Coastal reservoirs as a source of nitrous oxide: spatio-temporal patterns and assessment strategy

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ABSTRACT

Coastal reservoirs are widely regarded as a viable solution to the water scarcity problem faced by coastal cities with growing populations. As a result of the accumulation of anthropogenic wastes and the alteration of hydroecological processes, these reservoirs may also become the emission hotspots of nitrous oxide (N₂O). Hitherto, accurate global assessment of N₂O emission suffers from the scarcity and low spatio-temporal resolution of field data, especially from small coastal reservoirs with high spatial heterogeneity and multiple water sources. In this study, we measured the surface water N₂O concentrations and emissions at a high spatial resolution across three seasons in a subtropical coastal reservoir in southeastern China, which was hydrochemically highly heterogeneous because of the combined influence of river runoff, aquacultural discharge, industrial discharge and municipal sewage. Both N₂O concentration and emission exhibited strong spatio-temporal variations, which were correlated with nitrogen loading from the river and wastewater discharge. The mean N₂O concentration and mission were found to be significantly higher in the summer than in spring and autumn. The results of redundancy analysis showed that NH₄⁺-N explained the greatest variance in N₂O emission, which implied that nitrification was the main microbial pathway for N₂O production in spite of the potentially increasing importance of denitrification of NO3-N in the summer. The mean N2O emission across the whole reservoir was 107 μ g m⁻² h⁻¹, which was more than an order of magnitude higher than that from global lakes and reservoirs. Based on our results of

Monte Carlo simulations, a minimum of 15 sampling points per km^2 would be needed to produce representative and reliable N₂O estimates in such a spatially heterogeneous aquatic system. Overall, coastal reservoirs could play an increasingly important role in future climate change via their N₂O emission to the atmosphere as water demand and anthropogenic pressure continue to rise.

Keywords: Greenhouse gases; Water management; itrous oxide; Spatial heterogeneity; Wastewater discharge; Coastal reservoir

1. Introduction

Over 30 % of the world's population currently live in water-scarce regions (United Nations, 2018). In many of these regions, water scarcity is caused not by water shortage but rather insufficient water storage, leading to the construction of reservoirs as one of the common solutions to help divert and retain surface runoff (Sitharam et al., 2020). Coastal reservoirs are particularly appealing because of their low cost, low ecological impacts on upstream areas, and close proximity to the coastal populations with growing water demand (Liu et al., 2013). Coastal reservoirs are considered a key water management solution in densely populated coastal cities around the world by providing portable water for irrigation, preventing flooding downstream, and treating the discharged wastewater (Sitharam et al., 2020).

By retaining surface runoff and lengthening the water residence time, coastal reservoirs may accumulate pollutants and alter various hydroecological processes, leading to the release of greenhouse gases (e.g., nitrous oxide or N₂O) into the atmosphere. This is of particular concern for reservoirs in densely populated areas with high loadings of nutrients and wastes. In general, high nutrient loadings will increase the microbial-mediated process of N₂O production (Davidson et al., 2015; Wang et al., 2017; Xiao et al., 2019a), subsequently triggering more N₂O emission from reservoirs to the atmosphere (Wang et al., 2017; He et al., 2017). In addition, high nutrient loads in the reservoirs can drive large oxygen depletion, further promoting denitrification as the primary pathway of N₂O production (Borges et al., 2015). Dammed rivers and

reservoirs are thus potentially N₂O emission hotspots.

N₂O is a highly potent greenhouse gas with a global warming potential almost 300 times higher than that of carbon dioxide (CO₂) over a 100-year period (IPCC, 2013), and is also one of the major ozone-depleting substances (Ravishankara et al., 2009; Shaaban et al., 2018). The average atmospheric N_2O concentration reached 331 ppbv in 2018, which was approximately 23% above the pre-industrial level (World Meteorological Organization, 2019) with a growth rate of 0.7–0.8 ppb yr^{-1} over the last three decades (Davidson, 2009; Saikawa et al., 2014; Xiao, et al., 2019a). Despite the climatic importance of N₂O, accurate assessment of N₂O emissions are limited by the low availability of field data and the use of emission factor (EF) that is associated with the background N stock only. While it may be appropriate to apply EF for agriculture-based emission with easily quantifiable N input from fertilizers, applying specific EF in aquatic systems may yield large errors in the estimation of N₂O emission (Maavara et al., 2019), because of the highly variable N dynamics in both quantity and quality (Ivens et al., 2011). Furthermore, the assessment of N₂O emission is often done at low spatial resolutions (e.g., $1^{\circ} \times 1^{\circ}$ grids), which is too crude for small coastal reservoirs with high hydrographical heterogeneity and multiple water sources.

To address these limitations and explore the potential importance of coastal reservoirs as a net N_2O source, we measured in this study the surface water N_2O concentrations at a high spatial resolution in a subtropical coastal reservoir in southeast China, from which we calculated N_2O emissions using the wind-based thin boundary

layer model. We also examined the influence of environmental factors on the spatial variations of N₂O emissions and determined the optimal sampling strategy for assess N₂O emissions from spatially heterogeneous water bodies.

2. Materials and methods

2.1. Study area

The Wenwusha Reservoir $(25^{\circ}49'36''-25^{\circ}54'00'' \text{ N}, 119^{\circ}35'12''-119^{\circ}38'11'' \text{ E})$ is a coastal reservoir located at the mouth of the Nangyangdong River in Fujian Province in southeast China (Fig. 1). It has a surface area of 5.2 km², a mean depth of 9.0 m, a total volume of $3.20 \times 10^8 \text{ m}^3$ and a catchment area of 275 km² (Yang et al., 2020a). Influenced by a humid subtropical monsoonal climate, this reservoir has an annual mean temperature of 19.3 °C and an annual precipitation of 1390 mm, with 75% of the precipitation occurring during the wet season (May - September).

The reservoir is divided into two basins, namely the north basin (NB) and the south basin (SB) (Fig. 1) (Yang et al., 2020a). NB (1.9 km^2 ; $1.40 \times 10^8 \text{ m}^3$) is located in an urbanized area and is heavily impacted by aquacultural, industrial and municipal waste discharges while also receiving input from the Nanyangdong River. On the other hand, the watershed of SB (3.3 km^2 ; $1.69 \times 10^8 \text{ m}^3$) is dominated by natural wetlands and forests. This study covered a total of 10 transects in NB and 11 transects in SB, with 3-10 sampling sites selected along each transect (Fig. 1). A total of 103 sampling

sites (56 in NB and 47 in SB) were included, which could be categorized into five different sections according to the types of discharge: industrial effluents (Section-I; n = 4), town sewage (Section-T; n = 6), river inputs (Section-R; n = 7), aquacultural waste (Section-A; n = 22) and non-wastewater discharge (Section-N; n = 64).

2.2. Dissolved N₂O concentrations

Three sampling surveys covering the entire reservoi were conducted in November 2018, March 2019, and June 2019, respectively. In each survey, sample collection was completed within two consecutive days. We collected water samples at 20 cm depth using a 100-mL syringe equipped with three-way stopcock, and quickly transferred them into 55-mL glass serum bottles. We added 0.2 mL of saturated HgCl₂ into each bottle to inhibit microbial activities (Taipale and Sonninen 2009; Zhang et al., 2013), and then immediately sealed the bottle with an open-topped screw cap equipped with a halobutyl rubber septum to exclude air bubbles (Borges et al., 2018; Xiao et al., 2019a; Yang et al., 2020b). All the water samples were stored in an ice box and analyzed within 48 hours of collection.

Dissolved N_2O concentrations were measured following the headspace equilibration method (Yu et al., 2013; Yu et al., 2017). In the laboratory, N_2 gas (>99.9% purity) was injected into each sample bottle to displace a 25-mL headspace. The bottles were then shaken vigorously for 10 minutes in an oscillator to gas achieve an equilibrium in N₂O between the headspace air and the liquid phase. After waiting for 30 minutes, 5 mL of headspace air sample was collected with a syringe and analyzed for N₂O concentration using a gas chromatograph (GC-2014, Shimadzu, Kyoto, Japan) equipped with an electron capture detector. The detection limit was 0.02 ppm and the relative standard deviations of the measurements were $\leq 5.0\%$ (Yang et al., 2020b). Dissolved N₂O concentration was then calculated based on the temperature, Henry's law constant (salinity-dependent) and the measured head space N₂O concentration according to Weiss and Price (1980).

2.3. N₂O emission

N₂O emission (F_{W-A} , µmol m⁻² h⁻¹ was calculated using the thin boundary layer model (Equation 1) that has been widely applied in the lentic ecosystem (e.g., Cole and Caraco, 1998; Musenze et al., 2014 Xiao et al., 2019b):

$$F_{W-A} = k \times (C_{obs} - C_{eq}) \tag{Eq1}$$

where k is the gas transfer velocity (m h⁻¹); C_{obs} is the measured N₂O concentration (nmol L⁻¹) in the surface water; C_{eq} is the dissolved N₂O concentration (nmol L⁻¹) in equilibrium with the atmospheric concentration at the prevailing *in situ* conditions, as a function of atmospheric pressure, water temperature, and the ambient atmospheric N₂O.

Because of negligible surface water flow at the reservoir, the N₂O transfer velocity is primarily controlled by wind speed (Xiao et al., 2019b). In this study, we calculated kvalue using a wind-dependent equation (Eq 2) for small shallow lakes (Cole and Caraco, 1998):

$$k = (Sc/660)^{-n}(2.07 + 0.215 \times U1^{1}0^{7})$$
(Eq2)

where *Sc* is the temperature-dependent Schmidt number for N₂O; *n* is a constant that varies between 0.66 (for wind speeds $\leq 3 \text{ m s}^{-1}$) and 0.50 (wind speeds $\geq 3 \text{ m s}^{-1}$); U_{10} is the frictionless wind speed (m s⁻¹) at 10 m height. *Sc* and U_{10} were calculated using the following equations (Crusius and Wanninkhof, 2003; Wanninkhof, 1992):

$$Sc = 2055.6 - 137.11T + 4.3173T^2 - 0.054350T^3$$
 (Eq 3)

$$U_{10} = U_z \left[1 + \frac{(C_{d10})^{1/2}}{K} \ln(\frac{10}{z})\right]$$
(Eq 4)

where U_z is the wind speed (m s⁻¹) at a ght of z above the water surface (2.0 m in this study); C_{d10} is the drag coefficient at 10 m above the water surface (0.0013 m s⁻¹); and K is the von Karman constant (0.41).

2.4. Environmental variables

Water samples were collected at a depth of 20 cm with a 5 L organic glass hydrophore and transferred into 150 mL polyethylene bottles. About 0.5 mL of saturated HgCl₂ solution was added to inhibit microbial activities (Zhang et al., 2013). The water samples were transported in an ice box back to the laboratory for analysis. In the laboratory, the water sample was filtered through a 0.45-µm cellulose acetate filter (BiotransTM) and then analyzed for the concentrations of total dissolved nitrogen (TDN), ammonium-nitrogen (NH₄⁺-N), and nitrate-nitrogen (NO₃⁻-N) using a flow injection analyzer (Skalar Analytical SAN⁺⁺, The Netherlands) (Yang et al., 2020b). The detection limit and relative standard deviations were 3.0 μ g L⁻¹ and \leq 2.0% for TDN, 0.6 μ g L⁻¹ and \leq 3.0% for NH₄⁺-N and 0.6 μ g L⁻¹ and \leq 2.0% for NO₃⁻-N.

Concurrent measurements of water temperature (T_W) and pH (IQ150 pH/mV/Temperature meter, IQ Scientific Instruments, USA), salinity (Eutech Instruments-Salt6 salinity meter, USA), dissolved oxygen (DO; 550A YSI sonde, USA) and electrical conductivity (EC; 2265FS electrical conductivity meter, Spectrum Technologies, USA) were also made at a water depth of 20 cm (Yang et al., 2020a). The relative standard deviations of the measurements were between $\leq 1.0\%$ and $\leq 2.0\%$. Air temperature (T_A), wind speed at 2 m height (W_S) and atmospheric pressure (P_{atm}) were measured by a portable weather meter (Kestrel-3500, USA).

2.5. Statistical analysis

Results were presented as mean ± 1 standard error. Significant differences in hydrographical properties, dissolved N₂O concentration and N₂O emission among reservoir basins and discharge sections were tested by analysis of variance (ANOVA) using the statistical software SPSS 17.0 (SPSS Inc., USA). Correlations between hydrographical parameters and N₂O concentration or emissions were investigated by Pearson correlation analysis. The extent to which environmental variables affected the spatiotemporal variations in N₂O emissions was tested by Redundancy Analysis (RDA) using the software CANOCO 5.0 (Microcomputer Power, Ithaca, USA), with *T*w, pH, DO, salinity, NO₃⁻-N and NH₄⁺-N, and TDN being inputted as environmental variables.

Monte Carlo analysis was used to determine the optimal sample size for assessing the mean N_2O emission at the whole-reservoir scale. The measured fluxes were resampled from the 103 sites with varying sample sizes (n = 10, 20, 30..., 100) without replacement for 10,000 times. The overall mean and standard deviation of N_2O emission were subsequently calculated for each sample size. Plots were created using OriginPro 7.5 (OriginLab Corp. USA). To illustrate the spatial variations in hydrographical parameters and N₂O concentration (or flux) in the reservoir, the Kriging method in ArcGIS 10.2 (ESRI Inc., Redlands, CA, USA) was employed for spatial interpolation.

3. Results

3.1. Hydrographical parameters

There were large differences in hydrographical parameters between NB and SB across the three campaigns. NB had lower DO, EC, and salinity (Fig. S1) and higher NO_3^- -N, NH_4^+ -N and TDN concentrations as compared to SB (Fig. 2). T_W and pH were comparable between the two basins (Fig. S2). Water chemistry also varied spatially depending on the discharge types. In Section-N without sewage discharge, the NO_3^- -N, NH_4^+ -N, and TDN concentrations were significantly lower (p<0.05) while DO and salinity were generally higher than in other sections.

The hydrographical parameters also showed strong temporal variation across the whole reservoir (p<0.001). Average water temperature was 23.6 °C during the study period, while the seasonal pattern was clear decreasing in the order: summer (29.3 °C) > spring (23.1 °C) > autumn (18.3 °C) (Fig. S2). The seasonal mean concentrations of NO₃⁻-N, NH₄⁺-N and TDN ranged from 0.95 to 1.47, 0.28 to 0.57, and 1.43 to 1.82 mg L⁻¹, respectively. The seasonal mean DO concentration, pH and salinity ranged from 4.60 to 8.87 mg L⁻¹, 6.88 to 10.70, and 0.41 to 1.49‰, respectively. We observed generally higher NO₃⁻-N, TDN, pH and salinity in autumn (Fig. 2 and Fig. S1), but higher DO and lower NH₄⁺-N in spring (Fig. 2 and Fig. S1). The seasonal mean water depth ranged from 6.0 to 8.6 m, with a maximum value occurring in summer.

3.2. Spatial variation in N₂O concentration

Surface water N₂O concentrati in the reservoir showed high spatial heterogeneity, varying by 9 to >30 folds between the two basins (Fig. 3a-3c). The mean N₂O concentrations were significantly higher in NB than in SB (Table 1), while the overall mean N₂O concentrations in November, March and June were 46.4 ± 2.0 , 58.9 ± 5.9 , and 43.9 ± 2.7 nmol L⁻¹, respectively.

N₂O concentration also varied spatially depending on the discharge types, with significantly lower values in Section-N (non-wastewater discharge) and generally higher values in Section-T (town sewage) (Table 1). Overall, N₂O concentration showed significant differences between the sections with and without wastewater

discharge (Fig. 4a-4c).

3.3. Spatial variation in N₂O emission

From November to June, N₂O emission varied spatially by from >100 to >300 folds between the two basins (Fig. 3d-3f). N₂O emissions were generally high in NB, ranging from 6.9 to 711.5 μ g m⁻² h⁻¹ with a mean of 153.1±8.6 μ g m⁻² h⁻¹. In contrast, N₂O emissions were significantly lower in SB (Table 1), with occasional occurrence of negative values (i.e. net uptake of N₂O). The reservoir-wide mean N₂O emissions were 110.4±9.3, 78.5±6.3, and 153.1±8.6 μ g m⁻² h⁻¹ in November, March, and June, respectively.

There were clear spatial differences in N_2O emission as a function of discharge types, with lower emissions in Section-N without wastewater discharge but higher emissions in Section-I (industrial effluent) and Section-R (river runoff) across the three seasons (Table 1). Overall, there were significant differences in mean N_2O emission between the sections with and without wastewater discharge (Fig. 4d-4f).

3.4. Effects of environmental variables on N₂O concentration and emission

Results of Pearson correlation analysis showed that the reservoir-wide N_2O concentration and emission were correlated negatively with DO, EC, and salinity (Table S2) and correlated positively with NO_3^--N , NH_4^+-N and TDN concentrations across all surveys (Table S2, Fig. S3 and S4). Over the whole study period, the N_2O concentration and flux in the reservoir were positively correlated with air temperature but negatively

correlated with wind speed (p<0.05 or < 0.01; Table S2). Atmospheric pressure had no significant correlation with either N₂O concentration or flux (p>0.05; Table S2).

Based on the result of RDA analysis, EC, TDN and NO_3^-N were the environmental variables that explained the most spatial variations in N₂O concentration and emission in November, March, and June, respectively (Fig. 5). Combining all the data together, NH_4^+-N had the largest explanatory power (50.4%) for the spatio-temporal variations in N₂O concentration and emission, followed by NO_3^--N (23.5%) and salinity (20.2%) (Fig. 5).

3.5. Determining the optimal sample size

Based on the results of Monte rlo analysis, the uncertainty of mean N_2O emission estimates across the reservoir decreased markedly with increasing number of sampling sites. For each sample size, we created the boxplots of N_2O resampling means after 10,000 times of resampling (Fig. 6a). The Highest Probability Density ranges converged with increasing sample size, indicating improved accuracy of N_2O emission. In addition, the bias of the resampling N_2O estimations declined and the standard deviations of resampling means decreased exponentially with sample size (Fig. 6b). Notably, an increase in sample size from 10 to 100 decreased the relative standard deviation of mean N_2O emission by nearly 20-fold, from 21% to 1.2%.

The results of Monte Carlo analysis could also be used to determine the optimal sample size for assessing the reservoir N_2O emission. To achieve an error of under 5%,

at least 80 samples over the whole reservoir (5.2 km²) or ca. 15 samples per km² would be needed to resolve the high spatial heterogeneity of N₂O emissions. With a sample size of only 10, we obtained a high relative standard deviation of estimated average N₂O emission of 22.6% after 10,000 times of resampling. Despite this large deviation, the median of the resampled means was 107 μ mol m⁻² h⁻¹ (Fig. 6a), which was close to the overall mean of all 103 samples combined. Therefore, the high spatial resolution measurements across multiple transects effectively reduced the estimation bias of N₂O emission in our study.

4. Discussion

4.1. N₂O production in reservoir systems

 N_2O production is driven by microbial transformation of N substrates (Herrman et al., 2008; Liu et al., 2011; Phanwilai et al., 2020; Xiao et al., 2019b). Previous studies have reported significant relationships between sewage input, N substrates and N_2O emission in lakes (e.g., Xiao et al., 2019a), coastal river network (Yu et al., 2013) and inland reservoirs (e.g., He et al., 2017; Liu et al., 2011, 2017). In this study, while several environmental variables such as air temperature, wind speed, atmospheric pressure and solar radiation were quite homogeneous across the reservoir, the availability of N substrates as reflected by the concentrations of NH_4^+ -N, NO_3^- -N and TDN showed high spatial variations (Fig. 2). Both the concentrations of N substrate and

 N_2O emission rates were much higher in the north basin than in the south basin (Table 1). Located in an urbanized region, our reservoir was heavily impacted by anthropogenic activities, particularly the discharge of domestic and industrial wastewaters and the nutrient-rich effluents from aquaculture ponds. The highest average NO_3^- -N and NH_4^+ -N concentrations were found in the section with river input, indicating the pronounced influence of the upstream area as a N source. In contrast, the N-section without wastewater discharge had better water quality as reflected by the lower NO_3^- -N and NH_4^+ -N concentrations and higher DO level (Table S1). Taken together, our results showed that the high spatial he erogeneity of N_2O emission was closely linked to the input of localized N subst ate into the reservoir.

 N_2O can be produced via denitrification of NO_3^- under oxygen-poor conditions, or nitrification of NH_4^+ -N under aer bic conditions. The DO levels obtained in this reservoir were generally high enough (ca. $\geq 25\%$ sat.) to favor nitrification process (Fig. S1), which was further supported by the results of RDA analysis in which NH_4^+ -N explained the largest overall variance in N₂O emission (ca. 50%) during the study period (Fig. 5d). However, it should be noted that DO was only measured in the surface water in this study, but could drop to lower levels in deeper water, especially during the summer, where denitrification might dominate. This hypothesis was supported by the results of RDA analysis in which NO_3^- -N explained >73% of the variance in N₂O emission in June (Fig. 5c) when the average DO level was slightly lower (Table S1) and was the second most important factor in accounting for the overall variance (ca. 24%) (Fig 5d). Interestingly, salinity also explained about 20% of the variance in N_2O emission (Fig. 5d). Because of the proximity of the south basin to the sea, salt intrusion and deposition has created spatial differences in salinity across the reservoir, which could adversely impact the abundance and activity of nitrifiers that are of freshwater in origin (Francis et al., 2003; Monteiro et al., 2017) and subsequently decrease N_2O production (Sun et al., 2013; Wang et al., 2018; Welti et al., 2017).

4.2. Spatial and temporal variations in N₂O emission

Large spatial variations in N₂O emission have been observed in freshwater lakes and reservoirs (Cheng et al., 2019; Musenze et al., 2014; Zhao et al., 2013). Based on field measurements with a high spatia resolution, we successfully characterized the spatial heterogeneity of N₂O emission from the Wenwusha Reservoir, which exhibited coefficients of variation ranging from 81% to 110% across the three seasons. Land use change in the catchment can disturb various biogeochemical processes on land and in adjacent waters (Hosen et al., 2014; Williams et al., 2016), with the consequence of changing the direction and magnitude of N₂O fluxes among different areas (He et al., 2017; Wang et al., 2017; Zhou et al., 2017). In our study, the concentrations of nitrogen substrate (NO₃⁻-N, NH₄⁺-N, and TDN) were substantially lower in sites without wastewater discharge than in those in close proximity with aquaculture ponds, municipal and agricultural lands with wastewater discharge (Table S1). The N₂O concentration and flux increased with nitrogen concentration across the whole reservoir. Most notably, the sites under the influence of wastewater discharge covered only 30% of the reservoir area but contributed to approximately 60% of the total N_2O emission from the whole reservoir. Our results demonstrated that the spatial variation in N_2O flux in the subtropical Wenwusha reservoirs was largely affected by various anthropogenic activities (e.g., urbanization and land use) in the catchment area.

N₂O emission also exhibited clear seasonal variation with significantly higher values in the summer (Table 1), similar to the observations made in other inland waters (Musenze et al., 2014; Xiao et al., 2019b; Zhu et al., 2013). Temperature is considered a dominant factor driving the temporal variation in N₂O emission by regulating microbial activity (Beaulieu et al. 2010; Harrison and Matson, 2003; Hinshaw and Dahlgren 2013). The higher N_2O emission observed in the summer coincided with the higher water temperature (Fig. S2c), which could have stimulated N₂O production. Interestingly, we observed that the verage N_2O emission in spring was substantially lower than that in autumn (Fig. 3 and Table 1), which were opposite to the trend with water temperature (Fig. S2a and S2b). There were several heavy rain events that occurred in spring. Previous studies have shown that heavy precipitation could transport a greater amount of nutrients and greenhouse gases from the surrounding watershed into the lentic ecosystems (Dinsmore et al., 2013; Sinha et al., 2017; Xiao et al., 2021), thereby increasing the saturation level of dissolved greenhouse gases and stimulating microbial greenhouse gas production (Stanley et al., 2016; Yu et al., 2017). However, we found no significant increase in dissolved N₂O concentrations and fluxes in association with the heavy spring precipitation, which was probably a result of the dilution effects of precipitation on the biogeochemical reactions in the reservoir (Outram and Hiscock, 2012; He et al. 2017).

Furthermore, the drainage of aquaculture ponds after harvest in autumn as a common management practice could introduce N-rich effluent into the reservoir and stimulate N₂O production. This was supported by the elevated concentrations of NO_3^--N , NH_4^+-N , and TDN found near the aquaculture ponds in November (Fig. 2) and the strong positive correlations detected between N₂O emission and the concentration of N substrates (Fig. S4).

4.3. Subtropical coastal reservoir as a net N₂O source

The average N₂O fluxes across the water-air interface from the Wenwusha Reservoir was 107.0±15.6 μ g m⁻² h⁻¹, which was about 3.9 and 1.9 times higher than that observed in the marsh (Wang et al., 2018) and aquaculture ponds (Yang et al., 2020b), respectively, in the same subtropical coastal region. This magnitude of N₂O emission was also substantially higher than the average rate reported in the lakes (12.2 μ g m⁻² h⁻¹) and reservoirs (42.3 μ g m⁻² h⁻¹) of China (Li et al., 2018). When compared with the N₂O emission rates globally (Table S3), the mean emission from our reservoir was lower than only a handful of tropical or subtropical reservoirs, but was over an order of magnitude greater than the global average of 7.2 μ g m⁻² h⁻¹ for lakes and reservoirs spanning across the tropical to polar regions (Hu et al., 2016). Our results, along with similar findings by others (Chen et al., 2014), suggested that subtropical

coastal reservoirs were strong N₂O sources and could play an increasingly important climatic role as countries seek to construct more coastal reservoirs (Yang and Kelly, 2015).

4.4. Implications for N₂O assessment and reservoir management

Over the past two decades, researchers have increasingly recognized the role of reservoirs as a potential source of greenhouse gases including N₂O (Deemer et al., 2016; Descloux et al., 2017; Maavara et al., 2019). However, reliable assessment of N₂O emission from global reservoirs remains a challenge due to the scarcity and low resolution of field data. Results from this study highlighted the high spatial and temporal heterogeneities of N₂O concentration and emission in a subtropical coastal reservoir, with significant implications to representative sampling for flux estimation. Our results of Monte Carlo analysis suggested that reliance of field data from only one or a few sampling points, as quite commonly done in the literature (Table S3), could yield large errors. Urbanized, coastal reservoirs such as the one studied here are often influenced by runoff and discharge from multiple sources. Our Monte Carlo simulation results (section 3.5) suggested that a minimum of 15 sampling points per km² would be needed to produce a reliable and representative estimate of spatially-averaged N₂O

Our data showed a strong correlation between N_2O emission and wastewater discharges, affirming the important role of wastewater input in promoting N_2O

emission. This problem of greenhouse gas release can be exacerbated by the race to modernization and urbanization, particularly in developing countries where wastewater treatment remains inadequate (Hosen et al., 2014; Williams et al., 2016; Yang et al., 2013). Strengthening environmental regulation, particularly for wastewater treatment and discharge, will be key to mitigate water pollution and greenhouse gas emission in reservoir management (Yang, 2014; Yang et al., 2015). To the best of our knowledge, this study was the first attempt to characterize N₂O concentration and emission from a coastal reservoir at a very high spatial resolution. Our data would be valuable for global biogeochemical modelling and prediction of N₂O emission from costal reservoirs that are under intense human disturbance.

Several aspects of our study could be improved further in future studies. Firstly, several studies in inland waters have reported diurnal variations with higher N₂O emissions at night (Baulch et al., 2012; Wu et al., 2018; Yang et al., 2011). Our measurements in the Wenwusha Reservoir were limited to daytime and therefore the emission might have been underestimated. Diurnal differences in N₂O emission from the reservoir should be compared to improve the reliability of flux estimates. Secondly, in a thermally stratified water column, nitrogen transformation can differ between the epilimnion and the hypolimnion (Beaulieu et al., 2015; Liang et al., 2019; Salk et al., 2016), which was not included in our study. Future studies might use novel methods such as isotopic tracing to examine the relative importance of different nitrogen biogeochemical processes throughout the water column (Sebilo et al., 2006). While we

aimed to resolve the spatial heterogeneity with high spatial resolution sampling, our data also showed considerable seasonal variations in N_2O concentration and emission (Fig. 4). However, resolving temporal variation at a higher frequency (monthly to daily) could be challenging logistically without the help of *in situ* automated sampling systems. Lastly, the microbial abundance and functional groups (e.g. nitrifier, denitrifier, and ammonifier) should be further characterized to shed light on the dominant microbial pathways of N_2O production in the reservoir.

5. Conclusions

We characterized the variation of N₂O concentration and emission in a coastal, urbanized reservoir in southeastern China with a high spatial resolution in this study. We found that N₂O concentration and emission were strongly influenced by localized runoff and discharges, resulting in very high spatial heterogeneity in addition to the seasonal variations. Results of Monte Carlo simulations showed that a minimum of 15 sampling points per km² would be required to produce a reliable estimate of the whole-reservoir N₂O emission. Based on the analysis of water chemistry data, nitrification appeared to be the main pathway for N₂O production, although denitrification might become increasingly important in the summer months with reducing DO level. Total N₂O emission from our subtropical coastal reservoir was estimated to be 4.8 Mg yr⁻¹, which was considerably higher than the global average for lakes and reservoirs. Continued construction of coastal reservoirs for water management may exacerbate future climate change by increasing N₂O emission.

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Table 1. Summary of ANOVAs of the effects of sampling location, sampling time and their interactions on surface water N₂O concentration and

2 emission in the Wenwusha Reservoir.

	Month	Reservoir basi	'n*	Discharge type	**			
		SB	NB	Section-I	Section-T	Section-R	Section-A	Section-N
N ₂ O concentration	Nov 2018	30.1±2.5	60.1±1.8	62.9±2.1 a	67.6±4.5 a	62.7±3.8 a	47.5±3.1 b	41.25±2.7 b
(nmol L ⁻¹)	May 2019	19.7±1.9	64.3±2.3	79.9±4.1ab	59.61±9.2 b	86.5±6.8 a	43.9±3.6 bc	35.62±3.2c
	Jun 2019	29.1±5.2	84.0 ± 8.6	85.5±15.4 a	43.97±11.9 b	123.7±37.2 a	63.1±13.2 b	50.21±6.7 c
	Average	26.3	69.5	76.1	57.1	91.0	51.5	42.4
N2O emission	Nov 2018	53.5±6.5	158.2±13.2	326.8±36.6 a	181.9±36.9 b	213.3±55.4 b	96.3±15.4c	83.77±8.2c
$(\mu g m^{-2} h^{-1})$	May 2019	30.3±5.8	118.9 ± 6.7	116.9±7.5 ab	96.9±20.7 bc	165.4±16.3 a	90.9±11.0 b	60.57±7.8 c
	Jun 2019	72.5±15.0	182.3±20.4	201.3±29.7 ab	105.2±43.6bc	269.1±104.9 a	127.1±27.4 b	117.2±16.00
	Average	52.1	153.1	215	128	215.9	104.8	87.2

letter are not significantly different from each other ($p \ge 0.05$) within the month. Average values are calculated across the three surveys for each column.

b b

 ω 4



Fig. 1. Sampling sites within the Wenwusha Reservoir in Fujian Province, southeastern China. A total of 1 21 transects were sampled, with 11 transects in the south basin (SB) (1–11), and 10 in the north basin 1 (NB) (12–21). Each circle represents a sampling site influenced by different discharge type: A: 2 aquacultural effluent (n = 4); I: industrial effluent (n = 6); R: river input (n = 7); T: municipal sewage (n = 22); N: non-discharge (n = 64).





8 Fig. 2. Spatial distributions of surface-water NO₃⁻-N (*a-c*), NH₄⁺-N (*d-f*) and TDN (*g-i*) in the

9 Wenwusha Reservoir.



10

11 Fig. 3. Spatial distributions of surface-water N_2O concentration (*a-c*) and fluxes (*d-f*) in the Wenwusha

12 Reservoir.









26 Fig. 6. (a) The box and whisker plots and (b) standard deviation of estimated N₂O fluxes from the Wenwusha 27 Reservoir. Monte Carlo analysis was used to evaluate the effect of sample size on the emission estimate. 28 Without replacements, the N₂O emission measurements were resampled from all sites (n = 10, 20, 30..., 29 100). The resampling process was repeated 10,000 times. For each sample size, the overall mean and 30 standard deviation of N₂O emission were calculated. The N₂O emission are a function of the number of 31 sampling sites selected from a total of 103 sites based on 1000 simulations. In panel a, the margins of the 32 boxes are the upper and lower quartiles, and the line in the box represents the median. The whiskers are the 33 extreme values that fall within 1.5 times the interquartile range and the circles represent extreme values that 34 fall outside 1.5 times the interquartile range. The red line are the overall average N₂O flux based on the 35 results from all 103 sites. The grey areas represent 70, 95 and 99% of the Highest Probability Density ranges.

1 Supporting Information

2 Coastal reservoirs as a source of the greenhouse gas nitrous oxide: spatial-

3 temporal patterns and assessment strategy

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S1

23	3 Supporting Information Summary
24	4 No. of pages: 13 No. of figures: 4 No. of tables: 4
25 26	Page S2: Fig. S1. Spatial distributions of surface-water DO $(a-c)$, electrical conductivity (EC) $(d-f)$, and salinity $(g-i)$ in the Wenwusha Reservoir.
27	 Page S3: Fig. S2. Spatial distributions of surface-water W_T (a-c), and pH (d-f) in the Wenwusha Reservoir.
30	 Page S4: Fig. S3. Relationships between surface-water N₂O concentration and NO₃-N, NH4⁺-N and TDN concentrations in the Wenwusha Reservoir. Parameter bounds on the
31	1 regression coefficients are at 95% confidence limits.
32	2 Page S5: Fig. S4. Relationships between N ₂ O fluxes and surface-water NO ₃ ⁻ -N, NH ₄ ⁺ -N
33	 and TDN concentrations in the Wenwusha Reservoir. Parameter bounds on the regression coefficients are at 95% confidence limits
35	5 Page S6: Table S1. Summary of F values of Two-way ANOVAs of the effect of reservoir
36	 basin, sampling month and their interaction on hydrographical parameters in the Wenwusha Reservoir.
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39	by types in the Wenwusha Reservoir.
40	0 Page S8: Table S3. Pearson correlation coefficients between N2O concentration /
41	l emission and different environmental variables in the Wenwusha Reservoir.
42	2 Page S9: Table S4. N ₂ O emission ($\mu g m^{-2} h^{-1}$) from reservoirs and lakes worldwide.
43	3 Page S10: References



S3



- 48 Fig. S2. Spatial distributions of surface-water T_W (a-c), and pH (d-f) in the Wenwusha
- 49 Reservoir.



concentrations in the Wenwusha Reservoir. Parameter bounds on the regression coefficients are at 95% confidence limits. 51 52 53



Fig. S4. Relationships between N_2O fluxes and surface-water NO_3^-N , NH_4^+-N and TDN concentrations in the Wenwusha Reservoir. Parameter bounds on the regression coefficients are at 95% confidence limits.

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Month	Discharge Area#	DO (mg L ⁻¹)*	Salinity (‰)*	NO ₃ N (mg L ⁻¹)	NH4 ⁺ -N (mg L ⁻¹)*	TDN (mg L ⁻¹)
Nov-2018	Section-I	4.50±0.28 ab	1.08±0.02 a	1.96±0.22 a	0.53±0.16 c	1.87±0.20 a
	Section-T	4.89±0.17 ab	1.10±0.00 a	1.55±0.08 ab	0.53±0.03 c	1.60±0.08 b
	Section-R	3.50±0.30 a	0.82±0.13 a	1.59±0.09 ab	1.28±0.16 a	2.00±0.14 ab
	Section-A	4.76±0.19 b	2.44±0.21 b	1.91±0.12 a	0.72±0.06 b	2.23±0.12 a
	Section-N	6.52±0.21c	2.01±0.12 b	1.27±0.10 b	0.41±0.03 d	1.67±0.07 b
	Average	4.83	1.49	1.66	0.69	1.87
Mar-2019	Section-I	8.69±0.21 b	1.08±0.02 b	1.69±0.12 a	0.41±0.05 b	2.31±0.15 ab
	Section-T	8.61±1.01 b	1.15±0.02 b	1.84±0.08 a	$0.40{\pm}0.08$ b	2.18±0.07 ab
	Section-R	4.32±0.61 a	0.66±0.13 a	1.98±0.29 a	0.97±0.23 a	2.93±0.13 a
	Section-A	9.71±0.69 c	1.27±0.02 b	1.40±0.08 a	$0.50{\pm}0.04\mathbf{b}$	1.75±0.09 b
	Section-N	13.04±0.31 d	$1.37{\pm}0.03\mathbf{c}$	1.01±0.10 b	$0.20{\pm}0.01\mathbf{c}$	1.36±0.11 c
	Average	8.87	1.11	1.58	0.50	2.11
Jun-2019	Section-I	4.02±0.23 a	0.45±0.03 a	1.14±0.23 ab	0.59±0.13 bc	1.24±0.38 bc
	Section-T	4.75±0.32 ab	0.35±0.02 a	1.08±0.16 b	0.48±0.07 c	1.22±0.11 bc
	Section-R	4.18±0.66 a	0.36±0.03 a	1.35±0.12 a	0.96±015 a	2.06±0.18 a

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 $\mathbf{S7}$

Sectio	n-A	4.44±0.20 a	0.45±0.02 a	0.99±0.06 b	0.68±0.05 b	1.56±0.11 b
Sectio	N-n	5.62±0.16 b	0.46±0.02 a	0.90±0.04 b	0.50±0.03 c	$1.08\pm0.04\mathbf{c}$
	Average	4.60	0.41	1.09	0.64	1.43
[#] Different sections impacted	by different disch	arge types: Section-I (in	dustrial effluents; $n = 4$), Se	sction-T (municipal sewage	n = 6, Section-R (river	input; $n = 7$), Section-A

 $\mathbf{S8}$

Environmental variables	N2O concentra	ıtion			N ₂ O emissio	U		
	Nov-2018	Mar-2019	Jun-2019	All data	Nov-2018	Mar-2019	Jun-2019	All data
Meteorologic variables								
Air temperature	0.602**	NS	NS	0.224*	0.373**	NS	NS	0.193**
Wind speed	NS	-0.607**	NS	-0.265*	NS	-0.300**	NS	-0.127*
Atmospheric pressure	NS	NS	NS	NS	NS	NS	NS	NS
Hydrographical parameters								
Water temperature	0.754**	NS	-0.495**	NS	0.453**	NS	-0.346**	NS
pH	NS	-0.508**	0.374**	NS	NS	-0.490**	0.274**	NS
Conductivity (EC)	-0.682**	-0.607**	-0.412**	-0.296**	-0.548**	-0.597**	-0.303**	-0.299**
Salinity	-0.705**	-0.738**	-0.442**	-0.322**	-0.543**	-0.715**	-0.343**	-0.320**
Dissolved oxygen (DO)	-0.763**	-0.358**	-0.495**	-0.294**	-0.522**	-0.219*	-0.304**	-0.302**
NO3 ⁻ -N	0.687**	0.859**	0.586**	0.445**	0.403**	0.716**	0.511**	0.369**
NH4 ⁺ -N	0.391*	0.487**	0.619**	0.452**	0.260**	0.551**	0.422**	0.402**
TDN	0.434^{**}	0.873**	0.596**	0.460^{**}	NS	0.796**	0.359**	0.316^{**}
The symbols * and ** denote significant correlati	ions at $p < 0.05$ an	d p < 0.01, resp	ectively. NS	means non-si	gnificant relation	onship.		

S9

Ecosystem Type	Location	Climate	N ₂ O emission	Reference
Reservoirs	Samuel Reservoir and Tucurui Reservoir, Brazil	Tropical	229.2	Lima et al., 2002
	Petit Saut Reservoir and Fortuna Reservoir	Tropical	-247.5 to 401.6 (167.2)	Guérin et al., 2008
	Serra de Mesa Reservoir, Brazil	Tropical	5.7	Sikar et al., 2005
	Manso Reservoir, Brazil	Tropical	5.7	Sikar et al., 2005
	Three Gorges Reservoir, China	Subtropical	-0.64 to 5.17 ()	Yu et al., 2018
	Chongqing, Southwest China	Subtropical	3.67 to 1221.0 (240.17)	Wang et al., 2017a
	Hongjiadu Reservoir, China	Subtropical	4.4 to 58.1 (19.8)	Liu et al., 2011a
	Wujiangdu Reservoir, China	Subtropical	3.5 to 77.4 (28.2)	Liu et al., 2011a
	Dongfeng Reservoir, China	Subtropical	0.4 to 26.8 (8.4)	Liu et al., 2017
	Jinghong Reservoir, China	Subtropical	0.6 to 1.1 ()	Shi et al., 2020
	Southeast Queensland, Australia	Subtropical	-0.2 to 3.5 (0.4)	Musenze et al., 2014
	Wenwusha Reservoir, China	Subtropical	-10.8 to 711.5 (107.0)	This study
	Xiaolangdi Reservoir, China	Temperate	24.2 to 25.1 (24.8)	Cheng et al., 2019
	Eguzon Reservoir, France	Temperate	4.9 to 142.1 (31.5)	Descloux et al., 2017
	Lokka Reservoir, Finland	Temperate	-3.5 to 11.4 ()	Huttunen et al., 2002
	Porttipahta Reservoir, Finland	Temperate	-0.9 to 10.6 ()	Huttunen et al., 2002

Table S3. Comparison of the N₂O fluxes (µg m⁻² h⁻¹) in reservoirs and lakes across different countries and climate zones.

S10

Wang et al., 2009	Wang et al., 2017b	Hirota et al., 2007	Liu et al., 2015	Huttunen et al., 2003	Liu et al., 2011b	Liu et al., 2011b	Liu et al., 2011b	Yan et al., 2018
18.04 to 25.5 ()	9.73 to 127.0 ()	31.2 to 78.8 (47.1)	-6.2 to 39.2 (7.9)	3.96 to 22.00	-8.2 to 68.5 (9.7)	-11.8 to 18.0 (7.9)	-35.6 to 45.2 (22.4)	0.1 to 38.3 (11.9)
Subtropical	Subtropical	Subtropical	Temperate	Temperate	Polar climate	Polar climate	Polar climate	Plateau-climate
Taihu Lake, China	Poyang Lake, China	Lake Nakaumi, Japan	Lake Yuqiao and 6 other lakes, Tianjin, China	Lake Kevätön, Finland	Lake Mochou, Antarctica	Lake Tuanjie, Antarctica	Lake Daming, Antarctica	Lakes on the Tibetan Plateau, China

Lakes

65 Numbers in brackets are averages. "---" indicates no data.

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