.4)		Equation (7.13)	
Item	RMSE (-)	Bias (-)	RMSE (-)	Bias (-)
$\cot \alpha = 3.0$				
All tests	0.00365	-7.46×10^{-6}	0.00271	6.10×10^{-6}
Unimodal	0.00284	8.82×10^{-4}	0.00221	6.34×10^{-4}
Bimodal	0.00463	-7.41×10^{-5}	0.00376	-4.95×10^{-5}
$\cot \alpha = 1.5$				
All tests	0.00336	-1.69×10^{-3}	0.00260	-1.36×10^{-3}
Unimodal	0.00191	1.36×10^{-4}	0.00135	1.06×10^{-4}
Bimodal	0.00304	-5.18×10^{-4}	0.00228	-3.78×10^{-4}
$\cot \alpha = 0.0$				
All tests	0.00972	-3.84×10^{-4}	0.00609	2.84×10^{-4}
Unimodal	0.00103	5.80×10^{-4}	0.000792	4.58×10^{-4}
Bimodal	0.00967	-3.65×10^{-4}	0.00628	3.22×10^{-4}

7.5 Sensitivity Analysis of the Fitted Coefficients

It was observed from Figure 7.14a that the coefficient A described for the vertical slope seems to pose certain uncertainties among other fitness parameters. To examine the effect of the uncertainty of that gradient on the overtopping predictions, sensitivity analysis were performed.

The Monte carlo sensitivity analysis are based on investigating various influences of the uncertainty of the input will pose on the output overtopping predictions. In this way the reliability and the robustness of the overtopping prediction can be improved. Different methods of performing sensitivity analysis have been fully described in Iooss and Lemaître (2015). In present study, the Monte Carlo analysis was adopted. This method allows the use of the uncertainty distributions obtained from different inputs into the model or determined from the model to establish effects of any of the uncertainties in the parameter. In this thesis both normal and exponential probability distributions are applied.

Figure 7.15(a–c) presents the graphical and statistical analysis obtained from the Monte Carlo sensitivity analysis. In all the three figures, the normal distributions are based on 3000 Monte Carlo simulations. Figure 7.15 shows the distribution of the sensitivities of all the parameters in the new formulation combined, while Figures 7.15(b & c) represent the normal distribution obtained using the parameters of coefficient A due to $\cot \alpha = 0$ alone (the lowest gradient in Figure 7.14a). The process performance capability Ppk accounts shown in Figure 7.15a accounts for the overall variation in the simulation results. As shown here the overall variation of 1.322 was recorded using all the combined parameters. Only 0.1 percent are out of specification. According to Verbeeck et al. (2006), the Ppk shows the drifts and shifts between each of the parameters. Ppk depends largely on the values of the standard error obtained during the simulation. A smaller value is obtained in this case. This result is similar to the effects of $\cot \alpha = 0$ (shown in Figures 7.15(b & c)). Smaller values of Ppk (1.343 and 1.35) are similarly obtained in both cases. These values are shown on the normal distributions.

The accurate reproduction of the normal distribution when effects due to $\cot \alpha = 0$ are applied to Equations (7.13) – (7.15) also shows that the distributions are not strongly sensitive to the effects of the vertical seawall significantly. Moreover, a low percentages of the distribution are out of specification with only 0.121 % and 0.118 respectively. Due to the similarities of the normal statistical distribution, and little variations observed in the Ppk and the percentages that are out of specification, it can be concluded that effects of $\cot \alpha = 0$ does not have significant effects on the fitness parameters applied in Equations (7.13) – (7.15).

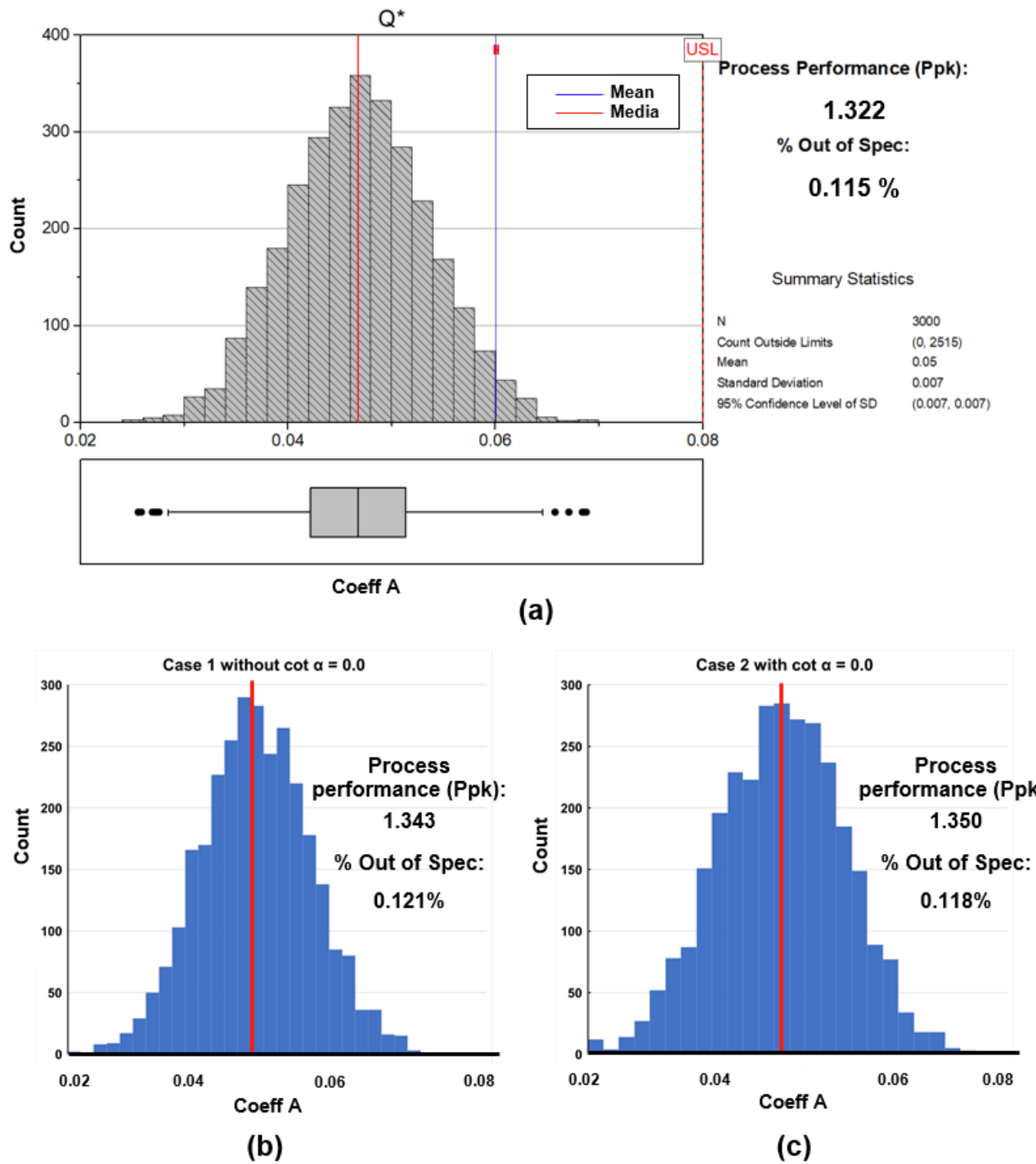


Figure 7.15 Sensitivity analysis of the distribution obtained from Monte Carlo simulation performed on (a) all the coefficient parameters, (b) due to the effect of $\cot \alpha = 0$ on both coefficients A and B and (c) due to the effect of $\cot \alpha = 0$ on both coefficients A alone within the new overtopping formulations in Equations (7.13) – (7.15)

7.6 Influence of Swell Spectral Peak Period

Figure 7.16 presents the influence of swell peak period that was investigated in this study. The box plots present similar information as described in Section 7.3. In the present Section, at prototype scales, only swells with peak periods of 11, 15, 20 and 25 seconds are respectively considered. As visible in Figure 7.16, non-dimensional wave overtopping are strongly influenced by the peak periods (or frequencies) of swell in the energy spectrum for the same sea-state. These effects could be significant in coastal seawall designs and costs for certain locations. Non-dimensional wave overtopping seems to respond strongly to the separation of spectral peaks between the wind wave and swell wave frequencies in the bimodal energy spectrum. There seems to be clearer wave overtopping being triggered by swell periods at 11 seconds than the rest for both $\cot \alpha = 0$ and 1.5. This effect does not seem to be very clear for $\cot \alpha = 3.0$ except for few extreme non-dimensional wave overtopping estimates shown by the outliers.

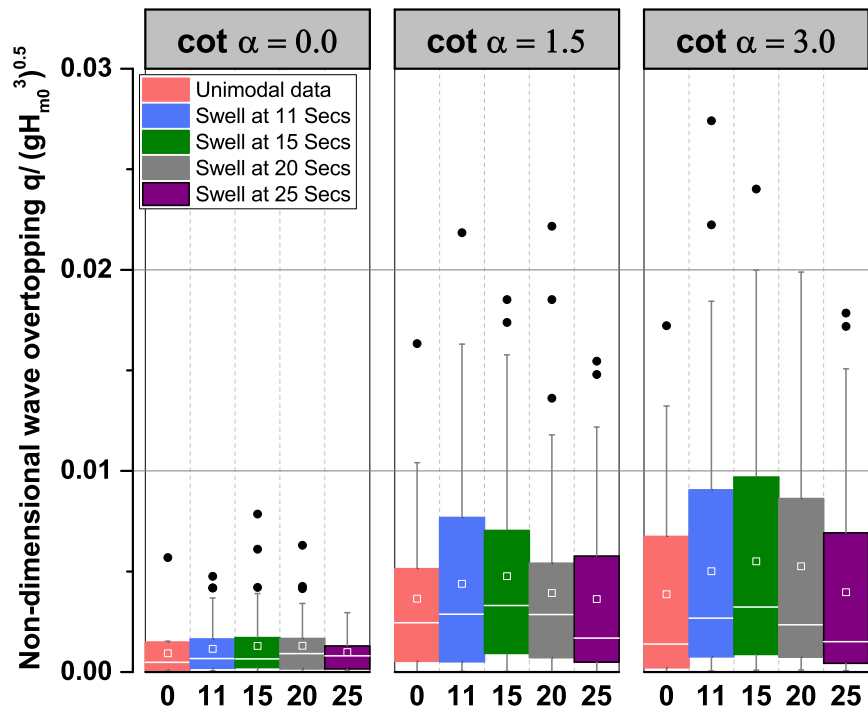


Figure 7.16 The relationship between the non-dimensional overtopping rate $q/\sqrt{H_{m0}^3}$ with Swell peak periods of $\cot \alpha = 0.0$ and $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$

In order to obtain a suitable empirical relationship between the non-dimensional wave overtopping and swell spectral peak periods, the non-linear regressions are applied to fit different dimensionless wave overtopping estimates for different swell periods.

Non-linear least squares regression analysis was obtained as a result of the fitting process on each sets of data. Plots of fitted results are presented in Figures 7.17–7.20.

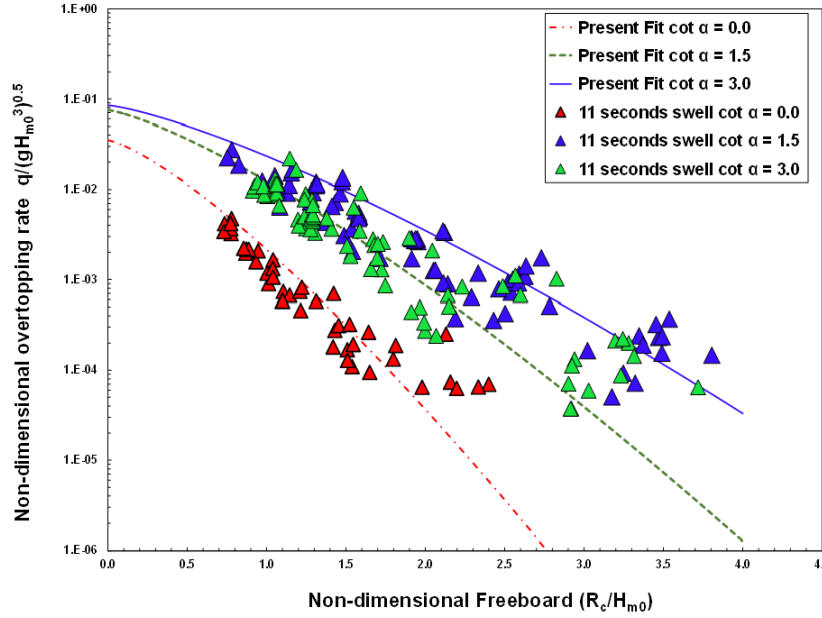


Figure 7.17 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for swell with spectral peak periods at 11 secs

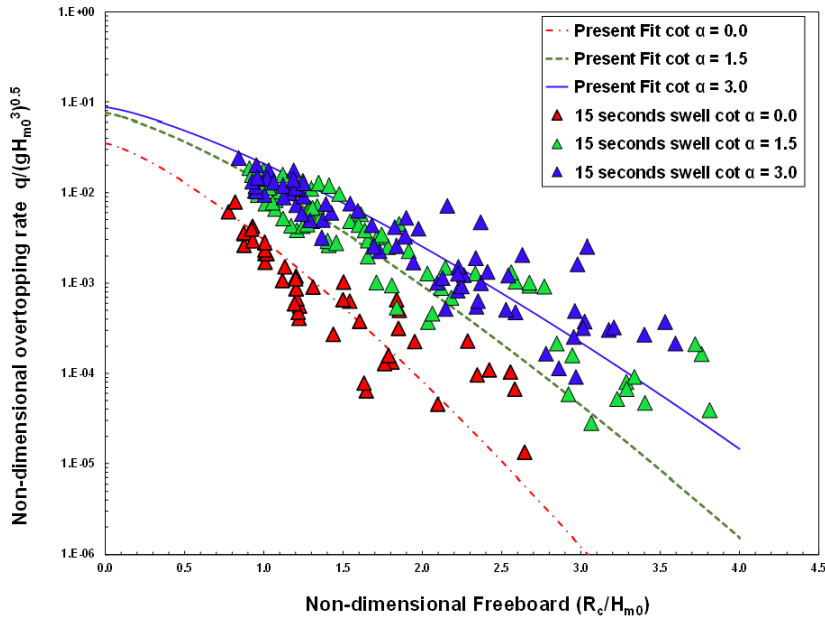


Figure 7.18 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for swell with spectral peak periods at 15 secs

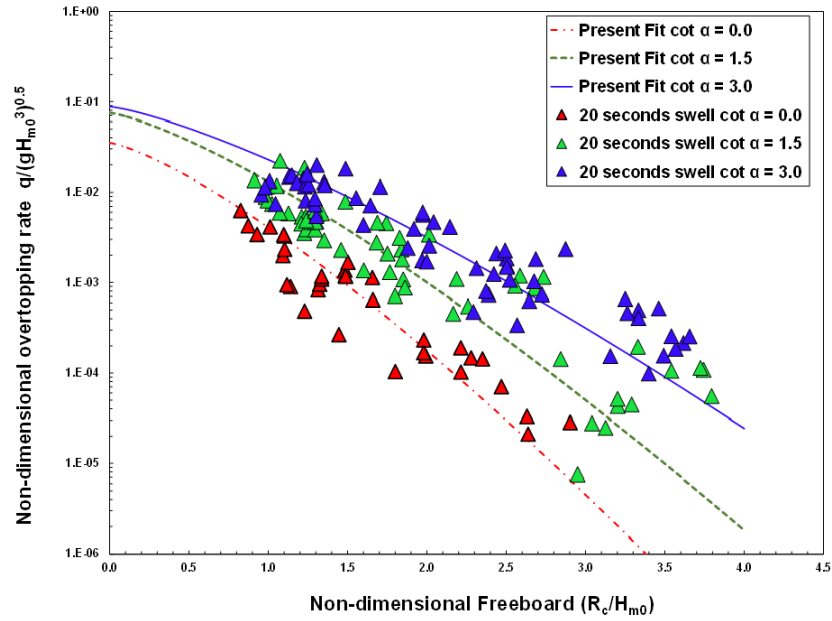


Figure 7.19 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for swell with spectral peak periods at 20 secs

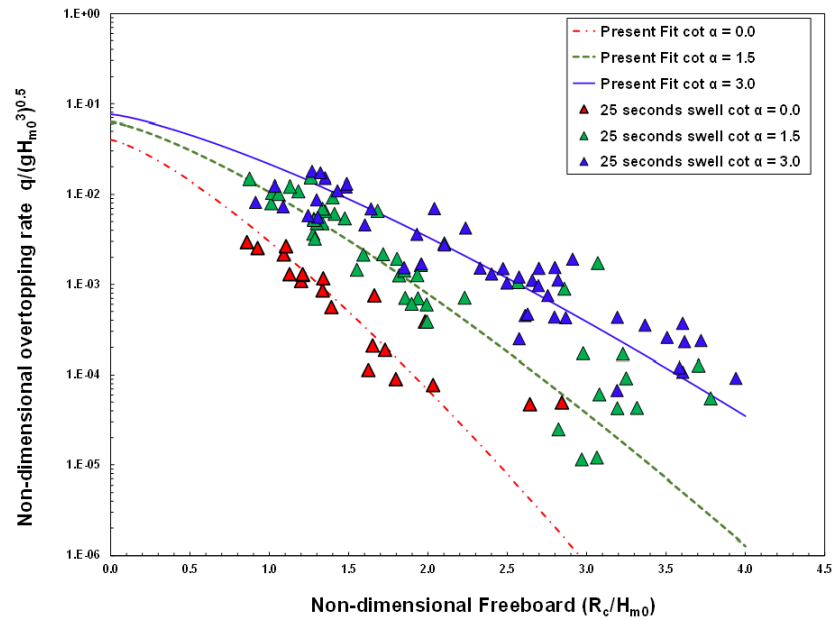
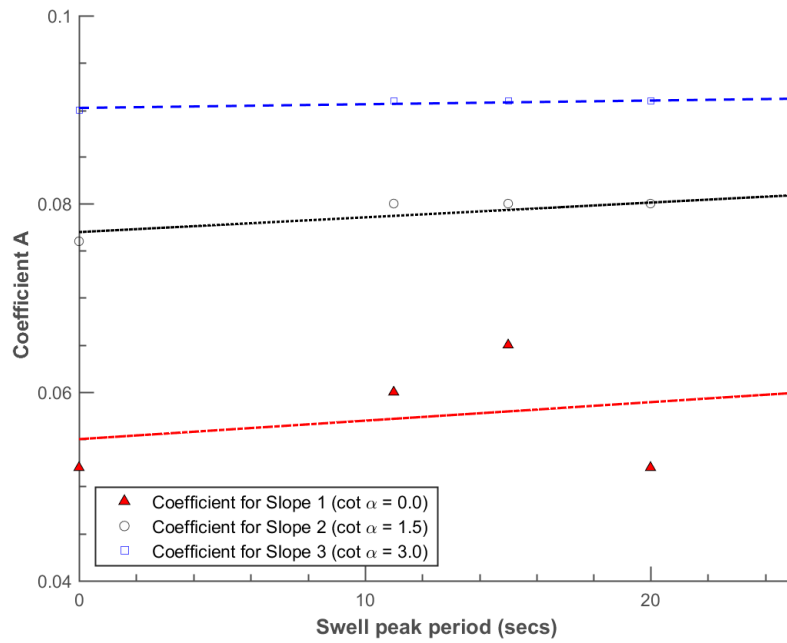
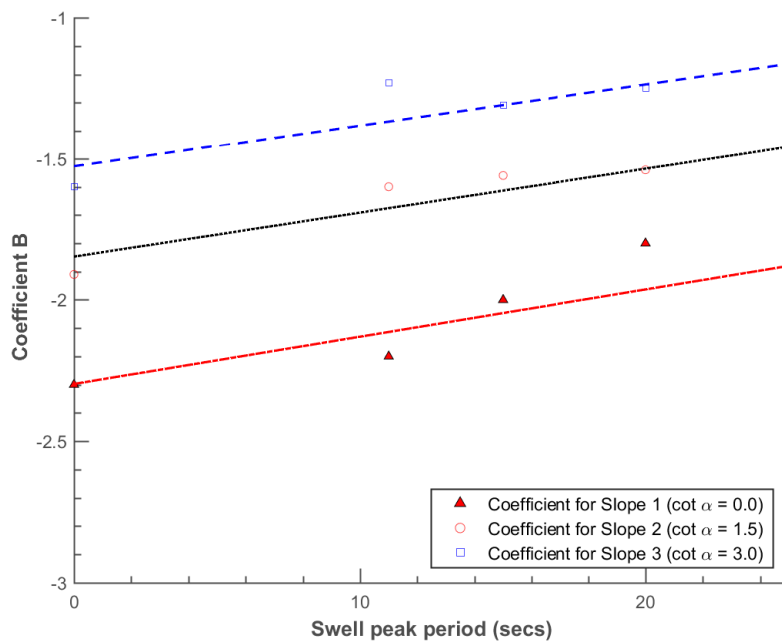


Figure 7.20 Results of non-linear regression performed on different non-dimensional wave overtopping obtained across three slopes for $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ for swell with spectral peak periods at 25 secs

Based on these fittings, individual coefficients A and B can be extracted as before. These are presented in Figure 7.21.



(a)



(b)

Figure 7.21 (a) Coefficient A as a function of swell peak period, varying from unimodal to 25 secs, (b) Coefficient B as a function of swell peak periods, varying from unimodal to 25 secs

Concerning effects of swell peak period $T_{Sm-1,0}$, as the peak separation increases, the spectral width of the resulting bimodal spectrum also increases and in turn, creating larger overall spectra wave period $T_{m-1,0}$ and the resulting wave length $L_{m-1,0}$. Also as described earlier, the sensitivity of overtopping is greater for the mildest slope. As shown in Figure 7.21, the highest wave overtopping performances were demonstrated by this sloping seawall $\cot \alpha = 3.0$. Unlike in the unimodal cases, lower frequency swells produce greater wave overtopping than swells at higher frequencies because the latter produces smaller overall spectra characteristics than the former.

7.7 Conclusions

In this chapter, different influences of the presence of swell in a sea-state have been described and analysed. Various parameters of spectra characteristics considered are the narrowness, the broadness, and peakedness parameters. Effects of swell percentages and swell peak periods on overtopping have also have been shown. Swell waves occurring at lower peak periods are producing the uppermost impact of wave overtopping on each slope investigated. However, it was hypothesise that the increased wave overtopping for longer waves/swell percentage is due to a change in wave behaviour, where more waves are surging than breaking on the structure. A simple modification to the existing formula for non-breaking waves is proposed that accounts for the presence of swell percentages in bimodal conditions. Further study is required on the non-linearities imposed by swell peak periods to have a more suitable formulation of this phenomena. Wave overtopping increases as the spectral values described by the narrowness, the broadness, and peakedness parameters reduces. Moreover, spectra values tend to be approaching asymptotic values for both the narrowness and the broadness parameters. Increasing trends of swell percentage in a bimodal sea could produce enormous wave overtopping with tendencies of more significant flooding. Broad-banded sea conditions yield less wave overtopping intensity than highly peaked bimodal seas.

Chapter 8

Reflection Analysis of Impermeable Slopes Under Bimodal Sea Conditions

Follow-up to the experimental test programme conducted in chapter 6, understanding of reflection characteristics of coastal seawalls is crucial for seawall designing and management. Wave reflection can cause difficulties to small vessel manoeuvring at the harbour entrance and cause damaging scouring at the toe of coastal structures. Previous studies have considered reflection characteristics of coastal seawalls, under wind-generated random waves, but have neglected to consider the effects of wave bimodality created by the presence of swell waves. Chapter 8 will focus on the influence of random wave bimodality on reflective characteristics of coastal seawalls. Reflection coefficients were computed from each storm in the experimental test. Analysis of results suggests that both unimodal and bimodal waves give similar reflection characteristics. However, a higher reflection coefficient in bimodal sea states than in the unimodal sea states. It was found that the reflection coefficient of coastal seawalls is strongly influenced by the seawall slope, the wave steepness, relative water depth, and the surf similarity parameters. A new empirical reflection equation to describe the influence of wave bimodality on the reflection characteristics of coastal seawalls has been formulated based on this study.

8.1 Reflection Characteristics of Smooth Impermeable Slopes

Battjes (1974) identified the surf similarity parameter ξ as a critical parameter affecting the reflection characteristics of sloping impermeable slopes with slope angle α under

incident monochromatic waves. It can be expressed mathematically as:

$$\xi = \frac{\tan \alpha}{\sqrt{2\pi H/gT^2}} \quad (8.1)$$

In this equation, H , and T represent the significant wave height and the wave period of the monochromatic wave, respectively. Values of $\xi \leq 2.3$ correspond to breaking waves while for $\xi \geq 2.3$ non-breaking waves occur. The equation showing a simplified relationship between the reflection coefficient K_r and the breaker parameter ξ of breaking monochromatic waves has also been described in the same study (Battjes (1974)), and is:

$$K_r = 0.1 \xi^2 \quad (8.2)$$

This expression is only valid for breaking monochromatic waves $\xi \leq 2.3$ as illustrated by Figure 8.1. In Seelig and Ahrens (1981), a modified version of Equation (8.2) has been presented. The modified relationship between the reflection coefficient K_r and the breaker parameter ξ is presented in Equation (8.3):

$$K_r = \tanh(0.1\xi^2) \quad (8.3)$$

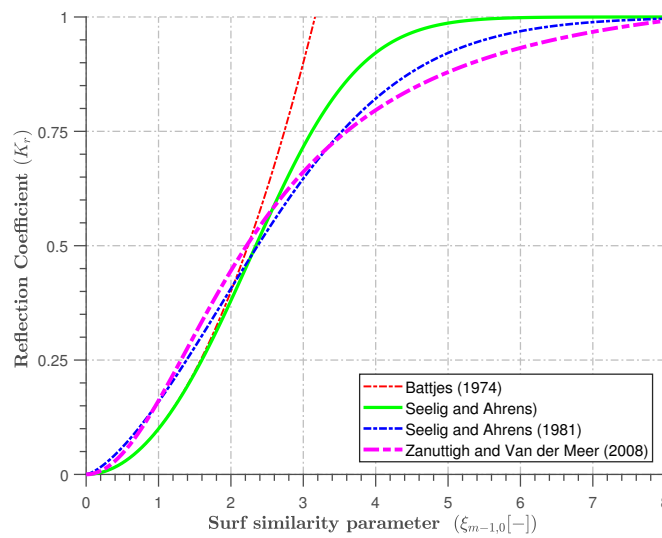


Figure 8.1 An illustration of the prediction of reflection characteristics of smooth impermeable slopes

It is worth noting that Equation (8.2) is a close approximation to Equation (8.3) for small values of the surf similarity parameter. For larger values of the parameter, Equation (8.3) tends asymptotically to 1. The mathematical expression from Seelig and Ahrens (1981) is valid for both breaking and non-breaking monochromatic waves. A more accurate expression was also proposed based on several experimental studies that were performed to describe the reflection behaviour of other sea defences, including revetments, beaches and breakwater. These expressions consider both the peak period of irregular waves and the spectrally determined breaker parameter $\xi_{m-1,0}$. These are described in Equation (8.4)a and (8.4)b below:

$$K_r = \frac{\xi_p^2}{\xi_p^2 + 5.5} \quad (8.4a)$$

$$K_r = \frac{\xi_{m-1,0}^2}{\xi_{m-1,0}^2 + 5.0} \quad (8.4b)$$

Equation (4b) is valid for $1.0 \leq \xi_{m-1,0} \leq 6.2$ and sloping seawalls with $1.5 \leq \cot \alpha \leq 2.5$ respectively.

Zanuttigh and Martinelli (2008) proposed a revised version of Equation (8.4)b from over 4000 reflection coefficient experimental test results. This can be generally expressed as:

$$K_r = \tanh 0.16(\xi^{1.43}) \quad (8.5)$$

This equation extends the range of applicability to structure slopes with $1.5 \leq \cot \alpha \leq 4.0$ and $1.0 \leq \xi_{m-1,0} \leq 4.1$ with dimensionless crest freeboard of $0.58 \leq R_c/H_{m0} \leq 4.5$.

The reflection analyses by Mansard and Funke (1980); Zelt and Skjelbreia (1992) apply strictly to linear waves. For the breaking wave cases, the non-linearity of the reflection performances cannot be accurately estimated by the method. The accuracy of the Zelt & Skjelbreia method was determined using the flume without any structure. The free reflection characteristics of the open flume were used to calibrate the performance of the method of Zelt and Skjelbreia (1992). These adjustments are applied to all wave cases studied in this thesis.

8.1.1 Estimation of Reflection Parameters

Some reflection parameters required for this study were estimated. These include the linear wave length, wave steepness and the dimensionless Iribarren number. Relationships obtained from these parameters are of special relevance to this study. The linear wave length $L_{m-1,0}$ applied here is calculated using the Newton-Raphson iteration technique on the dispersion relationship, for example in Reeve et al. (2015).

$$L_{m-1,0} = \frac{gT_{m-1,0}}{2\pi} \tanh\left(\frac{2\pi h}{L_{m-1,0}}\right) \quad (8.6)$$

In Equation (8.6), h is the offshore water depth, and $T_{m-1,0}$ represents the spectra wave period. Also, the wave steepness $S_{m-1,0}$ can be defined in terms of the dimensionless ratio of the spectral wave height H_{m0} and the wave length $L_{m-1,0}$ obtained from Equation (8.7).

$$S_{m-1,0} = \frac{H_{m0}}{L_{m-1,0}} \quad (8.7)$$

Similarly, the surf similarity parameter $\xi_{m-1,0}$ can be estimated from Equation (8.8):

$$\xi_{m-1,0} = \frac{\tan\alpha}{\sqrt{S_{m-1,0}}} \quad (8.8)$$

Other relevant parameters are the non-dimensional wave height H_{m0}/h and non-dimensional water depth $d/L_{m-1,0}$. Full details of equations solved to determine the reflection coefficient K_r can be found in Zelt and Skjelbreia (1992) and Zelt and Skjelbreia (1992). Table 6.3 presents the wave conditions tested in this study. As shown in the table for each test numbers, there are four different swell peak periods across three different sloping seawalls.

8.2 Results and Discussions

To derive a functional improvement from the reflection coefficient K_r of impermeable walls under bimodal sea conditions, effects of various secondary factors influencing K_r will be considered in this section. Such factors includes the wall slope, water depth, wave steepness and the crest freeboard.

8.2.1 Influence of Wall Slope on Reflection Characteristics

Reflection characteristics of a coastal seawall can be determined by the steepness of the wall slope. Combined plots of the relationship between the reflection coefficient and the breaker index parameters are presented in Figures 8.2 and 8.3. Figure 8.2 presented results for unimodal sea-states while reflection results due to bimodal sea-states are represented in Figure 8.3. In general, for unimodal and bimodal sea conditions, K_r varies between 0.4 and 1.038. As expected, lower values of K_r are observed for the gentle slope of 1:3, while higher of K_r are clearly observed for the case of vertical seawall. These results are consistent with the findings of Seelig and Ahrens (1981) and Allsop and Hettiarachchi (1989).

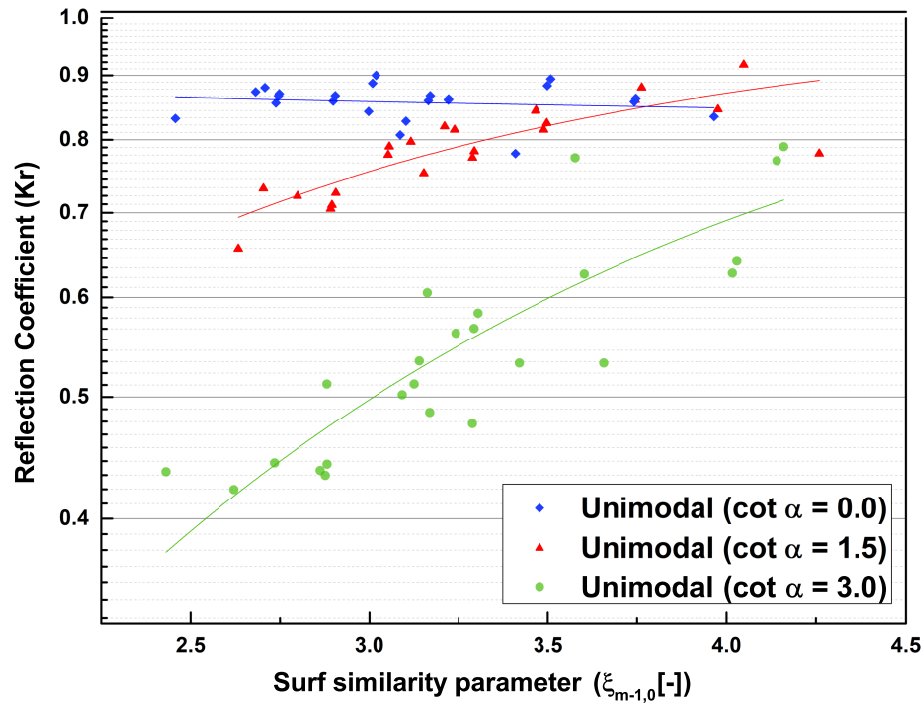


Figure 8.2 Comparison of the reflection coefficient K_r and breaker parameters $\xi_{m-1,0}$ across three slopes ($\cot \alpha = 1.5$, 3.0 and 0.0) as shown under unimodal sea states

The range in K_r for a vertical wall depends largely on the degree of wave overtopping, and it increases as the crest freeboard increases (Figures 6.14, 6.15 and 6.16). For sloped seawalls, the K_r is increases as the $\xi_{m-1,0}$ increases for the incident waves. This increments tends towards an asymptotic values for waves with longer periods. In bimodal seas, higher bound of K_r is created which corresponds to swell of long period exhibited by the bimodal seas. The case is different with a vertical seawall as K_r values are generally around 100 percent as expected.

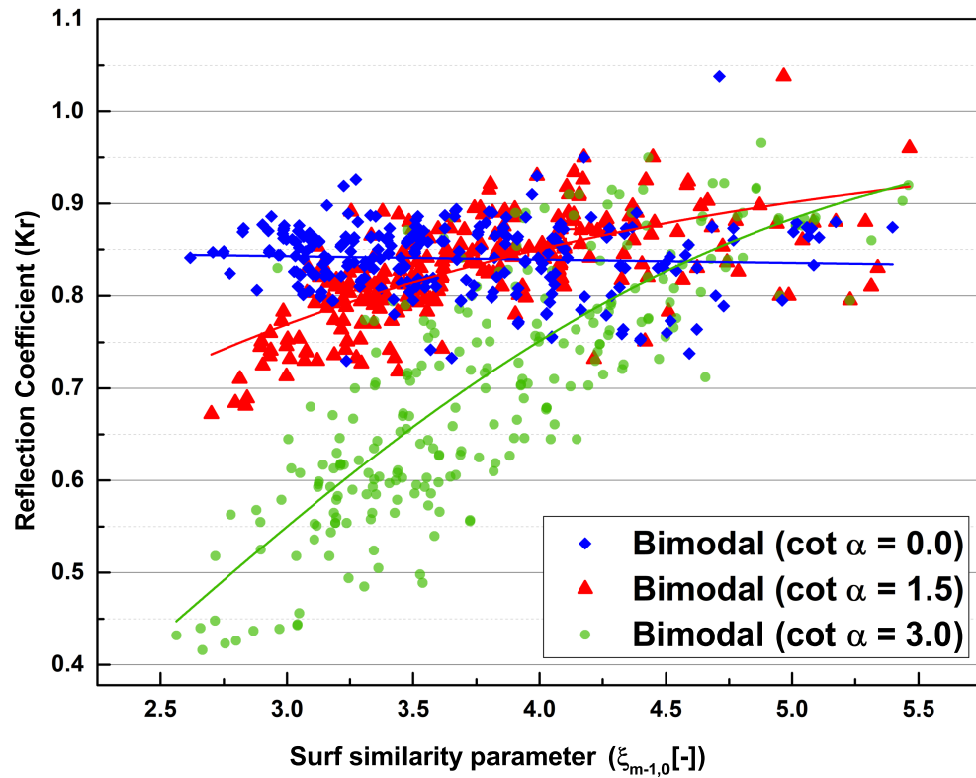


Figure 8.3 Comparison of the reflection coefficient K_r and breaker parameters $\xi_{m-1,0}$ across three slopes (cot $\alpha = 1.5$, 3.0 and 0.0) as shown under bimodal sea conditions

Figure 8.4 shows a combined plot of the best-fit curves of K_r against $\xi_{m-1,0}$ for both unimodal (solid lines) and bimodal (in dotted lines) across different slopes. It is obtained by applying the non-linear fit algorithm to all the datasets under unimodal and bimodal seas. Unimodal sea states exhibited similar trends as the bimodal sea conditions under the same test conditions. However, the relationship between K_r and $\xi_{m-1,0}$ are more extended in the bimodal sea states than in the unimodal sea conditions. This is because of the longer periods in the bimodal seas (Figure 7.8a). As the wave period increases, there is a further reduction in wave steepness while values of $\xi_{m-1,0}$ increases. It follows that values of K_r are inversely proportional to the wave steepness as reported in previous studies, for example, Seelig and Ahrens (1981).

Least reflections are observed for the gentle slope (cot $\alpha = 3.0$), which is in accordance with the observations of Neelamani and Sandhya (2003); Van Gent (1994), and is a result of the wave energy dissipation associated with energy losses due to wave overtopping.

Further examination of these results shows that the relationships between K_r and $\xi_{m-1,0}$ for unimodal state obtained in this study are within the bounds of previous

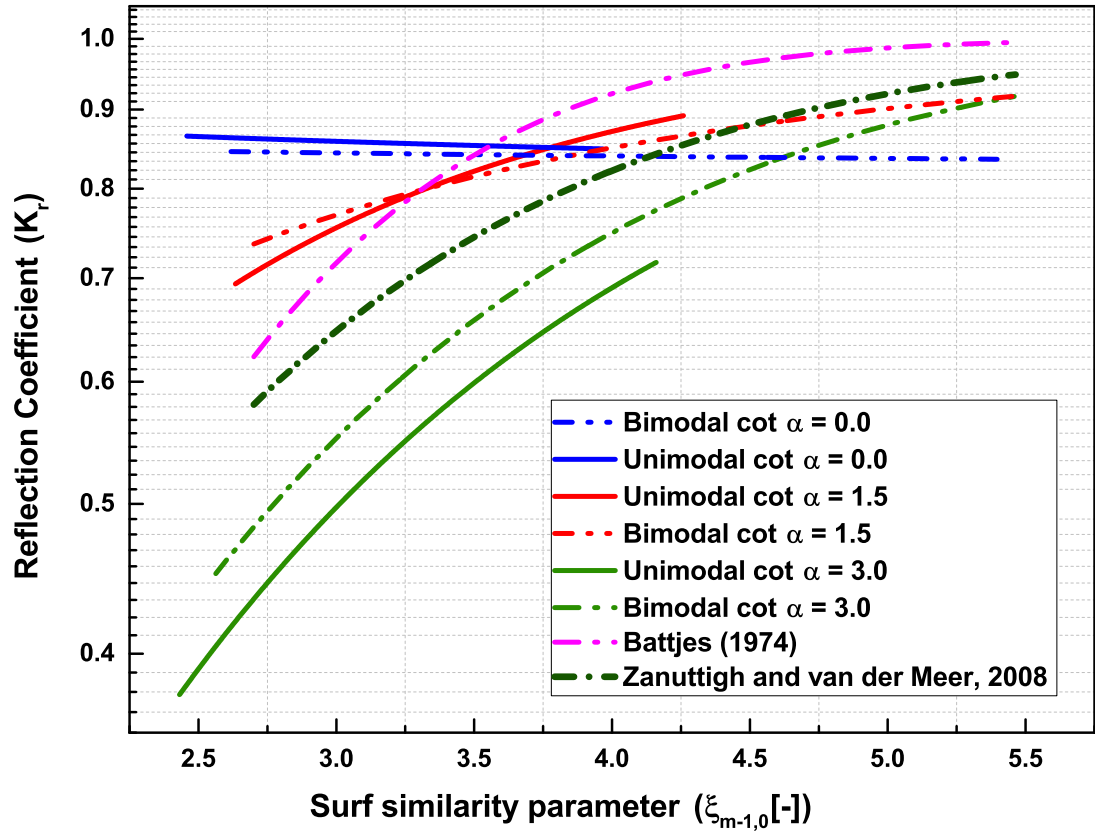


Figure 8.4 Variations of reflection coefficient K_r with breaker index $\xi_{m-1,0}$ across three different slopes

studies derived by Battjes (1974) and Zanuttigh and Martinelli (2008). However, a better fit is obtained from the modified K_r formulation derived by Zanuttigh and Martinelli (2008). The fitted coefficients of Zanuttigh and Martinelli (2008) does not accurately fit for the relationship between K_r and $\xi_{m-1,0}$ in bimodal seas.

8.2.2 Influence of Water Depth Variations

In this section the influence of water depth, h/L , on reflection coefficient, K_r , was investigated across three slopes (i.e., $\cot \alpha = 0.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$). Figure 8.5 presents these results for all the datasets acquired during this study. Generally, for the steep ($\cot \alpha = 1.5$) and the mild sloping ($\cot \alpha = 3.0$) seawalls, it is found that the value of K_r decreases with increasing relative depth h/L . This behaviour agrees with the findings of Neelamani and Sandhya (2003) and Nassar and Negm (2016). These observations are peculiar to plunging wave breaking phenomena as previously described in cited literatures. For the vertical seawall, the reflection coefficient K_r is almost independent of water depth. This is expected because of the reflection of waves by vertical walls irrespective of the depth limiting values. Standing waves are formed during these tests, and more energy is reflected than for sloping walls. These observations are similar for both the unimodal and bimodal sea states.

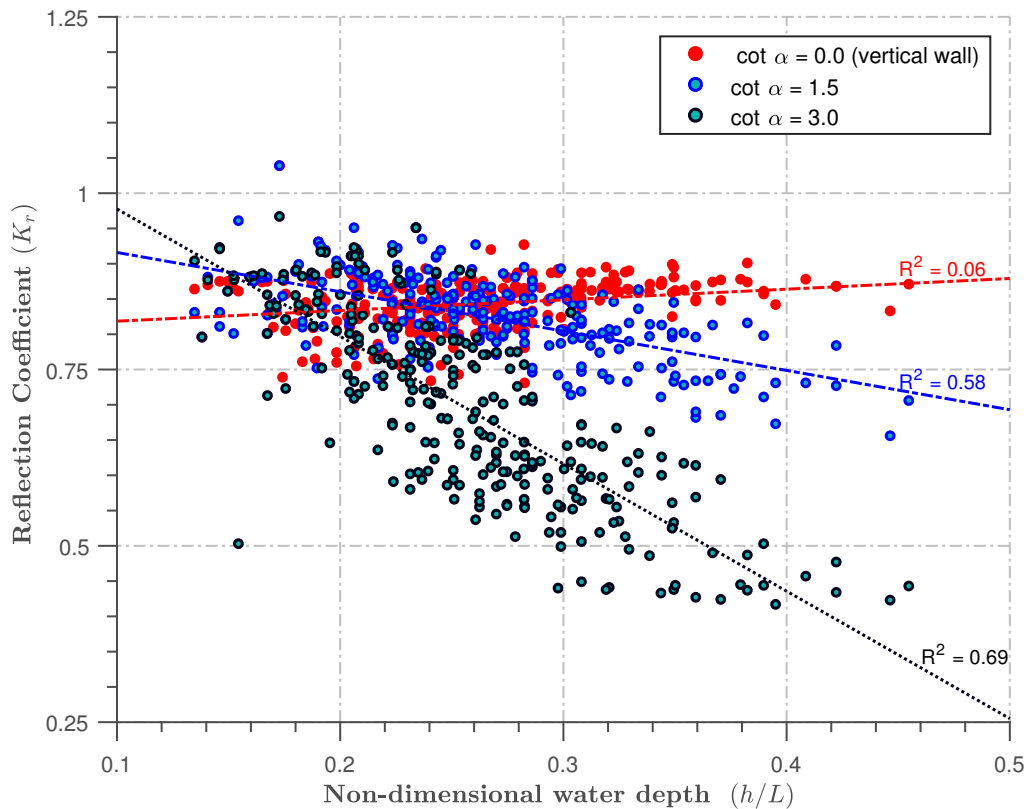


Figure 8.5 Variations of reflection coefficient K_r with non-dimensional water depth h/L across three different slopes investigated

8.2.3 Influence of Wave Steepness

Many sloping and vertical impermeable seawalls are built solely to dissipate wave energies that are directly incident on them. Reflected waves are produced whenever the waves are interacting with the plain seawalls. In sloping seawalls, it has been suggested in Goda (2010) that values of reflection coefficient K_r are inversely proportional to the incident wave steepness $S_{m-1,0}$. The suggestion is true for observations recorded in this study as shown in Figure 8.6. It forms completely similar correlation with the influence of water depth observed in previous section (8.3.2). However, the vertical seawall totally deviated from this theory in both cases. In the case of vertical seawalls, standing waves are mostly formed during propagation. This will eventually lead to cancelling out incident waves leading to destructive interference and higher K_r up to 100 percent in most cases (Figure 8.5).

The K_r performance of vertical seawall as shown in Figures 8.5 and 8.6 suggests that vertical seawalls would provide a more valuable protection to cities, harbours or ports as previously prescribed in previous studies. It is because of the high reflective characteristics that made them suitable for stopping waves from overtopping structures. However, this could in turn lead to more pronounced standing waves that can exacerbated scouring of sediments near the toe of the structure, and eventually to its failure and collapse (Zanuttigh and Martinelli, 2008).

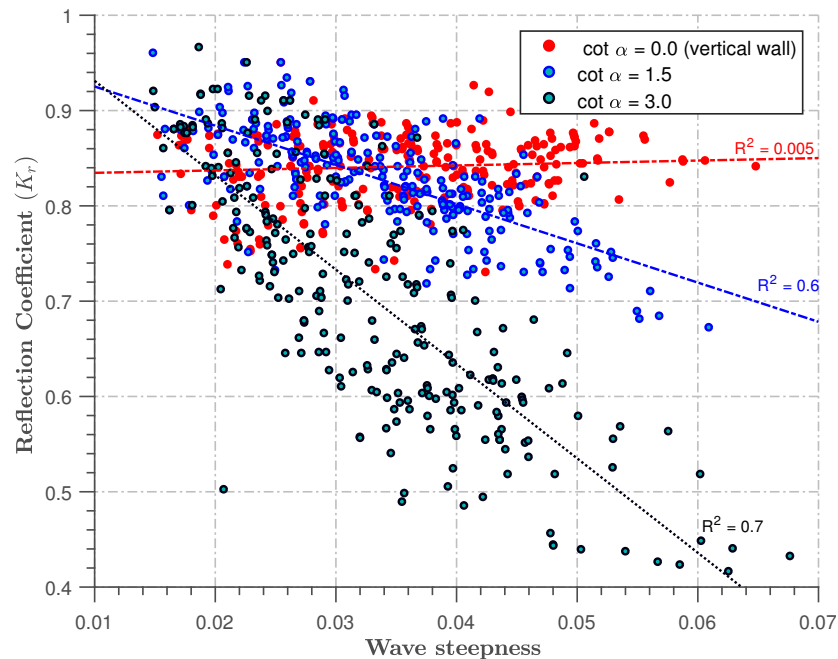


Figure 8.6 Relationships between wave steepness and the reflection coefficient K_r across three different slopes investigated

8.2.4 Effects of Crest Freeboard

Accurate selection of sizes of crest freeboard R_c/H_{m0} is an essential requirement for designing coastal seawalls against wave overtopping and to serve as flood barriers. Reliable prediction of reflection characteristics K_r of a coastal seawall suitable for selected R_c/H_{m0} is also key. In Figure 4.7, K_r is presented in terms of only R_c/H_{m0} across three seawall slopes. It can be seen that K_r is directly proportional to values of R_c/H_{m0} . As the crest freeboard increases, values of K_r also gradually increases. It implies that as crest freeboard reduces, there are more wave overtopping and tendencies for reflection reduces altogether.

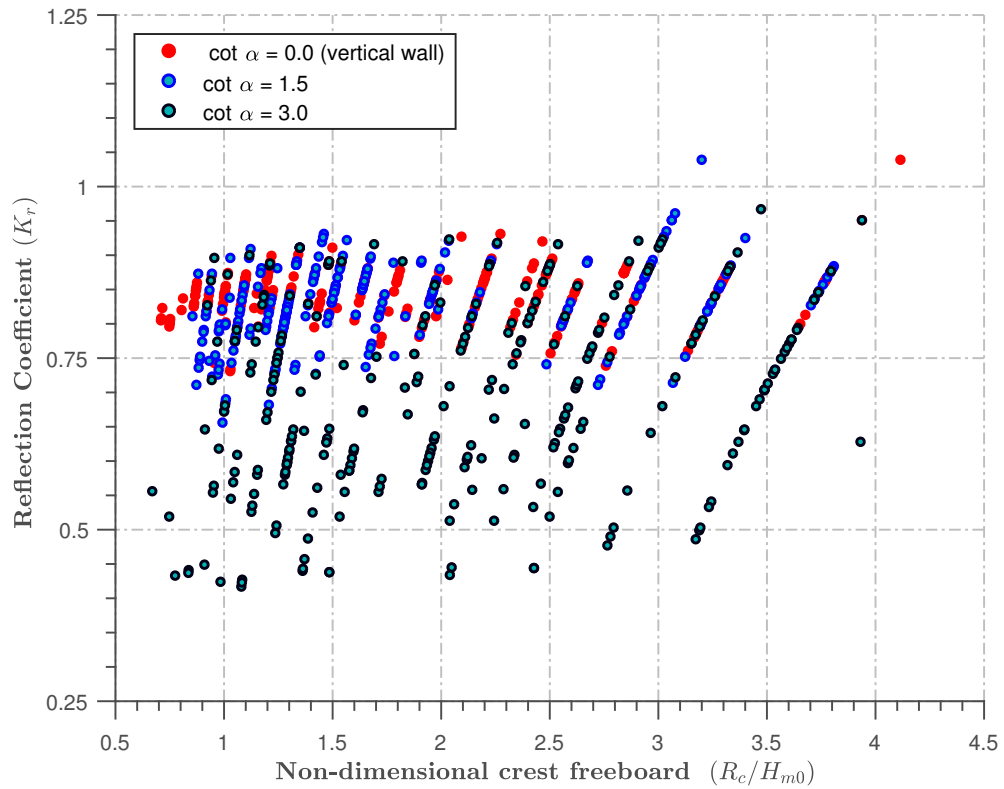


Figure 8.7 Influence of dimensionless crest freeboard R_c/H_{m0} with the reflection coefficient K_r across three different slopes investigated

8.3 Reflection Coefficients of Steep Slopes Under Bimodal Waves

In order to establish an improved formulation for the reflection coefficient K_r for bimodal sea conditions, a more detailed analysis of the results were performed. Figure 8.8 presents a more detailed relationship achieved between K_r and $\xi_{m-1,0}$ for each slope for the bimodal cases. The residuals obtained from the non-linear multi-regression fit algorithm are also presented in Figure 8.9. The residual plots show that the equation is suitable at predicting K_r under bimodal seas.

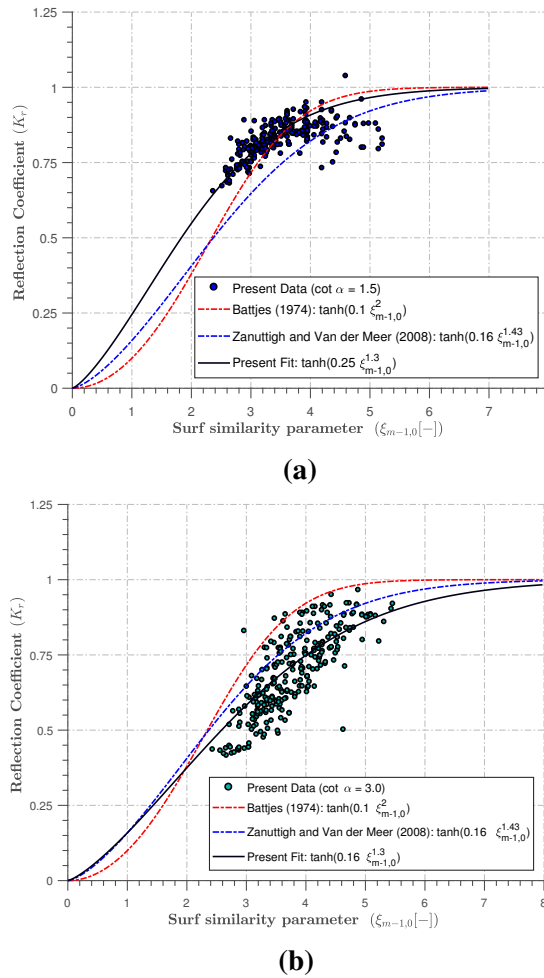


Figure 8.8 A representation of the non-linear multi-regression fit between K_r and $\xi_{m-1,0}$ for sloping seawalls

Figures 8.9(a-c) present the performance of the new formulation for the prediction of K_r of the impermeable slopes under the influence of bimodal seas. The residuals are plotted from the difference between the observed values and the predicted ones. A

positive residual values implies that the K_r predictions made by the new formulations are while on the other hand, negative values means that the new model is over predicting the K_r values. As visible from these Figure 8.9(a-c), the new equation performed very well under the bimodal sea conditions for across the three slopes. The residual plot for the mildest slope with $\cot \alpha = 3.0$ seems to give the best performance as shown in the residual plots. This may due to the possibility of obtaining the most extreme wave overtopping for this slope. The spread of the residual plots of the steeper slopes ($\cot \alpha = 1.5$) and vertical wall ($\cot \alpha = 0.0$) are symmetrically distributed showing more clustered properties around 3.5 as shown in the Figure 8.9(b-c).

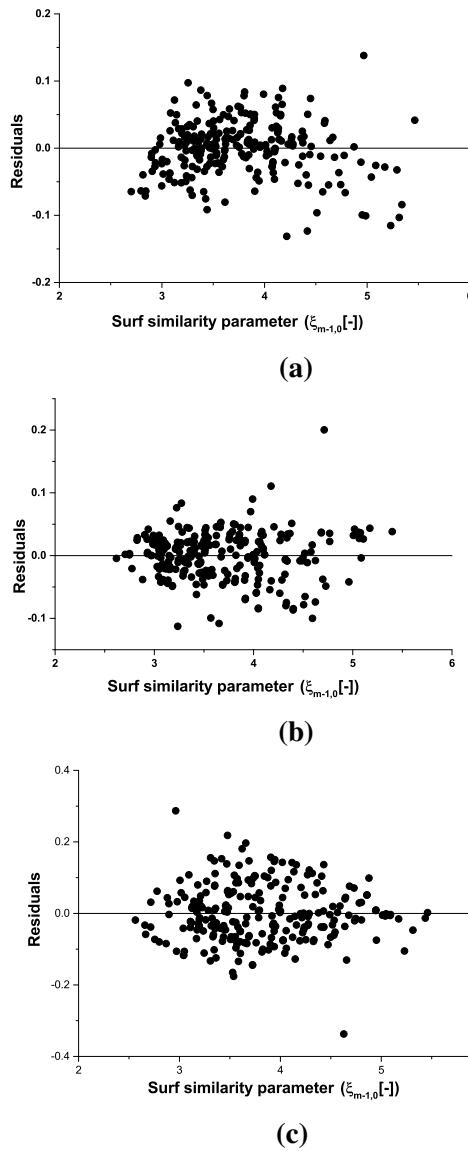


Figure 8.9 A residual plots between the residuals of K_r with respect to $\xi_{m-1,0}$ for sloping seawall with: (a) $\cot \alpha = 0.0$; (b) $\cot \alpha = 1.5$, (c) $\cot \alpha = 3.0$

Based on multiple regression analysis of the observed bimodal datasets, empirical relationship that considers effects of reflection due to bimodal waves is proposed. As suggested in Zanuttigh and Martinelli (2008), this equation contains two corresponding calibration coefficients applied in defining K_r of different impermeable slopes. It can only be applied for sloping seawalls under bimodal sea conditions. The coefficients are slightly modified from Zanuttigh and Martinelli (2008). For simplicity, a more general form of Equation (8.5) can be written in terms of coefficients a and b presented in Equation (8.9):

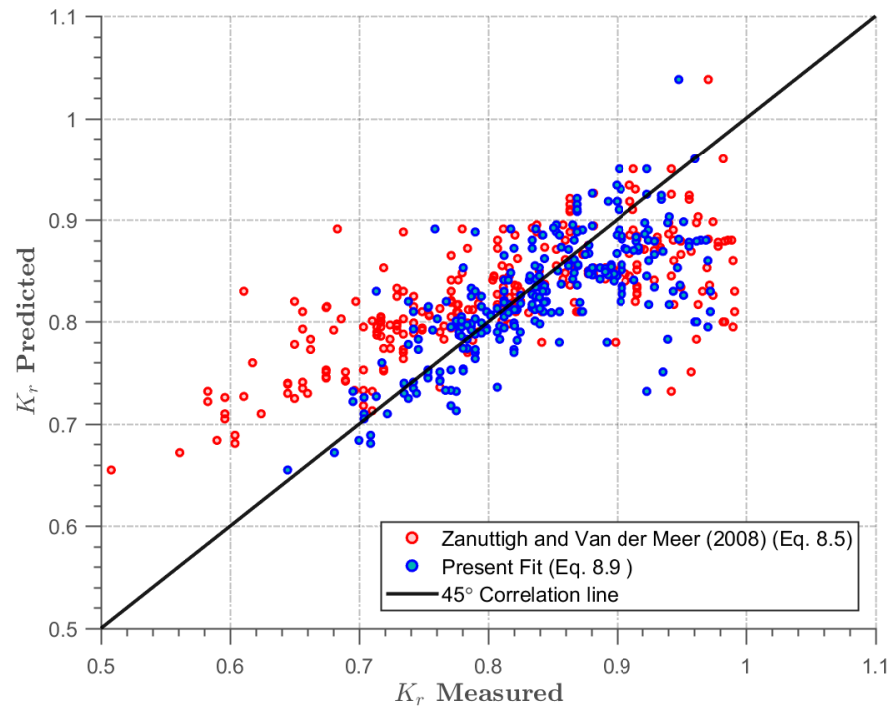
$$K_r(a, b) = \tanh a(\xi_{m-1,0}^b) \quad (8.9)$$

As a general rule of conditions for bimodal seas, values of a and b can be simplified:

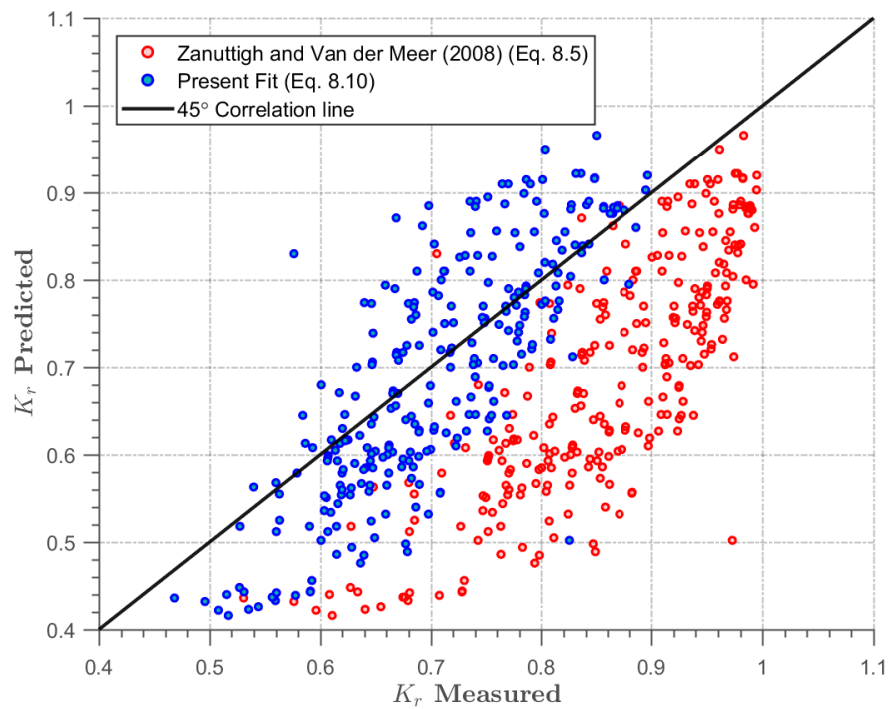
$$\left. \begin{array}{l} K_r = \tanh \left(a \xi_{m-1,0}^b \right) \\ \xi_{m-1,0} = (2.4 \leq \xi_{m-1,0} \leq 5.5) \\ \vdots \\ \frac{R_c}{H_{m0}} = \left(0.8 \leq \frac{R_c}{H_{m0}} \leq 4.0 \right) \end{array} \right\} \begin{array}{l} a = 0.25 \text{ (for } 1 \leq \cot \alpha \leq 1.5) \\ b = 1.3 \text{ (Bimodal seas)} \end{array} \quad (8.10)$$

$$\left. \begin{array}{l} K_r = \tanh \left(a \xi_{m-1,0}^b \right) \\ \xi_{m-1,0} = (2.4 \leq \xi_{m-1,0} \leq 5.5) \\ \vdots \\ \frac{R_c}{H_{m0}} = \left(0.8 \leq \frac{R_c}{H_{m0}} \leq 4.0 \right) \end{array} \right\} \begin{array}{l} a = 0.16 \text{ (for } 1.5 \leq \cot \alpha \leq 3.0) \\ b = 1.3 \text{ (Bimodal seas)} \end{array} \quad (8.11)$$

To assess the suitability and conformance of the new formulation, a correlation assessment of the new equation with Zanuttigh & van der Meer Zanuttigh and Martinelli (2008) is made. Figure 8.10 shows a comparison of the verification of Equation (8.5) from the Zanuttigh & van der Meer Zanuttigh and Martinelli (2008) and Equations (8.10) and (8.11) obtained in present study. The observed K_r and the predicted K_r are fully described in this figure for the two slopes ((a) with $\cot \alpha = 1.5$; and (b) with $\cot \alpha = 3.0$).



(a)



(b)

Figure 8.10 Comparison between the measured K_r and the predicted K_r of wave reflection coefficient (a) $\cot \alpha = 1.5$; and (b) $\cot \alpha = 3.0$

As shown in these figures, observed data points are well distributed either side within 45° correlation line. It can be deduced from the slight convergence of the agreement between observed and predicted values that Equations (8.10) – (8.11) works better in defining reflection characteristics of sloping seawalls under the influence bimodal waves than Equation (8.5). The prediction of prolonged breaker index imposed longer periods in swell waves are more accurately predicted with Equations (8.10) – (8.11) than Equation (8.5). It is under-predicting K_r under the bimodal cases most especially with increasing gentle slopes. The present formulation (represented by Equations (8.10) – (8.11)) will be applicable at locations exposed to local storm waves and open oceans under bimodal wave conditions.

The equations are valid within the following ranges of smooth impermeable slopes:

$$0 \leq \cot \alpha \leq 1.5 \text{ for Equation (8.9), and } 0 \leq \cot \alpha \leq 1.5$$

range of applicability of structure slopes to $1.5 \leq \cot \alpha \leq 4.0$ serving intervals of breaker indices within $1.0 \leq \xi_{m-1,0} \leq 4.1$ with dimensionless crest freeboard of $0.58 \leq R_c/H_{m0} \leq 4.5$.

8.4 Conclusions

This chapter has examined the reflection performance of smooth sloping impermeable aluminium seawall under bimodal sea states. Three different sloping seawalls were investigated at three different water levels to conduct 823 storm tests. An array of four wave gauges positioned around the centre (with constant water depth) were applied to effectively capture both incident and reflected wave elevations. The FFT were applied to decompose analysed wave signals from four number wave gauges into frequency components in the frequency domain. The analysed reflection response of the studied coastal seawall is highly dependent on the seawall slope and wave bimodality. The resultant reflection coefficient also increases with swell peak periods and swell percentages. From the results of the reflection tests are presented and analysed in this paper yielding an improved empirical formula to determining reflection under bimodal sea conditions. New expressions for the reflection coefficient to take into account swell driven seas with wave bimodality have been proposed.

Chapter 9

Analysis of Wave Overtopping in a Numerical Flume

The capability of the IH2VOF computational model to simulate wave overtopping of coastal seawalls is presented in Chapter. A method of managing the effect of bimodal wave overtopping is to test how more complex geometries will perform under these wave conditions in the numerical domain. Results are compared for calibration and validation of the numerical performance. Storm walls and parapets are then incorporated into the original structure geometry. Firstly I present the validated wave overtopping due to simple structures. This is followed by results of the simulated storm wall and parapet in order to reduce excess wave overtopping induced by bimodal waves. This chapter concludes by comparing the overtopping reduction capabilities of storm-walls and parapets under bimodal waves.

9.1 Tested Geometries and Cases

Figure 9.1 shows sets of possible seawall geometries that were selected for study within the numerical domain. Overall, different structure geometries are investigated to illustrate how best wave overtopping can be reduced by the incorporation of storm walls and parapets. Different geometries tested include a replication of the plain seawall

slopes already tested experimentally for validation and incorporations of both storm walls and parapets at the start of the crest with crest width $B = 0$.

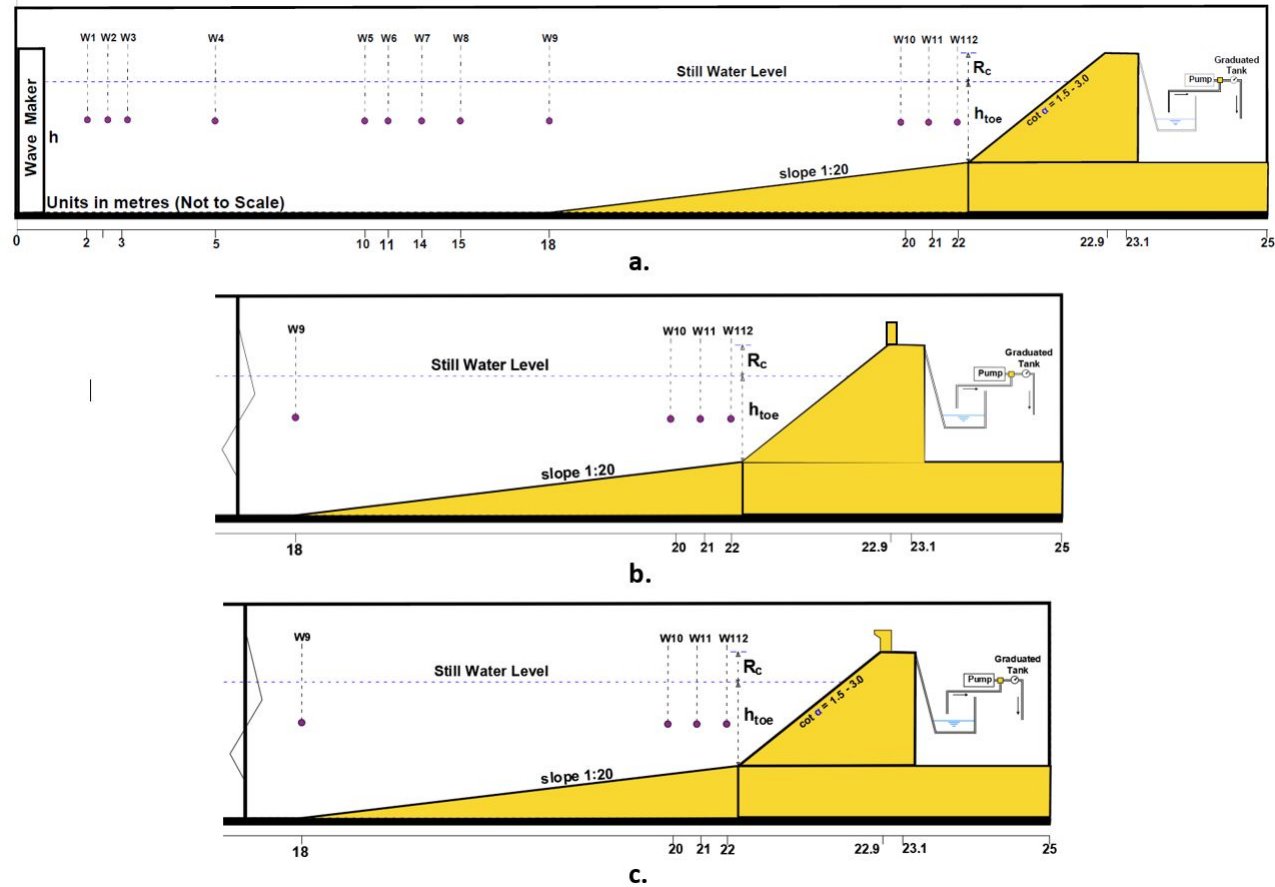


Figure 9.1 Tested geometries for the wave overtopping simulation for plain sloping seawall with: (a) No storm wall and parapet; (b) Storm wall at $B = 0$ and (c) storm wall at $B = 0$

The model set-up was similar to the descriptions we have in chapter 4. In the numerical tests conducted here, additional wave gauges are incorporated at the top of the crest to detect wave overtopping. Table 9.1 contains different wave conditions and water levels that are investigated. Cases with higher overtopping volumes were numerically simulated to reduce excess overtopping volumes using storm walls and parapets. Different numbers are selected for each seawall geometries: $\cot \alpha = 1.5$ having 86 tests, $\cot \alpha = 3.0$ having a total of 100 tests, while for the vertical seawall ($\cot \alpha = 0.0$) has a total of 57 tests. Altogether for the numerical modelling, a total of number of 243 tests were conducted.

Table 9.1 Selected extreme wave conditions simulated in the numerical model

Test No	H_{m0} (m)	T_{pW} (secs)	T_{pS1} (secs)	T_{pS2} (secs)	T_{pS3} (secs)	T_{pS4} (secs)	$h(m)$	$\cot\alpha = 0.0$	$\cot\alpha = 1.5$	$\cot\alpha = 3.0$
T001	0.125	1.107	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T002	0.125	1.264	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T003	0.125	1.423	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T004	0.125	1.581	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T005	0.1	1.107	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T006	0.125	1.264	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T007	0.1	1.423	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T008	0.125	1.581	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T009	0.1	1.107	1.739	2.3712	3.162	3.953	0.6	0	1.5	3
T010	0.1	1.264	1.739	2.3712	3.162	3.953	0.6	0	1.5	3
T011	0.125	1.423	1.739	2.3712	3.162	3.953	0.6	0	1.5	3
T012	0.125	1.581	1.739	2.3712	3.162	3.953	0.6	0	1.5	3
T013	0.075	1.107	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T014	0.075	1.264	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T015	0.075	1.423	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T016	0.075	1.581	1.739	2.3712	3.162	3.953	0.65	0	1.5	3
T017	0.1	1.107	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T018	0.1	1.264	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T019	0.1	1.423	1.739	2.3712	3.162	3.953	0.7	0	1.5	3
T020	0.1	1.581	1.739	2.3712	3.162	3.953	0.7	0	1.5	3

9.2 Validation of Results

In order to rely on the numerical system to replicate other physical geometries or scenarios, model calibration and validation is necessary. This was achieved by applying how well the surface elevation time series and wave energy (spectra) were obtained from the numerical model, and how they conform to the expected theoretical and experimental values. The validated results are based on carefully selected vertical mesh resolution, horizontal mesh resolution, and domain sizes that are compatible with the experimental set-up. Moreover, the physical parameters have to be carefully chosen to

ensure suitable turbulence dissipation and avoid destructive interference of waves. This is because zero values are obtained for numerical wave gauges placed at those positions.

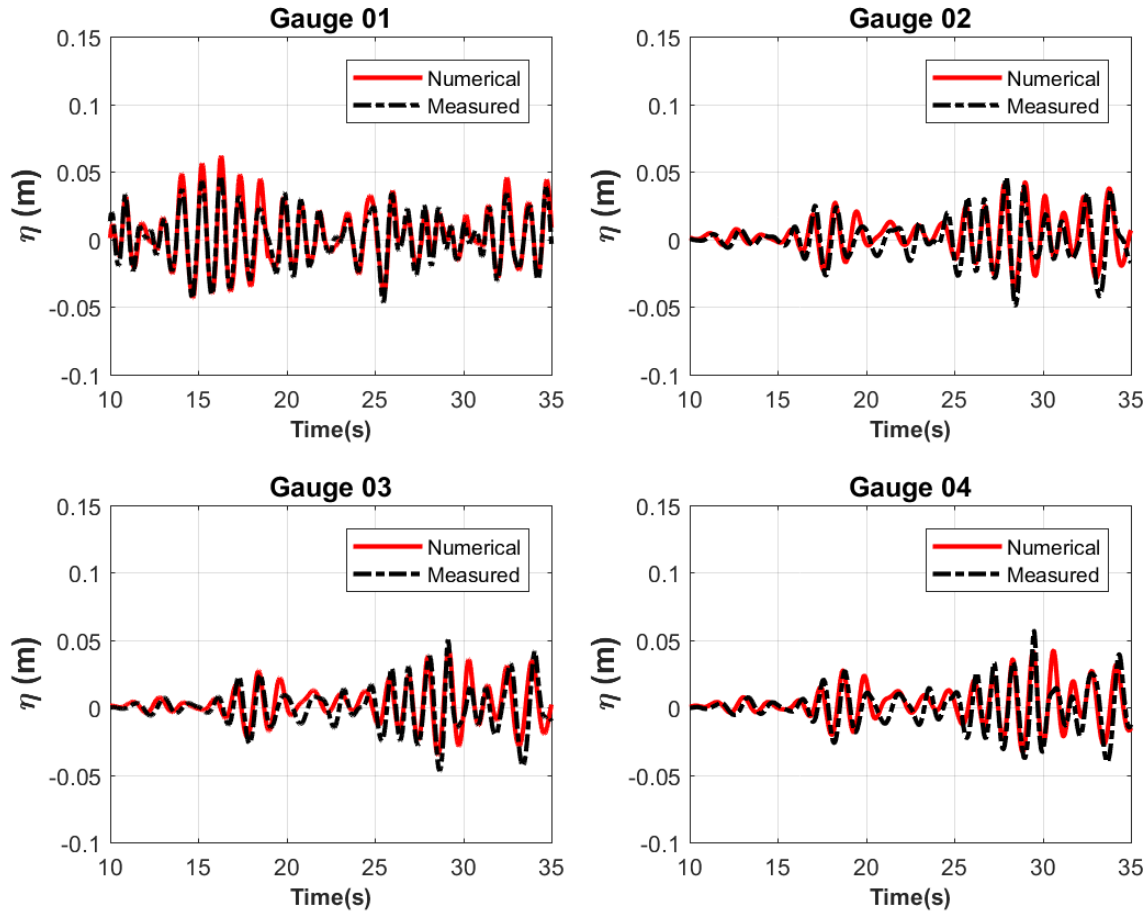


Figure 9.2 An example of free surface validation between the numerical and experimental model obtained during this study for a wave with $H_{m0} = 0.1$ metres at a water depth 0.65 metres approaching a slope with $\cot \alpha = 1.5$

Numerically reproduced wave conditions are compared with the experimental and theoretical values. At first, the output time series from the numerical domain are mostly out of phase in several cases due to incompatibility of generation sequences. It is because the numerical wave paddle start at different times with the physical model. This problem was resolved by reinforcing the boundary conditions with the same time series from the physical model. Wave gauge positions were severally calibrated until a compatible time series were obtained in the computational domain. Figures 9.2 and 9.3 show the validation of the reproduction of surface elevations and wave spectra respectively from the numerical domain.

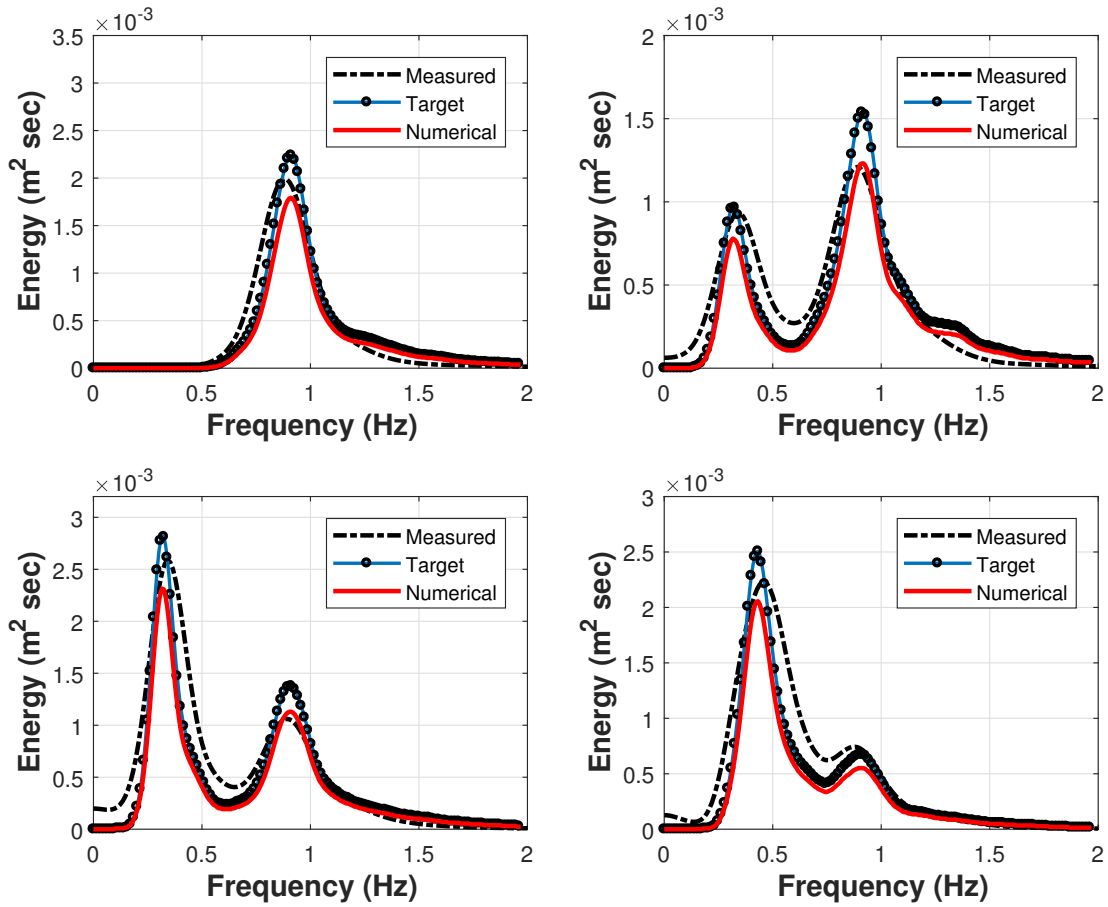


Figure 9.3 An example of reproduction of different spectra types that are validated between the theoretical, numerical and experimental model for a wave with $H_{m0} = 0.1$ metres at a water depth 0.65 metres approaching a slope with $\cot \alpha = 1.5$

As expected, good agreement was obtained in terms of the wave propagation pattern and the energy distribution within the spectra. The peak frequencies of the modal values obtained in the numerical model were well captured. The spectral values are under-predicted in this case due to extreme turbulence dissipation in the model. This may be due to the numerical inconsistencies that arise from numerical assumptions and scaling effects. However, the overall spectra energy were sufficient to describe both the theoretical and the physical wave-forcing across various frequencies for both spectra types.

9.2.1 Spectra Evolution Across Numerical Gauges

Generally across the numerical domain, the evolution of bimodal wave spectra were accurately described. Examples of the reproduction of a wave condition with a bimodal spectra of 75 % swell are presented in Figure 9.4 across the 6 wave gauges.

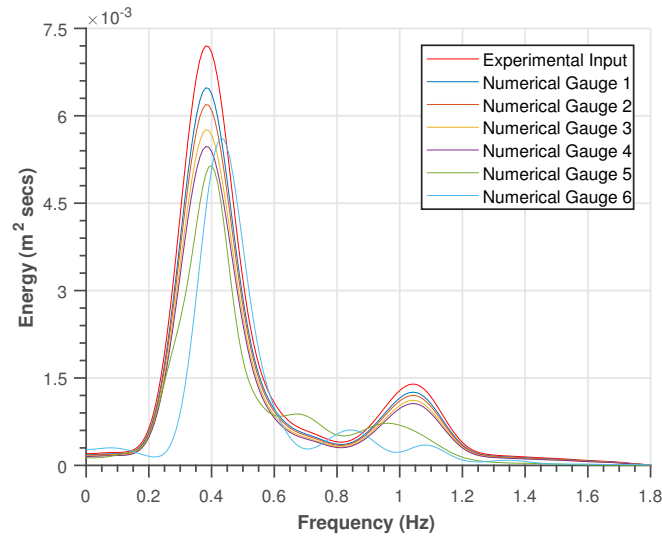


Figure 9.4 Evolution of bimodal spectrum across different wave gauges in the numerical domain for $\cot \alpha = 3.0$ and test T001 – 10 and 75% swell

As expected, similar energy patterns were maintained between the experiment and the numerical output at the peaks. Most deformations and transformations occurred with wave gauges 5 and 6 placed in sloping positions. More energy is shifted from lower to higher frequencies at these wave gauges than observed in other ones. Within the numerical domain as waves propagates, there's numerical dissipation going on in the tank as the energy spectra are decreasing with distance from the paddle from Gauge 1 near the wave paddle to Gauge 8 by the structure slope (Figure 9.4).

9.2.2 Flow Patterns in the Numerical System

Accurate modelling of the free surface and flow velocity is an essential component of a wave hydrodynamics processes. Within the IH2VOF model, the free surface and the velocity fields of wave can be accurately modelled and tracked. Figures 9.5 shows the free surface elevation of a propagating bimodal wave leading to overtopping. Figure 9.6 show examples of the flow magnitude obtained from the horizontal and vertical velocity vectors of a leading wave. The wave velocity pattern were very well described with the IH2VOF model. The velocity pattern within the crest portrays accurately the non-breaking surging wave overtopping as can be seen in the velocity magnitude as captured in the numerical model at the times shown.

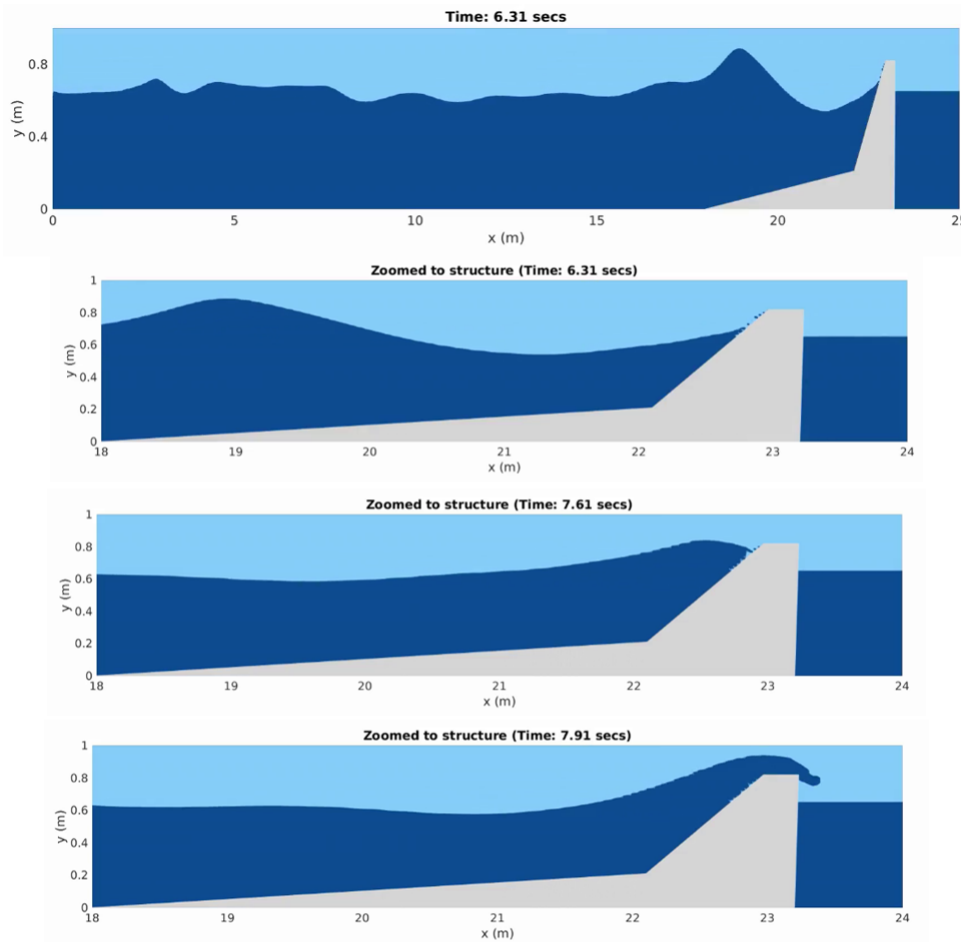


Figure 9.5 An example of free surface by non-breaking (surging) overtopping waves with $H_{m0} = 0.1$ metres at a water depth 0.65 metres approaching a slope with $\cot \alpha = 1.5$ in a numerical domain IH2VOF

Sufficient optimization of the meshes were performed to accurately visualise this behaviour. In the case shown in the following figures, the leading overtopping waves

are characterised by the highest velocity values (Figure 9.6). In some other cases these waves can be nullified by any returning reflected waves of equal or higher magnitude.

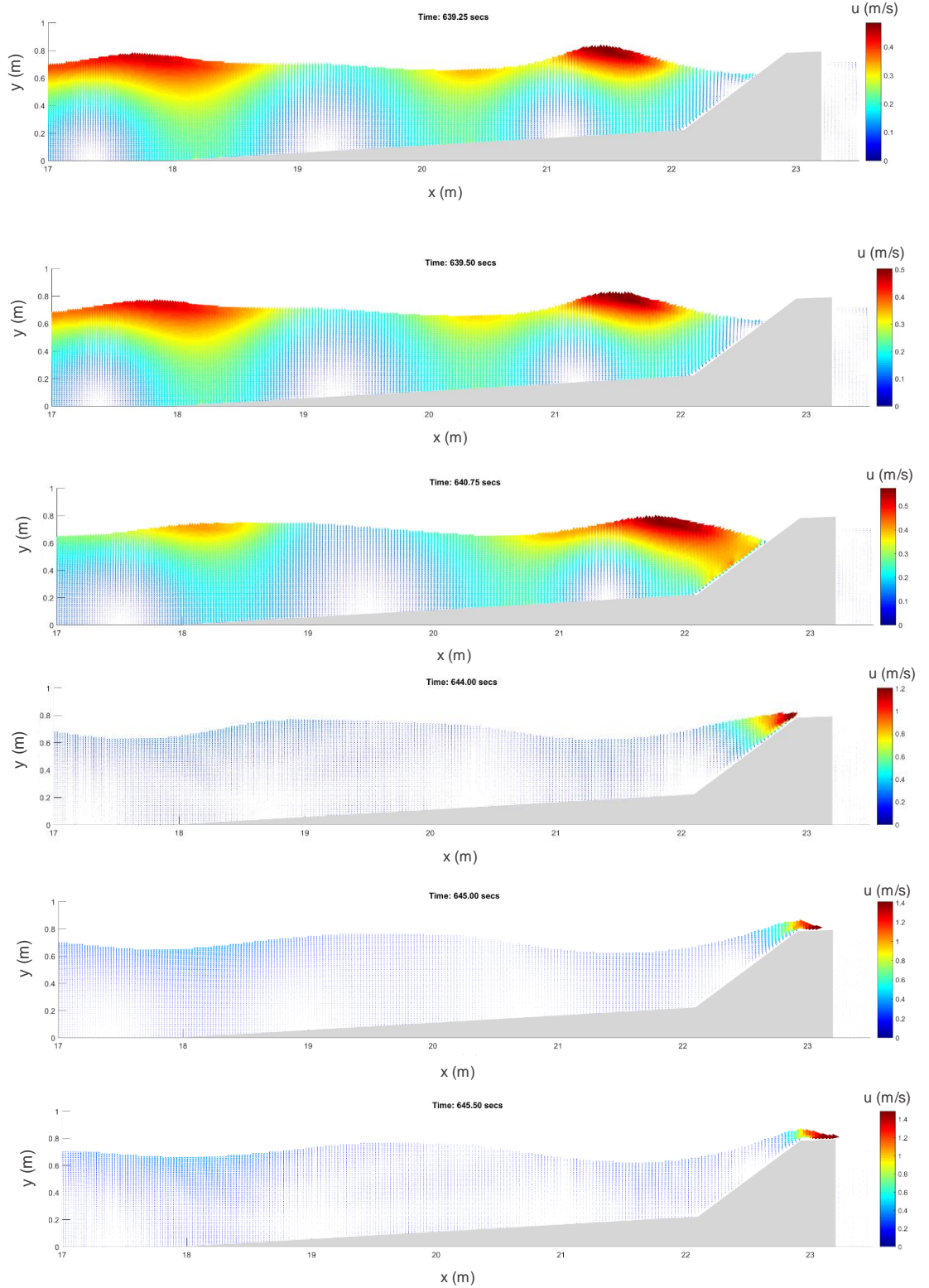


Figure 9.6 An example of flow velocity showing non-breaking (surging) overtopping waves with $H_{m0} = 0.1$ metres at a water depth 0.65 metres approaching a slope with $\cot\alpha = 1.5$ in a numerical domain IH2VOF

9.3 Determination of Wave Overtopping

It is important to derive wave overtopping estimate from the numerical results in order to compare with the experimental results. Recorded times series of the wave gauge placed on top of the crest has been applied to estimate wave overtopping. In this case, wave overtopping is computed using the method suggested in Schüttrumpf and Oumeraci (2005). Using this method, wave overtopping is computed using the product of the layer thickness $h(t)$ of the wave on top of the crest, and the depth-averaged velocity $V(t)$ at the same position. This can be expressed mathematically as shown in Equation (9.1):

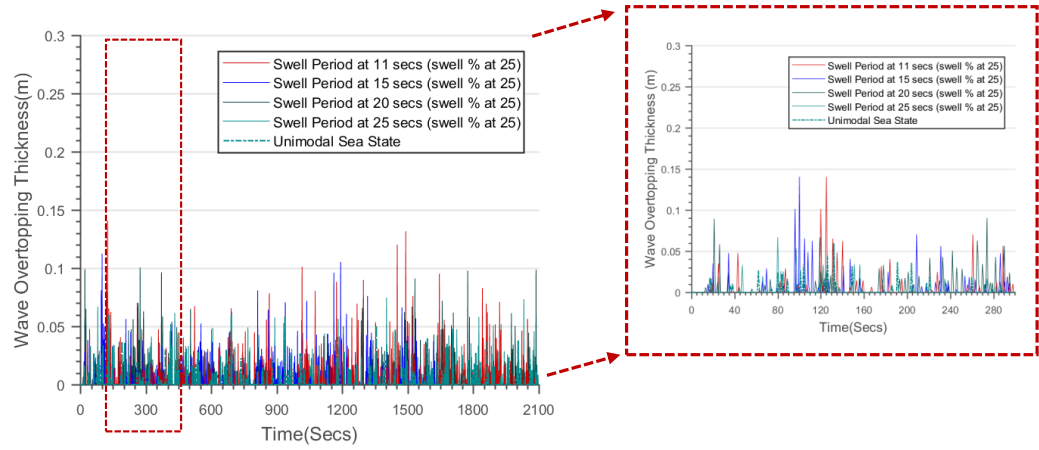
$$q(t) = h(t) * v(t) \quad (9.1)$$

Figures 9.7 and 9.8 give a representation of wave overtopping thicknesses and velocities on top of the structure's crest for different percentages of the swell that are present in the sea-state.

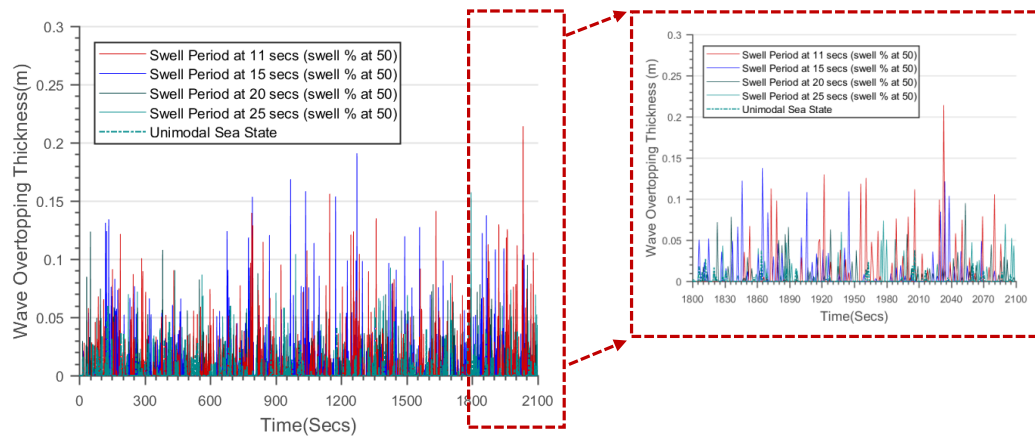
As expected, the numerical domain accurately reproduced the times series of thicknesses and velocities of wave overtopping for different percentages of swell. The thicknesses and velocities of overtopping increases with increasing swell percentages. There is also a direct proportionality in the relationship between overtopping thicknesses and the velocity. It is the increment in the flow velocity on top of the crest that increases the thicknesses of the flow leading to more non-breaking/ flushing wave overtopping.

Moreover, this observation becomes more evident with results of the derived wave overtopping from the two figures by Equation (9.1). These are presented in Figure 9.9, which shows the mean overtopping discharge, and Figure 9.10, which shows the cumulative wave overtopping estimates.

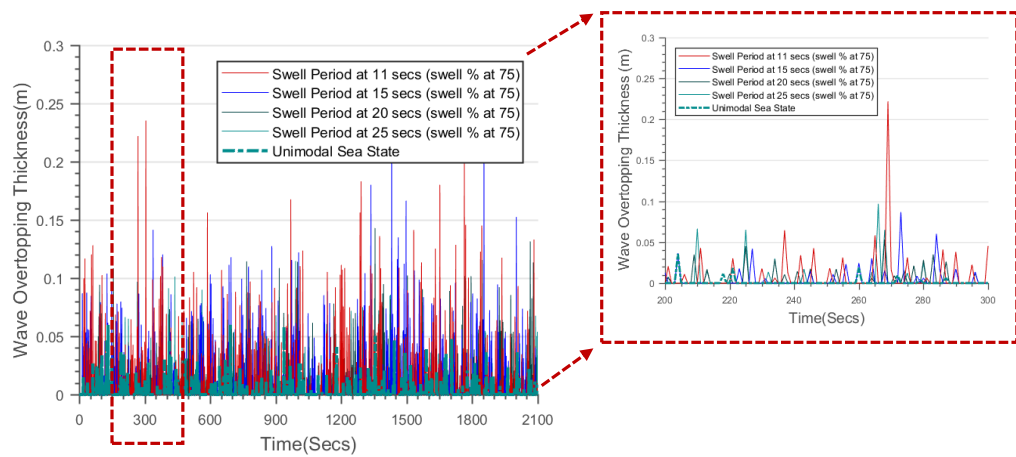
The unimodal sea states without swell shows least overtopping thicknesses and velocities out of all cases. It is the presence of the swell in the sea-state that increases the flushing severities of the flow and causes more overtopping effects.



(a)



(b)



(c)

Figure 9.7 An example of the wave overtopping thickness computed for (a) 25 % swell, (b) 50 % swell, and (c) 75 % swell

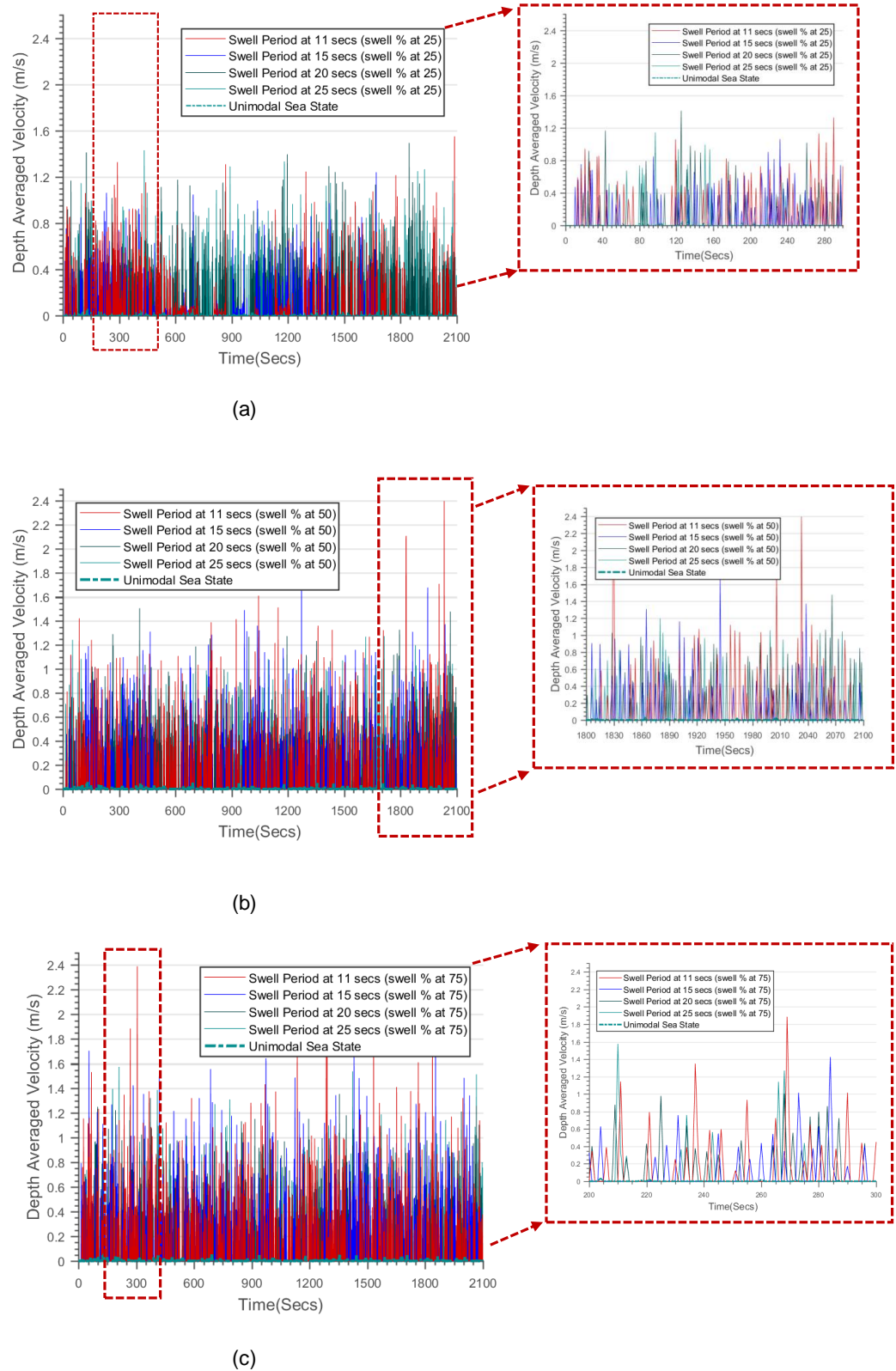


Figure 9.8 An example of the depth-averaged velocity computed for (a) 25 % swell, (b) 50 % swell, and (c) 75 % swell

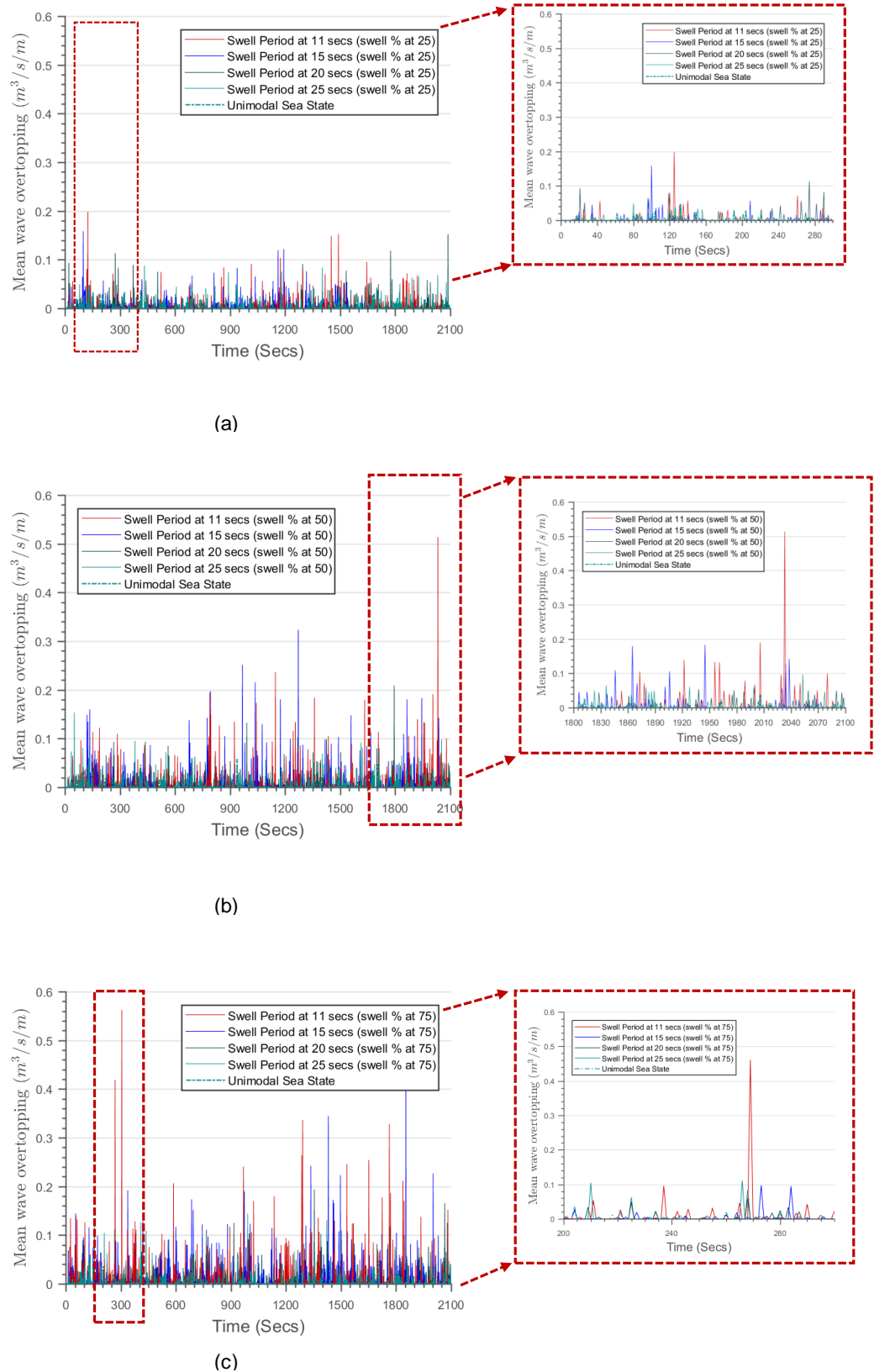
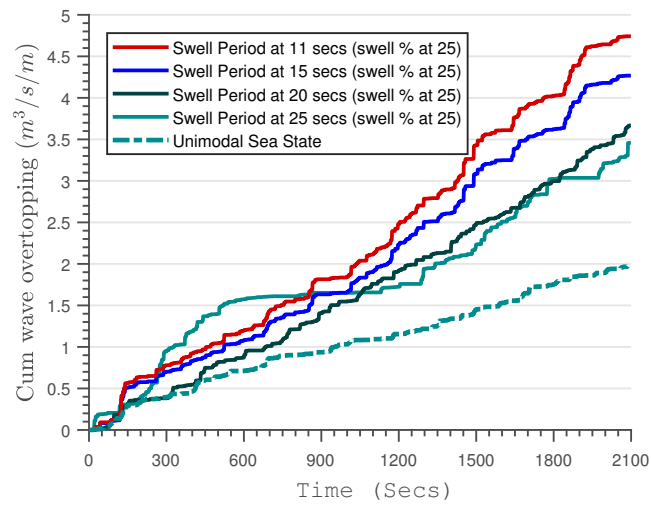
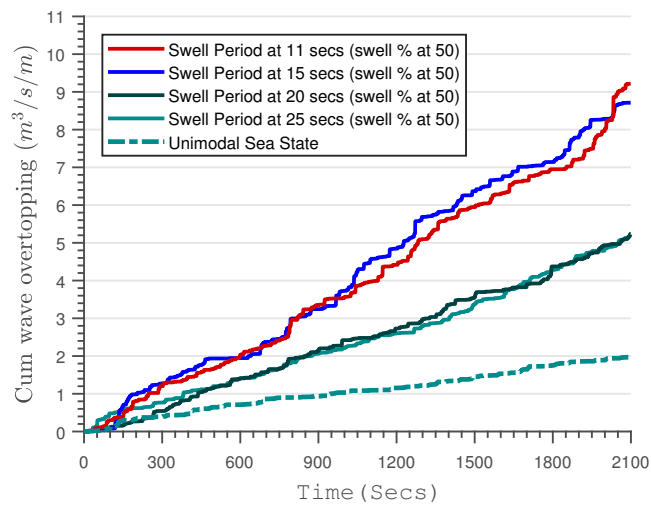


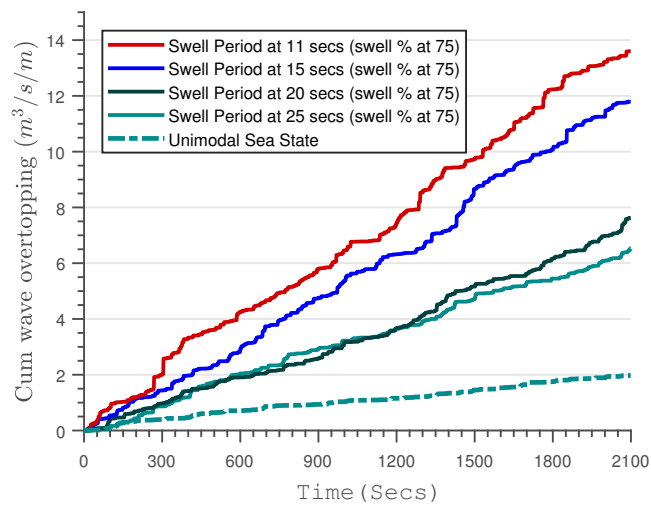
Figure 9.9 Estimated mean wave overtopping volume from the numerical model for (a) 25 % swell, (b) 50 % swell, and (c) 75 % swell



(a)



(b)



(c)

Figure 9.10 Estimated cumulative wave overtopping volume from the numerical model for (a) 25 % swell, (b) 50 % swell, and (c) 75 % swell

9.4 Validation of Wave overtopping Results

Figures 9.11 – 9.13 present a replot of the experimental overtopping results already reported in Chapter 6 across the three slopes investigated in this study. The validation are based on the impermeable geometry without parapet or storm wall. In this section, the dimensionless overtopping determined computationally are shown in Figures 9.11 – 9.13. Generally in the three figures, similar trends of wave overtopping are observed for both the numerical and the experimental outcomes each for different slopes. The combination of numerical and experimental overtopping for both sloping and vertical seawalls are represented by Figures 9.11 and 9.12. It can be observed that overtopping (determined from the numerical model) are lesser than values that are empirically determined. It is the loss in the energy of the computational domain that is responsible for this event. However, similar trends are observed in both numerical and experimental domains.

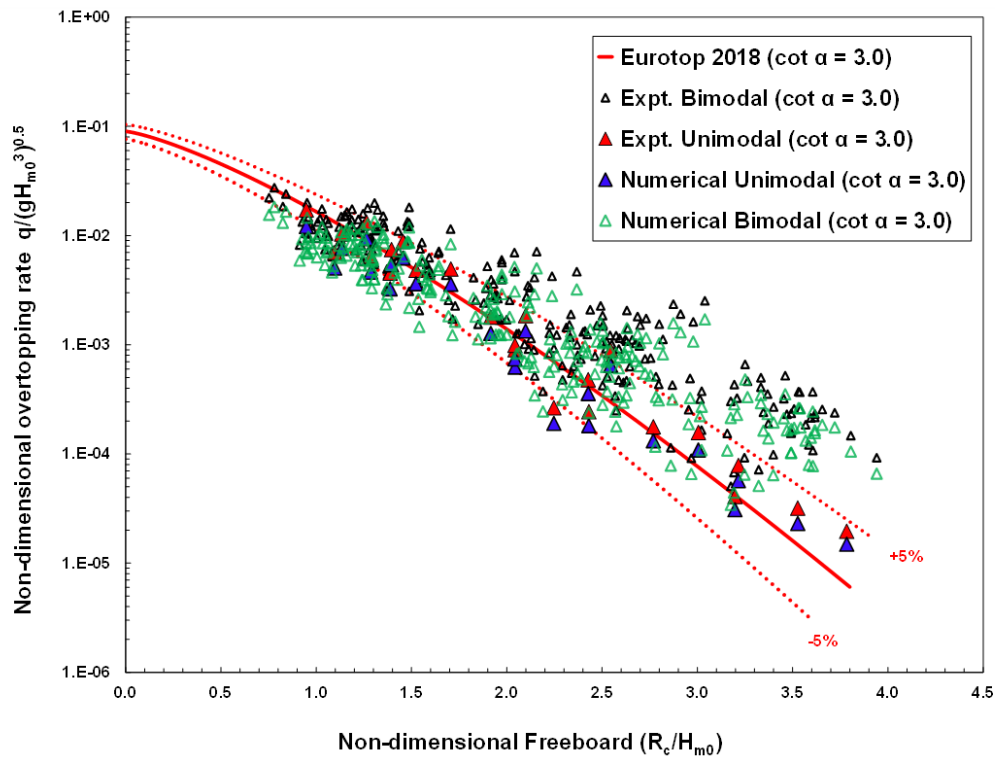


Figure 9.11 Plots of computational and experimental non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$ plotted within the corresponding prediction band of EurOTop (2018) obtained for a sloping seawall with $\cot \alpha = 3.0$

Moreover, numerically determined overtopping values from unimodal sea states are accurately predicted by the corresponding 90 % confidence limits of EurOTop (2018).

However, bimodal seas were under-predicted across slopes with $\cot \alpha = 1.5$ and $\cot \alpha = 3.0$ and $\cot \alpha = 0.0$.

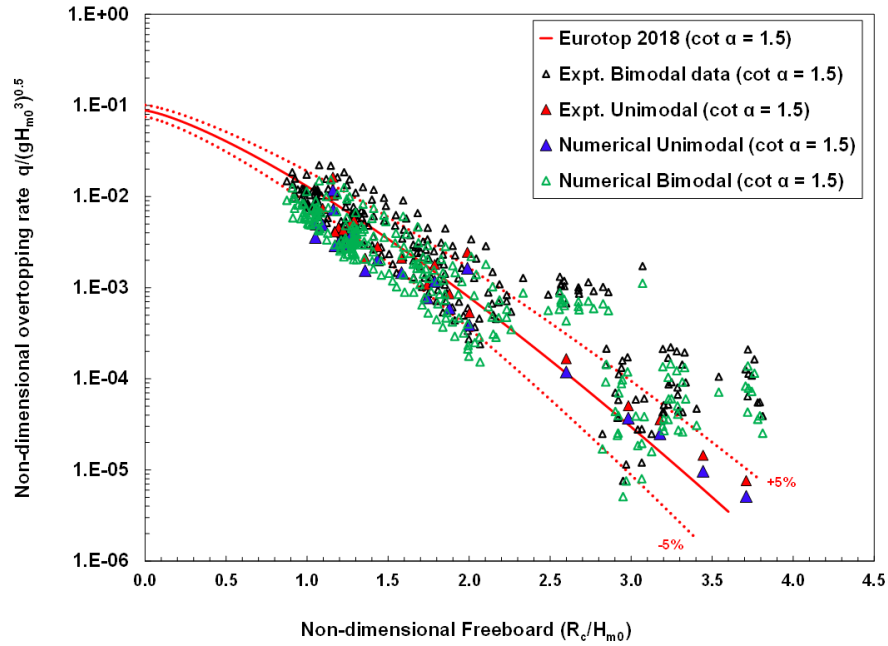


Figure 9.12 Plots of computational and experimental non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$ plotted within the corresponding prediction band of EurOTop (2018) obtained for a sloping seawall with $\cot \alpha = 1.5$.

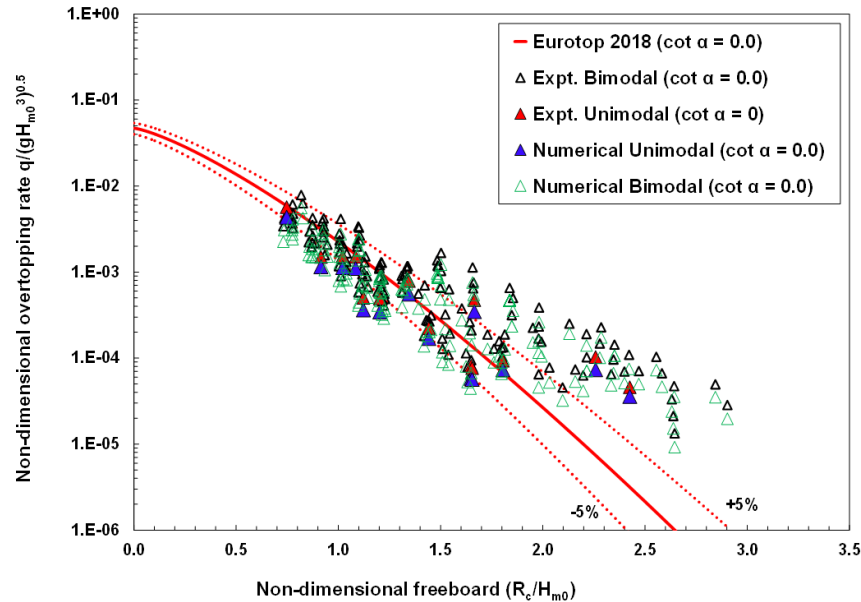


Figure 9.13 Plots of computational and experimental non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$ plotted within the corresponding prediction band of EurOTop (2018) obtained for a sloping seawall with $\cot \alpha = 3.0$

9.5 Verification of Numerical and Experimental Results

It is necessary to verify the conformance of the numerical model with the experimental results. To achieve this purpose, two methods of coefficient of determination (R^2) and Nash–Sutcliffe efficiency (NSE) suggested by Krause et al. (2005) are applied. Basically R^2 can be computed using Equation (9.2).

$$R^2 = \left(\frac{\sum_{i=1}^n (E_i - \bar{E})(N_i - \bar{N})}{\sqrt{\sum_{i=1}^n (E_i - \bar{E})^2} \sqrt{\sum_{i=1}^n (N_i - \bar{N})^2}} \right)^2 \quad (9.2)$$

The equation describes the linear dependence between the numerical and the experimental variables expressed in the form of simple correlation coefficients existing between the datasets. In Equation (9.2), E_i represents the recorded dimensionless overtopping from the experimental results, N_i are numerical non-dimensional overtopping. Also, \bar{E} and \bar{N} are the mean values of n number wave overtopping datasets recorded in both experimental and numerical models. Values of R^2 lies between 0 and 1. It is the value of R^2 that describes the strength of linear association between E and N . The value of 1 shows the strongest positive correlation, while $R^2 = 0$ shows that no agreement exists between the two variables compared.

To validate the model using the efficiency, Equation (9.3) is applied. The equation was first proposed by Nash and Sutcliffe (1970). Values of NSE ranges from negative infinity ($-\infty$) to 1, a perfectly matched model between experimental and computational values using the NSE criteria should be greater than 80 %. In general, there is a strong agreement if the value of NSE is zero. Negative values of NSE shows that the experimental results presents a better predictor of the captured data.

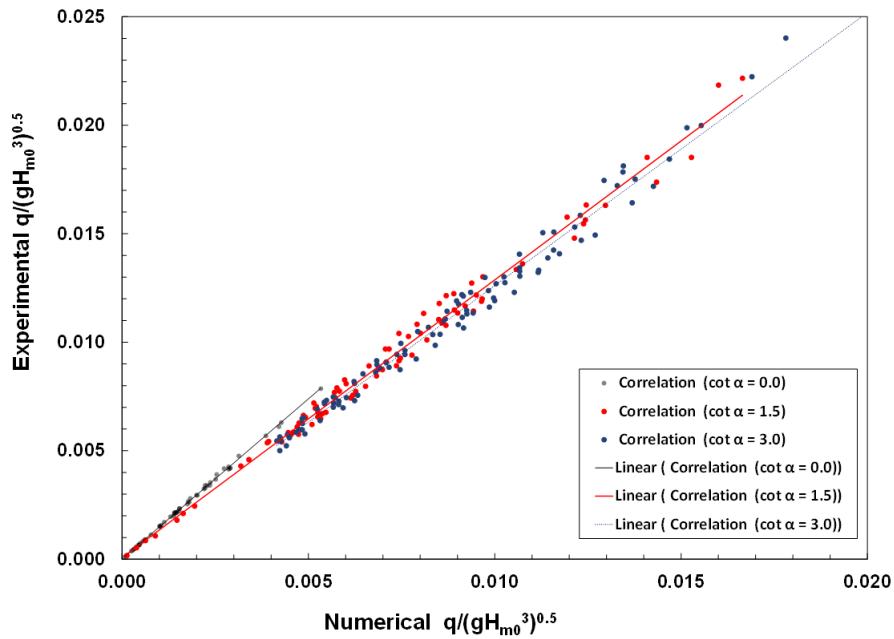
$$NSE = 1 - \left(\frac{\sum_{i=1}^n (E_i - N_i)^2}{\sum_{i=1}^n (E_i - \bar{E})^2} \right) \quad (9.3)$$

Table 9.2 Values of R^2 and NSE computed between the numerical and experimental values for the smooth impermeable seawalls investigated during present study

Parameters		Equations (9.1 & 9.2)	
Item	R^2 (-)	NSE (-)	
$\cot \alpha = 3.0$	0.93	0.701	
$\cot \alpha = 1.5$	0.91	0.723	
$\cot \alpha = 0.0$	0.89	0.699	

Figure 9.14 presents the comparison to verify the relationship between E_i and the N_i across the three with $\cot \alpha = 3.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 0.0$ investigated. As visible in the figure, strong correlation (90 %) exists between numerical and experimental values for the three slopes. Table 9.2 presents summary of R^2 values across the three slopes.

In the present study, as shown in Table 9.2, for the three slopes, simulated results from the computational domain closely agree with the observed data with an efficiency of approximately 70 %. In this case, it is a high degree of closeness because of the numerical complexities of resolving bimodal wave patterns.

**Figure 9.14** Verification of computational and experimental non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ across the three slopes with $\cot \alpha = 3.0$, $\cot \alpha = 1.5$ and $\cot \alpha = 0.0$

9.6 Bimodal Wave Overtopping Due to Storm Walls and Parapets

To reduce the excessive overtopping emerging from bimodal waves, the seawall geometry was modified using a storm wall and a parapet (Figure 9.15). The storm wall and the parapet were placed at a position with berm width $B = 0$ (in both cases), and selected test series (Table 9.1) were simulated using the validated model for the most vulnerable slope ($\cot \alpha = 3.0$). Also these falls within the lower dimensionless freeboards ($0.5 \leq \frac{R_c}{H_{m0}} \leq 2.5$) with excessive overtopping.

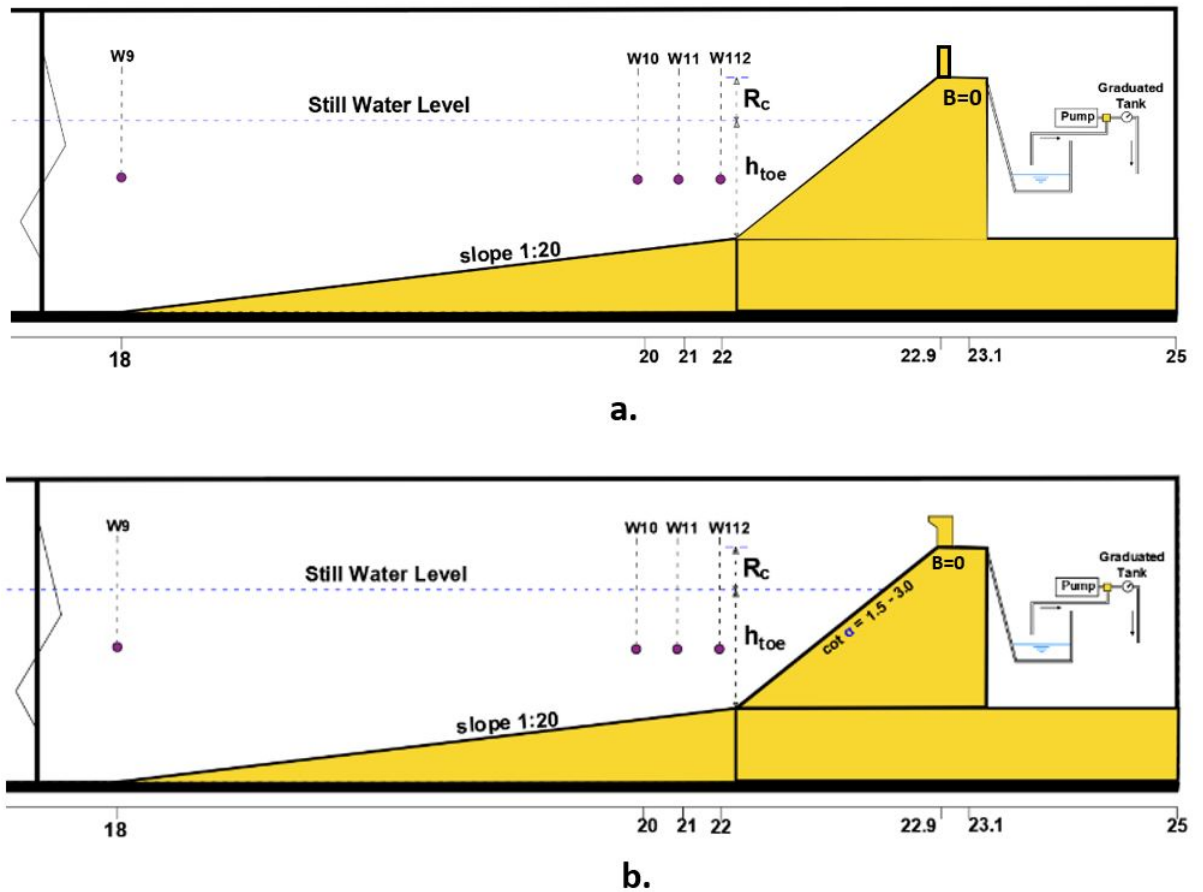


Figure 9.15 Tested geometries for the wave overtopping simulation for plain sloping seawall with: (a) Storm wall at $B = 0$; (b) Parapet at $B = 0$

Figure 9.16 shows the relationship between the non-breaking non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ against the non-dimensional crest freeboard $\frac{R_c}{H_{m0}}$. It was plotted within the corresponding prediction band of EurOTop (2018) and present fit using Equation (7.13) obtained for a sloping seawall with $\cot \alpha = 3.0$ numerically obtained between the storm walls and the parapet for $\cot \alpha = 3.0$. Results obtained for the

original sea-wall geometry with $\cot \alpha = 3.0$ were also presented for comparison. As clearly shown in Figure 9.16, bimodal overtopping rates were significantly reduced by the incorporation of both the storm wall and the parapets. However, with more scattering still experienced at the dimensionless crest freeboard above.

Generally as expected, the parapet clearly provides more overtopping reduction at $B = 0$ position. The parapet provides a more efficiency at reduction of overtopping volumes in up to 13.4 % at that portion of the crest (i.e., $B = 0$). Most overtopping estimates are far lower than the predicted overtopping and corresponding 90 % confidence limits of EurOTop (2018) for the bimodal cases are presented in each plot. The fitness obtained for present dataset with Equation (7.13) present a better prediction of bimodal overtopping than the For these cases, there is no means of checking any evidence of heteroscedasticity between data groups. This is because no bimodal dataset is available for to determine such analysis in this type of seawall.

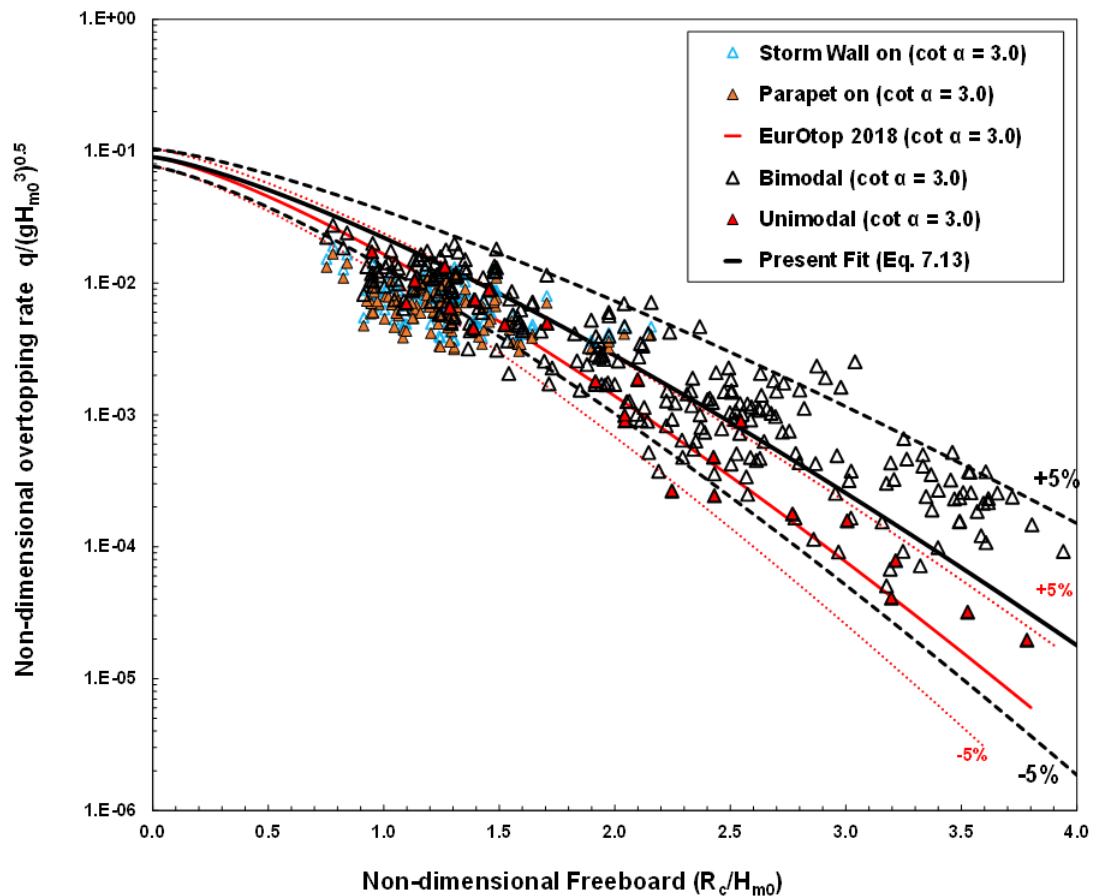


Figure 9.16 Verification of computational and experimental non-dimensional overtopping rate $\frac{q}{\sqrt{H_{m0}^3}}$ for with $\cot \alpha = 3.0$: (a) Storm wall at crest width $B = 0$; (b) Parapet at crest width $B = 0$

9.7 Conclusion

A numerical approach to validate numerical wave overtopping has been presented in this chapter. It has also been complimented with the use of more complex geometries to reduce overtopping occurrences by bimodal waves. Significantly and similarly to the experiment, numerically determined overtopping values from unimodal sea states are more accurately predicted within the corresponding 90 % confidence limits of EurOTop (2018) than in bimodal seas. The computational model resolved the numerical complexities of propagating bimodal wave patterns. The agreement between simulated results from the computational domain and the observed empirical data is with a correlation R^2 of 90 % and efficiency greater than 70 %. Moreover, bimodal overtopping rates were drastically reduced by the incorporation of the storm wall and the parapets. However, as suggested in previous literatures, the parapets provided more reduction of up to 13 % than the implemented storm wall in this case.

Chapter 10

Conclusions and Future Research

In this chapter, the main conclusions that are drawn from each aspects investigated are presented. The overview of the objectives and methodology of the thesis was first presented. It was followed by summaries of results obtained to fulfil each objective.

10.1 Overview of the Thesis Contribution

The main problem studied in this thesis concerns the modelling of wave overtopping of coastal defences under storm-induced bimodal waves. This has been assessed by developing a specific form of energy-conserved bimodal spectrum that have varying proportions of swell, while at the same time containing a fixed amount of energy with the unimodal wind seas for easy comparison. Numerical and experimental methods have been applied to selected bimodal spectra and their corresponding unimodal sea states to compare the overtopping tendencies of smooth, sloping and vertical impermeable seawalls. The main conclusions are presented in each section below.

10.2 Effects of Swell in Energy-Conserved Bimodal Seas

Presence of swell in a bimodal sea state will significantly alter the non-linearity of the sea-state despite the fact that the energy content of the sea-state remain the same as the unimodal sea. This would in turn, effect the characteristics of waves derived from the unimodal and bimodal spectra, whether near a wave maker, or near a structure. There is an increasingly smooth and continuous exponential relationship between the skewness and kurtosis as the swell percentage or the peak period grows in an energy

conserved bimodal sea. The Weibull distributions produced the best fitness for wave patterns of high swell percentages in this study. As the swell frequency and percentages changes in a bimodal sea, the relationship between the skewness and the kurtosis are tending towards a unimodal sea. The distribution of the waveheight on the Rayleigh model shows that the non-linearity increases with the distance away from the wave maker. This could be caused by the development of other physical processes like reflection, wave breaking, and other shoaling processes arising from wave-structure interactions. In agreement with previous studies, the unimodal sea states showed the greatest non-linear behaviour however, non-linearity patterns vary as the swell grows. The Rayleigh distribution could not properly represent bimodal seas with wider spectral width and also overpredicts waveheight at low exceedances in these conditions.

10.3 Effects of Swell Percentage on Wave Overtopping

The wave overtopping performance of sloping impermeable seawalls under the influence of bimodal sea conditions have been investigated using 823 wave overtopping tests. Different percentages of swell components were derived for each sea state. Both unimodal and equivalent bimodal waves were tested in order to compare effects of bimodality. With wave overtopping derived from a minimum of 1000 waves, results show that relative wave overtopping increases with increasing swell percentages. It was hypothesised that the increased wave overtopping for longer waves/swell percentage is due to a change in wave behaviour, where more waves are surging than breaking on the structure. Also, for fixed wave conditions, more wave overtopping was observed with increasing slope mildness as established in previous studies. The unimodal tests agree closely with predictions using the EurOtop formulae. EurOtop formulae underpredict the observed wave overtopping rates for bimodal conditions, (even when using the $T_{m-1,0}$ period for such cases). A simple modification to the existing formula for non-breaking waves was proposed that accounts for the presence of swell percentages under bimodal conditions.

10.4 Effects of Swell Peak Periods on Wave Overtopping

Based on the same experiments, effects of swell appearance at different peak periods on wave overtopping were equally analysed. It was observed that swell occurrence at

lower peak periods produce the uppermost impact of wave overtopping on each slope investigated than those at higher peak periods. These observations are characterised by more non-linearities when compared to effects posed by swell percentages. Modification to the existing formula for non-breaking waves under the presence of swell peak periods under bimodal conditions have not been provided in this thesis. Further analysis would be necessary to establish better relationship between swell peak periods with overtopping estimates.

10.5 Effects of Spectral Characteristics on Wave Overtopping

Influences of different spectra characteristics such as the narrowness parameter, the broadness parameter, and the peakedness parameter on wave overtopping were analysed for different spectrum types in this thesis. It was observed that spectra characteristics of each parameter were strongly correlated with swell percentages in the sea for different spectra types. Overtopping increases as spectra values of each parameter reduces. However, spectra values tend to be approaching asymptotic values for both the narrowness and the broadness parameters. Increasing trends of swell percentage in a bimodal sea could produce enormous wave overtopping with tendencies of more significant flooding. Broad-banded sea conditions yield less wave overtopping intensity than highly peaked bimodal seas. Overtopping is highly sensitive to the succession of wave groups and wave patterns in the random wave sequence than overtopping caused by individual peaks. However, bimodal waves occurring from different percentages of wind waves while the swell content of the sea remain constant were not analysed in this study.

10.6 Reflection Analysis of Coastal Seawalls Under Bimodal seas

At locations exposed to local storm waves and open oceans, long period swell waves can be present leading to bimodal wave conditions. Existing literature provides little guidance on reflection characteristics in this situation. The experimental results from this thesis was applied to fill the gap in our understanding of seawall performance under bimodal wave conditions. As in previous literatures, the analysed reflection response of the studied coastal seawall is highly dependent on the seawall slope and wave bimodality. However, the resultant reflection coefficient increases with swell peak

periods and swell percentages. From these analyses, an improved formulation for the reflection coefficient to take into account swell driven seas with wave bimodality have been proposed.

10.7 Wave Overtopping in a Numerical Flume

The overtopping from the experiment was applied to validate the wave overtopping from the numerical model. More geometries involving a storm wall and a parapet were also simulated computationally to test their reductive capability on bimodal waves. Although simulating the storm walls and parapets in the computational domain is a very challenging task because of the complexity of numerically describing the structures' geometry. The challenges of multiple sampling of double-peak frequencies described by bimodal waves were resolved. There is a great agreement between simulated results from the computational domain and the observed empirical data. Moreover, bimodal overtopping rates were drastically reduced by the incorporation of the storm wall and the parapets as expected. However, a further reduction of up to 13 % was obtained in the overtopping of parapet simulation than storm walls.

10.8 Recommendations for Further Research

Having extensively studied bimodal waves in this present study, it would be essential in future to consider research in the following areas:

1. Large sets of experiments on the individual wave overtopping threats of bimodal waves. Emphasis of this study should be based on maximum wave overtopping occurrences.
2. Considering using rubble mounds (Figure 10.1) in model experiments and using extended experimental studies to examine structural responses to wave bimodality.

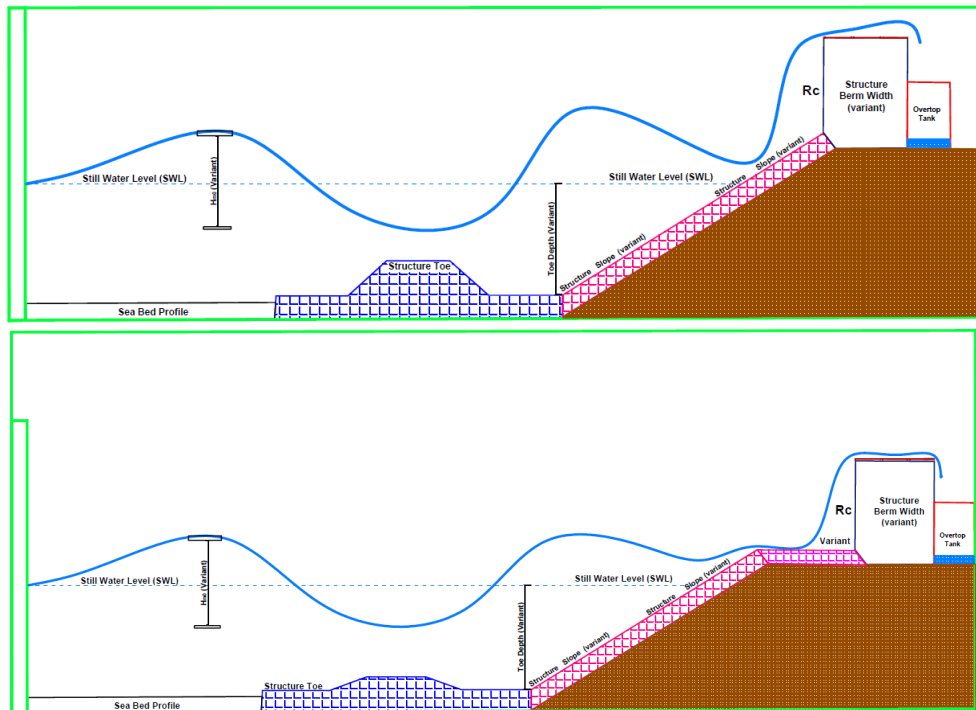


Figure 10.1 *Further modifications to coastal seawalls in future research.*

3. Examining the driven phenomena of force effects on different coastal structures under bimodal storms.
4. It is essential to replicate this entire study in a three-dimensional flume for a more realistic representation of the overtopping effects on coastal seawalls. Worst effects of wave overtopping usually occur in a normally incident wave-structure interaction, however, incorporating the directional component of the spectra into the energy-conserved bimodal spectra would be a useful aspect of research in a wave basin.

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Appendix A

Outcome of Current Research

Orimoloye, S., Karunarathna, H., and Reeve, D. E. (2019). "Effects of swell on wave height distribution of energy-conserved bimodal seas". *Journal of Marine Science and Engineering*, 7(3):79.

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Appendix B

Bimodal Wave Overtopping Test Conditions and Results

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
1	T001 – 1	2	25	7	11	0.70	0.10	0.125	1.08	0.006396
2	T001 – 2	2	25	7	15	0.70	0.10	0.125	1.03	0.017520
3	T001 – 3	2	25	7	20	0.70	0.10	0.125	1.05	0.007447
4	T001 – 4	2	25	7	25	0.70	0.10	0.125	1.09	0.007323
5	T001 – 5	2	50	7	11	0.70	0.10	0.125	1.13	0.009150
6	T001 – 6	2	50	7	15	0.70	0.10	0.125	1.00	0.009440
7	T001 – 7	2	50	7	20	0.70	0.10	0.125	1.22	0.014079
8	T001 – 9	2	75	7	11	0.70	0.10	0.125	1.06	0.009861
9	T001 – 10	2	75	7	15	0.70	0.10	0.125	0.94	0.010494
10	T001 – 13	1	0	7	-	0.70	0.10	0.125	1.10	0.007010
11	T002 – 1	2	25	8	11	0.70	0.10	0.125	0.83	0.018440
12	T002 – 2	2	25	8	15	0.70	0.10	0.125	0.84	0.024020
13	T002 – 3	2	25	8	20	0.70	0.10	0.125	0.96	0.009442
14	T002 – 4	2	25	8	25	0.70	0.10	0.125	0.91	0.008193
15	T002 – 5	2	50	8	11	0.70	0.10	0.125	0.78	0.027410
16	T002 – 6	2	50	8	15	0.70	0.10	0.125	0.95	0.019990
17	T002 – 7	2	50	8	20	0.70	0.10	0.125	1.01	0.013278

($\cot \alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
18	T002 – 9	2	75	8	11	0.70	0.10	0.125	0.75	0.022240
19	T002 – 10	2	75	8	15	0.70	0.10	0.125	1.02	0.014251
20	T002 – 13	1	0	8	-	0.70	0.10	0.125	0.95	0.017220
21	T003 – 1	2	25	9	11	0.70	0.10	0.125	1.14	0.011150
22	T003 – 2	2	25	9	15	0.70	0.10	0.125	0.95	0.010370
23	T003 – 3	2	25	9	20	0.70	0.10	0.125	0.98	0.011350
24	T003 – 4	2	25	9	25	0.70	0.10	0.125	1.04	0.012380
25	T003 – 5	2	50	9	11	0.70	0.10	0.125	1.16	0.015050
26	T003 – 6	2	50	9	15	0.70	0.10	0.125	0.95	0.011760
27	T003 – 7	2	50	9	20	0.70	0.10	0.125	1.13	0.014940
28	T003 – 9	2	75	9	11	0.70	0.10	0.125	1.05	0.014060
29	T003 – 10	2	75	9	15	0.70	0.10	0.125	0.93	0.013320
30	T003 – 13	1	0	9	-	0.70	0.10	0.125	1.13	0.010359
31	T004 – 1	2	25	10	11	0.70	0.10	0.125	1.05	0.010660
32	T004 – 2	2	25	10	15	0.70	0.10	0.125	1.06	0.012990
33	T004 – 3	2	25	10	20	0.70	0.10	0.125	1.15	0.015310
34	T004 – 4	2	25	10	25	0.70	0.10	0.125	1.27	0.017850
35	T004 – 5	2	50	10	11	0.70	0.10	0.125	0.97	0.012300
36	T004 – 6	2	50	10	15	0.70	0.10	0.125	1.19	0.017460
37	T004 – 9	2	75	10	11	0.70	0.10	0.125	1.16	0.016430
38	T004 – 10	2	75	10	15	0.70	0.10	0.125	0.96	0.014700
39	T004 – 13	1	0	10	-	0.70	0.10	0.125	1.27	0.013223
40	T005 – 1	2	25	7	11	0.65	0.15	0.100	2.05	0.001260
41	T005 – 2	2	25	7	15	0.65	0.15	0.100	1.94	0.001700
42	T005 – 3	2	25	7	20	0.65	0.15	0.100	1.97	0.001750
43	T005 – 4	2	25	7	25	0.65	0.15	0.100	1.96	0.001690
44	T005 – 5	2	50	7	11	0.65	0.15	0.100	1.72	0.001720
45	T005 – 6	2	50	7	15	0.65	0.15	0.100	1.84	0.002570
46	T005 – 7	2	50	7	20	0.65	0.15	0.100	1.88	0.002410
47	T005 – 8	2	50	7	25	0.65	0.15	0.100	1.93	0.003610
48	T005 – 9	2	75	7	11	0.65	0.15	0.100	1.95	0.002640
49	T005 – 10	2	75	7	15	0.65	0.15	0.100	1.83	0.004120

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
50	T005 – 11	2	75	7	20	0.65	0.15	0.100	1.98	0.005640
51	T005 – 13	1	0	7	-	0.65	0.15	0.100	2.04	0.000899
52	T006 – 1	2	25	8	11	0.65	0.15	0.125	1.49	0.003090
53	T006 – 2	2	25	8	15	0.65	0.15	0.125	1.37	0.003170
54	T006 – 3	2	25	8	20	0.65	0.15	0.125	1.60	0.004350
55	T006 – 4	2	25	8	25	0.65	0.15	0.125	1.60	0.004590
56	T006 – 5	2	50	8	11	0.65	0.15	0.125	1.54	0.003640
57	T006 – 6	2	50	8	15	0.65	0.15	0.125	1.37	0.004940
58	T006 – 7	2	50	8	20	0.65	0.15	0.125	1.55	0.008630
59	T006 – 9	2	75	8	11	0.65	0.15	0.125	1.56	0.005750
60	T006 – 10	2	75	8	15	0.65	0.15	0.125	1.55	0.007560
61	T006 – 13	1	0	8	-	0.65	0.15	0.125	1.52	0.004820
62	T007 – 1	2	25	9	11	0.65	0.15	0.100	1.91	0.002850
63	T007 – 2	2	25	9	15	0.65	0.15	0.100	1.73	0.002270
64	T007 – 3	2	25	9	20	0.65	0.15	0.100	2.02	0.002560
65	T007 – 4	2	25	9	25	0.65	0.15	0.100	2.10	0.002750
66	T007 – 5	2	50	9	11	0.65	0.15	0.100	1.93	0.002720
67	T007 – 6	2	50	9	15	0.65	0.15	0.100	1.68	0.004320
68	T007 – 7	2	50	9	20	0.65	0.15	0.100	1.92	0.003940
69	T007 – 9	2	75	9	11	0.65	0.15	0.100	1.95	0.002840
70	T007 – 10	2	75	9	15	0.65	0.15	0.100	1.69	0.002550
71	T007 – 11	2	75	9	20	0.65	0.15	0.100	1.97	0.005990
72	T007 – 12	2	75	9	25	0.65	0.15	0.100	2.04	0.006980
73	T007 – 13	1	0	9	-	0.65	0.15	0.100	1.92	0.001790
74	T008 – 1	2	25	10	11	0.65	0.15	0.125	1.58	0.005450
75	T008 – 2	2	25	10	15	0.65	0.15	0.125	1.59	0.006270
76	T008 – 3	2	25	10	20	0.65	0.15	0.125	1.64	0.007170
77	T008 – 4	2	50	10	25	0.65	0.15	0.125	1.64	0.006920
78	T008 – 5	2	50	10	11	0.65	0.15	0.125	1.59	0.005010
79	T008 – 6	2	50	10	15	0.65	0.15	0.125	1.43	0.005970
80	T008 – 7	2	50	10	20	0.65	0.15	0.125	1.71	0.011470
81	T008 – 9	2	75	10	11	0.65	0.15	0.125	1.58	0.004790

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
82	T008 – 10	2	75	10	15	0.65	0.15	0.125	1.39	0.007480
83	T008 – 13	1	0	10	-	0.65	0.15	0.125	1.71	0.004940
84	T009 – 1	2	25	7	11	0.60	0.20	0.100	2.43	0.000357
85	T009 – 2	2	25	7	15	0.60	0.20	0.100	2.59	0.000475
86	T009 – 3	2	25	7	20	0.60	0.20	0.100	2.65	0.000630
87	T009 – 4	2	25	7	25	0.60	0.20	0.100	2.61	0.000452
88	T009 – 5	2	50	7	11	0.60	0.20	0.100	2.54	0.000723
89	T009 – 6	2	50	7	15	0.60	0.20	0.100	2.22	0.000828
90	T009 – 7	2	50	7	20	0.60	0.20	0.100	2.52	0.001070
91	T009 – 8	2	50	7	25	0.60	0.20	0.100	2.57	0.001210
92	T009 – 9	2	75	7	11	0.60	0.20	0.100	2.59	0.001150
93	T009 – 10	2	75	7	15	0.60	0.20	0.100	2.98	0.001620
94	T009 – 11	2	75	7	20	0.60	0.20	0.100	2.87	0.002350
95	T009 – 13	1	0	7	-	0.60	0.20	0.100	2.43	0.000244
96	T010 – 1	2	25	8	11	0.60	0.20	0.100	2.50	0.000424
97	T010 – 2	2	25	8	15	0.60	0.20	0.100	2.34	0.000550
98	T010 – 3	2	25	8	20	0.60	0.20	0.100	2.39	0.000738
99	T010 – 4	2	25	8	25	0.60	0.20	0.100	2.66	0.001110
100	T010 – 5	2	50	8	11	0.60	0.20	0.100	2.29	0.000643
101	T010 – 6	2	50	8	15	0.60	0.20	0.100	2.09	0.001010
102	T010 – 7	2	50	8	20	0.60	0.20	0.100	2.31	0.001440
103	T010 – 8	2	50	8	25	0.60	0.20	0.100	2.33	0.001520
104	T010 – 9	2	75	8	11	0.60	0.20	0.100	2.14	0.000895
105	T010 – 10	2	75	8	15	0.60	0.20	0.100	2.63	0.002060
106	T010 – 13	1	0	8	-	0.60	0.20	0.100	2.25	0.000265
107	T011 – 1	2	25	9	11	0.60	0.20	0.125	2.06	0.001270
108	T011 – 2	2	25	9	15	0.60	0.20	0.125	2.34	0.001900
109	T011 – 3	2	25	9	20	0.60	0.20	0.125	2.00	0.001700
110	T011 – 4	2	25	9	25	0.60	0.20	0.125	1.85	0.001540
111	T011 – 5	2	50	9	11	0.60	0.20	0.125	1.91	0.001700
112	T011 – 6	2	50	9	15	0.60	0.20	0.125	1.89	0.003290
113	T011 – 7	2	50	9	20	0.60	0.20	0.125	2.15	0.004130

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
114	T011 – 9	2	75	9	11	0.60	0.20	0.125	1.54	0.002060
115	T011 – 10	2	75	9	15	0.60	0.20	0.125	2.37	0.004670
116	T011 – 13	1	0	9	-	0.60	0.20	0.125	2.04	0.000990
117	T0012 – 1	2	25	10	11	0.60	0.20	0.125	1.94	0.002820
118	T0012 – 2	2	25	10	15	0.60	0.20	0.125	1.97	0.003950
119	T0012 – 3	2	25	10	20	0.60	0.20	0.125	2.04	0.004680
120	T0012 – 4	2	25	10	25	0.60	0.20	0.125	2.24	0.004240
121	T0012 – 5	2	50	10	11	0.60	0.20	0.125	2.12	0.003330
122	T0012 – 6	2	50	10	15	0.60	0.20	0.125	1.90	0.005230
123	T0012 – 9	2	75	10	11	0.60	0.20	0.125	2.11	0.003470
124	T0012 – 10	2	75	10	15	0.60	0.20	0.125	2.16	0.007130
125	T0012 – 13	1	0	10	-	0.60	0.20	0.125	2.10	0.001850
126	T013 – 1	2	25	7	11	0.65	0.15	0.075	2.78	0.000510
127	T013 – 2	2	25	7	15	0.65	0.15	0.075	2.53	0.000508
128	T013 – 3	2	25	7	20	0.65	0.15	0.075	2.57	0.000339
129	T013 – 4	2	25	7	25	0.65	0.15	0.075	2.58	0.000250
130	T013 – 5	2	50	7	11	0.65	0.15	0.075	2.19	0.000374
131	T013 – 6	2	50	7	15	0.65	0.15	0.075	2.35	0.000634
132	T013 – 7	2	50	7	20	0.65	0.15	0.075	2.73	0.000742
133	T013 – 8	2	50	7	25	0.65	0.15	0.075	2.87	0.000427
134	T013 – 9	2	75	7	11	0.65	0.15	0.075	2.62	0.001120
135	T013 – 10	2	75	7	15	0.65	0.15	0.075	2.54	0.001220
136	T013 – 11	2	75	7	20	0.65	0.15	0.075	2.51	0.001530
137	T013 – 12	2	75	7	25	0.65	0.15	0.075	2.40	0.001310
138	T013 – 13	1	0	7	-	0.65	0.15	0.075	2.77	0.000177
139	T014 – 1	2	25	8	11	0.65	0.15	0.075	2.46	0.000786
140	T014 – 2	2	25	8	15	0.65	0.15	0.075	2.15	0.000518
141	T014 – 3	2	25	8	20	0.65	0.15	0.075	2.29	0.000477
142	T014 – 4	2	25	8	25	0.65	0.15	0.075	2.63	0.000468
143	T014 – 5	2	50	8	11	0.65	0.15	0.075	2.52	0.000924
144	T014 – 6	2	50	8	15	0.65	0.15	0.075	2.13	0.001140
145	T014 – 7	2	50	8	20	0.65	0.15	0.075	2.37	0.000803

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
146	T014-8	2	50	8	25	0.65	0.15	0.075	2.76	0.000754
147	T014-9	2	75	8	11	0.65	0.15	0.075	2.63	0.001080
148	T014-10	2	75	8	15	0.65	0.15	0.075	2.26	0.001230
149	T014-11	2	75	8	20	0.65	0.15	0.075	2.50	0.001850
150	T014-12	2	75	8	25	0.65	0.15	0.075	2.50	0.001040
151	T014-13	1	0	8	-	0.65	0.15	0.075	2.43	0.000478
152	T015-1	2	25	9	11	0.65	0.15	0.075	2.55	0.001010
153	T015-2	2	25	9	15	0.65	0.15	0.075	2.25	0.000926
154	T015-3	2	25	9	20	0.65	0.15	0.075	2.68	0.001030
155	T015-4	2	25	9	25	0.65	0.15	0.075	2.70	0.000973
156	T015-5	2	50	9	11	0.65	0.15	0.075	2.59	0.000915
157	T015-6	2	50	9	15	0.65	0.15	0.075	2.37	0.001000
158	T015-7	2	50	9	20	0.65	0.15	0.075	2.42	0.001240
159	T015-8	2	50	9	25	0.65	0.15	0.075	2.82	0.001120
160	T015-9	2	75	9	11	0.65	0.15	0.075	2.73	0.001740
161	T015-10	2	75	9	15	0.65	0.15	0.075	2.23	0.001520
162	T015-11	2	75	9	20	0.65	0.15	0.075	2.50	0.001480
163	T015-12	2	75	9	25	0.65	0.15	0.075	2.47	0.001500
164	T015-13	1	0	9	-	0.65	0.15	0.075	2.55	0.000893
165	T016-1	2	25	10	11	0.65	0.15	0.075	2.11	0.000912
166	T016-2	2	25	10	15	0.65	0.15	0.075	3.04	0.002528
167	T016-3	2	25	10	20	0.65	0.15	0.075	2.69	0.001815
168	T016-4	2	25	10	25	0.65	0.15	0.075	2.70	0.001511
169	T016-5	2	50	10	11	0.65	0.15	0.075	2.33	0.001175
170	T016-6	2	50	10	15	0.65	0.15	0.075	2.41	0.001337
171	T016-7	2	50	10	20	0.65	0.15	0.075	2.44	0.002092
172	T016-8	2	50	10	25	0.65	0.15	0.075	2.80	0.001545
173	T016-9	2	75	10	11	0.65	0.15	0.075	2.63	0.001410
174	T016-10	2	75	10	15	0.65	0.15	0.075	2.22	0.001275
175	T016-11	2	75	10	20	0.65	0.15	0.075	2.49	0.002271
176	T016-12	2	75	10	25	0.65	0.15	0.075	2.91	0.001911
177	T016-13	1	0	10	-	0.65	0.15	0.075	3.01	0.000157

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
178	T017-1	2	25	7	11	0.70	0.10	0.100	1.37	0.004340
179	T017-2	2	25	7	15	0.70	0.10	0.100	1.29	0.004920
180	T017-3	2	25	7	20	0.70	0.10	0.100	1.30	0.005440
181	T017-4	2	25	7	25	0.70	0.10	0.100	1.31	0.005600
182	T017-5	2	50	7	11	0.70	0.10	0.100	1.43	0.007200
183	T017-6	2	50	7	15	0.70	0.10	0.100	1.20	0.007240
184	T017-7	2	50	7	20	0.70	0.10	0.100	1.24	0.008110
185	T017-8	2	50	7	25	0.70	0.10	0.100	1.48	0.012130
186	T017-9	2	75	7	11	0.70	0.10	0.100	1.46	0.008970
187	T017-10	2	75	7	15	0.70	0.10	0.100	1.21	0.010690
188	T017-11	2	75	7	20	0.70	0.10	0.100	1.18	0.012700
189	T017-13	1	0	7	-	0.70	0.10	0.100	1.39	0.004540
190	T018-1	2	25	8	11	0.70	0.10	0.100	1.41	0.006520
191	T018-2	2	25	8	15	0.70	0.10	0.100	1.24	0.005780
192	T018-3	2	25	8	20	0.70	0.10	0.100	1.29	0.007310
193	T018-4	2	25	8	25	0.70	0.10	0.100	1.25	0.005850
194	T018-5	2	50	8	11	0.70	0.10	0.100	1.28	0.007280
195	T018-6	2	50	8	15	0.70	0.10	0.100	1.25	0.009070
196	T018-7	2	50	8	20	0.70	0.10	0.100	1.23	0.011620
197	T018-8	2	50	8	25	0.70	0.10	0.100	1.43	0.011000
198	T018-9	2	75	8	11	0.70	0.10	0.100	1.31	0.011910
199	T018-10	2	75	8	15	0.70	0.10	0.100	1.22	0.010820
200	T018-11	2	75	8	20	0.70	0.10	0.100	1.35	0.013020
201	T018-13	1	0	8	-	0.70	0.10	0.100	1.40	0.007440
202	T019-1	2	25	9	11	0.70	0.10	0.100	1.28	0.009620
203	T019-2	2	25	9	15	0.70	0.10	0.100	1.16	0.009950
204	T019-3	2	25	9	20	0.70	0.10	0.100	1.30	0.008550
205	T019-4	2	25	9	25	0.70	0.10	0.100	1.30	0.008740
206	T019-5	2	50	9	11	0.70	0.10	0.100	1.29	0.009230
207	T019-6	2	50	9	15	0.70	0.10	0.100	1.12	0.008850
208	T019-7	2	50	9	20	0.70	0.10	0.100	1.25	0.012200
209	T019-8	2	50	9	25	0.70	0.10	0.100	1.35	0.015080

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
210	T019 – 9	2	75	9	11	0.70	0.10	0.100	1.31	0.011080
211	T019 – 10	2	75	9	15	0.70	0.10	0.100	1.12	0.011920
212	T019 – 11	2	75	9	20	0.70	0.10	0.100	1.49	0.018130
213	T019 – 13	1	0	9	-	0.70	0.10	0.100	1.29	0.006480
214	T020 – 1	2	25	10	11	0.70	0.10	0.100	1.48	0.012300
215	T020 – 2	2	25	10	15	0.70	0.10	0.100	1.20	0.011310
216	T020 – 3	2	25	10	20	0.70	0.10	0.100	1.36	0.012040
217	T020 – 4	2	25	10	25	0.70	0.10	0.100	1.49	0.013050
218	T020 – 5	2	50	10	11	0.70	0.10	0.100	1.48	0.013430
219	T020 – 6	2	50	10	15	0.70	0.10	0.100	1.25	0.012750
220	T020 – 7	2	50	10	20	0.70	0.10	0.100	1.25	0.015850
221	T020 – 8	2	50	10	25	0.70	0.10	0.100	1.32	0.017190
222	T020 – 9	2	75	10	11	0.70	0.10	0.100	1.32	0.011430
223	T020 – 10	2	75	10	15	0.70	0.10	0.100	1.19	0.013890
224	T020 – 11	2	75	10	20	0.70	0.10	0.100	1.31	0.019890
225	T020 – 13	1	0	10	-	0.70	0.10	0.100	1.46	0.008820
226	T021 – 1	2	25	7	11	0.60	0.20	0.075	3.18	0.000051
227	T021 – 2	2	25	7	15	0.60	0.20	0.075	2.97	0.000091
228	T021 – 3	2	25	7	20	0.60	0.20	0.075	3.40	0.000098
229	T021 – 4	2	25	7	25	0.60	0.20	0.075	3.19	0.000068
230	T021 – 5	2	50	7	11	0.60	0.20	0.075	3.32	0.000072
231	T021 – 6	2	50	7	15	0.60	0.20	0.075	2.78	0.000165
232	T021 – 7	2	50	7	20	0.60	0.20	0.075	3.16	0.000154
233	T021 – 8	2	50	7	25	0.60	0.20	0.075	3.94	0.000092
234	T021 – 9	2	75	7	11	0.60	0.20	0.075	3.47	0.000231
235	T021 – 10	2	75	7	15	0.60	0.20	0.075	3.02	0.000375
236	T021 – 11	2	75	7	20	0.60	0.20	0.075	3.46	0.000521
237	T021 – 12	2	75	7	25	0.60	0.20	0.075	2.80	0.000434
238	T021 – 13	1	0	7	-	0.60	0.20	0.075	3.20	0.000041
239	T022 – 1	2	25	8	11	0.60	0.20	0.075	3.25	0.000092
240	T022 – 2	2	25	8	15	0.60	0.20	0.075	2.86	0.000114
241	T022 – 3	2	25	8	20	0.60	0.20	0.075	3.49	0.000156

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
242	T022-4	2	25	8	25	0.60	0.20	0.075	3.61	0.000108
243	T022-5	2	50	8	11	0.60	0.20	0.075	3.35	0.000240
244	T022-6	2	50	8	15	0.60	0.20	0.075	3.18	0.000303
245	T022-7	2	50	8	20	0.60	0.20	0.075	3.62	0.000216
246	T022-8	2	50	8	25	0.60	0.20	0.075	3.61	0.000370
247	T022-9	2	75	8	11	0.60	0.20	0.075	3.45	0.000321
248	T022-10	2	75	8	15	0.60	0.20	0.075	3.01	0.000318
249	T022-11	2	75	8	20	0.60	0.20	0.075	3.33	0.000500
250	T022-13	1	0	8	-	0.60	0.20	0.075	3.22	0.000079
251	T023-1	2	25	9	11	0.60	0.20	0.075	3.37	0.000190
252	T023-2	2	25	9	15	0.60	0.20	0.075	3.40	0.000269
253	T023-3	2	25	9	20	0.60	0.20	0.075	3.57	0.000185
254	T023-4	2	25	9	25	0.60	0.20	0.075	3.59	0.000121
255	T023-5	2	50	9	11	0.60	0.20	0.075	3.02	0.000165
256	T023-6	2	50	9	15	0.60	0.20	0.075	3.60	0.000214
257	T023-7	2	50	9	20	0.60	0.20	0.075	3.66	0.000256
258	T023-8	2	50	9	25	0.60	0.20	0.075	3.72	0.000238
259	T023-9	2	75	9	11	0.60	0.20	0.075	3.54	0.000368
260	T023-10	2	75	9	15	0.60	0.20	0.075	2.96	0.000491
261	T023-11	2	75	9	20	0.60	0.20	0.075	3.33	0.000408
262	T023-12	2	75	9	25	0.60	0.20	0.075	3.19	0.000432
263	T023-13	1	0	9	-	0.60	0.20	0.075	3.78	0.000020
264	T024-1	2	25	10	11	0.60	0.20	0.075	3.49	0.000154
265	T024-2	2	25	10	15	0.60	0.20	0.075	3.53	0.000372
266	T024-3	2	25	10	20	0.60	0.20	0.075	3.54	0.000257
267	T024-4	2	25	10	25	0.60	0.20	0.075	3.62	0.000232
268	T024-5	2	50	10	11	0.60	0.20	0.075	3.49	0.000235
269	T024-6	2	50	10	15	0.60	0.20	0.075	3.21	0.000324
270	T024-7	2	50	10	20	0.60	0.20	0.075	3.26	0.000463
271	T024-8	2	50	10	25	0.60	0.20	0.075	3.51	0.000258
272	T024-9	2	75	10	11	0.60	0.20	0.075	3.81	0.000147
273	T024-10	2	75	10	15	0.60	0.20	0.075	2.95	0.000254

(cot $\alpha = 3.0$) Continue on the next page

Table B.1 Dimensionless Wave Overtopping Results for $\cot \alpha = 3.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
274	T024 – 11	2	75	10	20	0.60	0.20	0.075	3.25	0.000661
275	T024 – 12	2	75	10	25	0.60	0.20	0.075	3.37	0.000353
276	T024 – 13	1	0	10	-	0.60	0.20	0.075	3.53	0.000032

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
1	T001 – 1	2	25	7	11	0.70	0.10	0.125	1.00	0.008450
2	T001 – 2	2	25	7	15	0.70	0.10	0.125	1.01	0.007560
3	T001 – 3	2	25	7	20	0.70	0.10	0.125	1.21	0.005370
4	T001 – 4	2	25	7	25	0.70	0.10	0.125	1.01	0.007970
5	T001 – 5	2	50	7	11	0.70	0.10	0.125	1.05	0.008760
6	T001 – 6	2	50	7	15	0.70	0.10	0.125	1.08	0.008270
7	T001 – 7	2	50	7	20	0.70	0.10	0.125	0.98	0.008910
8	T001 – 9	2	75	7	11	0.70	0.10	0.125	1.07	0.011490
9	T001 – 10	2	75	7	15	0.70	0.10	0.125	0.91	0.018520
10	T001 – 13	1	0	7	-	0.70	0.10	0.125	1.10	0.007429
11	T002 – 1	2	25	8	11	0.70	0.10	0.125	0.91	0.009690
12	T002 – 2	2	25	8	15	0.70	0.10	0.125	0.96	0.009410
13	T002 – 3	2	25	8	20	0.70	0.10	0.125	1.02	0.007450
14	T002 – 4	2	25	8	25	0.70	0.10	0.125	1.02	0.010420
15	T002 – 5	2	50	8	11	0.70	0.10	0.125	0.92	0.010780
16	T002 – 6	2	50	8	15	0.70	0.10	0.125	0.97	0.012240
17	T002 – 7	2	50	8	20	0.70	0.10	0.125	1.23	0.018520
18	T002 – 8	2	50	8	25	0.70	0.10	0.125	0.88	0.014800
19	T002 – 9	2	75	8	11	0.70	0.10	0.125	0.94	0.012000
20	T002 – 10	2	75	8	15	0.70	0.10	0.125	1.00	0.017380
21	T002 – 11	2	75	8	20	0.70	0.10	0.125	0.91	0.013620
22	T002 – 13	1	0	8	-	0.70	0.10	0.125	1.04	0.007203

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
23	T003 – 1	2	25	9	11	0.70	0.10	0.125	1.06	0.009150
24	T003 – 2	2	25	9	15	0.70	0.10	0.125	1.28	0.006480
25	T003 – 3	2	25	9	20	0.70	0.10	0.125	1.24	0.005410
26	T003 – 4	2	25	9	25	0.70	0.10	0.125	1.26	0.015460
27	T003 – 5	2	50	9	11	0.70	0.10	0.125	0.98	0.008740
28	T003 – 6	2	50	9	15	0.70	0.10	0.125	0.93	0.015770
29	T003 – 7	2	50	9	20	0.70	0.10	0.125	1.00	0.011670
30	T003 – 9	2	75	9	11	0.70	0.10	0.125	0.98	0.010830
31	T003 – 10	2	75	9	15	0.70	0.10	0.125	1.12	0.015640
32	T003 – 13	1	0	9	-	0.70	0.10	0.125	1.16	0.010406
33	T004 – 1	2	25	10	11	0.70	0.10	0.125	1.05	0.013020
34	T004 – 2	2	25	10	15	0.70	0.10	0.125	0.98	0.012150
35	T004 – 3	2	25	10	20	0.70	0.10	0.125	1.05	0.011790
36	T004 – 4	2	25	10	25	0.70	0.10	0.125	1.06	0.010110
37	T004 – 5	2	50	10	11	0.70	0.10	0.125	1.05	0.011350
38	T004 – 6	2	50	10	15	0.70	0.10	0.125	1.09	0.010270
39	T004 – 9	2	75	10	11	0.70	0.10	0.125	1.06	0.011440
40	T004 – 10	2	75	10	15	0.70	0.10	0.125	1.03	0.011330
41	T004 – 13	1	0	10	-	0.70	0.10	0.125	1.16	0.016328
42	T005 – 1	2	25	7	11	0.65	0.15	0.100	1.55	0.006280
43	T005 – 2	2	25	7	15	0.65	0.15	0.100	1.54	0.004800
44	T005 – 3	2	25	7	20	0.65	0.15	0.100	1.46	0.002290
45	T005 – 4	2	25	7	25	0.65	0.15	0.100	1.68	0.006590
46	T005 – 5	2	50	7	11	0.65	0.15	0.100	1.31	0.003290
47	T005 – 6	2	50	7	15	0.65	0.15	0.100	1.41	0.002630
48	T005 – 7	2	50	7	20	0.65	0.15	0.100	1.74	0.004550
49	T005 – 8	2	50	7	25	0.65	0.15	0.100	1.40	0.009270
50	T005 – 9	2	75	7	11	0.65	0.15	0.100	1.53	0.001820
51	T005 – 10	2	75	7	15	0.65	0.15	0.100	1.85	0.004470
52	T005 – 11	2	75	7	20	0.65	0.15	0.100	1.35	0.002940
53	T005 – 13	1	0	7	-	0.65	0.15	0.100	1.44	0.002830
54	T006 – 1	2	25	8	11	0.65	0.15	0.125	1.15	0.021850

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
55	T006-2	2	25	8	15	0.65	0.15	0.125	1.30	0.011050
56	T006-3	2	25	8	20	0.65	0.15	0.125	1.13	0.005840
57	T006-4	2	25	8	25	0.65	0.15	0.125	1.18	0.010890
58	T006-5	2	50	8	11	0.65	0.15	0.125	1.19	0.016310
59	T006-6	2	50	8	15	0.65	0.15	0.125	1.07	0.006560
60	T006-7	2	50	8	20	0.65	0.15	0.125	1.00	0.008090
61	T006-9	2	75	8	11	0.65	0.15	0.125	1.20	0.004660
62	T006-10	2	75	8	15	0.65	0.15	0.125	1.25	0.013360
63	T006-13	1	0	8	-	0.65	0.15	0.125	1.05	0.004880
64	T007-1	2	25	9	11	0.65	0.15	0.100	1.25	0.003690
65	T007-2	2	25	9	15	0.65	0.15	0.100	1.61	0.005750
66	T007-3	2	25	9	20	0.65	0.15	0.100	1.49	0.007900
67	T007-4	2	25	9	25	0.65	0.15	0.100	1.48	0.005430
68	T007-5	2	50	9	11	0.65	0.15	0.100	1.41	0.003660
69	T007-6	2	50	9	15	0.65	0.15	0.100	1.40	0.002920
70	T007-7	2	50	9	20	0.65	0.15	0.100	1.21	0.004510
71	T007-9	2	75	9	11	0.65	0.15	0.100	1.51	0.002350
72	T007-10	2	75	9	15	0.65	0.15	0.100	1.46	0.002780
73	T007-11	2	75	9	20	0.65	0.15	0.100	1.23	0.003530
74	T007-12	2	75	9	25	0.65	0.15	0.100	1.41	0.006100
75	T007-13	1	0	9	-	0.65	0.15	0.100	1.36	0.002110
76	T008-1	2	25	10	11	0.65	0.15	0.125	1.25	0.008920
77	T008-2	2	25	10	15	0.65	0.15	0.125	1.21	0.003810
78	T008-3	2	25	10	20	0.65	0.15	0.125	1.07	0.022170
79	T008-4	2	25	10	25	0.65	0.15	0.125	1.13	0.012180
80	T008-5	2	50	10	11	0.65	0.15	0.125	1.20	0.003950
81	T008-6	2	50	10	15	0.65	0.15	0.125	1.35	0.012730
82	T008-7	2	50	10	20	0.65	0.15	0.125	1.07	0.005950
83	T008-9	2	75	10	11	0.65	0.15	0.125	1.38	0.004800
84	T008-10	2	75	10	15	0.65	0.15	0.125	1.17	0.004280
85	T008-13	1	0	10	-	0.65	0.15	0.125	1.20	0.004593
86	T009-1	2	25	7	11	0.60	0.20	0.100	2.49	0.000841

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
87	T009 – 2	2	25	7	15	0.60	0.20	0.100	2.77	0.000921
88	T009 – 3	2	25	7	20	0.60	0.20	0.100	2.74	0.001150
89	T009 – 4	2	25	7	25	0.60	0.20	0.100	3.07	0.001730
90	T009 – 5	2	50	7	11	0.60	0.20	0.100	2.83	0.001020
91	T009 – 6	2	50	7	15	0.60	0.20	0.100	2.59	0.001040
92	T009 – 7	2	50	7	20	0.60	0.20	0.100	2.68	0.000861
93	T009 – 8	2	50	7	25	0.60	0.20	0.100	2.86	0.000898
94	T009 – 9	2	75	7	11	0.60	0.20	0.100	2.60	0.000673
95	T009 – 10	2	75	7	15	0.60	0.20	0.100	2.68	0.000937
96	T009 – 11	2	75	7	20	0.60	0.20	0.100	2.59	0.001190
97	T009 – 13	1	0	7	-	0.60	0.20	0.100	2.60	0.000167
98	T010 – 1	2	25	8	11	0.60	0.20	0.100	2.57	0.001100
99	T010 – 2	2	25	8	15	0.60	0.20	0.100	2.56	0.001320
100	T010 – 3	2	25	8	20	0.60	0.20	0.100	2.56	0.000919
101	T010 – 4	2	25	8	25	0.60	0.20	0.100	2.57	0.001060
102	T010 – 5	2	50	8	11	0.60	0.20	0.100	2.14	0.000671
103	T010 – 6	2	50	8	15	0.60	0.20	0.100	2.18	0.000690
104	T010 – 7	2	50	8	20	0.60	0.20	0.100	2.26	0.000545
105	T010 – 8	2	50	8	25	0.60	0.20	0.100	2.23	0.000714
106	T010 – 9	2	75	8	11	0.60	0.20	0.100	2.15	0.000506
107	T010 – 10	2	75	8	15	0.60	0.20	0.100	2.68	0.001010
108	T011 – 1	2	25	9	11	0.60	0.20	0.125	2.23	0.000833
109	T011 – 2	2	25	9	15	0.60	0.20	0.125	2.15	0.001490
110	T011 – 3	2	25	9	20	0.60	0.20	0.125	2.17	0.000455
111	T011 – 4	2	25	9	25	0.60	0.20	0.125	1.94	0.000702
112	T011 – 5	2	50	9	11	0.60	0.20	0.125	1.72	0.001290
113	T011 – 6	2	50	9	15	0.60	0.20	0.125	1.91	0.002270
114	T011 – 7	2	50	9	20	0.60	0.20	0.125	1.84	0.002310
115	T011 – 9	2	75	9	11	0.60	0.20	0.125	1.73	0.002610
116	T011 – 10	2	75	9	15	0.60	0.20	0.125	1.78	0.002520
117	T011 – 13	1	0	9	-	0.60	0.20	0.125	1.99	0.002450
118	T0012 – 1	2	25	10	11	0.60	0.20	0.125	1.67	0.002810

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
119	T0012-2	2	25	10	15	0.60	0.20	0.125	1.66	0.002970
120	T0012-3	2	25	10	20	0.60	0.20	0.125	1.68	0.002770
121	T0012-4	2	25	10	25	0.60	0.20	0.125	1.72	0.002170
122	T0012-5	2	50	10	11	0.60	0.20	0.125	1.68	0.002410
123	T0012-6	2	50	10	15	0.60	0.20	0.125	1.72	0.003210
124	T0012-9	2	75	10	11	0.60	0.20	0.125	1.70	0.002490
125	T0012-10	2	75	10	15	0.60	0.20	0.125	1.74	0.003410
126	T0012-13	1	0	10	-	0.60	0.20	0.125	1.79	0.001800
127	T013-1	2	25	7	11	0.65	0.15	0.075	1.90	0.002880
128	T013-2	2	25	7	15	0.65	0.15	0.075	1.71	0.001020
129	T013-3	2	25	7	20	0.65	0.15	0.075	1.83	0.003130
130	T013-4	2	25	7	25	0.65	0.15	0.075	1.99	0.000597
131	T013-5	2	50	7	11	0.65	0.15	0.075	1.97	0.000494
132	T013-6	2	50	7	15	0.65	0.15	0.075	1.84	0.000537
133	T013-7	2	50	7	20	0.65	0.15	0.075	1.85	0.001080
134	T013-8	2	50	7	25	0.65	0.15	0.075	1.96	0.001650
135	T013-9	2	75	7	11	0.65	0.15	0.075	1.59	0.003470
136	T013-10	2	75	7	15	0.65	0.15	0.075	1.66	0.001980
137	T013-11	2	75	7	20	0.65	0.15	0.075	1.86	0.000874
138	T013-12	2	75	7	25	0.65	0.15	0.075	1.93	0.001270
139	T013-13	1	0	7	-	0.65	0.15	0.075	1.74	0.001080
140	T014-1	2	25	8	11	0.65	0.15	0.075	1.66	0.001300
141	T014-2	2	25	8	15	0.65	0.15	0.075	1.60	0.004340
142	T014-3	2	25	8	20	0.65	0.15	0.075	1.75	0.002090
143	T014-4	2	25	8	25	0.65	0.15	0.075	2.10	0.002870
144	T014-5	2	50	8	11	0.65	0.15	0.075	1.75	0.000857
145	T014-6	2	50	8	15	0.65	0.15	0.075	1.64	0.003780
146	T014-7	2	50	8	20	0.65	0.15	0.075	2.01	0.003400
147	T014-8	2	50	8	25	0.65	0.15	0.075	1.86	0.000709
148	T014-9	2	75	8	11	0.65	0.15	0.075	2.00	0.000273
149	T014-10	2	75	8	15	0.65	0.15	0.075	2.12	0.000886
150	T014-11	2	75	8	20	0.65	0.15	0.075	1.60	0.001360

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
151	T014 – 12	2	75	8	25	0.65	0.15	0.075	1.55	0.001460
152	T014 – 13	1	0	8	-	0.65	0.15	0.075	1.58	0.002110
153	T015 – 1	2	25	9	11	0.65	0.15	0.075	2.04	0.002090
154	T015 – 2	2	25	9	15	0.65	0.15	0.075	1.81	0.000944
155	T015 – 3	2	25	9	20	0.65	0.15	0.075	1.80	0.000715
156	T015 – 4	2	25	9	25	0.65	0.15	0.075	1.82	0.001260
157	T015 – 5	2	50	9	11	0.65	0.15	0.075	1.99	0.000336
158	T015 – 6	2	50	9	15	0.65	0.15	0.075	2.04	0.000371
159	T015 – 7	2	50	9	20	0.65	0.15	0.075	1.69	0.004600
160	T015 – 8	2	50	9	25	0.65	0.15	0.075	1.81	0.001930
161	T015 – 9	2	75	9	11	0.65	0.15	0.075	1.70	0.001730
162	T015 – 10	2	75	9	15	0.65	0.15	0.075	2.06	0.000459
163	T015 – 11	2	75	9	20	0.65	0.15	0.075	1.84	0.001800
164	T015 – 12	2	75	9	25	0.65	0.15	0.075	1.90	0.000609
165	T015 – 13	1	0	9	-	0.65	0.15	0.075	1.88	0.000861
166	T016 – 1	2	25	10	11	0.65	0.15	0.075	2.07	0.000240
167	T016 – 2	2	25	10	15	0.65	0.15	0.075	2.33	0.001280
168	T016 – 3	2	25	10	20	0.65	0.15	0.075	2.19	0.001090
169	T016 – 4	2	25	10	25	0.65	0.15	0.075	1.99	0.000387
170	T016 – 5	2	50	10	11	0.65	0.15	0.075	1.91	0.000440
171	T016 – 6	2	50	10	15	0.65	0.15	0.075	2.03	0.001280
172	T016 – 7	2	50	10	20	0.65	0.15	0.075	1.77	0.001300
173	T016 – 8	2	50	10	25	0.65	0.15	0.075	1.85	0.001420
174	T016 – 12	2	75	10	25	0.65	0.15	0.075	1.59	0.002150
175	T016 – 13	1	0	10	-	0.65	0.15	0.075	2.00	0.000532
176	T017 – 1	2	25	7	11	0.70	0.10	0.100	1.24	0.005010
177	T017 – 2	2	25	7	15	0.70	0.10	0.100	1.25	0.004210
178	T017 – 3	2	25	7	20	0.70	0.10	0.100	1.24	0.003900
179	T017 – 4	2	25	7	25	0.70	0.10	0.100	1.28	0.005370
180	T017 – 5	2	50	7	11	0.70	0.10	0.100	1.28	0.003620
181	T017 – 6	2	50	7	15	0.70	0.10	0.100	1.31	0.004800
182	T017 – 7	2	50	7	20	0.70	0.10	0.100	1.30	0.003870

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
183	T017-8	2	50	7	25	0.70	0.10	0.100	1.34	0.004780
184	T017-9	2	75	7	11	0.70	0.10	0.100	1.29	0.004480
185	T017-10	2	75	7	15	0.70	0.10	0.100	1.12	0.005140
186	T017-13	1	0	7	-	0.70	0.10	0.100	1.22	0.004300
187	T018-1	2	25	8	11	0.70	0.10	0.100	1.28	0.004780
188	T018-2	2	25	8	15	0.70	0.10	0.100	1.27	0.004410
189	T018-3	2	25	8	20	0.70	0.10	0.100	1.27	0.004680
190	T018-4	2	25	8	25	0.70	0.10	0.100	1.28	0.003670
191	T018-5	2	50	8	11	0.70	0.10	0.100	1.28	0.005290
192	T018-6	2	50	8	15	0.70	0.10	0.100	1.41	0.011890
193	T018-7	2	50	8	20	0.70	0.10	0.100	1.30	0.004760
194	T018-8	2	50	8	25	0.70	0.10	0.100	1.29	0.003240
195	T018-9	2	75	8	11	0.70	0.10	0.100	1.30	0.004980
196	T018-10	2	75	8	15	0.70	0.10	0.100	1.35	0.005160
197	T018-13	1	0	8	-	0.70	0.10	0.100	1.17	0.004100
198	T019-1	2	25	9	11	0.70	0.10	0.100	1.29	0.007690
199	T019-2	2	25	9	15	0.70	0.10	0.100	1.29	0.006770
200	T019-3	2	25	9	20	0.70	0.10	0.100	1.30	0.004690
201	T019-4	2	25	9	25	0.70	0.10	0.100	1.30	0.004820
202	T019-5	2	50	9	11	0.70	0.10	0.100	1.29	0.005250
203	T019-6	2	50	9	15	0.70	0.10	0.100	1.48	0.009690
204	T019-7	2	50	9	20	0.70	0.10	0.100	1.31	0.005760
205	T019-8	2	50	9	25	0.70	0.10	0.100	1.35	0.006540
206	T019-9	2	75	9	11	0.70	0.10	0.100	1.08	0.006630
207	T019-10	2	75	9	15	0.70	0.10	0.100	1.34	0.007040
208	T019-11	2	75	9	20	0.70	0.10	0.100	1.34	0.005860
209	T019-13	1	0	9	-	0.70	0.10	0.100	1.29	0.005550
210	T020-1	2	25	10	11	0.70	0.10	0.100	1.24	0.007750
211	T020-2	2	25	10	15	0.70	0.10	0.100	1.06	0.007740
212	T020-3	2	25	10	20	0.70	0.10	0.100	1.24	0.004880
213	T020-4	2	25	10	25	0.70	0.10	0.100	1.28	0.005230
214	T020-5	2	50	10	11	0.70	0.10	0.100	1.59	0.009090

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
215	T020 – 6	2	50	10	15	0.70	0.10	0.100	1.31	0.006770
216	T020 – 7	2	50	10	20	0.70	0.10	0.100	1.30	0.006210
217	T020 – 8	2	50	10	25	0.70	0.10	0.100	1.34	0.006960
218	T020 – 9	2	75	10	11	0.70	0.10	0.100	1.29	0.006710
219	T020 – 13	1	0	10	-	0.70	0.10	0.100	1.27	0.005140
220	T021 – 1	2	25	7	11	0.60	0.20	0.075	3.20	0.000213
221	T021 – 2	2	25	7	15	0.60	0.20	0.075	2.85	0.000214
222	T021 – 3	2	25	7	20	0.60	0.20	0.075	2.84	0.000143
223	T021 – 4	2	25	7	25	0.60	0.20	0.075	2.98	0.000172
224	T021 – 5	2	50	7	11	0.60	0.20	0.075	3.23	0.000087
225	T021 – 6	2	50	7	15	0.60	0.20	0.075	3.29	0.000080
226	T021 – 7	2	50	7	20	0.60	0.20	0.075	3.04	0.000028
227	T021 – 8	2	50	7	25	0.60	0.20	0.075	3.07	0.000012
228	T021 – 9	2	75	7	11	0.60	0.20	0.075	3.03	0.000059
229	T021 – 10	2	75	7	15	0.60	0.20	0.075	3.07	0.000028
230	T021 – 11	2	75	7	20	0.60	0.20	0.075	3.20	0.000043
231	T021 – 12	2	75	7	25	0.60	0.20	0.075	3.08	0.000061
232	T021 – 13	1	0	7	-	0.60	0.20	0.075	3.18	0.000036
233	T022 – 1	2	25	8	11	0.60	0.20	0.075	3.28	0.000200
234	T022 – 2	2	25	8	15	0.60	0.20	0.075	3.76	0.000164
235	T022 – 3	2	25	8	20	0.60	0.20	0.075	3.74	0.000109
236	T022 – 4	2	25	8	25	0.60	0.20	0.075	3.78	0.000055
237	T022 – 5	2	50	8	11	0.60	0.20	0.075	2.94	0.000132
238	T022 – 6	2	50	8	15	0.60	0.20	0.075	3.34	0.000091
239	T022 – 7	2	50	8	20	0.60	0.20	0.075	2.95	0.000008
240	T022 – 8	2	50	8	25	0.60	0.20	0.075	2.97	0.000012
241	T022 – 9	2	75	8	11	0.60	0.20	0.075	2.92	0.000037
242	T022 – 10	2	75	8	15	0.60	0.20	0.075	3.40	0.000047
243	T022 – 11	2	75	8	20	0.60	0.20	0.075	3.80	0.000055
244	T022 – 13	1	0	8	-	0.60	0.20	0.075	3.44	0.000015
245	T023 – 1	2	25	9	11	0.60	0.20	0.075	2.92	0.000112
246	T023 – 2	2	25	9	15	0.60	0.20	0.075	3.29	0.000067

(cot $\alpha = 1.5$) Continue on the next page

Table B.2 Dimensionless Wave Overtopping Results for $\cot \alpha = 1.5$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
247	T023 – 3	2	25	9	20	0.60	0.20	0.075	3.29	0.000045
248	T023 – 4	2	25	9	25	0.60	0.20	0.075	3.32	0.000043
249	T023 – 5	2	50	9	11	0.60	0.20	0.075	2.92	0.000037
250	T023 – 6	2	50	9	15	0.60	0.20	0.075	3.81	0.000039
251	T023 – 7	2	50	9	20	0.60	0.20	0.075	3.13	0.000025
252	T023 – 8	2	50	9	25	0.60	0.20	0.075	2.82	0.000025
253	T023 – 9	2	75	9	11	0.60	0.20	0.075	3.72	0.000065
254	T023 – 10	2	75	9	15	0.60	0.20	0.075	3.23	0.000052
255	T023 – 11	2	75	9	20	0.60	0.20	0.075	3.20	0.000052
256	T023 – 12	2	75	9	25	0.60	0.20	0.075	3.19	0.000043
257	T023 – 13	1	0	9	-	0.60	0.20	0.075	3.71	0.000008
258	T024 – 1	2	25	10	11	0.60	0.20	0.075	2.90	0.000071
259	T024 – 2	2	25	10	15	0.60	0.20	0.075	2.95	0.000159
260	T024 – 3	2	25	10	20	0.60	0.20	0.075	3.54	0.000106
261	T024 – 4	2	25	10	25	0.60	0.20	0.075	3.23	0.000170
262	T024 – 5	2	50	10	11	0.60	0.20	0.075	3.24	0.000222
263	T024 – 6	2	50	10	15	0.60	0.20	0.075	3.72	0.000210
264	T024 – 7	2	50	10	20	0.60	0.20	0.075	3.72	0.000114
265	T024 – 8	2	50	10	25	0.60	0.20	0.075	3.71	0.000127
266	T024 – 9	2	75	10	11	0.60	0.20	0.075	3.32	0.000144
267	T024 – 10	2	75	10	15	0.60	0.20	0.075	2.92	0.000059
268	T024 – 11	2	75	10	20	0.60	0.20	0.075	3.33	0.000195
269	T024 – 12	2	75	10	25	0.60	0.20	0.075	3.25	0.000092
270	T024 – 13	1	0	10	-	0.60	0.20	0.075	2.98	0.000051

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
1	T001 – 1	2	25	7	11	0.10	0.70	0.125	0.93	0.001570

(cot α = 0.0) Continue on the next page

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
2	T001 – 2	2	25	7	15	0.10	0.70	0.125	1.01	0.001710
3	T001 – 3	2	25	7	20	0.10	0.70	0.125	1.09	0.001980
4	T001 – 4	2	25	7	25	0.10	0.70	0.125	1.09	0.002160
5	T001 – 5	2	50	7	11	0.10	0.70	0.125	0.87	0.001960
6	T001 – 6	2	50	7	15	0.10	0.70	0.125	1.01	0.002140
7	T001 – 7	2	50	7	20	0.10	0.70	0.125	1.10	0.003280
8	T001 – 9	2	75	7	11	0.10	0.70	0.125	0.78	0.003240
9	T001 – 10	2	75	7	15	0.10	0.70	0.125	0.88	0.002630
10	T001 – 13	1	0	7	-	0.10	0.70	0.125	1.08	0.001504
11	T002 – 1	2	25	8	11	0.10	0.70	0.125	0.95	0.002080
12	T002 – 2	2	25	8	15	0.10	0.70	0.125	1.02	0.002130
13	T002 – 3	2	25	8	20	0.10	0.70	0.125	1.10	0.002330
14	T002 – 4	2	25	8	25	0.10	0.70	0.125	1.11	0.002660
15	T002 – 5	2	50	8	11	0.10	0.70	0.125	0.87	0.002180
16	T002 – 6	2	50	8	15	0.10	0.70	0.125	1.01	0.002340
17	T002 – 7	2	50	8	20	0.10	0.70	0.125	1.10	0.003390
18	T002 – 9	2	75	8	11	0.10	0.70	0.125	0.77	0.003680
19	T002 – 10	2	75	8	15	0.10	0.70	0.125	1.01	0.002800
20	T002 – 13	1	0	8	-	0.10	0.70	0.125	1.02	0.001531
21	T003 – 1	2	25	9	11	0.10	0.70	0.125	0.89	0.002160
22	T003 – 2	2	25	9	15	0.10	0.70	0.125	0.87	0.003540
23	T003 – 3	2	25	9	20	0.10	0.70	0.125	0.87	0.004260
24	T003 – 4	2	25	9	25	0.10	0.70	0.125	0.93	0.002540
25	T003 – 5	2	50	9	11	0.10	0.70	0.125	0.86	0.002220
26	T003 – 6	2	50	9	15	0.10	0.70	0.125	0.93	0.002950
27	T003 – 7	2	50	9	20	0.10	0.70	0.125	1.01	0.004150
28	T003 – 9	2	75	9	11	0.10	0.70	0.125	0.78	0.004760
29	T003 – 10	2	75	9	15	0.10	0.70	0.125	0.93	0.004210
30	T003 – 13	1	0	9	-	0.10	0.70	0.125	0.92	0.001497
31	T004 – 1	2	25	10	11	0.10	0.70	0.125	0.74	0.004180
32	T004 – 2	2	25	10	15	0.10	0.70	0.125	0.78	0.006110
33	T004 – 3	2	25	10	20	0.10	0.70	0.125	0.83	0.006300

(cot $\alpha = 0.0$) Continue on the next page

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
34	T004 – 4	2	25	10	25	0.10	0.70	0.125	0.86	0.002950
35	T004 – 5	2	50	10	11	0.10	0.70	0.125	0.73	0.003410
36	T004 – 6	2	50	10	15	0.10	0.70	0.125	0.82	0.007860
37	T004 – 7	2	50	10	20	0.10	0.70	0.125	0.93	0.003410
38	T004 – 9	2	75	10	11	0.10	0.70	0.125	0.77	0.004180
39	T004 – 10	2	75	10	15	0.10	0.70	0.125	0.92	0.003900
40	T004 – 13	1	0	10	-	0.10	0.70	0.125	0.75	0.005690
41	T005 – 1	2	25	7	11	0.15	0.65	0.100	1.81	0.000188
42	T005 – 2	2	25	7	15	0.15	0.65	0.100	1.81	0.000134
43	T005 – 3	2	25	7	20	0.15	0.65	0.100	1.98	0.000233
44	T005 – 4	2	25	7	25	0.15	0.65	0.100	1.65	0.000210
45	T005 – 5	2	50	7	11	0.15	0.65	0.100	1.80	0.000132
46	T005 – 6	2	50	7	15	0.15	0.65	0.100	1.95	0.000225
47	T005 – 7	2	50	7	20	0.15	0.65	0.100	1.99	0.000155
48	T005 – 8	2	50	7	25	0.15	0.65	0.100	1.73	0.000187
49	T005 – 9	2	75	7	11	0.15	0.65	0.100	1.51	0.000168
50	T005 – 10	2	75	7	15	0.15	0.65	0.100	1.76	0.000128
51	T005 – 11	2	75	7	20	0.15	0.65	0.100	2.28	0.000147
52	T005 – 13	1	0	7	-	0.15	0.65	0.100	1.81	0.000093
53	T006 – 1	2	25	8	11	0.15	0.65	0.125	1.42	0.000699
54	T006 – 2	2	25	8	15	0.15	0.65	0.125	1.54	0.000631
55	T006 – 3	2	25	8	20	0.15	0.65	0.125	1.66	0.000646
56	T006 – 4	2	25	8	25	0.15	0.65	0.125	1.66	0.000753
57	T006 – 5	2	50	8	11	0.15	0.65	0.125	1.31	0.000578
58	T006 – 6	2	50	8	15	0.15	0.65	0.125	1.50	0.001030
59	T006 – 7	2	50	8	20	0.15	0.65	0.125	1.66	0.001130
60	T006 – 9	2	75	8	11	0.15	0.65	0.125	1.01	0.000898
61	T006 – 10	2	75	8	15	0.15	0.65	0.125	1.50	0.000655
62	T006 – 13	1	0	8	-	0.15	0.65	0.125	1.67	0.000472
63	T007 – 1	2	25	9	11	0.15	0.65	0.100	1.55	0.000193
64	T007 – 2	2	25	9	15	0.15	0.65	0.100	1.65	0.000064
65	T007 – 3	2	25	9	20	0.15	0.65	0.100	1.80	0.000104

(cot $\alpha = 0.0$) Continue on the next page

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
66	T007-4	2	25	9	25	0.15	0.65	0.100	1.80	0.000090
67	T007-5	2	50	9	11	0.15	0.65	0.100	1.54	0.000109
68	T007-6	2	50	9	15	0.15	0.65	0.100	1.63	0.000078
69	T007-7	2	50	9	20	0.15	0.65	0.100	1.98	0.000168
70	T007-9	2	75	9	11	0.15	0.65	0.100	1.51	0.000128
71	T007-10	2	75	9	15	0.15	0.65	0.100	1.79	0.000157
72	T007-11	2	75	9	20	0.15	0.65	0.100	2.21	0.000190
73	T007-12	2	75	9	25	0.15	0.65	0.100	1.62	0.000114
74	T007-13	1	0	9	-	0.15	0.65	0.100	1.65	0.000081
75	T008-1	2	25	10	11	0.15	0.65	0.125	1.11	0.000738
76	T008-2	2	25	10	15	0.15	0.65	0.125	1.22	0.000408
77	T008-3	2	25	10	20	0.15	0.65	0.125	1.23	0.000487
78	T008-4	2	50	10	25	0.15	0.65	0.125	1.39	0.000559
79	T008-5	2	50	10	11	0.15	0.65	0.125	1.21	0.000456
80	T008-6	2	50	10	15	0.15	0.65	0.125	1.23	0.000559
81	T008-7	2	50	10	20	0.15	0.65	0.125	1.31	0.000828
82	T008-9	2	75	10	11	0.15	0.65	0.125	1.15	0.000670
83	T008-10	2	75	10	15	0.15	0.65	0.125	1.31	0.000906
84	T008-13	1	0	10	-	0.15	0.65	0.125	1.20	0.000491
85	T009-1	2	25	7	11	0.20	0.60	0.100	2.40	0.000070
86	T009-2	2	25	7	15	0.20	0.60	0.100	2.65	0.000013
87	T009-3	2	25	7	20	0.20	0.60	0.100	2.64	0.000021
88	T009-4	2	25	7	25	0.20	0.60	0.100	2.64	0.000047
89	T009-5	2	50	7	11	0.20	0.60	0.100	2.16	0.000074
90	T009-6	2	50	7	15	0.20	0.60	0.100	2.58	0.000067
91	T009-7	2	50	7	20	0.20	0.60	0.100	2.90	0.000028
92	T009-8	2	50	7	25	0.20	0.60	0.100	2.84	0.000050
93	T009-9	2	75	7	11	0.20	0.60	0.100	1.98	0.000065
94	T009-10	2	75	7	15	0.20	0.60	0.100	2.56	0.000104
95	T009-11	2	75	7	20	0.20	0.60	0.100	2.63	0.000033
96	T009-13	1	0	7	-	0.20	0.60	0.100	2.43	0.000046
97	T010-1	2	25	8	11	0.20	0.60	0.100	2.33	0.000065

(cot $\alpha = 0.0$) Continue on the next page

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
98	T010-2	2	25	8	15	0.20	0.60	0.100	2.10	0.000046
99	T010-3	2	25	8	20	0.20	0.60	0.100	2.47	0.000071
100	T010-4	2	25	8	25	0.20	0.60	0.100	2.03	0.000078
101	T010-5	2	50	8	11	0.20	0.60	0.100	2.20	0.000063
102	T010-6	2	50	8	15	0.20	0.60	0.100	2.35	0.000096
103	T010-7	2	50	8	20	0.20	0.60	0.100	2.22	0.000102
104	T010-8	2	50	8	25	0.20	0.60	0.100	2.23	0.000000
105	T010-9	2	75	8	11	0.20	0.60	0.100	2.13	0.000251
106	T010-10	2	75	8	15	0.20	0.60	0.100	2.29	0.000228
107	T010-13	1	0	8	-	0.20	0.60	0.100	2.26	0.000103
108	T011-1	2	25	9	11	0.20	0.60	0.125	1.65	0.000095
109	T011-2	2	25	9	15	0.20	0.60	0.125	2.42	0.000109
110	T011-3	2	25	9	20	0.20	0.60	0.125	2.35	0.000143
111	T011-4	2	25	9	25	0.20	0.60	0.125	1.74	0.000000
112	T011-5	2	50	9	11	0.20	0.60	0.125	1.64	0.000263
113	T011-6	2	50	9	15	0.20	0.60	0.125	1.85	0.000316
114	T011-7	2	50	9	20	0.20	0.60	0.125	1.71	0.000000
115	T011-9	2	75	9	11	0.20	0.60	0.125	1.43	0.000276
116	T011-10	2	75	9	15	0.20	0.60	0.125	1.85	0.000499
117	T011-13	1	0	9	-	0.20	0.60	0.125	1.65	0.000077
118	T0012-1	2	25	10	11	0.20	0.60	0.125	1.45	0.000313
119	T0012-2	2	25	10	15	0.20	0.60	0.125	1.44	0.000271
120	T0012-3	2	25	10	20	0.20	0.60	0.125	1.45	0.000267
121	T0012-4	2	25	10	25	0.20	0.60	0.125	1.98	0.000386
122	T0012-5	2	50	10	11	0.20	0.60	0.125	1.52	0.000322
123	T0012-6	2	50	10	15	0.20	0.60	0.125	1.60	0.000377
124	T0012-9	2	75	10	11	0.20	0.60	0.125	1.42	0.000181
125	T0012-10	2	75	10	15	0.20	0.60	0.125	1.84	0.000654
126	T0012-13	1	0	10	-	0.20	0.60	0.125	1.44	0.000224
127	T013-1	2	25	7	11	0.15	0.65	0.075	2.21	0.000000
128	T013-2	2	25	7	15	0.15	0.65	0.075	2.49	0.000000
129	T013-3	2	25	7	20	0.15	0.65	0.075	2.49	0.000000

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Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
130	T013-4	2	25	7	25	0.15	0.65	0.075	2.49	0.000000
131	T013-5	2	50	7	11	0.15	0.65	0.075	2.20	0.000000
132	T013-6	2	50	7	15	0.15	0.65	0.075	2.43	0.000000
133	T013-7	2	50	7	20	0.15	0.65	0.075	2.85	0.000000
134	T013-8	2	50	7	25	0.15	0.65	0.075	2.49	0.000000
135	T013-9	2	75	7	11	0.15	0.65	0.075	1.93	0.000000
136	T013-10	2	75	7	15	0.15	0.65	0.075	2.40	0.000000
137	T013-11	2	75	7	20	0.15	0.65	0.075	2.85	0.000000
138	T013-12	2	75	7	25	0.15	0.65	0.075	2.74	0.000000
139	T013-13	1	0	7	-	0.15	0.65	0.075	2.21	0.000000
140	T014-1	2	25	8	11	0.15	0.65	0.075	2.24	0.000000
141	T014-2	2	25	8	15	0.15	0.65	0.075	2.51	0.000000
142	T014-3	2	25	8	20	0.15	0.65	0.075	2.86	0.000000
143	T014-4	2	25	8	25	0.15	0.65	0.075	2.85	0.000000
144	T014-5	2	50	8	11	0.15	0.65	0.075	2.20	0.000000
145	T014-6	2	50	8	15	0.15	0.65	0.075	2.43	0.000000
146	T014-7	2	50	8	20	0.15	0.65	0.075	2.85	0.000000
147	T014-8	2	50	8	25	0.15	0.65	0.075	2.85	0.000000
148	T014-9	2	75	8	11	0.15	0.65	0.075	1.93	0.000000
149	T014-10	2	75	8	15	0.15	0.65	0.075	2.35	0.000000
150	T014-11	2	75	8	20	0.15	0.65	0.075	2.85	0.000000
151	T014-12	2	75	8	25	0.15	0.65	0.075	2.85	0.000000
152	T014-13	1	0	8	-	0.15	0.65	0.075	2.52	0.000000
153	T015-1	2	25	9	11	0.15	0.65	0.075	1.98	0.000000
154	T015-2	2	25	9	15	0.15	0.65	0.075	2.18	0.000000
155	T015-3	2	25	9	20	0.15	0.65	0.075	2.19	0.000000
156	T015-4	2	25	9	25	0.15	0.65	0.075	2.45	0.000000
157	T015-5	2	50	9	11	0.15	0.65	0.075	1.96	0.000000
158	T015-6	2	50	9	15	0.15	0.65	0.075	2.13	0.000000
159	T015-7	2	50	9	20	0.15	0.65	0.075	2.45	0.000000
160	T015-8	2	50	9	25	0.15	0.65	0.075	2.47	0.000000
161	T015-9	2	75	9	11	0.15	0.65	0.075	1.90	0.000000

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Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
162	T015 – 10	2	75	9	15	0.15	0.65	0.075	2.11	0.000000
163	T015 – 11	2	75	9	20	0.15	0.65	0.075	2.85	0.000000
164	T015 – 12	2	75	9	25	0.15	0.65	0.075	2.85	0.000000
165	T015 – 13	1	0	9	-	0.15	0.65	0.075	2.20	0.000000
166	T016 – 1	2	25	10	11	0.15	0.65	0.075	1.95	0.000000
167	T016 – 2	2	25	10	15	0.15	0.65	0.075	1.72	0.000000
168	T016 – 3	2	25	10	20	0.15	0.65	0.075	1.95	0.000000
169	T016 – 4	2	25	10	25	0.15	0.65	0.075	1.95	0.000000
170	T016 – 5	2	50	10	11	0.15	0.65	0.075	1.91	0.000000
171	T016 – 6	2	50	10	15	0.15	0.65	0.075	2.10	0.000000
172	T016 – 7	2	50	10	20	0.15	0.65	0.075	2.17	0.000000
173	T016 – 8	2	50	10	25	0.15	0.65	0.075	2.44	0.000000
174	T016 – 9	2	75	10	11	0.15	0.65	0.075	1.88	0.000000
175	T016 – 10	2	75	10	15	0.15	0.65	0.075	2.10	0.000000
176	T016 – 11	2	75	10	20	0.15	0.65	0.075	2.50	0.000000
177	T016 – 12	2	75	10	25	0.15	0.65	0.075	2.49	0.000000
178	T016 – 13	1	0	10	-	0.15	0.65	0.075	1.95	0.000000
179	T017 – 1	2	25	7	11	0.10	0.70	0.100	1.21	0.000738
180	T017 – 2	2	25	7	15	0.10	0.70	0.100	1.21	0.000671
181	T017 – 3	2	25	7	20	0.10	0.70	0.100	1.32	0.000954
182	T017 – 4	2	25	7	25	0.10	0.70	0.100	1.20	0.001090
183	T017 – 5	2	50	7	11	0.10	0.70	0.100	1.10	0.000581
184	T017 – 6	2	50	7	15	0.10	0.70	0.100	1.20	0.000864
185	T017 – 7	2	50	7	20	0.10	0.70	0.100	1.34	0.001090
186	T017 – 8	2	50	7	25	0.10	0.70	0.100	1.21	0.001280
187	T017 – 9	2	75	7	11	0.10	0.70	0.100	1.02	0.001220
188	T017 – 10	2	75	7	15	0.10	0.70	0.100	1.20	0.001200
189	T017 – 11	2	75	7	20	0.10	0.70	0.100	1.48	0.001370
190	T017 – 13	1	0	7	-	0.10	0.70	0.100	1.21	0.000000
191	T018 – 1	2	25	8	11	0.10	0.70	0.100	1.22	0.000818
192	T018 – 2	2	25	8	15	0.10	0.70	0.100	1.22	0.000474
193	T018 – 3	2	25	8	20	0.10	0.70	0.100	1.34	0.001170

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Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
194	T018-4	2	25	8	25	0.10	0.70	0.100	1.34	0.000858
195	T018-5	2	50	8	11	0.10	0.70	0.100	1.10	0.000583
196	T018-6	2	50	8	15	0.10	0.70	0.100	1.19	0.000591
197	T018-7	2	50	8	20	0.10	0.70	0.100	1.49	0.001190
198	T018-8	2	50	8	25	0.10	0.70	0.100	1.34	0.001170
199	T018-9	2	75	8	11	0.10	0.70	0.100	1.04	0.001650
200	T018-10	2	75	8	15	0.10	0.70	0.100	1.20	0.000867
201	T018-11	2	75	8	20	0.10	0.70	0.100	1.50	0.001670
202	T018-13	1	0	8	-	0.10	0.70	0.100	1.35	0.000788
203	T019-1	2	25	9	11	0.10	0.70	0.100	1.00	0.001170
204	T019-2	2	25	9	15	0.10	0.70	0.100	1.12	0.001070
205	T019-3	2	25	9	20	0.10	0.70	0.100	1.14	0.000898
206	T019-4	2	25	9	25	0.10	0.70	0.100	1.13	0.001300
207	T019-5	2	50	9	11	0.10	0.70	0.100	1.04	0.001320
208	T019-6	2	50	9	15	0.10	0.70	0.100	1.13	0.001520
209	T019-7	2	50	9	20	0.10	0.70	0.100	1.12	0.000942
210	T019-8	2	50	9	25	0.10	0.70	0.100	1.21	0.001300
211	T019-9	2	75	9	11	0.10	0.70	0.100	1.04	0.001060
212	T019-10	2	75	9	15	0.10	0.70	0.100	1.20	0.001120
213	T019-11	2	75	9	20	0.10	0.70	0.100	1.49	0.001170
214	T019-13	1	0	9	-	0.10	0.70	0.100	1.12	0.000501
215	T020-1	2	25	10	11	0.10	0.70	0.100	1.03	0.000000
216	T020-2	2	25	10	15	0.10	0.70	0.100	0.96	0.000000
217	T020-3	2	25	10	20	0.10	0.70	0.100	1.03	0.000000
218	T020-4	2	25	10	25	0.10	0.70	0.100	1.07	0.000000
219	T020-5	2	50	10	11	0.10	0.70	0.100	1.16	0.000000
220	T020-6	2	50	10	15	0.10	0.70	0.100	1.01	0.000000
221	T020-7	2	50	10	20	0.10	0.70	0.100	1.18	0.000000
222	T020-8	2	50	10	25	0.10	0.70	0.100	1.48	0.000000
223	T020-9	2	75	10	11	0.10	0.70	0.100	1.08	0.000000
224	T020-10	2	75	10	15	0.10	0.70	0.100	1.18	0.000000
225	T020-11	2	75	10	20	0.10	0.70	0.100	1.44	0.000000

(cot $\alpha = 0.0$) Continue on the next page

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
226	T020 – 13	1	0	10	-	0.10	0.70	0.100	1.06	0.000000
227	T021 – 1	2	25	7	11	0.20	0.60	0.075	3.30	0.000000
228	T021 – 2	2	25	7	15	0.20	0.60	0.075	3.78	0.000000
229	T021 – 3	2	25	7	20	0.20	0.60	0.075	3.77	0.000000
230	T021 – 4	2	25	7	25	0.20	0.60	0.075	3.77	0.000000
231	T021 – 5	2	50	7	11	0.20	0.60	0.075	3.28	0.000000
232	T021 – 6	2	50	7	15	0.20	0.60	0.075	3.66	0.000000
233	T021 – 7	2	50	7	20	0.20	0.60	0.075	3.77	0.000000
234	T021 – 8	2	50	7	25	0.20	0.60	0.075	3.94	0.000000
235	T021 – 9	2	75	7	11	0.20	0.60	0.075	2.56	0.000000
236	T021 – 10	2	75	7	15	0.20	0.60	0.075	3.17	0.000000
237	T021 – 11	2	75	7	20	0.20	0.60	0.075	4.12	0.000000
238	T021 – 12	2	75	7	25	0.20	0.60	0.075	3.77	0.000000
239	T021 – 13	1	0	7	-	0.20	0.60	0.075	3.77	0.000000
240	T022 – 1	2	25	8	11	0.20	0.60	0.075	2.97	0.000000
241	T022 – 2	2	25	8	15	0.20	0.60	0.075	3.33	0.000000
242	T022 – 3	2	25	8	20	0.20	0.60	0.075	3.32	0.000000
243	T022 – 4	2	25	8	25	0.20	0.60	0.075	3.31	0.000000
244	T022 – 5	2	50	8	11	0.20	0.60	0.075	3.30	0.000000
245	T022 – 6	2	50	8	15	0.20	0.60	0.075	3.68	0.000000
246	T022 – 7	2	50	8	20	0.20	0.60	0.075	3.77	0.000000
247	T022 – 8	2	50	8	25	0.20	0.60	0.075	3.31	0.000000
248	T022 – 9	2	75	8	11	0.20	0.60	0.075	2.56	0.000000
249	T022 – 10	2	75	8	15	0.20	0.60	0.075	3.13	0.000000
250	T022 – 11	2	75	8	20	0.20	0.60	0.075	3.80	0.000000
251	T022 – 13	1	0	8	-	0.20	0.60	0.075	3.34	0.000000
252	T023 – 1	2	25	9	11	0.20	0.60	0.075	2.93	0.000000
253	T023 – 2	2	25	9	15	0.20	0.60	0.075	2.90	0.000000
254	T023 – 3	2	25	9	20	0.20	0.60	0.075	3.28	0.000000
255	T023 – 4	2	25	9	25	0.20	0.60	0.075	3.27	0.000000
256	T023 – 5	2	50	9	11	0.20	0.60	0.075	2.90	0.000000
257	T023 – 6	2	50	9	15	0.20	0.60	0.075	3.18	0.000000

(cot α = 0.0) Continue on the next page

Table B.3 Dimensionless Wave Overtopping Results for $\cot \alpha = 0.0$ (cont.)

ID	Test No	Mode		T_{pW} (secs)	T_{pS} (secs)	h (m)	R_c (m)	H_{m0} (m)	R^* (-)	Q^* (-)
		Uni = 1 Bi = 2	Swell (%)							
258	T023 – 7	2	50	9	20	0.20	0.60	0.075	3.18	0.000000
259	T023 – 8	2	50	9	25	0.20	0.60	0.075	3.74	0.000000
260	T023 – 9	2	75	9	11	0.20	0.60	0.075	2.54	0.000000
261	T023 – 10	2	75	9	15	0.20	0.60	0.075	3.14	0.000000
262	T023 – 11	2	75	9	20	0.20	0.60	0.075	3.78	0.000000
263	T023 – 12	2	75	9	25	0.20	0.60	0.075	3.65	0.000000
264	T023 – 13	1	0	9	-	0.20	0.60	0.075	3.29	0.000000
265	T024 – 1	2	25	10	11	0.20	0.60	0.075	2.61	0.000000
266	T024 – 2	2	25	10	15	0.20	0.60	0.075	2.53	0.000000
267	T024 – 3	2	25	10	20	0.20	0.60	0.075	2.89	0.000000
268	T024 – 4	2	25	10	25	0.20	0.60	0.075	2.86	0.000000
269	T024 – 5	2	50	10	11	0.20	0.60	0.075	2.56	0.000000
270	T024 – 6	2	50	10	15	0.20	0.60	0.075	2.79	0.000000
271	T024 – 7	2	50	10	20	0.20	0.60	0.075	3.25	0.000000
272	T024 – 8	2	50	10	25	0.20	0.60	0.075	3.24	0.000000
273	T024 – 9	2	75	10	11	0.20	0.60	0.075	2.51	0.000000
274	T024 – 10	2	75	10	15	0.20	0.60	0.075	2.76	0.000000
275	T024 – 11	2	75	10	20	0.20	0.60	0.075	3.31	0.000000
276	T024 – 12	2	75	10	25	0.20	0.60	0.075	3.30	0.000000
277	T024 – 13	1	0	10	-	0.20	0.60	0.075	2.61	0.000000

Appendix C

Reflection Characteristics Under Bimodal Seas

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$

ID	Test No	Swell (%)	Reff Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
1	T001 – 1	25	0.416	2.665
2	T001 – 2	25	0.423	2.755
3	T001 – 3	25	0.568	2.879
4	T001 – 4	25	0.426	2.799
5	T001 – 5	50	0.525	2.896
6	T001 – 6	50	0.671	3.182
7	T001 – 7	50	0.700	3.269
8	T001 – 9	75	0.608	3.052
9	T001 – 10	75	0.862	3.621
10	T001 – 13	1	0.422	2.621
11	T002 – 1	25	0.432	2.563
12	T002 – 2	25	0.440	2.658
13	T002 – 3	25	0.563	2.778
14	T002 – 4	25	0.448	2.715
15	T002 – 5	50	0.518	2.716
16	T002 – 6	50	0.645	3.005
17	T002 – 7	50	0.680	3.093
18	T002 – 9	75	0.555	2.895

($\cot \alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
19	T002 – 10	75	0.871	3.476
20	T002 – 13	1	0.436	2.431
21	T003 – 1	25	0.551	3.118
22	T003 – 2	25	0.553	3.109
23	T003 – 3	25	0.617	3.225
24	T003 – 4	25	0.544	3.175
25	T003 – 5	50	0.579	3.196
26	T003 – 6	50	0.717	3.479
27	T003 – 7	50	0.740	3.682
28	T003 – 9	75	0.583	3.201
29	T003 – 10	75	0.826	3.841
30	T003 – 13	1	0.534	3.140
31	T004 – 1	25	0.774	3.308
32	T004 – 2	25	0.790	3.469
33	T004 – 3	25	0.773	3.543
34	T004 – 4	25	0.782	3.714
35	T004 – 5	50	0.773	3.345
36	T004 – 6	50	0.841	3.689
37	T004 – 9	75	0.794	3.418
38	T004 – 10	75	0.895	4.020
39	T004 – 13	1	0.774	3.576
40	T005 – 1	25	0.444	3.041
41	T005 – 2	25	0.599	3.127
42	T005 – 3	25	0.630	3.198
43	T005 – 4	25	0.618	3.213
44	T005 – 5	50	0.554	3.192
45	T005 – 6	50	0.706	3.350
46	T005 – 7	50	0.755	3.560
47	T005 – 8	50	0.810	3.591
48	T005 – 9	75	0.611	3.438
49	T005 – 10	75	0.890	3.945
50	T005 – 11	75	0.856	4.078
51	T005 – 13	1	0.433	2.877

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
52	T006 – 1	25	0.437	2.868
53	T006 – 2	25	0.439	2.971
54	T006 – 3	25	0.613	3.017
55	T006 – 4	25	0.617	3.143
56	T006 – 5	50	0.518	3.037
57	T006 – 6	50	0.643	3.358
58	T006 – 7	50	0.739	3.353
59	T006 – 9	75	0.554	3.236
60	T006 – 10	75	0.890	3.905
61	T006 – 13	1	0.437	2.862
62	T007 – 1	25	0.565	3.426
63	T007 – 2	25	0.573	3.560
64	T007 – 3	25	0.679	3.667
65	T007 – 4	25	0.770	3.790
66	T007 – 5	50	0.586	3.506
67	T007 – 6	50	0.720	3.728
68	T007 – 7	50	0.797	4.022
69	T007 – 9	75	0.606	3.669
70	T007 – 10	75	0.915	4.287
71	T007 – 11	75	0.855	4.347
72	T007 – 12	75	0.922	4.730
73	T007 – 13	1	0.567	3.293
74	T008 – 1	25	0.585	3.315
75	T008 – 2	25	0.603	3.457
76	T008 – 3	25	0.673	3.460
77	T008 – 4	50	0.670	3.458
78	T008 – 5	50	0.593	3.339
79	T008 – 6	50	0.725	3.604
80	T008 – 7	50	0.751	3.803
81	T008 – 9	75	0.585	3.352
82	T008 – 10	75	0.828	3.875
83	T008 – 13	1	0.562	3.242
84	T009 – 1	25	0.443	3.041

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
85	T009 – 2	25	0.596	3.125
86	T009 – 3	25	0.646	3.209
87	T009 – 4	25	0.618	3.213
88	T009 – 5	50	0.554	3.192
89	T009 – 6	50	0.703	3.348
90	T009 – 7	50	0.769	3.572
91	T009 – 8	50	0.810	3.750
92	T009 – 9	75	0.600	3.430
93	T009 – 10	75	0.910	4.158
94	T009 – 11	75	0.820	4.428
95	T009 – 13	1	0.443	2.735
96	T010 – 1	25	0.518	3.168
97	T010 – 2	25	0.608	3.323
98	T010 – 3	25	0.653	3.448
99	T010 – 4	25	0.656	3.473
100	T010 – 5	50	0.558	3.332
101	T010 – 6	50	0.760	3.585
102	T010 – 7	50	0.786	3.681
103	T010 – 8	50	0.800	3.730
104	T010 – 9	75	0.622	3.285
105	T010 – 10	75	0.854	3.910
106	T010 – 13	1	0.512	2.881
107	T011 – 1	25	0.536	3.108
108	T011 – 2	25	0.604	3.366
109	T011 – 3	25	0.830	2.963
110	T011 – 4	25	0.667	3.262
111	T011 – 5	50	0.565	3.186
112	T011 – 6	50	0.714	3.476
113	T011 – 7	50	0.810	3.745
114	T011 – 9	75	0.579	2.978
115	T011 – 10	75	0.854	4.205
116	T011 – 13	1	0.512	3.126
117	T0012 – 1	25	0.593	3.172

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
118	T0012-2	25	0.635	3.345
119	T0012-3	25	0.708	3.489
120	T0012-4	25	0.717	3.513
121	T0012-5	50	0.601	3.345
122	T0012-6	50	0.722	3.792
123	T0012-9	75	0.590	3.259
124	T0012-10	75	0.828	3.953
125	T0012-13	1	0.605	3.162
126	T013-1	25	0.489	3.537
127	T013-2	25	0.625	3.759
128	T013-3	25	0.661	3.820
129	T013-4	25	0.666	3.911
130	T013-5	50	0.603	3.554
131	T013-6	50	0.756	4.001
132	T013-7	50	0.786	4.230
133	T013-8	50	0.890	4.364
134	T013-9	75	0.705	3.953
135	T013-10	75	0.915	4.410
136	T013-11	75	0.886	4.782
137	T013-12	75	0.800	4.952
138	T013-13	1	0.476	3.289
139	T014-1	25	0.566	3.604
140	T014-2	25	0.557	3.725
141	T014-3	25	0.704	3.927
142	T014-4	25	0.710	3.932
143	T014-5	50	0.619	3.826
144	T014-6	50	0.793	4.285
145	T014-7	50	0.775	4.429
146	T014-8	50	0.808	4.533
147	T014-9	75	0.710	4.057
148	T014-10	75	0.917	4.858
149	T014-11	75	0.884	4.947
150	T014-12	75	0.880	5.172

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
151	T014 – 13	1	0.532	3.656
152	T015 – 1	25	0.646	3.899
153	T015 – 2	25	0.661	4.059
154	T015 – 3	25	0.748	4.239
155	T015 – 4	25	0.763	4.278
156	T015 – 5	50	0.677	4.026
157	T015 – 6	50	0.772	4.403
158	T015 – 7	50	0.818	4.490
159	T015 – 8	50	0.855	4.591
160	T015 – 9	75	0.790	4.282
161	T015 – 10	75	0.886	4.646
162	T015 – 11	75	0.885	5.092
163	T015 – 12	75	0.860	5.314
164	T015 – 13	1	0.641	4.029
165	T016 – 1	25	0.780	4.199
166	T016 – 2	25	0.721	4.327
167	T016 – 3	25	0.756	4.498
168	T016 – 4	25	0.766	4.531
169	T016 – 5	50	0.740	4.231
170	T016 – 6	50	0.808	4.378
171	T016 – 7	50	0.831	4.736
172	T016 – 8	50	0.841	4.804
173	T016 – 9	75	0.715	4.252
174	T016 – 10	75	0.883	5.048
175	T016 – 11	75	0.876	5.014
176	T016 – 12	75	0.920	5.460
177	T016 – 13	1	0.770	4.140
178	T017 – 1	25	0.456	3.048
179	T017 – 2	25	0.593	3.123
180	T017 – 3	25	0.613	3.185
181	T017 – 4	25	0.616	3.212
182	T017 – 5	50	0.560	3.196
183	T017 – 6	50	0.670	3.485

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
184	T017-7	50	0.725	3.534
185	T017-8	50	0.885	3.658
186	T017-9	75	0.608	3.436
187	T017-10	75	0.884	3.939
188	T017-11	75	0.827	4.049
189	T017-13	1	0.442	2.882
190	T018-1	25	0.524	3.344
191	T018-2	25	0.494	3.244
192	T018-3	25	0.597	3.405
193	T018-4	25	0.505	3.362
194	T018-5	50	0.565	3.337
195	T018-6	50	0.750	3.751
196	T018-7	50	0.717	3.782
197	T018-8	50	0.810	3.906
198	T018-9	75	0.628	3.603
199	T018-10	75	0.887	4.134
200	T018-11	75	0.910	4.317
201	T018-13	1	0.486	3.169
202	T019-1	25	0.579	3.436
203	T019-2	25	0.586	3.570
204	T019-3	25	0.598	3.600
205	T019-4	25	0.604	3.646
206	T019-5	50	0.595	3.513
207	T019-6	50	0.728	3.917
208	T019-7	50	0.757	3.984
209	T019-8	50	0.910	4.116
210	T019-9	75	0.628	3.687
211	T019-10	75	0.899	4.270
212	T019-11	75	0.890	4.803
213	T019-13	1	0.583	3.305
214	T020-1	25	0.627	3.602
215	T020-2	25	0.659	3.656
216	T020-3	25	0.700	3.785

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
217	T020 – 4	25	0.646	3.940
218	T020 – 5	50	0.632	3.690
219	T020 – 6	50	0.749	4.094
220	T020 – 7	50	0.742	4.159
221	T020 – 8	50	0.645	4.147
222	T020 – 9	75	0.635	3.579
223	T020 – 10	75	0.838	4.240
224	T020 – 11	75	0.840	4.678
225	T020 – 13	1	0.626	3.602
226	T021 – 1	25	0.485	3.306
227	T021 – 2	25	0.640	3.529
228	T021 – 3	25	0.644	3.559
229	T021 – 4	25	0.498	3.528
230	T021 – 5	50	0.593	3.546
231	T021 – 6	50	0.751	3.997
232	T021 – 7	50	0.771	4.214
233	T021 – 8	50	0.950	4.430
234	T021 – 9	75	0.689	3.938
235	T021 – 10	75	0.922	4.685
236	T021 – 11	75	0.966	4.878
237	T021 – 12	75	0.502	4.629
238	T021 – 13	1	0.502	3.092
239	T022 – 1	25	0.540	3.584
240	T022 – 2	25	0.556	3.724
241	T022 – 3	25	0.703	3.926
242	T022 – 4	25	0.770	3.989
243	T022 – 5	50	0.610	3.818
244	T022 – 6	50	0.783	4.275
245	T022 – 7	50	0.776	4.430
246	T022 – 8	50	0.793	4.516
247	T022 – 9	75	0.679	4.028
248	T022 – 10	75	0.916	4.857
249	T022 – 11	75	0.882	4.945

(cot $\alpha = 3.0$) Continue on the next page

Table C.1 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 3.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
250	T022 – 13	1	0.532	3.420
251	T023 – 1	25	0.627	3.883
252	T023 – 2	25	0.645	4.044
253	T023 – 3	25	0.748	4.239
254	T023 – 4	25	0.758	4.273
255	T023 – 5	50	0.679	4.028
256	T023 – 6	50	0.764	4.143
257	T023 – 7	50	0.845	4.520
258	T023 – 8	50	0.834	4.567
259	T023 – 9	75	0.730	4.221
260	T023 – 10	75	0.881	4.640
261	T023 – 11	75	0.882	5.088
262	T023 – 12	75	0.795	5.230
263	T023 – 13	1	0.627	4.016
264	T024 – 1	25	0.702	4.122
265	T024 – 2	25	0.726	4.332
266	T024 – 3	25	0.732	4.472
267	T024 – 4	25	0.776	4.542
268	T024 – 5	50	0.703	4.194
269	T024 – 6	50	0.804	4.639
270	T024 – 7	50	0.839	4.746
271	T024 – 8	50	0.712	4.656
272	T024 – 9	75	0.876	4.422
273	T024 – 10	75	0.876	5.039
274	T024 – 11	75	0.876	5.014
275	T024 – 12	75	0.903	5.437
276	T024 – 13	1	0.790	4.160

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$

ID	Test No	Swell (%)	Reff Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
1	T001 – 1	25	0.672	2.811
2	T001 – 2	25	0.684	2.910
3	T001 – 3	25	0.681	3.235
4	T001 – 4	25	0.689	2.959
5	T001 – 5	50	0.760	3.054
6	T001 – 6	50	0.814	3.292
7	T001 – 7	50	0.796	3.345
8	T001 – 9	75	0.803	3.325
9	T001 – 10	75	0.872	3.380
10	T001 – 13	0	0.655	2.634
11	T002 – 1	25	0.710	2.815
12	T002 – 2	25	0.735	2.932
13	T002 – 3	25	0.746	2.998
14	T002 – 4	25	0.753	3.117
15	T002 – 5	50	0.773	2.979
16	T002 – 6	50	0.825	3.246
17	T002 – 7	50	0.810	3.296
18	T002 – 8	50	0.790	3.462
19	T002 – 9	75	0.787	3.161
20	T002 – 10	75	0.872	3.599
21	T002 – 11	75	0.810	3.852
22	T002 – 13	0	0.722	2.799
23	T003 – 1	25	0.792	3.296
24	T003 – 2	25	0.793	3.518
25	T003 – 3	25	0.727	3.295
26	T003 – 4	25	0.770	3.811
27	T003 – 5	50	0.798	3.347
28	T003 – 6	50	0.848	3.460
29	T003 – 7	50	0.820	3.754
30	T003 – 9	75	0.788	3.357
31	T003 – 10	75	0.895	4.067
32	T003 – 13	0	0.778	3.178

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
33	T004 – 1	25	0.774	3.444
34	T004 – 2	25	0.790	3.469
35	T004 – 3	25	0.773	3.688
36	T004 – 4	25	0.782	3.714
37	T004 – 5	50	0.773	3.482
38	T004 – 6	50	0.841	3.840
39	T004 – 9	75	0.794	3.558
40	T004 – 10	75	0.895	4.020
41	T004 – 13	0	0.774	3.424
42	T005 – 1	25	0.739	2.600
43	T005 – 2	25	0.751	2.841
44	T005 – 3	25	0.753	2.886
45	T005 – 4	25	0.732	2.788
46	T005 – 5	50	0.804	2.761
47	T005 – 6	50	0.865	3.087
48	T005 – 7	50	0.842	3.223
49	T005 – 8	50	0.780	3.188
50	T005 – 9	75	0.830	3.176
51	T005 – 10	75	0.921	3.657
52	T005 – 11	75	0.930	3.845
53	T005 – 13	0	0.726	2.549
54	T006 – 1	25	0.725	2.494
55	T006 – 2	25	0.741	2.520
56	T006 – 3	25	0.745	2.508
57	T006 – 4	25	0.751	2.451
58	T006 – 5	50	0.783	2.710
59	T006 – 6	50	0.833	2.869
60	T006 – 7	50	0.817	2.790
61	T006 – 9	75	0.795	2.963
62	T006 – 10	75	0.880	3.273
63	T006 – 13	0	0.732	2.321
64	T007 – 1	25	0.817	2.872
65	T007 – 2	25	0.827	3.084

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
66	T007 – 3	25	0.834	3.108
67	T007 – 4	25	0.836	3.045
68	T007 – 5	50	0.826	3.135
69	T007 – 6	50	0.871	3.428
70	T007 – 7	50	0.853	3.188
71	T007 – 9	75	0.830	3.392
72	T007 – 10	75	0.918	3.802
73	T007 – 11	75	0.880	3.788
74	T007 – 12	75	0.830	4.004
75	T007 – 13	0	0.825	2.856
76	T008 – 1	25	0.792	2.922
77	T008 – 2	25	0.813	3.145
78	T008 – 3	25	0.797	2.948
79	T008 – 4	25	0.799	2.949
80	T008 – 5	50	0.803	3.037
81	T008 – 6	50	0.853	2.933
82	T008 – 7	50	0.837	3.212
83	T008 – 9	75	0.810	3.307
84	T008 – 10	75	0.873	3.426
85	T008 – 13	0	0.790	2.959
86	T009 – 1	25	0.740	3.077
87	T009 – 2	25	0.743	3.408
88	T009 – 3	25	0.718	3.440
89	T009 – 4	25	0.713	3.672
90	T009 – 5	50	0.786	3.549
91	T009 – 6	50	0.821	3.613
92	T009 – 7	50	0.891	3.860
93	T009 – 8	50	0.810	4.146
94	T009 – 9	75	0.830	3.621
95	T009 – 10	75	0.888	4.135
96	T009 – 11	75	0.820	4.428
97	T009 – 13	0	0.710	3.052
98	T010 – 1	25	0.805	3.383

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
99	T010 – 2	25	0.799	3.475
100	T010 – 3	25	0.796	3.567
101	T010 – 4	25	0.805	3.599
102	T010 – 5	50	0.807	3.222
103	T010 – 6	50	0.845	3.504
104	T010 – 7	50	0.915	3.800
105	T010 – 8	50	0.890	3.814
106	T010 – 9	75	0.814	3.437
107	T010 – 10	75	0.888	4.135
108	T011 – 1	25	0.891	3.388
109	T011 – 2	25	0.810	3.225
110	T011 – 3	25	0.830	3.084
111	T011 – 4	25	0.888	3.316
112	T011 – 5	50	0.820	3.024
113	T011 – 6	50	0.891	3.497
114	T011 – 7	50	0.810	3.609
115	T011 – 9	75	0.830	3.158
116	T011 – 10	75	0.888	3.631
117	T011 – 13	0	0.820	3.085
118	T0012 – 1	25	0.754	3.048
119	T0012 – 2	25	0.736	3.188
120	T0012 – 3	25	0.770	3.296
121	T0012 – 4	25	0.812	3.346
122	T0012 – 5	50	0.764	3.108
123	T0012 – 6	50	0.809	3.531
124	T0012 – 9	75	0.792	3.181
125	T0012 – 10	75	0.842	3.693
126	T0012 – 13	0	0.752	3.153
127	T013 – 1	25	0.812	2.906
128	T013 – 2	25	0.812	3.001
129	T013 – 3	25	0.813	2.906
130	T013 – 4	25	0.862	3.380
131	T013 – 5	50	0.835	3.357

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
132	T013 – 6	50	0.894	3.379
133	T013 – 7	50	0.867	3.385
134	T013 – 8	50	0.890	3.424
135	T013 – 9	75	0.846	3.209
136	T013 – 10	75	0.925	3.545
137	T013 – 11	75	0.881	3.974
138	T013 – 12	75	0.800	4.120
139	T013 – 13	0	0.783	2.947
140	T014 – 1	25	0.848	3.233
141	T014 – 2	25	0.855	3.133
142	T014 – 3	25	0.850	3.191
143	T014 – 4	25	0.875	3.342
144	T014 – 5	50	0.851	3.549
145	T014 – 6	50	0.881	3.435
146	T014 – 7	50	0.870	3.427
147	T014 – 8	50	0.897	3.503
148	T014 – 9	75	0.849	3.753
149	T014 – 10	75	0.920	4.349
150	T014 – 11	75	0.878	4.110
151	T014 – 12	75	0.880	4.147
152	T014 – 13	0	0.844	2.883
153	T015 – 1	25	0.850	3.343
154	T015 – 2	25	0.856	3.473
155	T015 – 3	25	0.856	3.553
156	T015 – 4	25	0.872	3.587
157	T015 – 5	50	0.845	3.838
158	T015 – 6	50	0.879	4.043
159	T015 – 7	50	0.869	3.937
160	T015 – 8	50	0.897	4.017
161	T015 – 9	75	0.835	3.396
162	T015 – 10	75	0.903	4.426
163	T015 – 11	75	0.879	4.403
164	T015 – 12	75	0.860	4.261

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
165	T015 – 13	0	0.861	3.243
166	T016 – 1	25	0.847	3.818
167	T016 – 2	25	0.841	3.984
168	T016 – 3	25	0.845	4.112
169	T016 – 4	25	0.860	4.147
170	T016 – 5	50	0.828	3.866
171	T016 – 6	50	0.879	3.858
172	T016 – 7	50	0.866	3.695
173	T016 – 8	50	0.898	4.054
174	T016 – 12	75	0.810	4.421
175	T016 – 13	0	0.846	3.773
176	T017 – 1	25	0.730	3.069
177	T017 – 2	25	0.742	3.232
178	T017 – 3	25	0.733	3.275
179	T017 – 4	25	0.801	3.354
180	T017 – 5	50	0.795	3.374
181	T017 – 6	50	0.841	3.630
182	T017 – 7	50	0.824	3.797
183	T017 – 8	50	0.885	4.007
184	T017 – 9	75	0.819	3.611
185	T017 – 10	75	0.908	3.794
186	T017 – 13	0	0.705	2.892
187	T018 – 1	25	0.800	3.378
188	T018 – 2	25	0.789	3.467
189	T018 – 3	25	0.783	3.556
190	T018 – 4	25	0.794	3.589
191	T018 – 5	50	0.806	3.529
192	T018 – 6	50	0.853	3.511
193	T018 – 7	50	0.836	4.082
194	T018 – 8	50	0.810	4.096
195	T018 – 9	75	0.827	3.777
196	T018 – 10	75	0.910	4.158
197	T018 – 13	0	0.815	3.238

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
198	T019 – 1	25	0.809	3.625
199	T019 – 2	25	0.817	3.768
200	T019 – 3	25	0.825	3.798
201	T019 – 4	25	0.830	3.846
202	T019 – 5	50	0.818	3.702
203	T019 – 6	50	0.855	3.688
204	T019 – 7	50	0.842	4.265
205	T019 – 8	50	0.910	4.509
206	T019 – 9	75	0.820	3.522
207	T019 – 10	75	0.885	4.462
208	T019 – 11	75	0.890	4.803
209	T019 – 13	0	0.815	3.489
210	T020 – 1	25	0.730	3.121
211	T020 – 2	25	0.742	2.951
212	T020 – 3	25	0.733	3.250
213	T020 – 4	25	0.801	3.379
214	T020 – 5	50	0.795	3.035
215	T020 – 6	50	0.841	3.654
216	T020 – 7	50	0.824	3.819
217	T020 – 8	50	0.885	3.899
218	T020 – 9	75	0.819	4.085
219	T020 – 13	0	0.780	4.259
220	T021 – 1	25	0.796	3.546
221	T021 – 2	25	0.801	3.456
222	T021 – 3	25	0.798	3.480
223	T021 – 4	25	0.892	3.643
224	T021 – 5	50	0.819	3.739
225	T021 – 6	50	0.853	4.097
226	T021 – 7	50	0.934	4.136
227	T021 – 8	50	0.950	4.177
228	T021 – 9	75	0.926	3.934
229	T021 – 10	75	0.950	4.448
230	T021 – 11	75	1.038	4.967

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
231	T021 – 12	75	0.960	5.152
232	T021 – 13	0	0.797	3.306
233	T022 – 1	25	0.850	3.852
234	T022 – 2	25	0.856	4.270
235	T022 – 3	25	0.847	4.346
236	T022 – 4	25	0.867	4.367
237	T022 – 5	50	0.866	3.826
238	T022 – 6	50	0.884	4.382
239	T022 – 7	50	0.873	4.277
240	T022 – 8	50	0.886	4.357
241	T022 – 9	75	0.852	3.959
242	T022 – 10	75	0.924	4.867
243	T022 – 11	75	0.874	5.596
244	T022 – 13	0	0.845	3.930
245	T023 – 1	25	0.855	3.502
246	T023 – 2	25	0.854	4.251
247	T023 – 3	25	0.856	4.352
248	T023 – 4	25	0.872	4.393
249	T023 – 5	50	0.849	3.956
250	T023 – 6	50	0.883	4.837
251	T023 – 7	50	0.751	4.685
252	T023 – 8	50	0.783	4.510
253	T023 – 9	75	0.833	4.626
254	T023 – 10	75	0.817	4.845
255	T023 – 11	75	0.800	5.289
256	T023 – 12	75	0.795	5.547
257	T023 – 13	0	0.880	4.267
258	T024 – 1	25	0.840	4.018
259	T024 – 2	25	0.870	4.230
260	T024 – 3	25	0.732	4.781
261	T024 – 4	25	0.817	4.588
262	T024 – 5	50	0.827	4.321
263	T024 – 6	50	0.834	4.996

(cot $\alpha = 1.5$) Continue on the next page

Table C.2 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 1.5$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
264	T024 – 7	50	0.836	5.377
265	T024 – 8	50	0.826	5.427
266	T024 – 9	75	0.871	4.417
267	T024 – 10	75	0.853	4.724
268	T024 – 11	75	0.880	5.611
269	T024 – 12	75	0.830	5.663
270	T024 – 13	0	0.918	4.048

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$

ID	Test No	Swell (%)	Reff Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
1	T001 – 1	25	0.841	2.618
2	T001 – 2	25	0.847	2.707
3	T001 – 3	25	0.848	2.750
4	T001 – 4	25	0.846	2.749
5	T001 – 5	50	0.824	2.775
6	T001 – 6	50	0.844	2.966
7	T001 – 7	50	0.864	3.066
8	T001 – 9	75	0.806	2.882
9	T001 – 10	75	0.850	3.361
10	T001 – 13	1	0.832	2.458
11	T002 – 1	25	0.869	2.826
12	T002 – 2	25	0.870	2.927
13	T002 – 3	25	0.867	2.983
14	T002 – 4	25	0.871	2.986
15	T002 – 5	50	0.847	2.930
16	T002 – 6	50	0.843	3.150
17	T002 – 7	50	0.857	3.228
18	T002 – 9	75	0.802	3.065
19	T002 – 10	75	0.841	3.450
20	T002 – 13	1	0.873	2.682

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
21	T003 – 1	25	0.853	3.097
22	T003 – 2	25	0.829	3.086
23	T003 – 3	25	0.840	3.166
24	T003 – 4	25	0.826	3.155
25	T003 – 5	50	0.826	3.028
26	T003 – 6	50	0.832	3.330
27	T003 – 7	50	0.854	3.412
28	T003 – 9	75	0.800	3.135
29	T003 – 10	75	0.830	3.466
30	T003 – 13	1	0.859	2.899
31	T004 – 1	25	0.822	3.116
32	T004 – 2	25	0.799	3.134
33	T004 – 3	25	0.836	3.243
34	T004 – 4	25	0.840	3.369
35	T004 – 5	50	0.804	3.138
36	T004 – 6	50	0.819	3.308
37	T004 – 7	50	0.832	3.467
38	T004 – 9	75	0.795	3.183
39	T004 – 10	75	0.817	3.554
40	T004 – 13	1	0.807	3.087
41	T005 – 1	25	0.877	2.904
42	T005 – 2	25	0.870	3.044
43	T005 – 3	25	0.862	3.085
44	T005 – 4	25	0.863	3.109
45	T005 – 5	50	0.859	3.129
46	T005 – 6	50	0.832	3.307
47	T005 – 7	50	0.875	3.510
48	T005 – 8	50	0.780	3.425
49	T005 – 9	75	0.815	3.293
50	T005 – 10	75	0.817	3.709
51	T005 – 11	75	0.930	3.990
52	T005 – 13	1	0.867	2.746
53	T006 – 1	25	0.873	2.829

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
54	T006 – 2	25	0.886	2.939
55	T006 – 3	25	0.873	2.988
56	T006 – 4	25	0.876	2.990
57	T006 – 5	50	0.849	2.932
58	T006 – 6	50	0.834	3.143
59	T006 – 7	50	0.864	3.234
60	T006 – 9	75	0.801	3.065
61	T006 – 10	75	0.829	3.440
62	T006 – 13	1	0.880	2.708
63	T007 – 1	25	0.855	3.216
64	T007 – 2	25	0.857	3.474
65	T007 – 3	25	0.863	3.499
66	T007 – 4	25	0.859	3.536
67	T007 – 5	50	0.845	3.402
68	T007 – 6	50	0.839	3.673
69	T007 – 7	50	0.861	3.912
70	T007 – 9	75	0.810	3.513
71	T007 – 10	75	0.846	4.034
72	T007 – 11	75	0.874	4.367
73	T007 – 12	75	0.830	4.624
74	T007 – 13	1	0.861	3.221
75	T008 – 1	25	0.826	3.038
76	T008 – 2	25	0.836	3.266
77	T008 – 3	25	0.848	3.253
78	T008 – 4	50	0.829	3.344
79	T008 – 5	50	0.823	3.052
80	T008 – 6	50	0.841	3.341
81	T008 – 7	50	0.852	3.514
82	T008 – 9	75	0.822	3.169
83	T008 – 10	75	0.846	3.508
84	T008 – 13	1	0.843	2.998
85	T009 – 1	25	0.861	3.021
86	T009 – 2	25	0.858	3.035

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
87	T009 – 3	25	0.848	3.074
88	T009 – 4	25	0.844	2.973
89	T009 – 5	50	0.823	3.101
90	T009 – 6	50	0.817	3.295
91	T009 – 7	50	0.866	3.502
92	T009 – 8	50	0.810	3.591
93	T009 – 9	75	0.810	3.289
94	T009 – 10	75	0.795	3.689
95	T009 – 11	75	0.856	4.078
96	T009 – 13	1	0.856	2.739
97	T010 – 1	25	0.805	3.088
98	T010 – 2	25	0.926	3.274
99	T010 – 3	25	0.919	3.225
100	T010 – 4	25	0.863	3.332
101	T010 – 5	50	0.859	3.264
102	T010 – 6	50	0.815	3.478
103	T010 – 7	50	0.875	3.763
104	T010 – 8	50	0.890	3.814
105	T010 – 9	75	0.796	3.422
106	T010 – 10	75	0.794	3.688
107	T010 – 13	1	0.887	3.009
108	T011 – 1	25	0.861	2.998
109	T011 – 2	25	0.880	3.281
110	T011 – 3	25	0.830	3.084
111	T011 – 4	25	0.840	3.065
112	T011 – 5	50	0.852	3.048
113	T011 – 6	50	0.821	3.215
114	T011 – 7	50	0.808	3.374
115	T011 – 9	75	0.810	3.043
116	T011 – 10	75	0.814	3.452
117	T011 – 13	1	0.867	2.905
118	T0012 – 1	25	0.843	3.033
119	T0012 – 2	25	0.821	3.151

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
120	T0012-3	25	0.832	3.240
121	T0012-4	25	0.811	3.740
122	T0012-5	50	0.822	3.052
123	T0012-6	50	0.804	3.487
124	T0012-9	75	0.794	3.182
125	T0012-10	75	0.809	3.547
126	T0012-13	1	0.828	3.103
127	T013-1	25	0.873	3.407
128	T013-2	25	0.874	3.519
129	T013-3	25	0.870	3.542
130	T013-4	25	0.873	3.626
131	T013-5	50	0.858	3.559
132	T013-6	50	0.820	3.832
133	T013-7	50	0.874	4.075
134	T013-8	50	0.872	4.096
135	T013-9	75	0.804	3.817
136	T013-10	75	0.800	4.286
137	T013-11	75	0.873	4.767
138	T013-12	75	0.800	4.698
139	T013-13	1	0.867	3.171
140	T014-1	25	0.894	3.671
141	T014-2	25	0.891	3.799
142	T014-3	25	0.885	3.869
143	T014-4	25	0.873	3.857
144	T014-5	50	0.865	3.825
145	T014-6	50	0.825	3.863
146	T014-7	50	0.873	4.277
147	T014-8	50	0.873	4.343
148	T014-9	75	0.808	3.916
149	T014-10	75	0.754	4.401
150	T014-11	75	0.875	4.683
151	T014-12	75	0.880	5.172
152	T014-13	1	0.894	3.508

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
153	T015 – 1	25	0.864	3.873
154	T015 – 2	25	0.841	3.995
155	T015 – 3	25	0.851	4.098
156	T015 – 4	25	0.841	4.111
157	T015 – 5	50	0.836	3.943
158	T015 – 6	50	0.799	4.179
159	T015 – 7	50	0.842	4.516
160	T015 – 8	50	0.855	4.591
161	T015 – 9	75	0.780	4.027
162	T015 – 10	75	0.773	4.519
163	T015 – 11	75	0.873	5.077
164	T015 – 12	75	0.874	5.059
165	T015 – 13	1	0.863	3.747
166	T016 – 1	25	0.829	4.007
167	T016 – 2	25	0.770	3.916
168	T016 – 3	25	0.833	4.321
169	T016 – 4	25	0.830	4.339
170	T016 – 5	50	0.793	4.040
171	T016 – 6	50	0.764	4.331
172	T016 – 7	50	0.835	4.498
173	T016 – 8	50	0.833	5.085
174	T016 – 9	75	0.755	4.047
175	T016 – 10	75	0.764	4.623
176	T016 – 11	75	0.879	5.018
177	T016 – 12	75	0.874	5.398
178	T016 – 13	1	0.835	3.966
179	T017 – 1	25	0.877	2.904
180	T017 – 2	25	0.876	3.049
181	T017 – 3	25	0.868	3.090
182	T017 – 4	25	0.862	2.986
183	T017 – 5	50	0.858	3.128
184	T017 – 6	50	0.864	3.333
185	T017 – 7	50	0.886	3.519

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
186	T017-8	50	0.882	3.512
187	T017-9	75	0.837	3.311
188	T017-10	75	0.866	3.754
189	T017-11	75	0.885	3.946
190	T017-13	1	0.870	2.748
191	T018-1	25	0.898	3.159
192	T018-2	25	0.889	3.244
193	T018-3	25	0.887	3.330
194	T018-4	25	0.886	3.351
195	T018-5	50	0.863	3.267
196	T018-6	50	0.848	3.506
197	T018-7	50	0.887	3.626
198	T018-8	50	0.892	3.666
199	T018-9	75	0.828	3.449
200	T018-10	75	0.869	3.757
201	T018-11	75	0.910	3.971
202	T018-13	1	0.900	3.019
203	T019-1	25	0.853	3.214
204	T019-2	25	0.859	3.475
205	T019-3	25	0.865	3.364
206	T019-4	25	0.864	3.402
207	T019-5	50	0.846	3.403
208	T019-6	50	0.848	3.682
209	T019-7	50	0.864	3.915
210	T019-8	50	0.878	4.083
211	T019-9	75	0.818	3.520
212	T019-10	75	0.862	4.050
213	T019-11	75	0.890	4.385
214	T019-13	1	0.862	3.222
215	T020-1	25	0.730	3.236
216	T020-2	25	0.742	3.569
217	T020-3	25	0.733	3.653
218	T020-4	25	0.801	3.731

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
219	T020 – 5	50	0.795	3.501
220	T020 – 6	50	0.841	3.822
221	T020 – 7	50	0.824	4.062
222	T020 – 8	50	0.885	4.206
223	T020 – 9	75	0.819	3.579
224	T020 – 10	75	0.820	4.042
225	T020 – 11	75	0.830	4.468
226	T020 – 13	1	0.780	3.409
227	T021 – 1	25	0.860	3.396
228	T021 – 2	25	0.865	3.511
229	T021 – 3	25	0.861	3.534
230	T021 – 4	25	0.859	3.613
231	T021 – 5	50	0.851	3.768
232	T021 – 6	50	0.799	3.812
233	T021 – 7	50	0.862	4.063
234	T021 – 8	50	0.950	4.177
235	T021 – 9	75	0.798	3.811
236	T021 – 10	75	0.779	4.265
237	T021 – 11	75	1.038	4.712
238	T021 – 12	75	0.860	4.767
239	T021 – 13	1	0.860	3.165
240	T022 – 1	25	0.886	3.664
241	T022 – 2	25	0.879	3.787
242	T022 – 3	25	0.874	3.858
243	T022 – 4	25	0.869	3.853
244	T022 – 5	50	0.859	3.623
245	T022 – 6	50	0.812	4.059
246	T022 – 7	50	0.863	4.267
247	T022 – 8	50	0.867	4.336
248	T022 – 9	75	0.801	3.910
249	T022 – 10	75	0.752	4.399
250	T022 – 11	75	0.874	4.682
251	T022 – 13	1	0.883	3.499

(cot $\alpha = 0.0$) Continue on the next page

Table C.3 Reflection Characteristics Under Bimodal Seas for $\cot \alpha = 0.0$ (cont.).

ID	Test No	Swell (%)	Ref. Coeff K_r (-)	Surf Similarity Parameter $\xi_{m-1,0}$ (-)
252	T023 – 1	25	0.856	3.866
253	T023 – 2	25	0.841	3.995
254	T023 – 3	25	0.850	4.097
255	T023 – 4	25	0.841	4.111
256	T023 – 5	50	0.838	3.945
257	T023 – 6	50	0.785	4.165
258	T023 – 7	50	0.789	4.729
259	T023 – 8	50	0.844	4.578
260	T023 – 9	75	0.781	4.028
261	T023 – 10	75	0.760	4.505
262	T023 – 11	75	0.864	5.065
263	T023 – 12	75	0.795	4.961
264	T023 – 13	1	0.858	3.742
265	T024 – 1	25	0.835	4.013
266	T024 – 2	25	0.772	3.918
267	T024 – 3	25	0.832	4.320
268	T024 – 4	25	0.809	4.317
269	T024 – 5	50	0.802	4.049
270	T024 – 6	50	0.759	4.326
271	T024 – 7	50	0.830	4.492
272	T024 – 8	50	0.827	4.542
273	T024 – 9	75	0.756	4.048
274	T024 – 10	75	0.738	4.594
275	T024 – 11	75	0.869	5.005
276	T024 – 12	75	0.863	5.107
277	T024 – 13	1	0.835	3.966

Appendix D

Examples of Wave Gauge Calibration

Accuracy of recorded surface elevation is an essential component of this research. As earlier mentioned, the wave gauges are regularly calibrated to ascertain better reliability of the outcome. Using the calibration routine of the HR Daq system, the linear relationship between raw voltages at zero water level and the voltages during wave motion are established. Recorded voltages are linearly converted into water surface elevation using Equation D.1.

$$Mod_{value} = (mod_{fac} * Dat_{volts}) - Zero_{volts} \quad (D.1)$$

where Mod_{value} is the recorded surface elevation at the model scale, mod_{fac} is the model factor of calibration, and $Zero_{volts}$ is the voltage values recorded at still water level. Appendix D present results of wave gauge calibration obtained for the first five tests conducted during this study. Respective channel numbers represent each wave gauge. Channel 1 is closest to the wavemaker while Channel 8 is positioned at the crest of the structure. Acceptability of calibration results depends on the value of the correlation R^2 factor at 95 per cent CI. Calibration values that are less than 0.999 are repeated to obtain a better result. Appendix D contains examples of outputs of selected wave gauge calibration results.

Table D.1 Test 1. wave gauge calibration

Name	Cal (m/V)	Zero (V)	R ²	95% CI
Channel 1	0.0335	0.029	1.00000	0.0000
Channel 2	0.0388	-0.115	0.99999	0.0002
Channel 3	0.0394	-0.084	0.99999	0.0002
Channel 5	0.0365	-0.024	1.00000	0.0001
Channel 6	0.0353	-0.109	0.99998	0.0003
Channel 7	0.0323	-0.092	0.99995	0.0004
Channel 8	0.0288	0.017	1.00000	0.0000

Table D.2 Test 2. wave gauge calibration

Name	Cal (m/V)	Zero (V)	R ²	95% CI
Channel 1	0.0281	0.093	0.99986	0.0005
Channel 2	0.0323	0.004	0.99994	0.0004
Channel 3	0.0331	-0.022	0.99985	0.0006
Channel 5	0.0316	-0.133	0.99995	0.0003
Channel 6	0.0299	-0.015	0.99984	0.0005
Channel 7	0.0298	-0.141	0.99996	0.0003
Channel 8	0.0216	-0.037	0.99979	0.0004

Table D.3 Test 3. wave gauge calibration

Name	Cal (m/V)	Zero (V)	R ²	95% CI
Channel 1	0.0274	-0.006	1.00000	0.0001
Channel 2	0.0314	-0.127	1.00000	0.0000
Channel 3	0.0319	-0.119	1.00000	0.0001
Channel 4	0.0292	-0.009	1.00000	0.0000
Channel 5	0.0302	0.003	1.00000	0.0000
Channel 6	0.0282	-0.093	0.99998	0.0002
Channel 7	0.0183	-0.367	1.00000	0.0000
Channel 8	0.0195	0.048	1.00000	0.0000

Table D.4 Test 4. wave gauge calibration

Name	Cal (m/V)	Zero (V)	R ²	95% CI
Channel 1	0.0259	0.014	0.99990	0.0004
Channel 2	0.0298	-0.105	0.99995	0.0003
Channel 3	0.0303	-0.123	1.00000	0.0000
Channel 4	0.0279	0.005	0.99999	0.0001
Channel 5	0.0291	-0.041	1.00000	0.0001
Channel 6	0.0273	-0.123	0.99999	0.0001
Channel 7	0.0234	-0.065	0.99994	0.0003
Channel 8	0.026	0.043	0.99996	0.0003

Table D.5 Test 5. wave gauge calibration

Name	Cal (m/V)	Zero (V)	R ²	95% CI
Channel 1	0.0256	0.034	1.00000	0.0001
Channel 2	0.0294	-0.066	1.00000	0.0000
Channel 3	0.0299	-0.106	1.00000	0.0001
Channel 4	0.0273	0.029	0.99999	0.0001
Channel 5	0.0285	-0.039	1.00000	0.0001
Channel 6	0.0267	-0.09	1.00000	0.0001
Channel 7	0.0229	-0.064	1.00000	0.0001
Channel 8	0.0261	-0.025	1.00000	0.0000