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1 **Environmental drivers of nitrous oxide emission factor for a coastal**
2 **reservoir and its catchment areas in southeastern China**

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18 **ABSTRACT**

19 Asia is projected to be a major contributor to nitrous oxide (N₂O) emission in the
20 coming decades, but assessment of N₂O budget and distribution has been hampered
21 by low data resolution and poorly constrained emission factor (EF). Urbanized coastal
22 reservoirs receive high nitrogen loads from diverse sources across a heterogeneous
23 landscape, and using a fixed EF may lead to large errors in N₂O assessment. We
24 conducted high spatial resolution sampling of dissolved N₂O, nitrate-nitrogen (NO₃⁻-
25 N) and related hydrographical parameters in Wenwusha Reservoir and its catchment
26 areas (river, drainage channels, and aquaculture ponds) in southeastern China in
27 November 2018, March 2019 and June 2019. The empirically derived EF (calculated
28 as N₂O-N:NO₃⁻-N) for the reservoir showed 10-fold spatial variations, ranging from
29 0.8×10^{-3} to 8.8×10^{-3} . The average EF varied significantly among the four water types
30 in the following descending order: aquaculture ponds > river > drainage channels >
31 reservoir. Across all water types, EF of the summer month was 1.8–3.5 and 1.7–2.8
32 fold higher on average than that of autumn and spring, respectively. EF was higher in
33 the summer likely due to elevated water temperature. Overall, the EF deviated
34 considerably from the Intergovernmental Panel on Climate Change (IPCC) default
35 value such that using the latter would result in over- or under-estimation of N₂O
36 emissions, sometimes by up to 42%. A new regression algorithm for EF based on
37 water temperature, dissolved organic carbon and nitrate-nitrogen had a high and
38 significant explanatory power ($r^2 = 0.82$; $p < 0.001$), representing an improvement

39 over the IPCC default EF for assessing N₂O emission from coastal reservoirs and
40 similar environments.

41 **Keywords:** N₂O; Greenhouse gas; IPCC; Spatio-temporal variation; Nitrate-nitrogen;
42 Inland waters

43 **1. Introduction**

44 As a greenhouse gas, nitrous oxide (N_2O) has approximately 270 times the
45 warming effect of carbon dioxide on a century timescale (Neubauer and Megonigal,
46 2015). The concentration of N_2O in the atmosphere increased at a rate of 0.7–0.8 ppbv
47 per year in the past five decades (Davidson, 2009; Saikawa et al., 2014), and reached
48 334 ppbv in 2021 (National Oceanic and Atmospheric, 2021). Biological conversion of
49 anthropogenic nitrogen to N_2O in soil and surface waters is estimated to contribute
50 approximately 55% of the N_2O emission globally (Ciais et al., 2014; Zhang et al., 2020).
51 Therefore, effective climate change mitigation will require better understanding of
52 spatiotemporal distribution of N_2O emission and its environmental drivers.

53 N_2O emission from soil has been studied extensively (e.g., Griffis et al., 2013;
54 Mosier et al., 1998) and is a relatively well-constrained component of the global N_2O
55 assessment (Quick et al., 2019; Zhang et al., 2020). Agriculture accounts for the largest
56 N_2O emission from soil (Tian et al., 2020). By comparison, N_2O emission from aquatic
57 systems including those influenced by nutrient runoff is poorly constrained (Outram
58 and Hiscock, 2012; Webb et al., 2021) and it remains a major source of uncertainty in
59 the global N_2O budget.

60 Because of the chemical stability of N_2O in surface waters, N_2O emission is
61 directly proportional to dissolved N_2O concentration. The production of N_2O itself,
62 however, is a rather complex biogeochemical process as it can involve multiple
63 precursors and redox reactions (Schreiber et al., 2012). In practice, the

64 Intergovernmental Panel on Climate Change (IPCC) estimates waterborne N₂O
65 emission as anthropogenic N loading multiplied by an emission factor (EF) and
66 leaching loss (IPCC, 2006). This approach is only suitable where the source and
67 amount of N input from the catchment can be readily defined and leaching loss is
68 known and invariant, such may be the case of fertilized farmland. Also, to initially
69 establish the EF one needs N₂O emission that is usually estimated from dissolved N₂O
70 concentration based on a wind-based gas transfer model, which itself is often not
71 calibrated for the local conditions such as river flowrate and sheltering effect from the
72 surrounding landscape (e.g., Qin et al. 2019). As such, application of the EF to running
73 water through a variable landscape can lead to significant errors.

74 For urbanized and artificial aquatic systems such as coastal reservoirs, N₂O
75 dynamics is compounded by N input from diverse and diffuse sources, such as
76 contributions from wastewater discharge, fertilization (such as fish feeds in aquaculture
77 ponds), leaching and runoff from surrounding landscape as well as atmospheric
78 deposition, in addition to in situ N₂O production. To assess N₂O emission from such
79 systems, an alternative EF can be calculated as the ratio between dissolved
80 concentrations of nitrous oxide nitrogen (N₂O-N) and nitrate-nitrogen (NO₃⁻-N)
81 (Hama-Aziz et al., 2017). This formulation is appealing because it uses parameters that
82 can be measured easily and accurately, without prior knowledge of the source of N.
83 Once the EF is established, it can be used to estimate N₂O-N from NO₃⁻-N, the latter of
84 which is a routine parameter in environmental surveys, thereby facilitating the

85 spatiotemporal assessment of N₂O. However, the conversion of nitrogenous input or
86 NO₃⁻-N to N₂O is mediated by a multitude of biological and physical factors (Schreiber
87 et al., 2012) and therefore, applying a fixed EF could be problematic (Cooper et al.,
88 2017; Fu et al. 2018; Turner et al., 2015). Indeed, some researchers have argued that the
89 IPCC default EF may have grossly overestimated N₂O emission from inland waters
90 (Maavara et al., 2019).

91 Asia is projected to account for the largest share of riverine and estuarine N₂O
92 emission in the coming decades under various Millennium Ecosystem Assessment
93 scenarios (Kroeze et al., 2010). As population size increases in many Asian coastal
94 cities, construction of coastal reservoirs is widely considered as the solution to water
95 scarcity (Yang et al., 2018). Due to increasing agricultural activities and rapid
96 urbanization in the catchment areas, coastal reservoirs often receive high N loading
97 from multiple sources, making them potential N₂O hotspots (Lacerda et al., 2008;
98 Páez-Osuna et al., 2017). The high heterogeneity in geography and hydrography around
99 and within coastal reservoirs may require a more nuanced approach to derive EF in
100 order to improve the assessment of N₂O emission from these inland water systems.

101 We conducted a field study, using high spatial resolution data of dissolved N₂O
102 and NO₃⁻-N, to assess the variability of EF across a coastal reservoir and its catchment
103 areas (river, drainage channels and aquaculture ponds) in southeastern China. We also
104 examined the relationships between EF and different hydrographical parameters, based
105 on which we derived algorithms to improve EF value for N₂O emission assessment.

106

107 **2. Materials and methods**

108 *2.1. Study area*

109 The Wenwusha Reservoir is at the mouth of the Min River Estuary in southeastern
110 China (25°49'36" to 25°54'00"N, 119°35'12" to 119°38'11"E), within a densely
111 populated area with approximately 1000 persons per km² (Fig. 1). The reservoir was
112 constructed primarily for irrigation and flood mitigation purposes, and it covers 5.2 km²
113 in surface area and 3.20×10^8 m³ in volume. The reservoir receives input from a
114 catchment area of 275 km² of diverse landscape features that include urban area (6.0%),
115 aquaculture ponds (9.8%), forest (15.0%), farmland (3.3%), sand (2.1%) and wetland
116 (14.4%) (Zhang et al., 2021). The reservoir is divided by dams into two basins (Fig. 1):
117 the north basin (NB) is heavily impacted by domestic, industrial and aquacultural waste
118 discharges as well as input from the Nanyangdong River; the south basin (SB) is
119 surrounded by farmlands (including aquaculture), wetland, forest and towns. The region
120 experiences subtropical monsoonal climate with an annual average temperature of
121 19.6 °C and precipitation of 1,390 mm (Yang et al., 2020).

122 *2.2. Sample collection and analysis*

123 High-resolution field sampling was conducted in November 2018 (autumn),
124 March 2019 (spring) and June 2019 (summer), where a total of 121 sites were sampled
125 each time, of which 103 sites were within the reservoir, seven in the river, four in

126 drainage channels, and seven in aquaculture ponds (Fig. 1). A total of 21 transects were
127 sampled within the reservoir, with 11 transects in the south basin and 10 in the north
128 basin (Fig. 1b).

129 To measure dissolved N₂O, water was collected from 20 cm below the surface
130 with a syringe and transferred into 55-mL glass serum bottles. Microbial activities in
131 the water sample were stopped by adding 0.2 mL saturated HgCl₂ solution (Zhang et al.,
132 2013); the bottle was then sealed immediately without headspace with a butyl rubber
133 stopper (Xiao et al., 2019a). Upon return to the laboratory, dissolved N₂O was
134 measured using the headspace equilibration technique (Davidson et al., 2015; Yu et al.,
135 2013). A headspace was created by displacing 25 mL of the water with nitrogen (N₂)
136 (>99.999% purity). The bottle was then shaken for 10 minutes to attain gas equilibrium
137 between the headspace and the liquid phase. 5 mL of the headspace gas was extracted
138 and injected into a gas chromatograph (GC-2014, Shimadzu, Kyoto, Japan) equipped
139 with an electron capture detector (ECD) for determining the N₂O concentration. A
140 calibration curve was produced with standard N₂O gas at 0.3, 0.4 and 1.0 ppm. The
141 N₂O measurements had a precision of ± 5% and a detection limit of 0.02 ppm.

142 Using the measured headspace N₂O concentrations, the original dissolved
143 concentrations of dissolved N₂O were calculated based on the Bunsen solubility
144 coefficient as a function of temperature and salinity (Brase et al., 2017; Weiss and Price,
145 1980). The N₂O emission factor (EF) was calculated as the ratio between N₂O-N and
146 NO₃⁻-N dissolved concentrations (Hama-Aziz et al., 2017) for the surface waters of the

147 reservoir, river, drainage channels and aquaculture ponds.

148 Additional water samples were collected into 150 mL polyethylene bottles at each
149 site and filtered through a 0.45 μm cellulose acetate filter (BiotransTM nylon
150 membranes). The filtrates were analyzed for nitrate-nitrogen (NO_3^- -N) and total
151 dissolved nitrogen (TDN) by a flow injection analyzer (Skalar Analytical SAN⁺⁺,
152 Netherlands), and dissolved organic carbon (DOC) by a TOC analyzer
153 (TOC-VCPH/CPN, Shimadzu, Japan). The detection limit and relative standard
154 deviations were 0.6 $\mu\text{g L}^{-1}$ and $\leq 2.0\%$ for NO_3^- -N, 3.0 $\mu\text{g L}^{-1}$ and $\leq 2.0\%$ for TDN,
155 and 0.4 $\mu\text{g L}^{-1}$ and $\leq 1.0\%$ for TOC, respectively.

156 In each sampling campaign, we also measured various hydrographical parameters
157 of surface water in situ at a depth of 20 cm at each site. Water temperature (T_w) and pH
158 were measured by a portable meter (IQ150, IQ Scientific Instruments, USA).
159 Conductivity (EC), dissolved oxygen (DO) and chemical oxygen demand (COD) were
160 determined by a EC meter (2265FS EC, Spectrum Technologies, USA), a water quality
161 checker (HORIBA, Japan) and a COD detector (LH-CM3H, China), respectively.

162 2.3. *Statistical analysis*

163 Statistical tests were conducted in SPSS version 17.0 (SPSS Inc., USA) at a
164 significance level of $p < 0.05$. Differences in hydrographical properties, dissolved N_2O
165 concentration and EF between water bodies (e.g., reservoir, river, drainage channels,
166 aquaculture ponds) and between reservoir basins (e.g., north basin and south basin)
167 were tested by one-way analysis of variance (ANOVA). Relationships between

168 hydrographical parameters and EF were tested by Pearson correlation analysis.
169 OriginPro version 7.5 (OriginLab Corp. USA) was used to generate the statistical plots.
170 The Kriging method in ArcGIS 10.2 (ESRI Inc., Redlands, CA, USA) was used for
171 spatial interpolation of EF in the reservoir. All results were presented as mean \pm 1
172 standard error (S.E.), unless otherwise stated.

173

174 **3. Results**

175 *3.1. Hydrographical conditions*

176 There was no significant difference in mean surface water T_W among the four
177 water types ($p > 0.05$; Fig. 2a). However, the surface water chemical properties varied
178 greatly. Overall, the mean pH (Fig. 2b) and EC (Fig. 2c) were highest in aquaculture
179 ponds (9.5 ± 0.6 and 3.49 ± 1.79 mS cm^{-1} , respectively), followed by reservoir ($8.7 \pm$
180 1.1 and 2.6 ± 1.0 mS cm^{-1}), drainage channels (8.1 ± 0.3 and 1.8 ± 0.5 mS cm^{-1}) and
181 river (7.7 ± 0.2 and 1.3 ± 0.3 mS cm^{-1}) ($p < 0.05$). The mean DOC concentration was
182 also highest in aquaculture ponds (8.4 ± 2.3 mg L^{-1}), followed by river (7.4 ± 2.0 mg
183 L^{-1}), drainage channels (7.2 ± 1.8 mg L^{-1}) and reservoir (4.1 ± 0.9 mg L^{-1}) (Fig. 2f). In
184 the river, the mean DO (4.0 ± 0.3 mg L^{-1}) (Fig. 2d) and COD (23.3 ± 2.4 mg L^{-1}) (Fig.
185 2e) concentrations were significantly lower ($p < 0.05$) while NO_3^- -N (1.64 ± 0.2 mg L^{-1})
186 (Fig. 2g) and TDN (2.2 ± 0.4 mg L^{-1}) (Fig. 2h) concentrations were generally higher
187 than in the other areas ($p < 0.05$).

188 The hydrographical conditions also showed strong temporal variations. Across the

189 four surface water types, the mean T_w (28.1 ± 0.5 °C) (Fig. 2a) and DOC (9.9 ± 1.4 mg
190 L^{-1}) (Fig. 2f) were highest in summer, whereas the mean pH (9.5 ± 0.7) (Fig. 2b) and
191 EC (3.8 ± 1.2 mS cm^{-1}) (Fig. 2c) were highest in autumn. We observed generally higher
192 DO (7.4 ± 1.2 mg L^{-1}) (Fig. 2d) in spring and higher COD (34.5 ± 3.0 mg L^{-1}) (Fig. 2e)
193 in autumn. The mean NO_3^- -N (Fig. 2g) and TDN (Fig. 2h) concentrations in river (2.0
194 ± 0.3 mg L^{-1} and 2.9 ± 0.1 mg L^{-1} , respectively) and drainage channels (1.4 ± 0.1 and
195 2.5 ± 0.1 mg L^{-1}) were highest in spring; in contrast, the mean NO_3^- -N (Fig. 2g) and
196 TDN (Fig. 2h) concentrations in aquaculture ponds (1.4 ± 0.1 and 1.4 ± 0.1 mg L^{-1}) and
197 reservoir (2.2 ± 0.1 and 1.9 ± 0.1 mg L^{-1}) were highest in autumn.

198 3.2. N_2O concentration in surface waters

199 The dissolved N_2O concentration in river, drainage channels and aquaculture ponds
200 varied seasonally in this decreasing order: summer > spring > autumn (Fig. 3). In the
201 reservoir, the mean N_2O concentration was slightly higher in spring (Fig. 3). Across all
202 sampling dates and sites, the N_2O concentration in river, drainage channels, aquaculture
203 ponds, and reservoir varied over the range of 49.6–261.3, 42.8–121.8, 42.2–198.9, and
204 6.1–261.2 nmol L^{-1} , respectively. Over the study period, there was a significant
205 difference in N_2O concentration between surface water types ($p < 0.001$; Table 1), with
206 the highest value in river (110.0 ± 36.1 nmol L^{-1}), followed by aquaculture ponds (91.7
207 ± 29.2 nmol L^{-1}), drainage channels (75.9 ± 8.2 nmol L^{-1}) and reservoir (49.8 ± 4.6
208 nmol L^{-1}).

209 3.3. Spatial and seasonal variations of EF

210 Across the reservoir, there were clear spatial differences in EF between the two
211 basins (ANOVA, $p < 0.05$; Fig. 4). Overall, EF in NB ranged from 0.3×10^{-3} to 8.8×10^{-3} ,
212 which were higher than that in SB (0.4×10^{-3} – 6.6×10^{-3}) (Fig. 4d). Significant differences
213 were also observed between sampling sites with respect to wastewater discharge
214 (ANOVA, $p < 0.05$). Generally, sites with wastewater discharge had a significantly
215 higher EF than those without wastewater discharge (Fig. 4f).

216 Over the study period, the EF value ranged from 1.8×10^{-3} – 6.0×10^{-3} in the river,
217 2.0×10^{-3} – 5.4×10^{-3} in the drainage channels, 2.1×10^{-3} – 7.4×10^{-3} in the aquaculture ponds
218 and 1.4×10^{-3} – 2.6×10^{-3} in the reservoir (Fig. 5). The EF value varied significantly
219 among the four water types (ANOVA, $p < 0.001$; Table 1) and the respective means
220 (\pm SEs) were $3.3 \times 10^{-3} \pm 1.4 \times 10^{-3}$ (river), $3.2 \times 10^{-3} \pm 1.1 \times 10^{-3}$ (drainage channels),
221 $4.5 \times 10^{-3} \pm 1.5 \times 10^{-3}$ (aquaculture ponds) and $1.8 \times 10^{-3} \pm 0.4 \times 10^{-3}$ (reservoir).

222 EF also varied significantly between seasons ($p < 0.001$; Table 1). The highest EF
223 was recorded in the summer months in all areas, in some cases more than two-fold
224 above the autumn and spring months (Fig. 5).

225 3.4. Relationships between EF and environmental variables

226 EF across all sampling sites was correlated linearly with T_w ($r^2 = 0.61$; $p < 0.01$; Fig.
227 6a). EF also correlated significantly with DOC but in an exponential manner ($r^2 = 0.93$;
228 $p < 0.01$; Fig. 6b). When plotted against DOC:NO₃⁻-N, a linear relationship is obtained
229 ($r^2 = 0.84$; $p < 0.01$; Fig. 6c). In contrast, EF was negatively but weakly correlated with
230 EC (Fig. 6d) and NO₃⁻-N concentration (Fig. 6e) ($r^2 = 0.26$ – 0.27 ; $p < 0.05$). We found

231 a significant and negative correlation between salinity and EF across the different
232 months ($p < 0.01$; Fig. 7a-c), and the combined data show an exponential decrease in EF
233 with increasing salinity ($r^2 = 0.41$; $p < 0.01$; Fig. 7d). No significant correlation was
234 found between EF and pH, DO, COD or TDN ($p > 0.05$).

235 **4. Discussion**

236 *4.1. N₂O distribution in coastal reservoir and its catchment*

237 Given its powerful warming effect, proper assessment and mitigation of N₂O is
238 critical for nations to meet their commitment to combat global warming (IPCC, 2018).
239 The “top-down” approach focuses on the outcome of increasing production and
240 emission of N₂O by measuring the changes in atmospheric N₂O level. However, in
241 order for nations to devise and implement policies to curb N₂O emission, it is necessary
242 to use “bottom-up” approach to identify and quantify the sources of N₂O (Del Grosso et
243 al., 2008). Notwithstanding its limitation, EF remains an indispensable tool for
244 bottom-up N₂O assessments because of its ability to encapsulate the complex aquatic
245 nitrogen biogeochemical processes into easily quantifiable parameters, such as NO₃⁻-N.
246 However, for urbanized coastal reservoirs that receive water and nutrient inputs from
247 multiple sources across a heterogeneous landscape, applying a fixed EF value may
248 translate to large errors in N₂O assessments. For improvement, we seek to refine and
249 constrain EF by considering the high-resolution distribution of N₂O and its
250 environmental driver(s).

251 We observed considerable variations in N₂O concentration in the four surface

252 water types. N₂O level was similarly low across all water types in the autumn month,
253 but it reached significantly higher levels in the summer month—more than double in
254 the river and aquaculture ponds (Fig. 3). Interestingly, both NO₃⁻-N and TDN
255 concentrations were lower in the summer (Fig. 2g). This suggests that the higher
256 summer-time N₂O was not caused by a higher amount of N substrates. Rather, the large
257 summer-time increase could be attributed to the higher water temperature (Fig. 2a)
258 increasing microbial transformation of N to N₂O (McMahon and Dennehy 1999; Tian
259 et al., 2017, 2018). Nevertheless, because the seasonal temperature change was nearly
260 identical across all sampling sites, temperature alone could not explain the spatial
261 differences in N₂O among the different water types. Also, the observation that N₂O
262 spatiotemporal variations were disconnected from NO₃⁻-N and TDN further highlights
263 the existence of moderating factors in N₂O production and the deficiency in using a
264 fixed EF value.

265 4.2. Environmental influence of EF

266 It is clear from the data that EF varied substantially in space and in time in the
267 reservoir. We observed considerably higher EF in the north basin than in the south basin
268 of the reservoir (Fig. 4), which may be attributed to the different hydrographical
269 properties between the two basins that influenced N₂O production. Salinity is higher in
270 SB due to its close proximity to the sea (Yang et al., 2020). Previous research has
271 shown that high salinity could adversely impact the abundance and activity of
272 ammonia-oxidizing bacteria (Li et al., 2021; Mosier and Francis, 2008; Neubauer et al.,

273 2019), thereby decreasing N₂O production and emission (Liu et al., 2015; Wang et al.,
274 2018; Welti et al., 2017). Similarly, we found a significant and negative correlation
275 between salinity and EF across the entire data set (Fig. 7), indicating that seawater
276 intrusion is an important factor moderating N₂O production. Alternatively, the
277 relationship could suggest low N₂O seawater mixing and diluting with higher N₂O
278 freshwater. Furthermore, the north basin is situated in an area with intensive human
279 activity and it receives sewage discharge and river inflows, leading to higher microbial
280 and substrate loading that could potentially increase N₂O production as compared to the
281 south basin. This is consistent with the observations that EF was higher in NB and in
282 sites impacted by wastewater (Fig. 4a and 4b).

283 The common formulations of EF are functions of nitrogen in the form of total N
284 input or NO₃⁻-N, based on the idea that increase in those should proportionally increase
285 N₂O. However, we found that EF correlated negatively (albeit weakly) with NO₃⁻-N
286 (Fig. 6e), as also reported for artificial surface waters (Webb et al., 2021), suggesting
287 that NO₃⁻-N alone is a poor predictor of N₂O in the system. The large spatial and
288 temporal variations in EF also mean that using a fixed EF would inevitably lead to
289 errors when assessing N₂O emission. For example, our empirically derived EF deviated
290 from the IPCC (2019) default EF value in all surface water types and all sampling
291 months (Fig. 5). On average, the IPCC default EF of 2.6×10^{-3} would overestimate N₂O
292 emission from the reservoir but underestimate N₂O emissions from the adjacent river,
293 drainage channels and aquaculture ponds, sometimes by as much as 42%.

294 4.3. Improving EF with hydrographical parameters

295 The application of emission factor (EF) offers important practical convenience in
296 greenhouse gas accounting at the national level. However, the biogeochemical
297 processes that lead to N₂O production and emission can be complex and variable at the
298 local level, which cannot be encapsulated realistically by a fixed EF value. Some
299 researchers have criticized the IPCC default EF for grossly overestimating N₂O
300 emissions (e.g. [Hu et al., 2017](#); [Kuang et al., 2021](#); [Maavara et al., 2019](#)). Likewise, our
301 results showed that the IPCC default EF tends to underestimate emission from waters
302 with a high DOC:NO₃⁻-N. With the DOC:NO₃⁻-N ratio trending upward in many
303 inland water bodies as a result of changing catchment processes ([Weyhenmeyer and](#)
304 [Jeppesen, 2010](#)), it is necessary to continuously revise and improve EF for more proper
305 N₂O emission assessments.

306 We examined the empirical relationships between EF and hydrographical
307 parameters in order to derive better algorithms for EF. Among the parameters we
308 investigated, EF correlated negatively with EC and NO₃⁻-N, but the corrections were
309 weak ($r^2 = 0.11-0.49$; [Fig. 6 d-e](#)), which we consider not suitable for practical use.
310 Water temperature is a most basic parameter in environmental monitoring and it affects
311 a multitude of biogeochemical processes; water temperature data are also widely
312 available in databases. Our study showed a significant linear relationship between EF
313 and surface water temperature (T_w) across a range (18-29 °C) typical for subtropical
314 environments ([Fig. 6a](#)). Scientists may therefore substitute the IPCC default EF with a

315 temperature-dependent EF to improve N₂O assessments, with the caveat that the
316 predictive model has only a moderate r^2 (0.54). This temperature-dependent algorithm
317 can be improved with more measurements including higher temperatures from the
318 tropical region and lower temperatures such as the winter month and higher latitudes.

319 While NO₃⁻-N is the precursor of N₂O, the biochemical transformation of N still
320 requires the work of microbes including denitrifying bacteria, which in turns rely on
321 dissolved organic carbon (DOC) as the energy source (Holmes et al., 1996; Stow et al.,
322 2005; Schipper et al., 2010). As such, one may derive an algorithm to incorporate DOC
323 as a surrogate indicator of microbial activity. Indeed, our observations showed that EF
324 correlated significantly ($p < 0.01$) with DOC as described by an exponential equation,
325 with a high r^2 of 0.84 (Fig. 6b). As DOC is a basic environmental monitoring parameter,
326 the new algorithm provides an easy way to adjust and improve EF with DOC data.
327 However, further consideration is needed when applying this algorithm elsewhere.
328 DOC concentration in inland waters varies widely and can reach up to 300 mg L⁻¹ in
329 hypereutrophic waters (Sobek et al., 2007), whereas our measurements cover a rather
330 narrow range of DOC. If we were to extrapolate the algorithm forward, it would give an
331 unrealistic EF at higher DOC concentrations. As such, despite the high r^2 (0.84),
332 application of this algorithm should be limited to coastal reservoirs with a similar DOC
333 concentration range.

334 A simple improvement to the algorithm is by adding NO₃⁻-N as an additional
335 constraining variable. In inland waters including ones impacted by anthropogenic

336 activities, DOC is often coupled to N input such that the DOC:NO₃⁻-N ratio tends to
337 vary within a relatively narrow range. In our study, we observed a significant linear
338 relationship between EF and DOC:NO₃⁻-N despite a slightly lower r^2 value (0.83) (Fig.
339 6c). The DOC:NO₃⁻-N (mg/mg) range in our data covers most of the DOC:NO₃⁻-N
340 values reported for inland waters (Weyhenmeyer and Jeppesen, 2010; Heppell et al.,
341 2017). As such, an algorithm based on DOC:NO₃⁻-N provides another useful tool to
342 derive EF for N₂O emission from coastal reservoirs and their catchments.

343 Lastly, considering that both T_W and DOC:NO₃⁻-N show good correlations with
344 EF, we derived a new algorithm with both independent variables in a multiple
345 regression analysis. The resultant equation is:

$$346 \quad EF = -0.001 + 0.00007 T_W + 0.00041 \text{ DOC:NO}_3^- \text{-N} \quad (\text{Eq. 1})$$

347 The two independent variables, when combined in Eq. 1, also give a significant and
348 high explanatory power ($F = 32.24$, $r^2 = 0.82$, $p < 0.001$). We therefore consider this
349 new algorithm most preferred and ecologically relevant because it incorporates the
350 effects of the three key environmental drivers by relating the variability of EF to the
351 influences of hydrographical condition (T_W) and N₂O precursor (NO₃⁻-N) as moderated
352 by microbial activity (with DOC as the surrogate indicator).

353

354 **5. Conclusions**

355 The trade-off between convenience and accuracy in N₂O assessment based on
356 IPCC-default EF has important ramifications as nations strive to improve greenhouse

357 gas mitigations and climate change policies. We found that EF varied temporally and
358 spatially in a complex and dynamic system such as a coastal reservoir that was
359 influenced by multiple input sources; therefore, applying the IPCC default EF could
360 lead to error in N₂O emission assessment, sometimes by as much as 42%. The observed
361 EF was significantly related to surface water temperature, dissolved organic carbon and
362 nitrate concentrations; therefore, an algorithm based on the three hydrographical
363 parameters combined could be used to derive a EF to improve N₂O emission from
364 urbanized coastal reservoirs. Further improvement will require more data to test its
365 application in other aquatic systems.

366 **Declaration of competing interest**

367 The authors declare that they have no known competing financial interests or
368 personal relationships that could have appeared to influence the work reported in this
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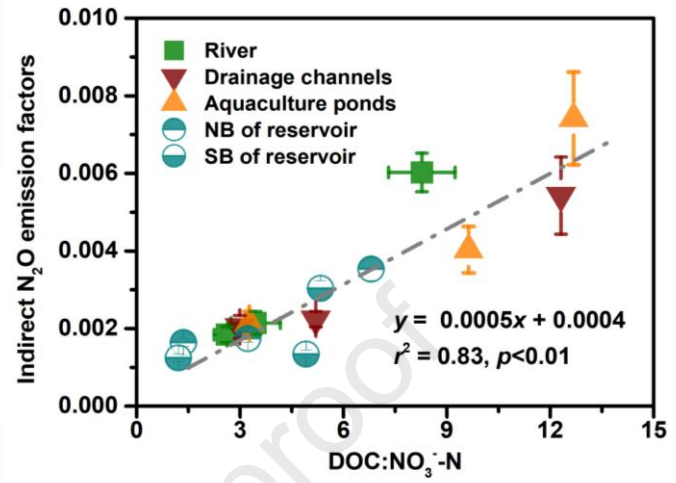
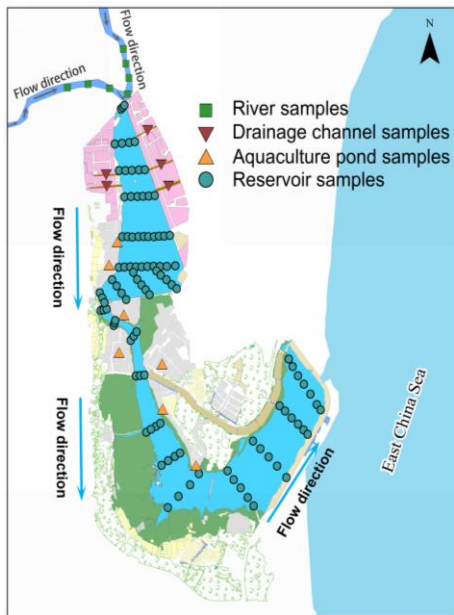
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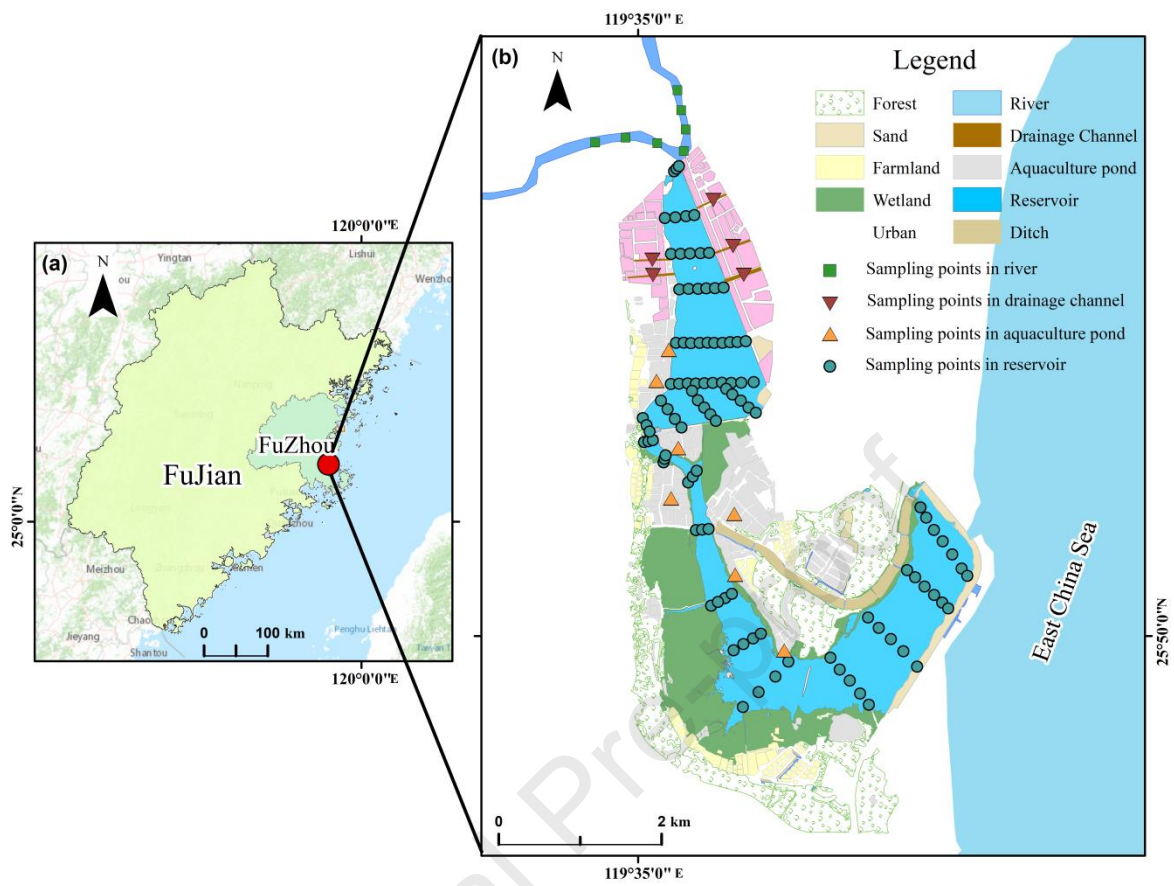
Graphical Abstract



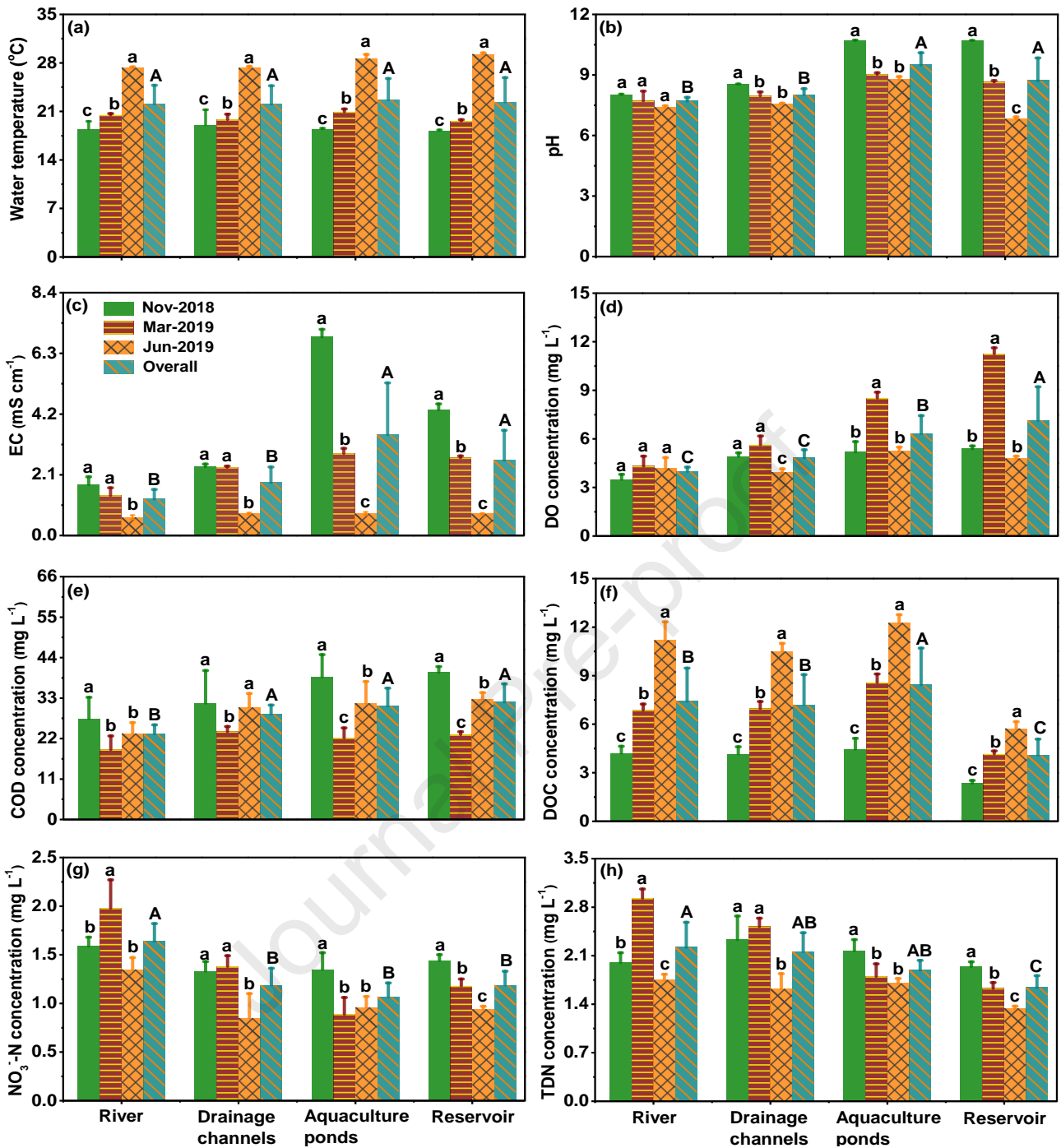
1 **Table 1** Results of two-way ANOVAs examining the effect of sampling area, season and their interactions on dissolved N₂O concentration and N₂O
 2 emission factor (EF).

	<i>df</i>	Dissolved N₂O concentration				N₂O emission factor (EF)			
		<i>Sum of squares</i>	<i>Mean square</i>	<i>F value</i>	<i>P value</i>	<i>Sum of squares</i>	<i>Mean square</i>	<i>F value</i>	<i>P value</i>
Sampling area	3	107882.38	35960.79	24.46	<0.01	1.88	0.63	36.30	<0.01
Season	2	61674.53	30837.27	20.98	<0.01	2.09	1.04	60.33	<0.01
Sampling area × Season	6	59233.93	9872.32	6.70	<0.01	0.96	0.16	9.26	<0.01
Residuals		515961.43	1469.98			6.07	0.02		

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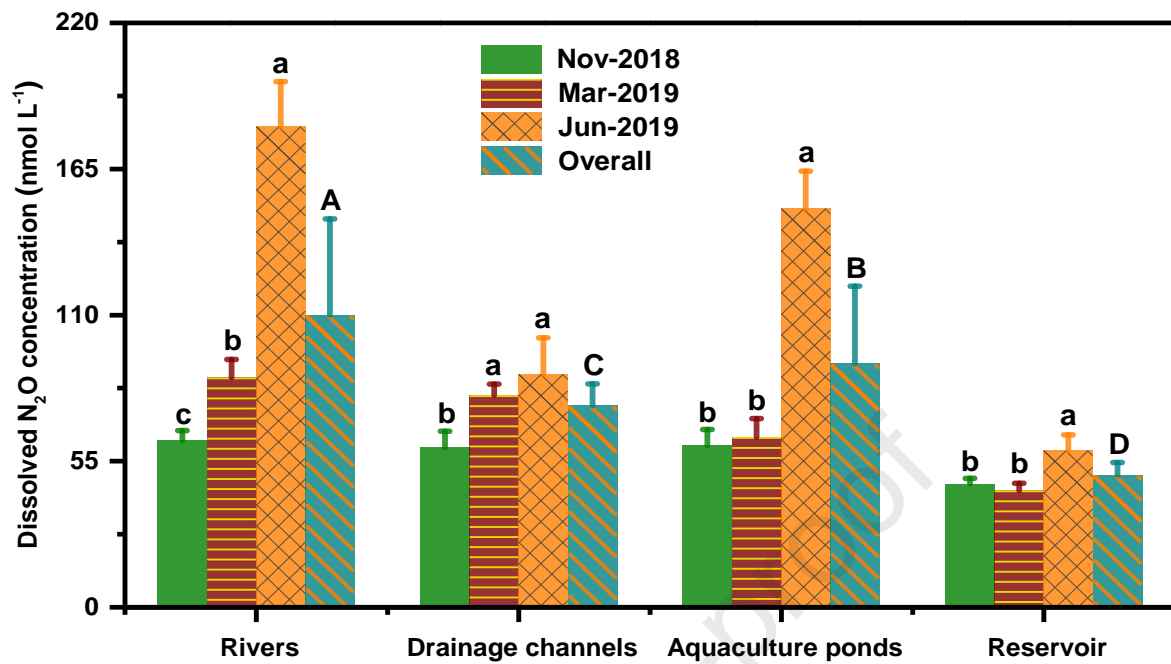


1
 2 **Fig. 1.** Location of (a) the study areas and (b) sampling sites in the Wenwusha Reservoir
 3 and the adjacent areas (river, drainage channels, aquaculture ponds) in Fujian Province,
 4 southeastern China.



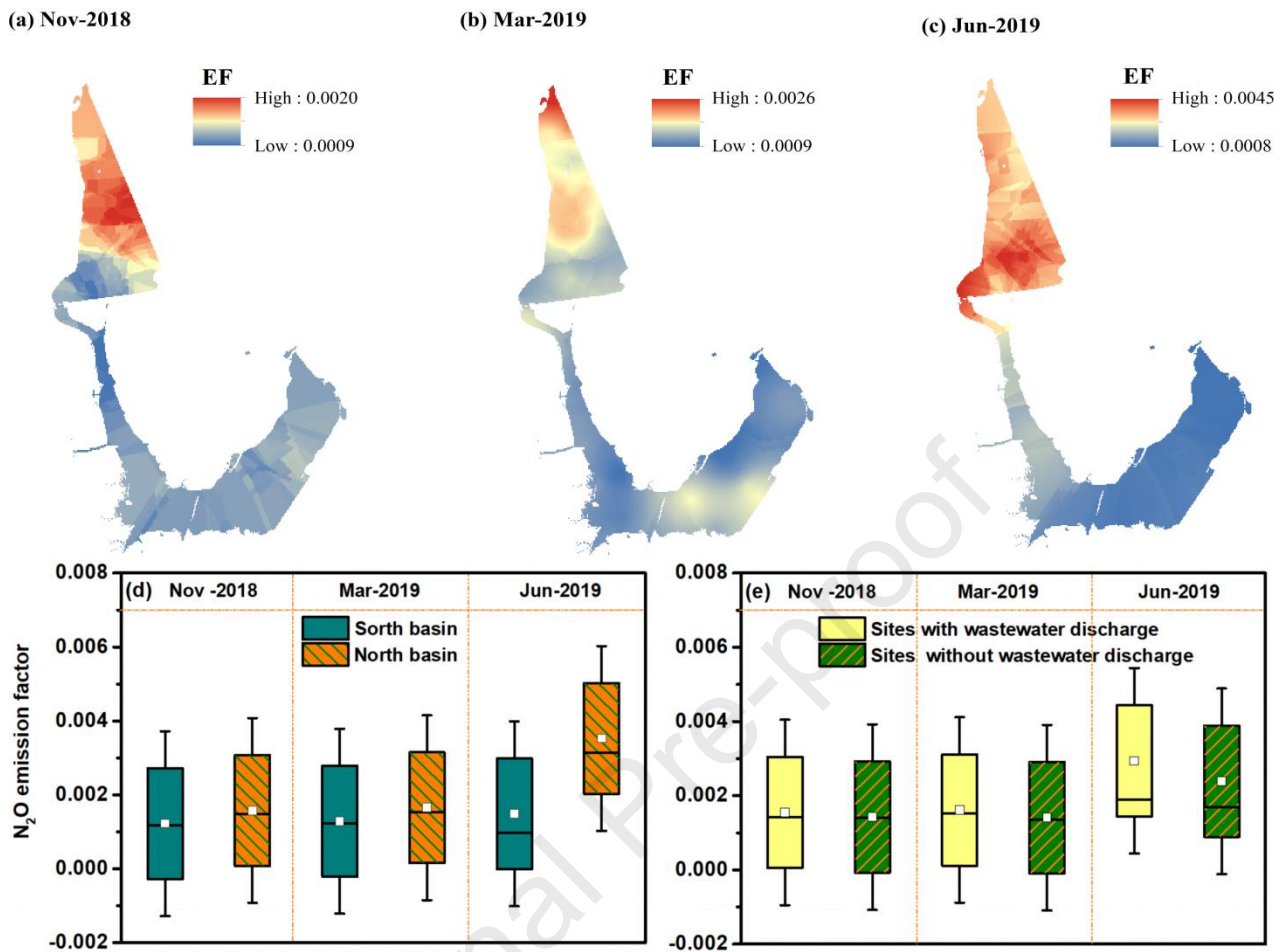
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6 **Fig. 2.** Variations in (a) water temperature (T_w), (b) pH, (c) conductivity (EC), (d) dissolved oxygen
 7 (DO), (e) chemical oxygen demand (COD), (f) dissolved organic carbon (DOC), (g) NO_3^- -N and (h)
 8 total dissolved nitrogen (TDN) in surface water (~ 0.2 m depth) among different sampling water
 9 bodies (river, drainage channels, aquaculture pond and reservoir) during each sampling campaign.
 10 Different lowercase letters above the bars indicate significant differences between months for
 11 specific sampling area ($p < 0.05$). Different uppercase letters above the bars indicate significant
 12 differences between sampling areas ($p < 0.05$). Differences in hydrographical properties between
 13 water bodies were tested by one-way analysis of variance (ANOVA). Bars represent mean \pm SE.



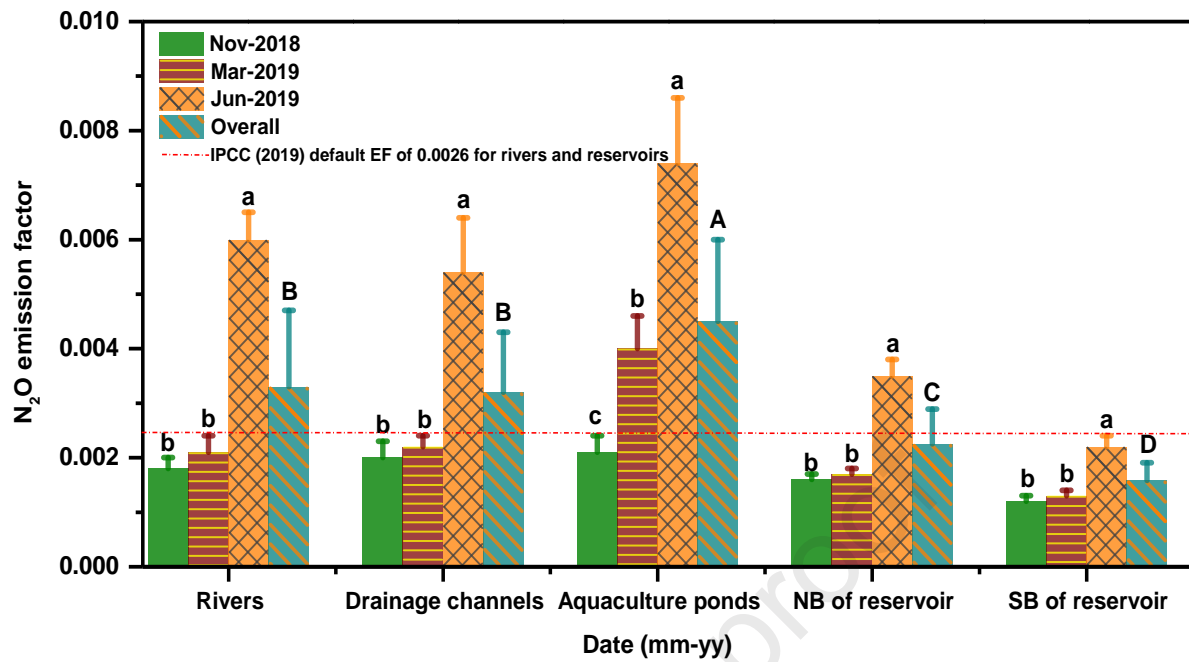
14

15 **Fig. 3.** Dissolved N₂O concentrations in surface water (~0.2 m depth) among different sampling
 16 water bodies (river, drainage channels, aquaculture ponds and reservoir) during each sampling
 17 campaign. Different lowercase letters above the bars indicate significant differences between
 18 months for specific sampling area ($p < 0.05$). Different uppercase letters above the bars indicate
 19 significant differences between sampling areas ($p < 0.05$). Bars represent mean \pm SE.

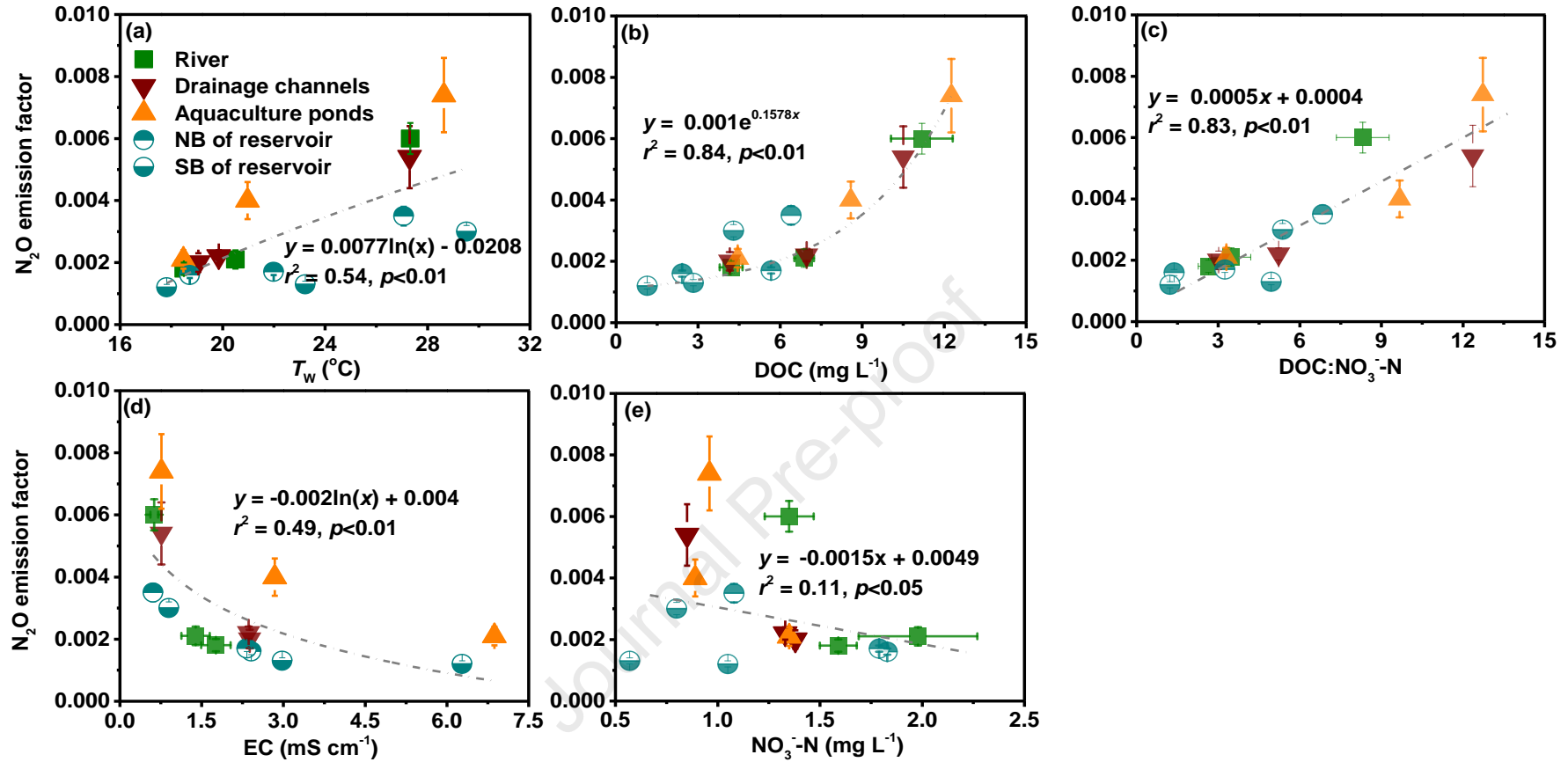


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21 **Fig. 4.** Spatial distributions of N₂O emission factor (EF) in the Wenwusha Reservoir during each
 22 sampling campaign (a-c), Boxplots of N₂O emission factors at the (d) different reservoir basins (south
 23 basin ($n = 47$) vs north basin ($n = 56$)), and (e) sampling sites with wastewater discharge ($n = 39$) or
 24 without wastewater discharge ($n = 64$). Each box shows the quartiles and median, while the square
 25 and whiskers represent the mean and values within 1.5 times of the interquartile range, respectively.



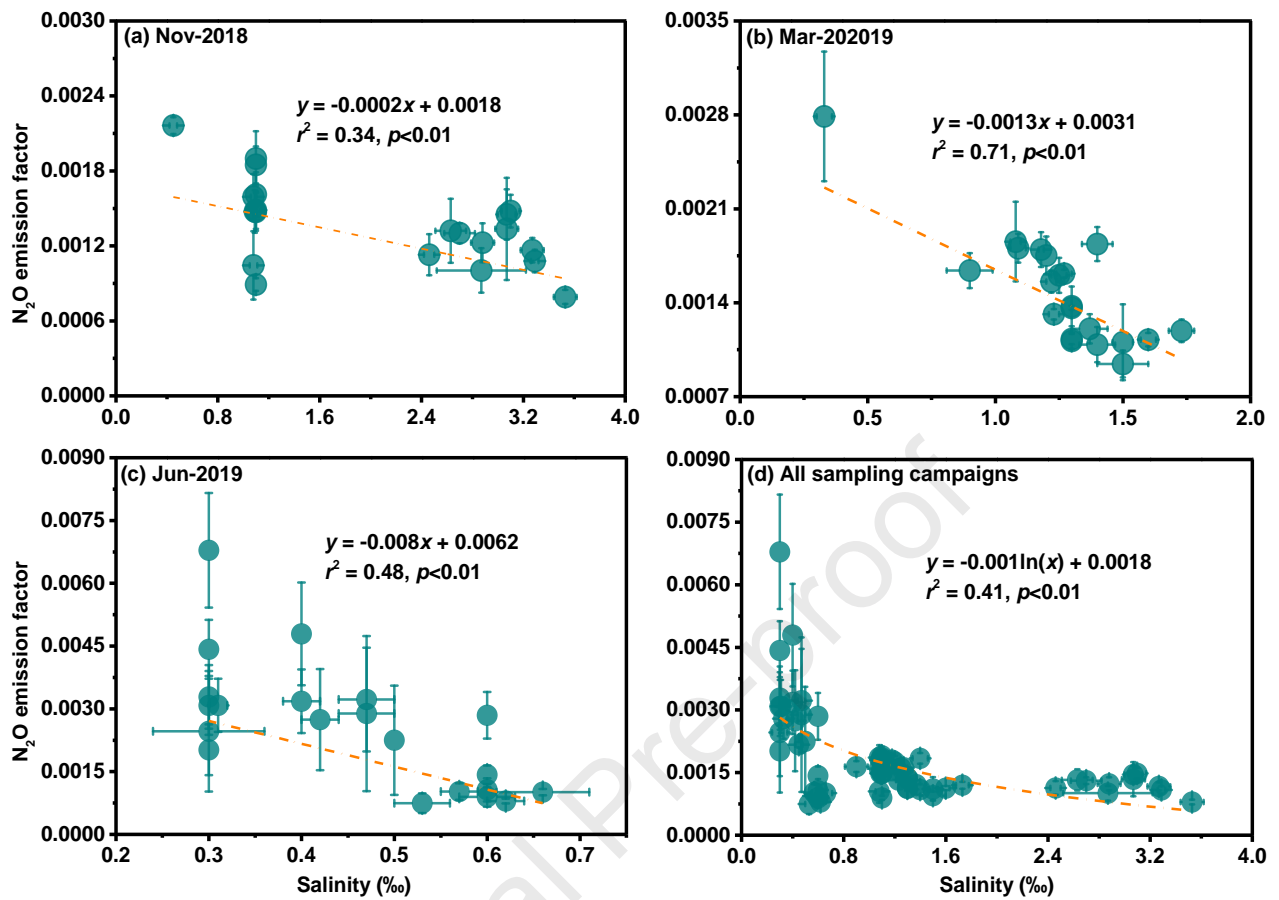
26
 27 **Fig. 5.** N₂O emission factor (EF) among different sampling water bodies (river, drainage channels,
 28 aquaculture ponds and reservoir) during each sampling campaign. Different lowercase letters
 29 above the bars indicate significant differences between months for specific sampling area
 30 ($p < 0.05$). Different uppercase letters above the bars indicate significant differences between
 31 sampling areas ($p < 0.05$). Bars represent mean \pm SE.



32

33 **Fig. 6.** Relationships between N_2O emission factor (EF) and (a) T_w , (b) DOC, (c) DOC: NO_3^- -N, (e) EC and (d) NO_3^- -N across all sampling

34 water bodies (reservoir, river, drainage channels and aquaculture ponds) during the study period.



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Fig. 7. Relationships between N_2O emission factor (EF) and surface-water salinity across 21 transects within the Wenwusha Reservoir during the Nov-2018 (a), Mar-2019 (b), Jun-2019 (c), and all sampling campaigns (d).

1 HIGHLIGHTS

- 2 ● Coastal reservoirs are a strong source of nitrous oxide (N_2O)
- 3 ● N_2O emission factor (EF) varied widely across different water types
- 4 ● Water temperature (T_w) affected seasonal variations in EF
- 5 ● Assessment of N_2O emissions using a fixed EF would introduce large errors
- 6 ● EF was best predicted by T_w , dissolved organic carbon and nitrate concentration

Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

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