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### **Paper:**

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1 **Tracing marine cryptotephra in the North Atlantic during the Last Glacial Period:**  
2 **Improving the North Atlantic marine tephrostratigraphic framework**

3

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18

19 ***Abstract***

20

21 Tephrochronology is increasingly being recognised as a key tool for the correlation of  
22 disparate palaeoclimatic archives, underpinning chronological models and facilitating  
23 climatically independent comparisons of climate proxies. Tephra frameworks integrating both  
24 distal and proximal tephra occurrences are essential to these investigations providing key  
25 details on their spatial distributions, geochemical signatures, eruptive sources as well as any  
26 available chronological and/or stratigraphic information. Frameworks also help to avoid mis-  
27 correlation of horizons and provide important information on volcanic history. Here we  
28 present a comprehensive chronostratigraphic framework of 14 tephra horizons from North  
29 Atlantic marine sequences spanning 60-25 cal ka BP. Horizons previously discovered as  
30 visible or coarse-grained deposits have been combined with 11 newly recognised volcanic  
31 deposits, identified through the application of cryptotephra identification and characterisation  
32 methods to a wide network of marine sequences. Their isochronous integrity has been  
33 assessed using their physical characteristics. All horizons originated from Iceland with the  
34 vast majority having a basaltic composition sourced from the Grímsvötn, Kverkfjöll,

35 Hekla/Vatnafjöll and Katla volcanic systems. New occurrences, improved stratigraphic  
36 placements and a refinement of the geochemical signature of the NAAZ II are reported and  
37 the range of the FMAZ IV has been extended. In addition, several significant geochemical  
38 populations that further investigations could show to be isochronous are reported. This tephra  
39 framework provides the foundation for the correlation and synchronisation of these marine  
40 records to the Greenland ice-cores and European terrestrial records to investigate the phasing,  
41 rate, timing and mechanisms controlling the rapid climate changes that characterised the last  
42 glacial period.

43

44 **Keywords:** Quaternary; palaeoceanography; tephrochronology; North Atlantic; tephra  
45 framework; marine cores

46

## 47 *1. Introduction*

48

49 Tephrochronology, the use of volcanic ash deposits as isochronous tie-lines between  
50 disparate palaeoclimatic records, is increasingly being utilised as a key geochronological tool  
51 for reconstructing the timing and phasing of past climatic events (e.g. Lowe, 2011; Lowe et  
52 al., 2012; Lane et al., 2013; Davies, 2015). This upsurge is directly linked to advances in  
53 cryptotephra analysis, which has dramatically increased the number of potential tie-lines and  
54 led to the compilation of regional tephra frameworks (e.g. Lowe et al., 2008; Tryon et al.,  
55 2009; Zanchetta et al., 2011; Davies et al., 2012; Abbott and Davies, 2012; Lowe et al.,  
56 2015). Tephrostratigraphical frameworks typically include a compilation of key information  
57 relating to the tephra horizons within them, including their spatial extent, based on  
58 preservation within palaeoclimate records, glass shard concentrations, glass shard  
59 composition and eruptive source alongside chronological and stratigraphic information (e.g.  
60 Lowe et al., 2008; Davies et al., 2014; Bourne et al., 2015; Matthews et al., 2015). The most  
61 comprehensive frameworks include both distal and proximal tephra findings, visible and  
62 cryptotephra occurrences and combine newly discovered data with previously published  
63 deposits. Integrating all this information can provide valuable frameworks for the volcanic  
64 history of a region and provide key reference tools for future studies. Distal archives are often  
65 more complete than proximal records, which are prone to removal or burial of deposits,  
66 although proximal archives can often record more information regarding eruptions, such as  
67 their full geochemical evolution. In addition, developing the most comprehensive tephra  
68 frameworks will help to reduce instances of mis-correlation which can occur if volcanic

69 regions produce multiple, closely-timed eruptions with similar geochemical compositions  
70 (e.g. Lowe, 2011; Bourne et al., 2013).

71

72 For the North Atlantic region, various detailed frameworks spanning a range of time-intervals  
73 are currently available. For example, Gudmundsdóttir et al. (2016) provides a proximal  
74 framework of Icelandic eruptions during the Holocene, Blockley et al. (2014) summarises the  
75 European tephra stratigraphy over the last glacial cycle and Davies et al. (2014) provides an  
76 integrated framework of MIS 5 tephra in Greenland ice-cores and North Atlantic marine  
77 records. The tephra framework for the Greenland ice-cores has significantly expanded in  
78 recent years (e.g. Mortensen et al., 2005; Abbott and Davies, 2012; Davies et al., 2014), in  
79 particular over the MIS 2-3 period (Bourne et al., 2015), highlighting the value of exploring  
80 these distal archives. In comparison, however, only a limited number of tephra horizons have  
81 been identified in North Atlantic marine records spanning MIS 2-3 (see Haflidason et al.,  
82 2000; Wastegård et al., 2006; Section 2). This relative paucity is despite considerable  
83 advances in distal tephrochronology and the high potential for a tephra framework from these  
84 sequences to be used to establish correlations to the Greenland ice-cores and European  
85 terrestrial records. Such correlations could help answer key questions regarding the relative  
86 timing of atmospheric and oceanic changes associated with the rapid climatic events, that  
87 punctuated the region during the last glacial period (e.g. NGRIP Members, 2004; Bond et al.,  
88 1993; Martrat et al., 2007; Hall et al., 2011; Zumaque et al., 2012; Henry et al., 2016).

89

90 Here we present a tephra framework for North Atlantic marine records spanning MIS 2-3,  
91 which is underpinned by our investigations of an extensive core network (Figure 1) using  
92 recently developed cryptotephra identification methods (Abbott et al., in revision). Prior  
93 studies are also reviewed (Section 2) and previously identified isochronous horizons are  
94 integrated with our new cryptotephra discoveries. This integration represents the most  
95 concerted attempt to improve the tephra framework for the North Atlantic, and overall a  
96 framework of 14 marine tephra or cryptotephra horizons from between 60-25 cal ka BP has  
97 been defined (Figure 2).

98

## 99 ***2. Prior North Atlantic Tephra Investigations between 25-60 ka BP***

100

101 It was highlighted earlier that tephra frameworks should integrate all isochronous tephra  
102 deposits from a region, so the framework presented in this work integrates our new

103 discoveries alongside previously published data from multiple cores sites from the North  
104 Atlantic (green sites on Figure 1). Within these prior tephrochronological studies of the MIS  
105 2-3 period, several isochronous tephra horizons have been identified, i.e. North Atlantic Ash  
106 Zone II (NAAZ II), Faroe Marine Ash Zone (FMAZ) II and FMAZ IV. Reviewing the  
107 literature does, however, highlight some of the challenges associated with determining the  
108 isochronous nature of deposits and the limitations of earlier studies that only focused on the  
109 coarse fraction ( $>150\ \mu\text{m}$ ) of the marine sediments. These were the major factors driving the  
110 development of a procedure for isolating fine-grained cryptotephra (down to  $25\ \mu\text{m}$   
111 diameter) and interpreting transportation and depositional processes (e.g. Abbott et al., 2011,  
112 in revision; Davies et al., 2014; Griggs et al., 2014). This is essential to determine the  
113 isochronous nature of fine-grained, cryptotephra deposits for which macro-sedimentary  
114 evidence cannot be utilised to determine the relative influence of primary and secondary  
115 processes. These methods were utilised by Abbott et al. (2016) to identify three previously  
116 undocumented MIS 2-3 volcanic events within a core retrieved from the Goban Spur (see  
117 Section 4 for details) and are more widely applied in this study.

118

119 The first MIS 3 tephra deposit to be recognised in the North Atlantic was NAAZ II, initially  
120 identified by Bramlette and Bradley (1941) and later described by Ruddiman and Glover  
121 (1972). NAAZ II is a complex ash zone composed of the products of several Icelandic  
122 eruptions (see Section 4.1.1) with rhyolitic material from one eruption (II-RHY-1) the most  
123 widespread, being traced into multiple marine cores and the Greenland ice-cores (e.g.  
124 Kvamme et al., 1989; Grönvold et al., 1995; Lacasse et al., 1996; Zielinski et al., 1997;  
125 Haflidason et al., 2000; Austin et al., 2004; Svensson et al., 2008). The widespread nature of  
126 II-RHY-1 gives rise to a key tie-line between North Atlantic marine records and the  
127 Greenland ice-cores within the North Atlantic tephra framework (Austin and Abbott, 2010).

128

129 The FMAZs comprise a series of ash zones identified in cores around the Faroe Islands  
130 region, and three, II, III and IV, were deposited during MIS 2-3. Two of these, FMAZ II and  
131 IV, have isochronous characteristics and are integrated within the framework (Figures 1 and  
132 2; Rasmussen et al., 2003; Wastegård et al., 2006; Wastegård and Rasmussen, 2014; Griggs  
133 et al., 2014). FMAZ II was described by Wastegård et al. (2006) as a visible horizon and was  
134 suggested to be a widespread primary fall deposit. The FMAZ II was subsequently traced into  
135 the NGRIP ice-core by Davies et al. (2008) (NGRIP 1848 m;  $26,740 \pm 390\ \text{yr b2k}$ ), providing  
136 a clear demonstration of the high potential for ice-marine correlations between the Greenland

137 ice-cores and North Atlantic marine sequences during the 60-25 cal ka BP period. FMAZ IV  
138 was first described by Wastegård and Rasmussen (2014) as a layer up to 20 cm thick  
139 deposited shortly after warming related to Dansgaard-Oeschger (DO) event 12. Due to its  
140 homogeneous composition and micro-sedimentary features (Griggs et al., 2014, 2015) it has  
141 been interpreted as a primary ashfall deposit.

142

143 FMAZ III, identified as a thick relatively dispersed zone of tephra spread over ~20 cm depth  
144 in the Faroes cores, was also thought to have a correlative in the NGRIP core (NGRIP  
145 2066.95 m;  $38,122 \pm 723$  yr b2k; Davies et al., 2010). However, Bourne et al. (2013)  
146 subsequently identified a series of closely-spaced tephra horizons in the NGRIP and NEEM  
147 ice-cores around NGRIP 2066.95 m, many with geochemical compositions that fall within  
148 the wide geochemical envelope of FMAZ III. This highlighted the complexity of the period  
149 and demonstrated that the suggested correlation was inappropriate and did not represent an  
150 ice-marine tie-line (Bourne et al., 2013). Bourne et al. (2013) and Griggs et al. (2014) both  
151 suggested that FMAZ III formed through the amalgamation of several separate tephra-fall  
152 events and low sedimentation rates at the core sites so the diachronous deposits are not  
153 incorporated in the marine tephra framework.

154

155 Early studies of North Atlantic tephra mainly focused on investigating visible tephra horizons  
156 or glass shards present within the coarse fraction of the sediment (i.e.  $>150 \mu\text{m}$  diameter).  
157 This may have created a bias towards the identification of horizons from large scale eruptions  
158 and/or horizons not deposited via primary ash-fallout (Brendryen et al., 2010; Abbott et al.,  
159 2011). The study of Lackschewitz and Wallrabe-Adams (1997) highlights the limitation of  
160 this approach. Several ash zones above NAAZ II were identified within and correlated  
161 between a series of cores from the Reykjanes Ridge, however, most of these deposits have  
162 heterogeneous geochemical compositions and in general coincide with distinct peaks in ice-  
163 rafted debris (IRD). Based on these factors Lackschewitz and Wallrabe-Adams (1997)  
164 concluded that this material was transported to the sites via iceberg rafting. This process  
165 could have significantly delayed the deposition of these deposits and, hence, they do not  
166 represent isochronous marker horizons and are not incorporated in the marine tephra  
167 framework. The only deposit with isochronous characteristics was the X peak, a discrete high  
168 concentration peak within VZ 1 in the SO82-5 core, with a homogeneous glass composition  
169 and no coeval IRD peak. This horizon was subsequently correlated to FMAZ II by Wastegård  
170 et al. (2006) (Figure 2).

171  
172 Voelker and Haflidason (2015) utilised the coarse sediment fraction to define a high-  
173 resolution tephrostratigraphy for the last 86 ka from the southern Greenland Sea PS2644 core.  
174 This sequence was interpreted as containing a record of 68 volcanic events between ~60-25  
175 cal ka BP based on the geochemical analysis of glass shards from 28 depths in the core. The  
176 volcanic events, however, are sometimes defined based on a limited number of geochemical  
177 analyses of deposits with multiple glass-based geochemical populations/events often  
178 identified at the same depth. According to protocols for assessing deposits this heterogeneity  
179 could be indicative of deposition via iceberg rafting and/or secondary depositional processes  
180 (Abbott et al., in revision), however, while these processes were acknowledged a distinction  
181 between tephra deposited via primary or secondary process is often not made. This may have  
182 led to the overreporting of the number of isochronous deposits present so the deposits from  
183 these volcanic events are not incorporated into the North Atlantic tephra framework presented  
184 here. However, it is important to note these findings as a reappraisal of these deposits  
185 together with IRD evidence may well reveal the presence of dominant populations and  
186 valuable isochrons in the future.

187

### 188 **3. Methodology**

189

#### 190 *3.1 Detecting, characterising and correlating cryptotephra deposits*

191

192 A widespread network of North Atlantic cores was investigated (Figure 1) and we applied the  
193 consistent methodological approach for cryptotephra identification outlined in Abbott et al.  
194 (in revision). Following preliminary low-resolution analysis, high-resolution glass shard  
195 concentration profiles were gained from the core deposits. The major element composition of  
196 peaks in glass shard concentrations were characterised using electron-probe micro-analysis  
197 (EPMA) with at least 20-40 individual shards from each deposit analysed (see Abbott et al.,  
198 in revision for full description). For all analysis and data comparison, the major element data  
199 were normalised to an anhydrous basis, i.e. 100 % total oxides, however, the raw  
200 geochemical data are provided in the Supplementary Data alongside secondary standard  
201 analyses (Table S12). Potential sources for geochemical populations and tephra or  
202 cryptotephra horizons were explored through graphical comparison of the composition of  
203 individual shards with glass and whole rock analyses from proximal Holocene Icelandic  
204 deposits from the three different rock suites and specific volcanic systems. We acknowledge

205 that some centres may have geochemically evolved or not been productive during the last  
206 glacial period, therefore, the potential sources proposed here may need to be revised.

207

208 Potential cross-correlations between all the isochronous horizons and significant glass shard  
209 derived geochemical populations in cores within the network and other marine records were  
210 explored using statistical comparisons of their average geochemical signature and graphical  
211 comparisons on bivariate plots. The similarity coefficient function (SC) of Borchardt et al.  
212 (1972) was utilised to construct a matrix for all these comparisons (Table S13). Twenty-five  
213 of the comparisons returned SC values greater than 0.97, which implies there are strong  
214 similarities in the geochemical signatures and further assessment was required to determine if  
215 they are correlatives. A combination of three main factors were used to rule out most of these  
216 comparisons as potential correlatives: large stratigraphic discrepancies, subtle geochemical  
217 differences, and occurrence at different depths in the same core sequence. Despite the  
218 majority being ruled out, upon further assessment two of the comparisons with high SC  
219 values were found to have very strong geochemical similarities and consistent stratigraphic  
220 positions and are suggested as correlatives between marine sequences in the network (see  
221 Section 4).

222

### 223 *3.2 Assessing the isochronous nature of cryptotephra deposits*

224

225 Several of the deposits reported here have been described in Abbott et al. (in revision) as  
226 illustrative examples for assessing the dominant controls on tephra deposition in the North  
227 Atlantic region. We synthesise these results in a framework of tephra deposits that represent  
228 isochronous marker horizons identified using protocols set out in Griggs et al. (2014) and  
229 Abbott et al. (in revision). The key characteristics used to define isochronous horizons are: (i)  
230 a clear peak in the shard concentration profile that can be used as the isochron position and  
231 (ii) a homogeneous geochemical population or distinct trend in glass shard analyses  
232 indicative of material deriving from a single volcanic eruption. Abbott et al. (in revision)  
233 outlines a tephra deposit type scheme that uses glass shard concentration profiles and  
234 geochemical homogeneity/heterogeneity to identify six North Atlantic marine tephra deposit  
235 types with common modes of tephra delivery and post-depositional reworking. Here that  
236 scheme is utilised to aid the assessment of the deposits identified in the marine records  
237 Although Type 1 and 3 deposits are typically characterised by single homogeneous  
238 populations there is greater variability and complexity in the geochemical signatures of Type



239 2 deposits. For the latter a larger number, typically >30 but on occasions up to 60, of single-  
240 grain major element analyses were acquired. These were graphically assessed to explore the  
241 relative homogeneity or heterogeneity of deposits, define homogeneous populations that may  
242 have derived from single eruptions, quantify their relative dominance within the deposits and  
243 categorise them as Type 2A or Type 2B deposits. Outliers were defined as analyses that were  
244 not consistently associated with a defined population. For some heterogeneous deposits  
245 where populations were not identifiable analyses were grouped based on affinities to the  
246 Icelandic rock suites (see Supplementary Figures).

247

### 248 *3.3 Age and stratigraphic constraints*

249

250 The timing of deposition for each tephra deposit is given based on the available  
251 climatostratigraphy for the specific core within which the horizons were isolated (Table 1).  
252 For some records, there is strong stratigraphic control based on proxy records from the cores  
253 that record the DO events which characterised the North Atlantic region during the last  
254 glacial period, e.g. MD04-2822 and MD04-2829CQ. However, for other cores, e.g. MD99-  
255 2251 and GIK23415-9, the stratigraphic frameworks are not as distinct with deposits from the  
256 Heinrich events providing the best stratigraphic control. Due to uncertainties in the relative  
257 timing of closely spaced horizons not identified in the same core sequence the stratigraphic  
258 relationships presented in Figure 2 should be treated with caution, e.g. the cluster of horizons  
259 that have been identified in various cores around the H4 event (Figure 2). Further  
260 investigations of these horizons, such as their tracing into other sequences, may help to refine  
261 the sequence of the volcanic events in the future.

262

### 263 *4. North Atlantic Tephra Framework*

264

265 An improved marine tephra framework for the North Atlantic between 60-25 cal kyr BP is  
266 presented in Figure 2 and Table 1. Overall, a framework of 14 isochronous horizons can be  
267 defined, including 8 new isochronous horizons presented for the first time, 3 cryptotephra  
268 deposits identified in MD04-2820CQ by Abbott et al. (2016) and 3 previously published  
269 deposits (NAAZ II, FMAZ IV and FMAZ II). This new framework represents a significant  
270 increase in the number of tephra marker horizons that could be utilised for the correlation of  
271 records during this period.

272

273 With the exception of NAAZ II (II-RHY-1) and MD04-2820CQ 497-498 cm, all tephra in  
274 the framework are basaltic in composition and originated from Iceland, specifically from the  
275 Grímsvötn, Kverkfjöll, Hekla/Vatnafjöll and Katla volcanic systems (Table 1). The most  
276 widespread isochronous horizon in the framework is the NAAZ II (II-RHY-1) (Figures 3 and  
277 4). The wide distribution and importance of this horizon had been established in prior studies,  
278 however, here we have isolated it in more sequences, gained greater control on the timing of  
279 deposition, with peaks in shard concentration determined at a 1 cm resolution, and provided  
280 an improved glass geochemical signature for the horizon (Section 4.1.1). The geographical  
281 range of the previously identified FMAZ IV can be expanded, to a limited extent, from the  
282 Faroe Islands region to the Norwegian Sea following its identification in MD95-2010 (Figure  
283 5; Section 4.1.2).

284

285 Within our network only two cores, MD04-2822 and MD04-2829CQ, exclusively preserved  
286 isochronous Type 1 deposits (Figures 6a and 6b). New isochronous horizons were also  
287 identified in two further cores, MD99-2251 and GIK23415-9, alongside other deposits  
288 without clear isochronous characteristics, i.e. Type 2B and Type 4 deposits (Figures 7a and  
289 8a), which can be attributed to temporal variability in the processes controlling tephra  
290 deposition at these sites (see Abbott et al., in revision). Further details regarding all the  
291 isochronous horizons are provided in Section 4.1 in chronological order from the oldest to the  
292 youngest horizon.

293

294 The Type 2B and Type 4 horizons are not overlooked though as analysis showed that within  
295 many of these deposits significant homogeneous geochemical populations could be isolated  
296 (Figures 7b and 8b; Table 1). These populations are presented alongside the framework of  
297 isochronous horizons as their geochemical homogeneity suggests that they were derived from  
298 single volcanic events, but, at present, questions remain over their depositional origin and  
299 isochronous nature. Further investigations, however, may permit their integration into the  
300 regional tephra framework and this is discussed further in Section 4.2.

301

#### 302 *4.1 Isochronous horizons*

303

##### 304 *4.1.1 NAAZ II*

305

306 NAAZ II is a crucial deposit within the North Atlantic marine tephra framework and it has  
307 been identified at nine sites within our network as a clear peak in rhyolitic material and at 6  
308 sites basaltic/intermediate material was also present. Based on occurrences of NAAZ II in  
309 several North Atlantic sites, this ash zone was defined as being composed of five  
310 geochemical populations, one rhyolitic (II-RHY-1) and four basaltic (II-THOL-1, II-THOL-  
311 2, II-THOL-3 and II-TAB-1) by Kvamme et al. (1989).

312

313 Shards from the peaks in rhyolitic material at the 9 sites have a consistent homogeneous  
314 transitional alkali rhyolitic composition (Figure 3a(i) and 4b; Table S2). In comparison to  
315 prior characterisations of NAAZ II from several North Atlantic marine cores, strong  
316 similarities can be observed for some oxides, e.g. FeO and CaO (Figure 3bi) but some offsets  
317 are apparent for other oxides, e.g. Na<sub>2</sub>O and SiO<sub>2</sub> (Figure 3bii). These differences are  
318 reflected in similarity coefficient comparisons (Table S2) and are consistent with sodium loss  
319 affecting the older EPMA analyses (Hunt and Hill, 1993; Kuehn et al., 2011), particularly for  
320 the analyses from Kvamme et al. (1989), and are highly unlikely to indicate a different source  
321 for the material. Therefore, the nine deposits in this network can be correlated to the II-RHY-  
322 1 component of NAAZ II. These new analyses provide an up-to-date composition for this  
323 component and highlight that data quality must be considered when assessing correlations  
324 between datasets, especially for rhyolitic material.

325

326 A peak in brown shards was isolated in direct association with the II-RHY-1 peak at 6 sites  
327 (Figure 4b; e.g. in MD99-2251 (Figure 4a)). Compositional analyses revealed a range of  
328 signatures with basaltic and intermediate material present (Figure 3a(ii)). Shards related to  
329 three of the basaltic populations of Kvamme et al. (1989) have been identified, but no shards  
330 related to the II-THOL-3 population were isolated (Figure 3c). Glass shards with an  
331 intermediate trachyandesite to trachydacite composition have been identified (Figure 3a(ii))  
332 and grouped as a new population, which we name II-INT-1. Some material with an  
333 intermediate composition was found in association with the proximal Icelandic deposit  
334 correlated to NAAZ II, the Thorsmörk ignimbrite (Jørgensen, 1980). However, this is less  
335 evolved than the material in these marine deposits with SiO<sub>2</sub> values of 56-58 % and is  
336 unlikely to be directly related. This additional intermediate population suggests that the  
337 basaltic material associated with NAAZ II derives from more individual eruptions than  
338 previously thought. This assertion is also supported by differences in the composition of  
339 material from this study attributed to the populations of Kvamme et al. (1989) which may

340 indicate they grouped material from multiple eruptions as single populations. For example,  
341 shards from M23485-1 and GIK23415-9 display geochemical differences, e.g. Figure 3cii,  
342 despite all falling into the II-THOL-2 field of Kvamme et al. (1989). At three of the sites the  
343 brown shards can be grouped as single populations: homogeneous populations within the II-  
344 THOL-2 geochemical field in M23485-1 and JM11-19PC and only shards from the  
345 intermediate population are present in MD01-2461 (Figure 4c). The remaining three sites  
346 preserve a mix of populations. MD04-2820CQ preserves three populations (II-THOL-1, II-  
347 THOL-2 and II-INT-1), each exceeding 24% of the shards present. GIK23415-9 and MD99-  
348 2251 are dominated by the II-THOL-1 and II-TAB-1 populations, respectively.

349

350 The contrast between the homogeneity of the rhyolitic material at all sites and the  
351 heterogeneity and inconsistent signatures of the basaltic/intermediate material may indicate  
352 that despite coeval deposition the two components were transported differentially. It has been  
353 suggested that NAAZ II was primarily transported from Iceland via sea-ice rafting and  
354 primary airfall (e.g. Ruddiman and Glover, 1972; Austin et al., 2004; Wastegård et al., 2006).  
355 Sea-ice rafting may have contributed towards the relatively higher rhyolitic shard  
356 concentrations at sites to the south and west of Iceland. The geochemical homogeneity and  
357 distinct peak with an upward tail in rhyolitic shard concentrations (i.e. Type 3 deposits; e.g.  
358 Figure 4a(i)), observed at all sites is consistent with these transport processes and supports  
359 the isochronous nature of the II-RHY-1 component.

360

361 The heterogeneity of the basaltic material and relative discreteness of the concentration  
362 peaks, e.g. Figure 4a(ii), are consistent with transport via iceberg rafting and the between-site  
363 contrasts in geochemical signatures highlights that icebergs calved from different margins of  
364 the Icelandic ice sheet could have transported and deposited material at the core sites. The  
365 absence of basaltic material associated with the rhyolitic peaks in the MD04-2822 and  
366 MD95-2010 sites is consistent with the findings of Abbott et al. (in revision) that ice rafting  
367 did not transport tephra to these sites during the last glacial period. Transportation via iceberg  
368 rafting can delay the deposition of tephra: therefore the peaks in basaltic material related to  
369 NAAZ II should not be utilised as isochronous markers. However, based on their dominance  
370 as homogeneous populations at some sites, II-THOL-2, II-TAB-1 and II-INT-1 are regarded  
371 as significant geochemical populations (Table 1). It cannot be ruled out that one or more of  
372 the basaltic populations were deposited coevally via primary fallout with the rhyolitic  
373 material, particularly at sites only containing one population. However, it is unlikely that this

374 process deposited all of the basaltic populations with subsequent amalgamation in the  
375 sediment column, as shard concentrations profiles for that type of deposit (Type 4) typically  
376 have a greater vertical spread within sequences and display multiple concentration peaks.

377

378 The coeval deposition of the two shard types may indicate that the volcanic eruption that  
379 produced the rhyolitic tephra horizon triggered an ice-rafting event which deposited the  
380 basaltic material, but the resolution of the marine records under investigation here is  
381 insufficient to resolve this temporal phasing.

382

#### 383 *4.1.2 FMAZ IV – MD95-2010 915-916 cm*

384

385 FMAZ IV was identified in the MD95-2010 core from the Norwegian Sea as a discrete  
386 deposit at 915-916 cm depth (Figure 5a). This deposit has a homogeneous basaltic glass  
387 composition with affinities to the Icelandic tholeiitic rock suite and the products of the  
388 Grímsvötn volcanic system. The glass composition of MD95-2010 915-916 cm is identical to  
389 the characterisation of the JM11-19PC 542-543 cm deposit of Griggs et al. (2014) (Figure 5b;  
390 SC – 0.985), previously correlated to the FMAZ IV of Wastegård and Rasmussen (2014).  
391 According to the age model and stratigraphy for MD95-2010 from Dokken and Jansen  
392 (1999), this layer has an age of ~44.45 cal ka BP and was deposited during the DO-12 event  
393 based on the magnetic susceptibility record. This stratigraphic position and age estimate are  
394 consistent with the work of Wastegård and Rasmussen (2014). This horizon has previously  
395 not been identified outside the Faroe Islands region and, therefore, this discovery expands its  
396 geographical range in a northeasterly direction to the Nordic Sea.

397

#### 398 *4.1.3 MD04-2820CQ 524-525 cm*

399

400 MD04-2820CQ 524-525 cm has previously been described by Abbott et al. (2016) where it  
401 was identified as a clear peak in shard concentrations spanning ~6 cm depth. Geochemical  
402 analyses of shards from this deposit form a homogeneous tholeiitic basaltic population  
403 sourced from either the Grímsvötn or Kverkfjöll Icelandic volcanic systems. These  
404 characteristics allow the deposit to be defined as Type 2A and, allied with a lack of direct  
405 covariance with IRD, this deposit is thought to have been deposited via primary fallout  
406 despite occurring during a period of elevated IRD concentrations (Abbott et al., 2016).

407

408 *4.1.4 MD04-2822 2017-2018 cm*

409

410 High-resolution analysis of MD04-2822 showed a well-constrained peak in brown glass  
411 shards in all grain-size fractions at 2017-2018 cm depth (Figure 6a). According to the core  
412 stratigraphy, this horizon was deposited during a stadial period prior to the warming  
413 transition into DO-9 (Figure 6a). Shards have a homogeneous basaltic composition with  
414 affinities to the Icelandic tholeiitic rock suite and the products of the Grímsvötn volcanic  
415 system (Figure 6c).

416

417 *4.1.5 MD04-2820CQ 497-498 cm*

418

419 MD04-2820CQ 497-498 cm was identified as a small peak in colourless glass shards, during  
420 a period of consistently elevated shard concentrations, deposited prior to DO-9 (Abbott et al.,  
421 2016). Shards from the peak have a transitional alkali rhyolitic composition and form a single  
422 population with affinities to a number of distal tephra deposits previously attributed to the  
423 Katla volcanic system (Abbott et al., 2016). This horizon is notable as it is the only other  
424 rhyolitic horizon within the marine tephra framework apart from the rhyolitic component of  
425 NAAZ II (Table 1). Due to its homogeneity and the prevalence of shards in the 25-80  $\mu\text{m}$   
426 fraction, this deposit was interpreted as an isochronous horizon deposited via primary ashfall  
427 (Abbott et al., 2016).

428

429 *4.1.6 MD04-2820CQ 487-488 cm*

430

431 Deposited just prior to Heinrich Event 4, MD04-2820CQ 487-488 cm was identified as a  
432 clear peak in brown glass shard concentrations across all grain size fractions spread over ~3  
433 cm depth (Abbott et al., 2016). While some transitional alkali outliers are present within  
434 shard analyses from this deposit, the vast majority of shards (~85 %) form a homogeneous  
435 geochemical population with a tholeiitic basaltic composition and affinities to the Grímsvötn  
436 volcanic system (Abbott et al., 2016). This homogeneous composition and a lack of  
437 covariance of shard concentrations with IRD suggests it was not deposited via iceberg rafting.  
438 Deposition is likely to have occurred via primary fall, however, the high proportion of shards  
439 in the coarser grain-size fractions (80-125  $\mu\text{m}$  and >125  $\mu\text{m}$ ) in comparison to the 25-80  $\mu\text{m}$   
440 fraction may also indicate transport via sea-ice rafting. Neither transport process would

441 impart a significant temporal delay in deposition, therefore, MD04-2820CQ 487-488 cm is  
442 viewed as an isochronous deposit (Abbott et al., 2016).

443

#### 444 *4.1.7 MD04-2829CQ 934-935 cm and 930-931 cm*

445

446 Two distinct and closely spaced peaks in brown glass shards were isolated in MD04-2829CQ  
447 with concentrations of ~35 shards per 0.5 g dws in the 25-80  $\mu\text{m}$  grain-size fraction (Figure  
448 6b). Only a limited number of shards were isolated in one of the three samples between these  
449 peaks. The stratigraphy for MD04-2829CQ indicates that these horizons were deposited  
450 during and just after the rapid warming into DO-8 (Figure 6b; Hall et al., 2011). Shards from  
451 both peaks were geochemically analysed and the analyses revealed two homogeneous  
452 basaltic populations with affinities to the Icelandic tholeiitic rock suite and the products of the  
453 Grímsvötn volcanic system. However, there are distinct differences in  $\text{Al}_2\text{O}_3$ , FeO, CaO and  
454 MgO between the two deposits (Figure 6c). These differences show that despite being  
455 separated by only 3 cm of sediment the horizons were produced by two separate volcanic  
456 eruptions and, coupled with their other characteristics, can both be considered as valuable  
457 isochronous marker horizons.

458

#### 459 *4.1.8 MD04-2822 2004-2005 cm*

460

461 High-resolution shard counts identified brown shards within the 25-80 and  $>125$   $\mu\text{m}$  grain-  
462 size fractions in the 2004-2005 cm sample of MD04-2822 (Figure 6a). While the shard  
463 concentrations are low the peaks are discrete as no further shards were identified in adjacent  
464 samples. Based on the stratigraphy of the core this material was deposited shortly after the  
465 warming transition into DO-8 (Figure 6a; Hibbert et al., 2010). Geochemical analysis shows  
466 that shards from the deposit have a homogeneous transitional alkali basaltic composition  
467 (Figure 6c). The shards are characterised by high  $\text{TiO}_2$  values of ~4.65 %wt and comparisons  
468 to proximal Icelandic deposits demonstrate that the deposit was most likely sourced from the  
469 Katla volcanic system (Figure 6c). The geochemical composition of the material in this peak  
470 is markedly distinct from the material in the underlying MD04-2822 2017-2018 cm horizon,  
471 indicating that they represent two discrete eruption events.

472

#### 473 *4.1.9 MD99-2251 1680-1681 cm*

474

475 The highest brown shard concentrations in MD99-2251 were identified as a peak centred  
476 around 1680-1681 cm depth (Figure 7a). Overall, high shard concentrations associated with  
477 this peak cover approximately 10 cm depth, typical of a Type 2 deposit, and glass shards  
478 from the main peak and a secondary peak at 1683-1684 cm were geochemically analysed.

479

480 Shards from 1680-1681 cm form a clear near-homogeneous population, with 76 % of the  
481 analyses in this population (Figure 7b). High TiO<sub>2</sub> concentrations in excess of 4.4 %wt  
482 strongly indicate an origin from the Katla volcanic system (Figure 7b). Within the remaining  
483 25 % of shards a minor population (6 %) of tholeiitic material, most likely sourced from the  
484 Kverkfjöll volcanic system, was also identified alongside several outlying shards (Figure 7b).  
485 The significant dominance of a single homogeneous population in the 1680-1681 cm peak,  
486 suggests that this material was deposited via primary ashfall and that this tephra deposit  
487 represents an isochronous marker horizon despite being deposited during a period of elevated  
488 IRD concentrations associated with Heinrich Event 3 (Figure 7a).

489

490 The glass-derived geochemical signature of material from the underlying 1683-1684 cm peak  
491 is the same as that of the major 1680-1681 cm peak suggesting that this does not represent an  
492 earlier and separate depositional event but instead represents downward reworking of  
493 material from the main concentration of glass. The slight deviation of the shard concentration  
494 profile from a gradational downward tail could imply that any reworking processes were not  
495 uniform across the core. Such variability was observed by Griggs et al. (2015) in 3D  
496 reconstructions of the structure of tephra deposits gained using X-ray microtomography.

497

#### 498 *4.1.10 MD04-2829CQ 800-801 cm*

499

500 The highest shard concentrations in core MD04-2829CQ were identified at 800-801 cm, with  
501 increases observed in all grain-size fractions (Figure 6b). This deposit is very discrete with  
502 limited shards identified in adjacent samples. Stratigraphic constraints indicate that this  
503 horizon was deposited in the cold period prior to DO-4 (Figure 6b; Hall et al., 2011).

504 Compositional analysis of individual shards shows that all material has a tholeiitic basaltic  
505 composition and can be grouped into two homogeneous populations, with clear bimodality  
506 observed for some oxides, including TiO<sub>2</sub>, FeO, CaO and MgO (Figure 6c). Analyses  
507 grouped into population THOL-1 were only derived from shards from the 25-80 µm grain-  
508 size fraction, whereas the majority of analyses in population THOL-2 are from shards >80



509  $\mu\text{m}$  in diameter. Based on comparisons to proximal Icelandic deposits, THOL-1 has a close  
510 affinity to products of Grímsvötn while THOL-2 is most likely derived from the Kverkfjöll  
511 volcanic system (Figure 6b; e.g. Óladóttir et al., 2011). This implies that the deposit was  
512 formed from the deposition of material from two coeval eruptions of these volcanic centres.

513

#### 514 *4.1.11 MD04-2822 1836-1837 cm - GIK23415-9 225-226 cm*

515

516 Within the MD04-2822 record the largest peak in brown shards was identified at 1836-1837  
517 cm depth with >40 shards per 0.5 g of dws present in the 25-80  $\mu\text{m}$  fraction (Figure 6a). The  
518 material is stratigraphically well constrained with only 2 shards present in the underlying  
519 sample. According to the stratigraphy this material was deposited during the cold stadial  
520 period shortly before the transition into DO-4 (Hibbert et al., 2010). Compositional analysis  
521 of glass shows that material from this peak has a transitional alkali basaltic composition and  
522 forms a homogeneous geochemical population (Figure 6c). Comparisons to proximal  
523 Icelandic deposits indicate that the horizon was sourced from either the Katla or  
524 Hekla/Vatnafjöll volcanic system (Figure 6c).

525

526 A discrete peak in shard concentrations, restricted to 1 cm and with the characteristics of a  
527 Type 1 deposit, was also isolated between 225-226 cm in GIK23415-9 (Figure 8a).  
528 Geochemical analysis of the shards from this deposit shows that all have a transitional alkali  
529 composition (Figure 8b). Within the analyses bimodality can be observed for some oxides,  
530 most notably  $\text{TiO}_2$ , and they can be split into two homogeneous populations. A dominant  
531 population (TAB-1) of 70 % of the shards with low  $\text{TiO}_2$  values and a smaller population  
532 (TAB-2) of 15 % of the analysed shards with  $\text{TiO}_2$  values  $\sim 0.35\%$  wt higher.  $\text{TiO}_2$  values  
533 have been identified as one of the primary oxides that can be used to discriminate between  
534 Icelandic basaltic eruptions from the last glacial period (e.g. Bourne et al., 2013, 2015). The  
535 remaining 15 % of analyses are classified as outliers. Comparisons to proximal deposits show  
536 that the populations have similarities to the products of both the Katla and Hekla/Vatnafjöll  
537 volcanic systems (Figure 8b). GIK23415-9 225-226 cm was deposited during Heinrich Event  
538 3 which could suggest it was deposited via iceberg rafting. However, the relative dominance  
539 of the TAB-1 population and a lack of direct covariance of shard concentrations with IRD,  
540 with the discrete shard peak contrasting with elevated IRD concentrations for  $\sim 25$  cm of core  
541 depth, do not support this interpretation. These indicators provide support for primary ashfall

542 deposition of glass shards from either a single chemically bimodal eruption or two eruption  
543 events very close in time.

544

545 Statistical analysis (SC of 0.987) and graphical comparisons support a correlation between  
546 MD04-2822 1836-1837 cm and GIK23415-9 225-226 cm (TAB-1) (Table S13; Figure 9a). In  
547 addition, there is a consistency in the stratigraphic position of the two horizons. MD04-2822  
548 1836-1837 cm was deposited between DO events 5 and 4 (Figure 6a), while GIK23415-9  
549 225-226 cm was deposited at the end of Heinrich Event 3 (Figure 8a), which, based on a  
550 comparison of ages for the Heinrich Events from Sanchez Goñi and Harrison (2010) and the  
551 Greenland ice-core chronology presented in Seierstad et al. (2014), occurred after Greenland  
552 Interstadial (GI) 5, the ice counterpart to DO-5. Based on the available information, we assert  
553 that these two deposits are the products of the same volcanic event and form a tie-line  
554 between the two relatively closely spaced sequences (Figure 2).

555

#### 556 *4.1.12 GIK23415-9 173-174 cm*

557

558 A peak in basaltic glass shard concentrations was identified in the GIK23415-9 core at a  
559 depth of 173-174 cm, following Heinrich Event 2 (Figure 8a). The shard concentration  
560 profile of this deposit is akin to a Type 1 deposit with a relatively discrete peak in shard  
561 concentrations restricted to ~1 cm (Figure 8a). Geochemical analysis of shards from this  
562 deposit show one clear homogeneous population, composed of 60 % of the analysed shards,  
563 with a basaltic tholeiitic composition and an affinity to the Kverkfjöll volcanic system  
564 (Figure 8b). The remaining 40 % are heterogeneous and can be regarded as outliers (Figure  
565 S8). Although the overall homogeneity of the deposit is not as distinct as most Type 1  
566 deposits, the occurrence of a homogeneous population deposited during a period of low IRD  
567 input does suggest that primary fall occurred to form an isochronous deposit. The outlying  
568 shards may derive from a low background of IRD input of ice-rafted shards during this  
569 period. In addition, the use of the percentage abundance of populations to assess this deposit  
570 has some limitations as only a low number of analyses, 15, were gained from shards within  
571 this deposit.

572

#### 573 *4.2 Significant geochemical populations and possible isochrons*

574

575 In addition to the isochronous deposits outlined in Section 4.1, six tephra deposits in the  
576 MD99-2251 core and four in the GIK23415-9 sequence were assessed as having non-  
577 isochronous characteristics and have been classified as Type 2B or Type 4 deposits (Figures  
578 7a and 8a). The main criterion underpinning this assessment was the geochemical  
579 heterogeneity of the deposits, indicative of the amalgamation of material from a number of  
580 volcanic eruptions. However, while only three deposits, MD99-2251 1654-1655 cm and  
581 1796-1797 cm and GIK23415-9 193-194 cm, have fully heterogeneous compositions the  
582 other deposits contain 16 significant homogeneous geochemical populations, in total, within  
583 their overall heterogeneity (Figure 2; Table 1). The significant geochemical populations may  
584 relate to single volcanic eruptions, but due to their occurrence within heterogeneous deposits  
585 further investigations are required to determine if they were deposited isochronously or  
586 otherwise. The full glass-based geochemical signatures of all MD99-2251 and GIK23415-9  
587 deposits and the populations identified within them are summarised in Figures S1-S14 and  
588 Tables S8 and S10.

589

590 The 16 populations all have a basaltic composition and were sourced from Iceland. In  
591 addition to the volcanoes which deposited isochronous horizons in the North Atlantic region,  
592 i.e. Grímsvötn, Kverkfjöll, Hekla/Vatnafjöll and Katla, homogeneous glass shard populations  
593 with geochemical similarities to the products of the Veidivötn-Bardarbunga and  
594 Vestmannaeyjar volcanic systems were identified (Table 1; Figures 7a and 7b). Their relative  
595 dominance within the deposits is variable, ranging from ~10 to 60 % of the total single-shard  
596 analyses used to characterise the deposits (Tables S8 and S10).

597

598 Co-variance of shard concentration profiles with IRD records was another variable used to  
599 assess the isochronous nature of the deposits (Abbott et al., in revision). Some of the deposits  
600 with heterogeneous signatures were deposited during periods of elevated or rising IRD  
601 concentration, which could indicate transport via iceberg rafting and a significant temporal  
602 delay between eruption and deposition. However, iceberg rafting is not the only process that  
603 can amalgamate the products of multiple eruptions. For example, for some deposits post-  
604 depositional mixing in the sediment column of the products of several closely-timed  
605 eruptions cannot be ruled out as some were isolated within periods of limited IRD deposition.  
606 In this later scenario, deposition would have been via primary ashfall with no temporal delay,  
607 however, determining the isochron position is challenging as complexity is often observed in  
608 the shard concentration profiles. Primary fallout could also have occurred during a period of

609 ice-rafting deposition resulting in the incorporation of a homogeneous ashfall population  
610 within a heterogeneous background rafted signal.

611

612 These differing scenarios and the uncertainty in the depositional processes implies that  
613 further investigations are required to assess whether these populations are isochronous.  
614 Consequentially we have reported the significant geochemical populations, but we have not  
615 incorporated them within the regional tephra framework until further evidence is gained.  
616 Such evidence may include their identification in other North Atlantic marine cores and/or  
617 the Greenland ice-core tephra framework in a similar stratigraphic position. In addition, for  
618 some records the covariance with IRD could not be fully explored because of the lower  
619 resolution in this dataset relative to the shard concentration profiles. Improved high-  
620 resolution IRD records would be highly advantageous for further assessing depositional  
621 processes. An example of how tracing these populations into other records could provide  
622 further insights into their isochronous nature is provided within our work.

623

624 The assessment of potential correlations (Table S13) highlighted a strong similarity between  
625 the glass-based geochemical signature of FMAZ II and the THOL-1 population in the  
626 GIK23415-9 202-203 cm deposit (Figure 7; Table 1). The SC comparison returned a high  
627 coefficient of 0.990, demonstrating that the signatures were nearly identical, and this  
628 observation is corroborated by graphical comparisons (Figure 9b). Stratigraphically, FMAZ II  
629 has been identified between Heinrich Events 3 and 2 in marine records and was deposited  
630 prior to an increase in IRD concentrations in the ENAM93-21 core (Rasmussen et al., 2003)  
631 and after GI-3 in the Greenland ice-core stratigraphy (Davies et al., 2010). GIK23415-9 202-  
632 203 cm was deposited during a period of increasing IRD concentrations related to the start of  
633 Heinrich Event 2 (Figure 8a). These stratigraphic juxtapositions are consistent and, coupled  
634 with the strong geochemical similarities, could imply isochronous deposition from the same  
635 volcanic event. GIK23415-9 202-203 cm (THOL-1) is one of 4 homogeneous geochemical  
636 populations within the deposit and, due to their co-occurrence, it was interpreted as being  
637 deposited by iceberg rafting. The proposed correlation does not contradict this interpretation  
638 but could demonstrate that GIK23415-9 202-203 cm (THOL-1) was deposited via primary  
639 ashfall during a period when tephra from other events was rafted by icebergs. Overall, this  
640 potential correlation highlights the complexity of some deposits, but demonstrates how these  
641 significant glass geochemical populations are important to consider as potential isochronous  
642 markers.

643

644 **5. Discussion**

645

646 *5.1 Future application of the North Atlantic marine tephra framework*

647

648 The North Atlantic marine tephra framework between MIS 2-3 has been significantly  
649 improved through the most extensive application of cryptotephra methods, comprehensive  
650 compositional analysis and rigorous and defined protocols to assess the isochronous nature of  
651 each deposit. For a long period only a limited number of horizons had been identified in this  
652 time period (Haflidason et al., 2000; Wastegård et al., 2006). Now this framework includes  
653 14 isochronous horizons that have considerable promise for correlating and synchronising  
654 palaeoclimatic records. There is also potential to add further isochronous markers given the  
655 significant geochemical populations identified in heterogeneous deposits also reported in this  
656 study.

657

658 NAAZ II remains a dominant tephra within this framework and our work has identified it in  
659 numerous additional cores with greater control on the timing of deposition derived from high-  
660 resolution shard counts and an improved geochemical signature for glass shards associated  
661 with the widespread rhyolitic component (II-RHY-1). This tephra, with an age of  $55,380 \pm$   
662  $1184$  yr b2k in the Greenland ice-core records (Svensson et al., 2008), represents a key  
663 marker horizon for the period providing an isochronous tie-line linking numerous widespread  
664 marine cores and the Greenland ice-core records beyond the radiocarbon window. The  
665 distribution of the FMAZ IV has been extended from the Faroe Islands region into the Nordic  
666 Seas and has the potential to be a key tie-line for DO 12. However, despite being found  
667 previously in several North Atlantic cores and the NGRIP ice-core (see summary map in  
668 Davies et al., 2012), the FMAZ II was only found in one additional core, GIK23415-9.  
669 Furthermore, most of the new cryptotephtras are single-core occurrences, highlighting  
670 challenges with cryptotephra tracing within the North Atlantic Ocean. The limited tracing of  
671 horizons may reflect the difficulties of detecting and isolating deposits that often only contain  
672 a low concentration of shards, but could also indicate the relatively constrained dispersal of  
673 the basaltic eruptions depositing material over the North Atlantic. Only one correlation has  
674 been made between newly identified isochronous horizons in the framework, MD04-2822  
675 1836-1837 cm and GIK23415-9 225-226 cm (Section 4.1.11; Figure 2). These cores are  
676 relatively closely spaced, supporting the suggestion of limited basaltic ash dispersal.

677

678 Assessing potential correlations between the records highlighted that while a range of factors  
679 demonstrated that there are few direct correlations, many of the horizons have similar major  
680 element geochemical signatures, especially eruptives from the Grímsvötn and Katla volcanic  
681 systems (Table 1). This conclusion corroborates the findings of Bourne et al. (2015) who  
682 observed similar repetition of major element glass geochemical signatures from these systems  
683 in tephra horizons in the Greenland ice-cores. This repetition is particularly notable for the  
684 period around H4 as a cluster of six closely spaced horizons has been identified in the marine  
685 cores (Figure 2). Of these, five horizons have similar tholeiitic glass major element basaltic  
686 compositions and are thought to be derived from the Grímsvötn volcanic system. However,  
687 subtle differences in geochemical signatures show they represent individual events, which  
688 may be further emphasised through trace element analysis (Lowe et al., 2017).

689

690 The observations that the new cryptotephra in the North Atlantic region may have limited  
691 dispersal and geochemical similarities do provide challenges for future correlation. There is,  
692 however, the potential to constraint a number of rapid climate events, such as H4 and DO 8  
693 and H3 as clusters of isochronous horizons are present around those events. Further  
694 investigations should initially focus on sites close to those preserving the isochronous  
695 horizons in this framework and/or re-evaluate previously explored sites (e.g. green sites on  
696 Figure 1), with adaptations to the methodological approach discussed in Section 5.3. It is  
697 imperative that potential correlations are rigorously assessed as correlating horizons or  
698 populations with close, but not identical, major element glass geochemical signatures, could  
699 lead to the establishment of incorrect tie-lines between records. Trace element analysis of the  
700 glass shards may aid this assessment as the additional signature may show greater differences  
701 between tephras (Lowe et al., 2017). Other supporting evidence such as broad stratigraphic  
702 constraints and independent age estimates can also be used to support and test correlations. A  
703 detailed assessment of possible correlations to the Greenland ice-cores will be discussed in a  
704 forthcoming publication whereby trace element signatures are also employed to assess and  
705 support correlations.

706

## 707 *5.2 Reconstructing Icelandic volcanic history*

708

709 This framework adds to our understanding of the volcanic history of Iceland during the last  
710 glacial period between 60-25 cal kyr BP. The dominance of basaltic over rhyolitic horizons

711 and the high productivity of the Grímsvötn/Kverkfjöll and Katla volcanic systems around  
712 Heinrich Event 4 and Heinrich Event 3, respectively, is consistent with the Greenland ice-core  
713 tephra framework for the same period (Bourne et al., 2015). The dominance of basaltic  
714 horizons in both sets of archives strongly suggests that differential dispersal of the products  
715 of rhyolitic eruptions was not occurring and their paucity reflects a relatively lower frequency  
716 of Icelandic rhyolitic eruptions during this period. Basaltic horizons potentially sourced from  
717 other volcanic centres were observed, including the Veidivötn-Bardabunga and  
718 Vestmannaeyjar volcanic systems. There are very few or no tephra in the Greenland  
719 framework with glass-based geochemical similarities to those horizons, potentially due to a  
720 bias in dispersal direction, a low number of eruptions from these sources and/or the nature of  
721 volcanic eruptions from these systems. This observation shows that a more complete  
722 reconstruction of Icelandic volcanism will be gained by integrating the two frameworks.  
723 There is, however, a notable difference between the number of tephra deposits identified  
724 between the marine and ice-core records. With 99 volcanic events recorded in the Greenland  
725 records in contrast to 33 events in the marine archives, if the homogeneous populations are  
726 assumed to derive from individual volcanic events. The lower resolution of the marine  
727 records, the potential for the amalgamation of airfall deposits, post-depositional reworking  
728 processes and the masking of low concentration glass shard deposits (see below) are the most  
729 likely causes of this disparity.

730

### 731 *5.3 Improving the marine tephra framework*

732

733 This work has demonstrated the potential of identifying isochronous cryptotephra in North  
734 Atlantic marine records of the last glacial period. However, the methodology employed to  
735 identify cryptotephra in this work most likely created a bias towards the identification of  
736 horizons depositing a high concentration of glass shards at core sites. As discussed by Timms  
737 et al. (2017), the process of completing low-resolution scans prior to a subjective peak  
738 selection for high-resolution (1 cm) analysis may introduce a bias as low concentration or  
739 discrete peaks might not have sufficient shard concentrations to be observed in the low-  
740 resolution record. The background of shards that is prevalent at some marine sites could mask  
741 individual eruptions that deposited a low concentration of shards. In the ice-core records,  
742 tephra events have been defined on the basis of as few as 3 shards (Bourne et al., 2015).  
743 Detecting deposits of this kind would be particularly challenging in the marine environment  
744 as they could be dismissed as “background” concentrations or hidden with the upward or

745 downward tail of a deposit or within an ash-rich deposit. We have attempted to explore the  
746 presence of such horizons in this study but agree with Timms et al. (2017) who advocate the  
747 use of more high-resolution shard concentration and glass-based chemical analyses to  
748 improve tephrostratigraphies, while acknowledging that this may be limited by sediment  
749 availability, time and financial considerations.

750

751 The marine tephra framework presented in this study should not be viewed as complete.  
752 However, by focusing on maximising the number and geographical range of sequences an  
753 initial framework has been produced that is a significant step towards a more comprehensive  
754 tephra-based synchronisation of North Atlantic marine records. Coupling the success of the  
755 methodology, the initial framework presented here and the insights into the spatial controls  
756 on tephra deposition discussed in Abbott et al. (in revision), there is huge potential to add to  
757 and refine the marine tephra framework. This can be achieved through focusing on new cores  
758 from areas with a high potential to preserve isochronous horizons and reassessing previously  
759 investigated cores at a high-resolution over key intervals during which isochronous horizons  
760 were identified in this work. In addition, innovative techniques for the identification and  
761 quantification of tephra that are currently being developed, for example X-ray fluorescence  
762 core scanning (e.g. Kolling and Bauch, 2017), hyperspectral core imaging (e.g. Aymerich et  
763 al., 2016) and automated flow cytometry and microscopy (e.g. D'Anjou et al., 2014), could  
764 be tested and incorporated into the methodological approach if appropriate.

765

## 766 **6. Conclusions**

767

768 A consistent methodology for the identification and characterisation of marine cryptotephra  
769 and the rigorous assessment of the influence of transportation and deposition processes on  
770 tephra deposits were used to build an enhanced North Atlantic marine tephra framework.  
771 Eleven isochronous deposits were identified in a wide network of marine sequences and have  
772 been integrated with prior data to create a marine tephra framework for the MIS 2-3 period.  
773 Key information for each deposit, such as their spatial extent, geochemical signature, eruptive  
774 source and timing of deposition, is synthesised. A number of significant geochemical  
775 populations are also reported that require further work to assess whether they originate from  
776 single volcanic eruptions and were deposited isochronously via primary tephra fallout.

777



778 There is considerable potential to improve this framework by tracing the deposits into other  
779 marine sequences, by identifying new deposits and/or gaining trace element characterisations  
780 to aid the differentiation of closely-spaced horizons. Combining this framework with  
781 knowledge of the processes controlling the deposition of tephra in the North Atlantic and the  
782 identification of key areas where isochronous horizons are preserved provided in Abbott et al.  
783 (in revision) these future investigations could be highly focussed, both temporally and  
784 spatially. The full potential of this framework will only be realised if attempts are made to  
785 trace these horizons into other archives such as the Greenland ice-cores and terrestrial  
786 records. If successful they can act as time-synchronous tie-lines to correlate and synchronise  
787 these palaeoclimatic records, providing insights into the phasing, rate, timing and  
788 mechanisms forcing the rapid climate changes that characterised this period.

789

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791

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808 stratigraphically, spatially, analytically, and temporally within the Quaternary), an INTAV-  
809 led project (International Focus Group on Tephrochronology and Volcanism) within the  
810 Stratigraphy and Chronology Commission (SACCOM) of the International Union for  
811 Quaternary research (INQUA).

812

813

814 **Figures**

815

816 **Figure 1:** Location map of cores within the marine network (red) and other cores referred to  
817 in the text (green). Location (1) includes cores SO82-2, SO82-5, LO09-23, LO09-21, SO82-7  
818 and SO82-4 described in Lackschewitz and Wallrabe-Adams (1997). Location (2) includes  
819 cores ENAM93-21 and ENAM93-20 and location (3) includes cores LINK16, LINK17,  
820 LINK15 and LINK04 described in Rasmussen et al. (2003), Wastegård et al. (2006) and  
821 Wastegård and Rasmussen (2014).

822

823 **Figure 2:** Schematic representation of the improved marine tephra framework for the North  
824 Atlantic between 60-25 cal kyr BP. Ages and the stratigraphic relationship of tephra horizons  
825 between cores are approximate should be treated with caution, see text for details. The ages  
826 utilised are based on either existing age models for sequences or estimates based on  
827 stratigraphic positions. Heinrich Events 2-5 are included as stratigraphic markers and their  
828 ages are based on those given in Sanchez Goñi and Harrison (2010).

829

830 **Figure 3:** (a) Total alkali v silica plot focusing on (i) rhyolitic material and (ii) basaltic and  
831 intermediate composition glass material from NAAZ II deposits in the marine network. (b)  
832 Comparison of new characterisations of NAAZ II rhyolitic glass to characterisations from  
833 prior studies. Geochemical fields based on analyses of glass from deposits in cores V23-23,  
834 V27-114, V23-82, V23-81 and V23-42 (Kvamme et al., 1989), MD95-2006 (Austin et al.,  
835 2004), ENAM93-20, ENAM33 and EW9302-2JPC (Wastegård et al., 2006) and MD99-2289  
836 (Brendryen et al., 2011). (c) Comparison of basaltic glass from newly characterised NAAZ II  
837 deposits to basaltic NAAZ II glass-based populations defined by Kvamme et al. (1989). All  
838 geochemical data plotted on a normalised anhydrous basis.

839

840 **Figure 4:** (a) Tephrostratigraphy of MD99-2251 between 1950-2030 cm covering the depth  
841 interval of NAAZ II. (i) Rhyolitic glass shards in the 25-80 µm grain-size fraction. (ii)  
842 Basaltic glass shards in the 25-80 µm grain-size fraction. (b) Peak concentrations of  
843 colourless (rhyolitic) and brown (basaltic) glass shards in tephra and cryptotephra deposits  
844 related to North Atlantic Ash Zone II. (c) Relative proportion of geochemical populations  
845 within analyses of basaltic glass tephra shards from NAAZ II deposits at six sites within the  
846 marine core network. Shard analyses not linked to the previously published populations or II-  
847 INT-1 were classified as uncorrelated.

848

849 **Figure 5:** (a) High-resolution concentration profiles of brown glass shards between 910-920  
850 cm in MD95-2010. (b) Comparison of the glass shard composition of MD95-2010 915-916  
851 cm to the glass-based characterisation of FMAZ IV (JM11-19PC 542-543 cm) from Griggs et  
852 al. (2014). All geochemical data plotted on a normalised anhydrous basis.

853

854 **Figure 6:** (a) (i) Percentage abundance of *Neogloboquadrina pachyderma* (sinistral) and (ii)  
855 brown glass shard tephrostratigraphy incorporating 5 cm and 1 cm counts for the MD04-2822  
856 core. (b) (i) Percentage abundance of *Neogloboquadrina pachyderma* (sinistral) and (ii)  
857 brown glass shard tephrostratigraphy incorporating 5 cm and 1 cm counts for the MD04-  
858 2829CQ core. Foram abundances and Dansgaard-Oeschger event numbering for MD04-2822  
859 and MD04-2829CQ from Hibbert et al. (2010) and Hall et al. (2011) respectively. (c)  
860 Geochemical characterisations of glass shards from Type 1 tephra deposits in the MD04-2822  
861 and MD04-2829CQ cores. (i) inset of total alkali vs. silica plot. Division line to separate  
862 alkaline and sub-alkaline material from MacDonald and Katsura (1964). Chemical  
863 classification and nomenclature after Le Maitre et al. (1989). (ii) FeO/TiO<sub>2</sub> vs. SiO<sub>2</sub> and (iii)  
864 TiO<sub>2</sub> vs. Al<sub>2</sub>O<sub>3</sub> compositional variations diagrams comparing the glass shard composition of  
865 MD04-2822 and MD04-2829CQ deposits to characterisations of proximal Icelandic material.  
866 Geochemical fields for Icelandic source volcanoes are based on normalised whole rock and  
867 glass shard analyses utilised in Bourne et al. (2015) and references within and additional data  
868 for the Kverkfjöll volcano from Gudmundsdóttir et al. (2016). All geochemical data plotted  
869 on a normalised anhydrous basis.

870

871 **Figure 7:** (a) (i) Percentage abundance of *Neogloboquadrina pachyderma* (sinistral), (ii) ash  
872 free IRD concentration and (iii) tephrostratigraphic record of the MD99-2251 marine core.  
873 Glass shard counts have been truncated for clarity. Shard counts in the 1686-1687 cm sample  
874 (\*) are 4991, 1862 and 507 shards per 0.5 g dws in the 25-80, 80-125 and >125 µm grain-size  
875 fractions, respectively. The shard counts for the 25-80 µm grain-size fraction from the 1904-  
876 1905 cm sample (\*\*) are 3776 shards per 0.5 g dws. Red bars denote samples depths from  
877 which glass shards were subsequently extracted for compositional characterisation. (b)  
878 Composition of significant geochemical populations identified in glass analyses of tephra  
879 deposits within the MD99-2251 core. (i) inset of total alkali vs. silica plot. Division line to  
880 separate alkaline and sub-alkaline material from MacDonald and Katsura (1964). Chemical  
881 classification and nomenclature after Le Maitre et al. (1989). (ii) FeO/TiO<sub>2</sub> vs. SiO<sub>2</sub> and (iii)

882 TiO<sub>2</sub> vs. Al<sub>2</sub>O<sub>3</sub> compositional variations diagrams comparing significant glass shard based  
883 geochemical populations from the MD99-2251 deposits to characterisations of proximal  
884 Icelandic material. Geochemical fields for Icelandic source volcanoes are based on  
885 normalised whole rock and glass shard analyses utilised in Bourne et al. (2015) and  
886 references within and additional data for the Kverkfjöll volcano from Gudmundsdóttir et al.  
887 (2016). All geochemical data plotted on a normalised anhydrous basis.

888

889 **Figure 8:** (a) (i) Percentage abundance of *Neogloboquadrina pachyderma* (sinistral), (ii)  
890 percentage IRD (>150 µm fraction) and (iii) tephrostratigraphic record of the GIK23415-9  
891 marine core. *Np*(s) and IRD data from Vogelsang et al. (2004) and Weinelt (2004),  
892 respectively. Labels for Heinrich Events from Weinelt et al. (2003) and Lu et al. (2007).  
893 Shard counts have been truncated for clarity. Shard counts in the 193-194 cm sample are  
894 5131 and 280 shards per 0.5 g dws in the 25-80 and >125 µm grain-size fractions  
895 respectively. Red bars denote samples depths from which glass shards were subsequently  
896 extracted for compositional characterisation. (b) Composition of significant glass-based  
897 geochemical populations identified in tephra deposits within the GIK23415-9 core. (i) inset  
898 of total alkali vs. silica plot. Division line to separate alkaline and sub-alkaline material from  
899 MacDonald and Katsura (1964). Chemical classification and nomenclature after Le Maitre et  
900 al. (1989). (ii) FeO/TiO<sub>2</sub> vs. SiO<sub>2</sub> and (iii) TiO<sub>2</sub> vs. Al<sub>2</sub>O<sub>3</sub> compositional variations diagrams  
901 comparing significant glass-based geochemical populations from the GIK23415-9 deposits to  
902 characterisations of proximal Icelandic material. Geochemical fields for Icelandic source  
903 volcanoes are based on normalised whole rock and glass shard analyses utilised in Bourne et  
904 al. (2015) and references within and additional data for the Kverkfjöll volcano from  
905 Gudmundsdóttir et al. (2016). All geochemical data plotted on a normalised anhydrous basis.

906

907 **Figure 9:** (a) Comparison of the glass analyses of the MD04-2822 1836-1837 cm tephra  
908 horizon and the GIK23415-9 225-226 cm (TAB-2) geochemical population. (b) Comparison  
909 of the glass analyses of the FMAZ II tephra horizon (JM11-19PC 202-203 cm from Griggs et  
910 al. (2014)) and that of the GIK23415-9 202-203 cm (THOL-1) geochemical population. All  
911 geochemical data plotted on a normalised anhydrous basis.

912 **Supplementary Information**

913

914 *Supplementary Figures*

915

916 **Figures S1-S13:** Graphical analysis of geochemical populations identified within single-  
917 shard major element glass analyses from tephra deposits within the MD99-2251 (S1-S7) and  
918 GIK23415-9 (S8-S13) cores.

919

920 *Supplementary Data*

921

922 **Table S1:** Original major oxide concentrations of glass shards from deposits related to the  
923 rhyolitic component of North Atlantic Ash Zone II (II-RHY-1). Deposits analysed are (i)  
924 MD04-2822 2168-2169 cm (ii) MD95-2024 1445-1446 cm (iii) MD99-2251 1974-1975 cm  
925 (supplementary peak) (iv) MD99-2251 2014-2015 cm (main peak) (v) M23485-1 622-623  
926 cm (vi) GIK23415-9 429-430 cm (vii) MD01-2461 942-943 cm (supplementary peak) (viii)  
927 MD01-2461 947-948 cm (main peak) (ix) MD04-2820CQ 610-611 cm (x) JM11-19PC 618-  
928 623 cm (xi) MD95-2010 996-1000 cm.

929

930 **Table S2:** Similarity coefficient comparisons of average concentrations of glass analyses of  
931 the II-RHY-1 component in deposits from cores analysed within this work and by Kvamme et  
932 al. (1989), Austin et al. (2004), Wastegård et al. (2006) and Brendryen et al. (2011).

933

934 **Table S3:** Original major oxide concentrations of glass shards from basaltic and intermediate  
935 shards directly associated with deposits of the rhyolitic component of North Atlantic Ash  
936 Zone II (II-RHY-1). Deposits analysed are (i) MD99-2251 2014-2015 cm (ii) M23485-1 622-  
937 623 cm (iii) GIK23415-9 429-430 cm (iv) MD01-2461 947-948 cm (v) MD04-2820CQ 610-  
938 611 cm (vi) JM11-19PC 618-623 cm.

939

940 **Table S4:** Original major oxide concentrations of glass shards from the MD95-2010 915-916  
941 cm tephra deposit.

942

943 **Table S5:** Original major oxide concentrations of glass shards from tephra deposits in the  
944 MD04-2822 core. Deposits analysed are from the depths of (i) 1836-1837 cm (ii) 2004-2005  
945 cm and (iii) 2017-2018 cm.

946

947 **Table S6:** Original major oxide concentrations of glass shards from tephra deposits in the  
948 MD04-2829CQ core. Deposits analysed are from the depths of (i) 800-801 cm (ii) 930-931  
949 cm and (iii) 934-935 cm.

950

951 **Table S7:** Original major oxide concentrations of glass shards from tephra deposits in the  
952 MD04-2820CQ core. Deposits analysed are from the depths of (i) 487-488 cm (ii) 497-498  
953 cm and (iii) 524-525 cm.

954

955 **Table S8:** Original major oxide concentrations of glass shards from tephra deposits in the  
956 MD99-2251 core. Deposits analysed are from the depths of (i) 1654-1655 cm (ii) 1680-1681  
957 cm (iii) 1683-1684 cm (iv) 1713-1714 cm (v) 1772-1773 cm (vi) 1796-1797 cm (vii) 1812-  
958 1813 cm and (viii) 1904-1905 cm.

959

960 **Table S9:** Analysis of glass-based geochemical populations present within tephra deposits  
961 identified in the MD99-2251 marine core. n = total number of analyses from deposits. Veid.-  
962 Bárd. = Veidivötn-Bárdarbunga.

963

964 **Table S10:** Original major oxide concentrations of glass shards from tephra deposits in the  
965 GIK23415-9 core. Deposits analysed are from the depths of (i) 173-174 cm (ii) 193-194 cm  
966 (iii) 202-203 cm (iv) 225-226 cm (v) 302-303 cm (vi) 305-306 cm and (vii) 375-376 cm.

967

968 **Table S11:** Analysis of glass-based geochemical populations present within tephra deposits  
969 identified in the GIK23415-9 marine core. n = total number of analyses from deposits. Veid.-  
970 Bárd. = Veidivötn-Bárdarbunga.

971

972 **Table S12a:** Original secondary standard analyses of the BCR2g standard made throughout  
973 analytical periods during which glass sample analyses presented in this work were analysed.

974

975 **Table S12b:** Original secondary standard analyses of the Lipari standard made throughout  
976 analytical periods during which glass sample analyses presented in this work were analysed.

977

978 **Table S13:** Similarity coefficient comparisons between the glass-based geochemical  
979 signatures of isochronous horizons and significant glass-based geochemical populations in

980 the marine tephra framework for the North Atlantic between 25-60 ka BP. Method of  
981 Borchartt et al. (1972) utilised. Red text shows SC values between 0.97 and 0.999 grey text  
982 shows SC values less than 0.95.



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**Table 1:** Summary of isochronous horizons and significant geochemical populations forming the marine tephra framework for the North Atlantic between 25-60 ka BP. The designation of climatic events is based on pre-existing stratigraphic frameworks for the cores. The stratigraphic ordering of horizons between cores is approximate. FMAZ II, FMAZ IV and NAAZ II have been identified in multiple cores. H = Heinrich Event; DO = Dansgaard-Oeschger Event. bas = basaltic; rhy = rhyolitic. Vat. = Vatnafjöll; Veid.-Bárd. = Veidivötn-Bárdarbunga. Deposit types based on the classification scheme outlined in Abbott et al. (submitted). References are as follows: 1: this study; 2: Rasmussen et al. (2003); 3: Wastegård et al. (2006); 4: Davies et al. (2008); 5: Griggs et al. (2014); 6: Abbott et al. (2016); 7: Wastegård and Rasmussen (2014); 8: Kvanne et al. (1989); 9: Austin et al. (2004); 10: Brendryen et al. (2010).

Tephra horizon/deposit (pop.)	Climatic event	Composition	Volcanic source	Deposit type	Ref(s)
<b><i>Isochronous horizons</i></b>					
GIK23415-9 173-174 cm (THOL-1)	Post H2	Tholeiitic bas	Kverkfjöll	1	1
FMAZ II	Post DO-3	Transitional alkali bas	Hekla/Vatnafjöll	2A	2, 3, 4, 5
MD04-2822 1836-1837 cm	Pre DO-4	Transitional alkali bas	Katla or Hekla/Vatnafjöll	1	1
GIK23415-9 225-226 cm (TAB-1)	H3	Transitional alkali bas	Katla or Hekla/Vat.	1	1
MD04-2829CQ 800-801 cm (THOL-1)	Pre DO-4	Tholeiitic bas	Grímsvötn	1	1
MD04-2829CQ 800-801 cm (THOL-2)	Pre DO-4	Tholeiitic bas	Kverkfjöll	1	1
MD99-2251 1680-1681 cm (TAB-1)	H3	Transitional alkali bas	Katla	2A	1
MD04-2822 2004-2005 cm	DO-8	Transitional alkali bas	Katla	1	1
MD04-2829CQ 930-931 cm	DO-8	Tholeiitic bas	Grímsvötn	1	1
MD04-2829CQ 934-935 cm	Pre DO-8	Tholeiitic bas	Grímsvötn	1	1
MD04-2820CQ 487-488 cm	Pre DO-8 (H4)	Tholeiitic bas	Grímsvötn	2A	6
MD04-2820CQ 497-498 cm	Pre DO-9	Transitional alkali rhy	Katla	2A	6
MD04-2822 2017-2018 cm	Pre DO-9	Tholeiitic bas	Grímsvötn	1	1
MD04-2820CQ 524-525 cm	Pre DO-11	Tholeiitic bas	Grímsvötn or Kverkfjöll	2A	6
FMAZ IV	Pre DO-12	Tholeiitic bas	Grímsvötn	2A	5, 7
NAAZ II (II-RHY-1)	End of DO-15	Transitional alkali rhy	Tindfjallajökull	3	1, 3, 8, 9, 10
<b><i>Significant geochemical populations</i></b>					
GIK23415-9 202-203 cm (TAB-1)	Pre H2	Transitional alkali bas	Katla	2B	1
GIK23415-9 202-203 cm (THOL-1)	Pre H2	Tholeiitic bas	Kverkfjöll	2B	1
GIK23415-9 202-203 cm (THOL-2)	Pre H2	Tholeiitic bas	Grímsvötn	2B	1
GIK23415-9 202-203 cm (THOL-3)	Pre H2	Tholeiitic bas	Veid. -Bárd.	2B	1
MD99-2251 1713-1714 cm (TAB-1)	Pre H3	Transitional alkali bas	Katla	2B	1
MD99-2251 1713-1714 cm (THOL-1)	Pre H3	Tholeiitic bas	Grímsvötn	2B	1
MD99-2251 1772-1773 cm (TAB-1)	Post H4	Transitional alkali bas	Katla	2B	1
MD99-2251 1772-1773 cm (TAB-2)	Post H4	Transitional alkali bas	Katla (?)	2B	1
GIK23415-9 302-306 cm (THOL-1)	H4	Tholeiitic bas	Grímsvötn	4	1
GIK23415-9 302-306 cm (THOL-2)	H4	Tholeiitic bas	Grímsvötn (?)	4	1
MD99-2251 1812-1813 cm (THOL-1)	H4	Tholeiitic bas	Grímsvötn	2B	1
MD99-2251 1812-1813 cm (THOL-2)	H4	Tholeiitic bas	Veid. -Bárd.	2B	1
MD99-2251 1812-1813 cm (THOL-3)	H4	Tholeiitic bas	Veid. -Bárd.	2B	1
MD99-2251 1812-1813 cm (TAB-1)	H4	Transitional alkali bas	Katla	2B	1
MD99-2251 1904-1905 cm (ALK-1)	Post H5	Alkali bas	Vestmannaeyjar	2B	1
GIK23415-9 375-376 cm (THOL-1)	Pre H5	Tholeiitic bas	Grímsvötn	2B	1
NAAZ II (II-THOL-2)	End of DO-15	Tholeiitic bas	Grímsvötn	2A, 2B	8, 1
NAAZ II (II-TAB-1)	End of DO-15	Transitional alkali bas	Katla	2A, 2B	8, 1
NAAZ II (II-INT-1)	End of DO-15	Trachyandesite-Trachydacite	Unknown	2A, 2B	1



Figure 1

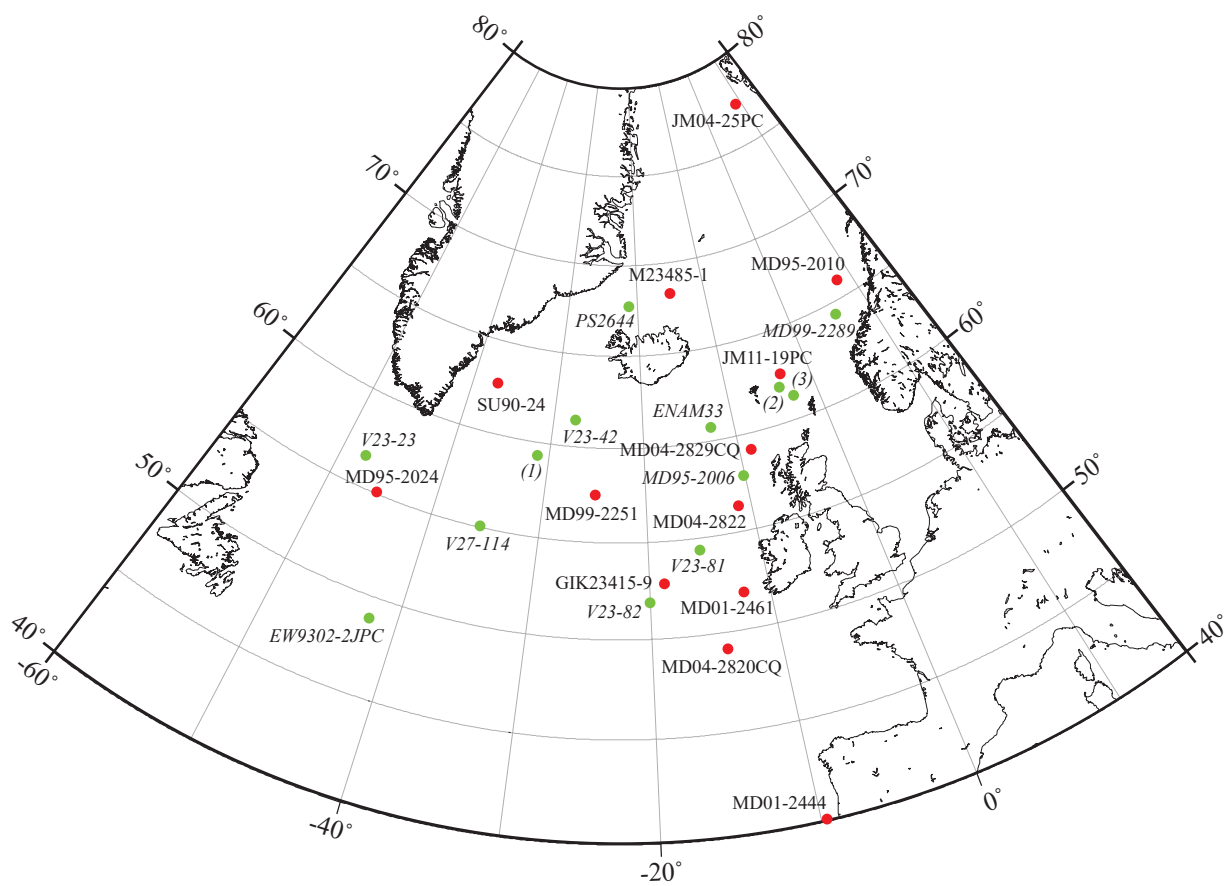
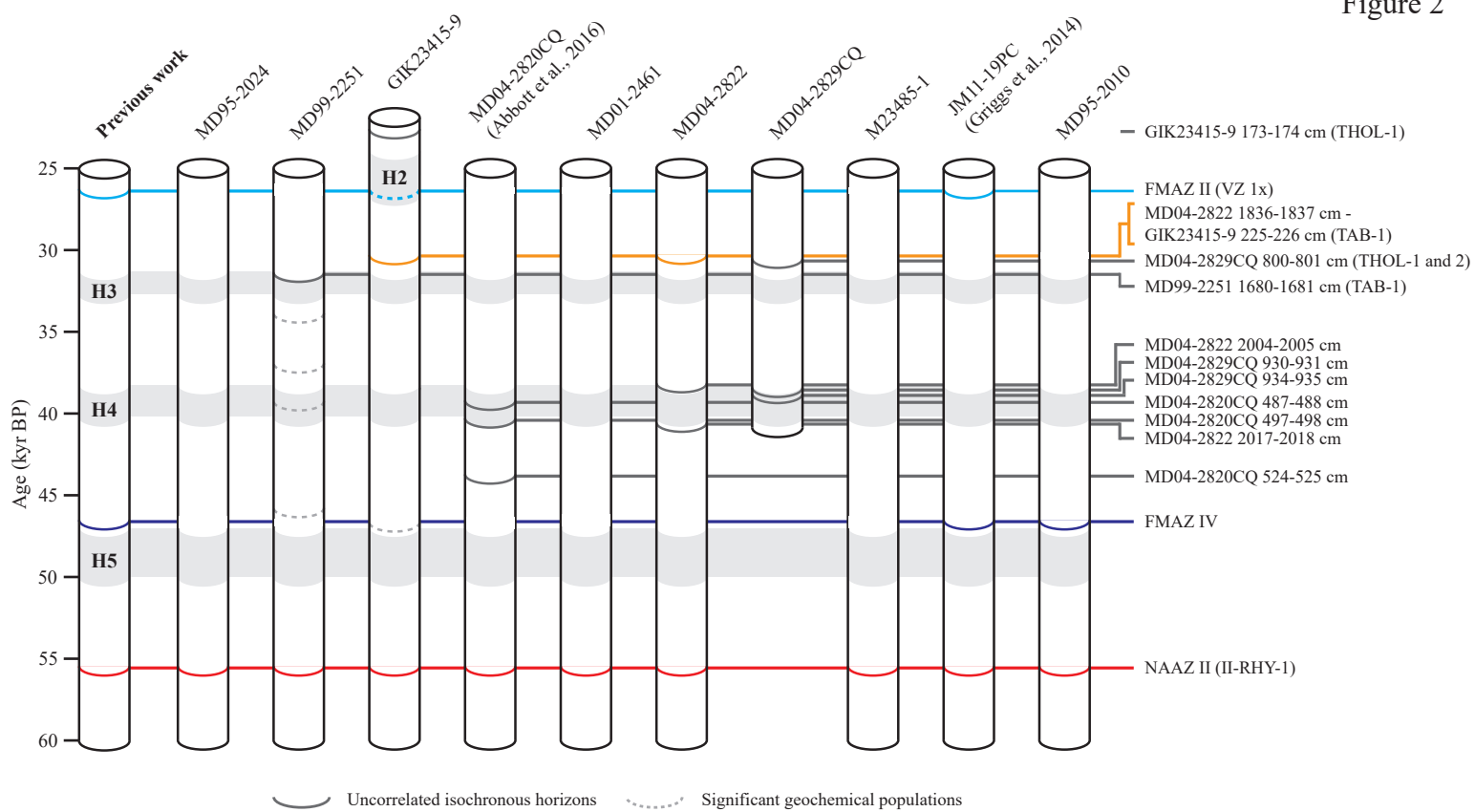
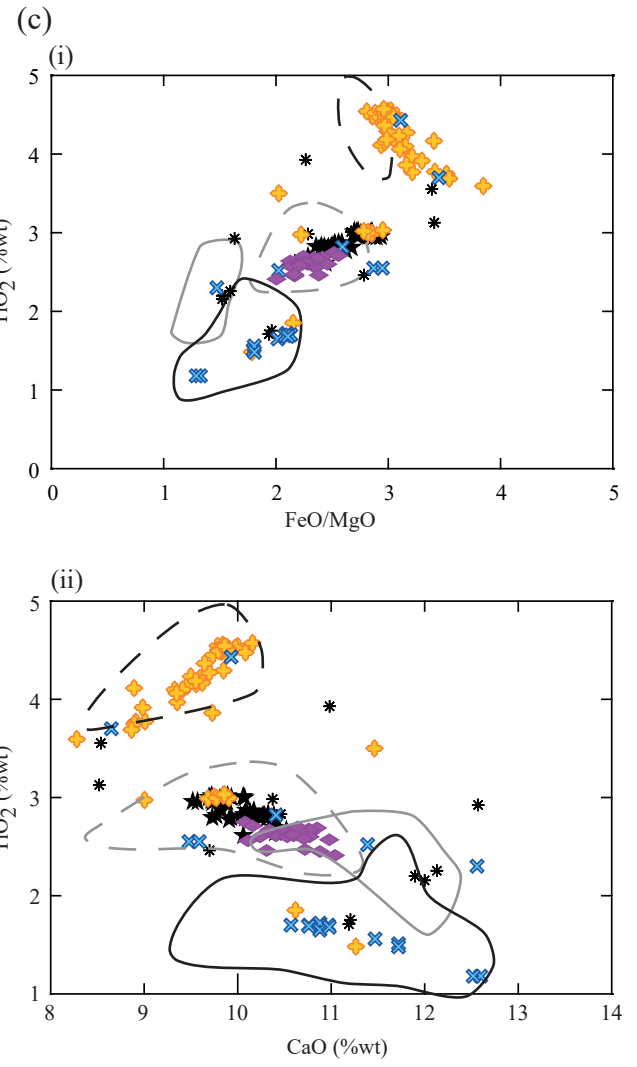
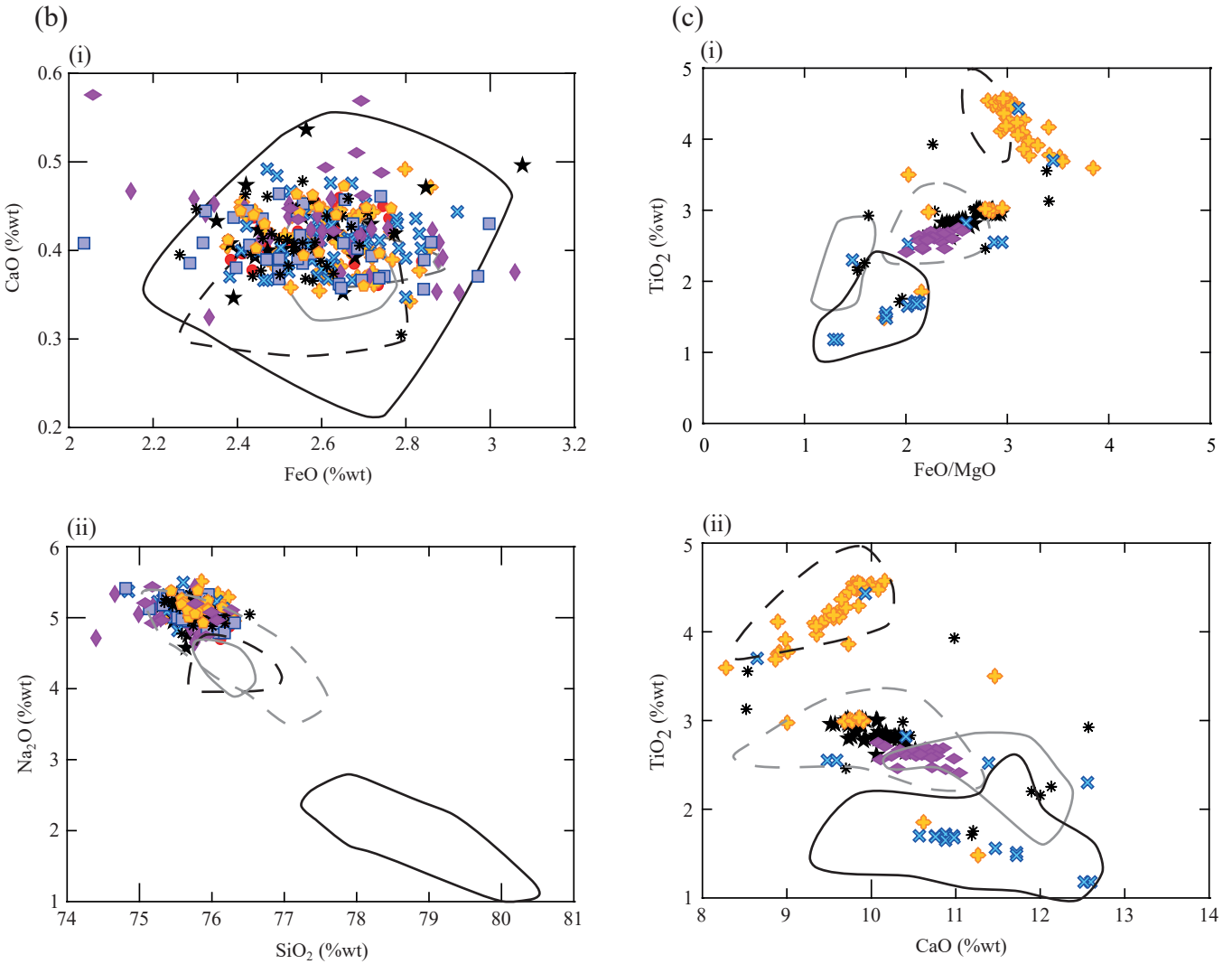
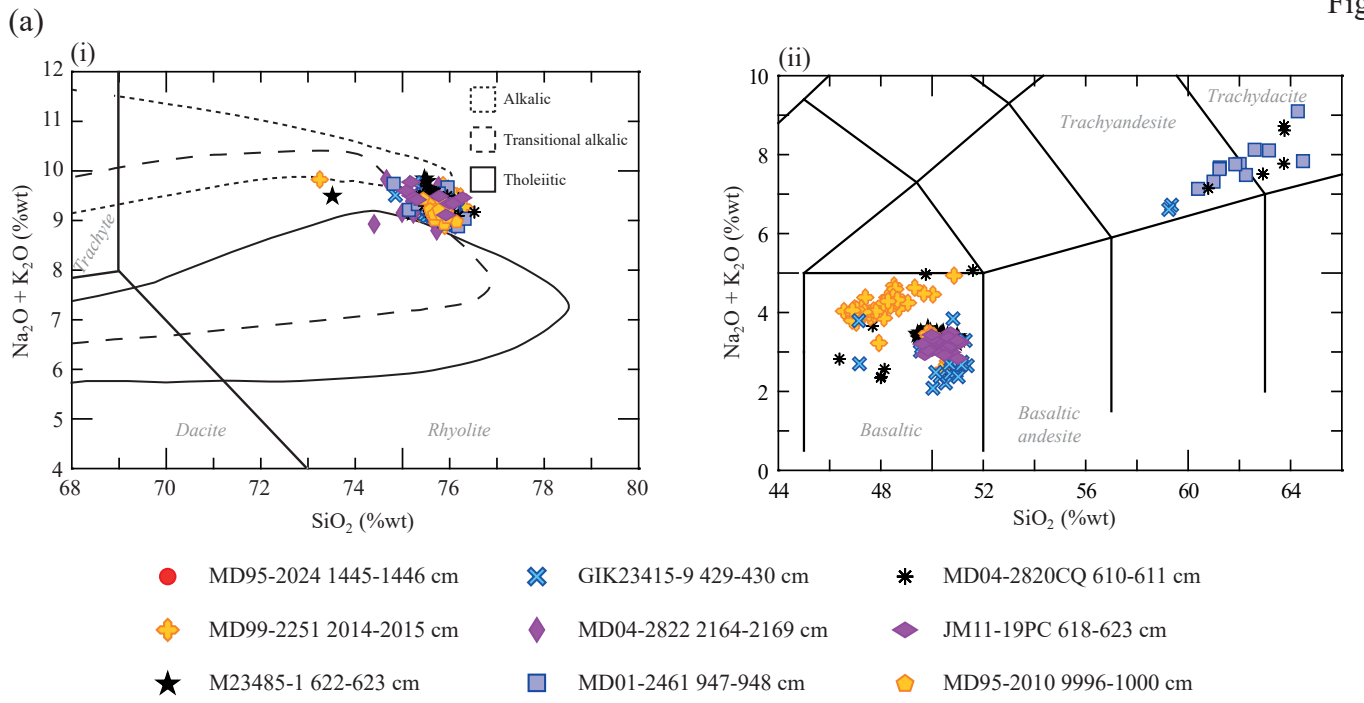


Figure 2





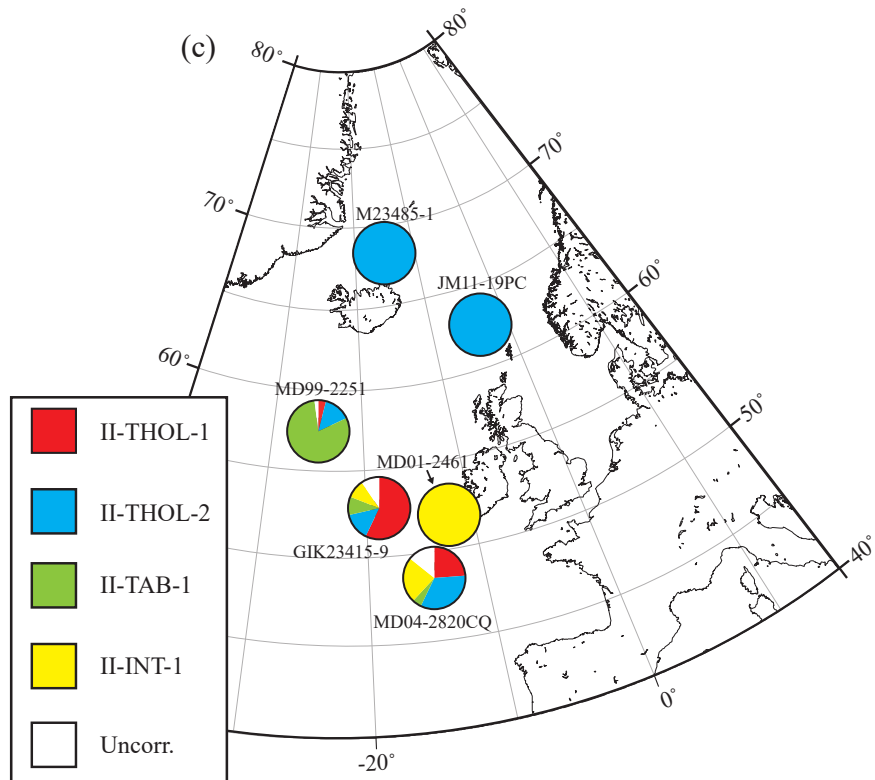
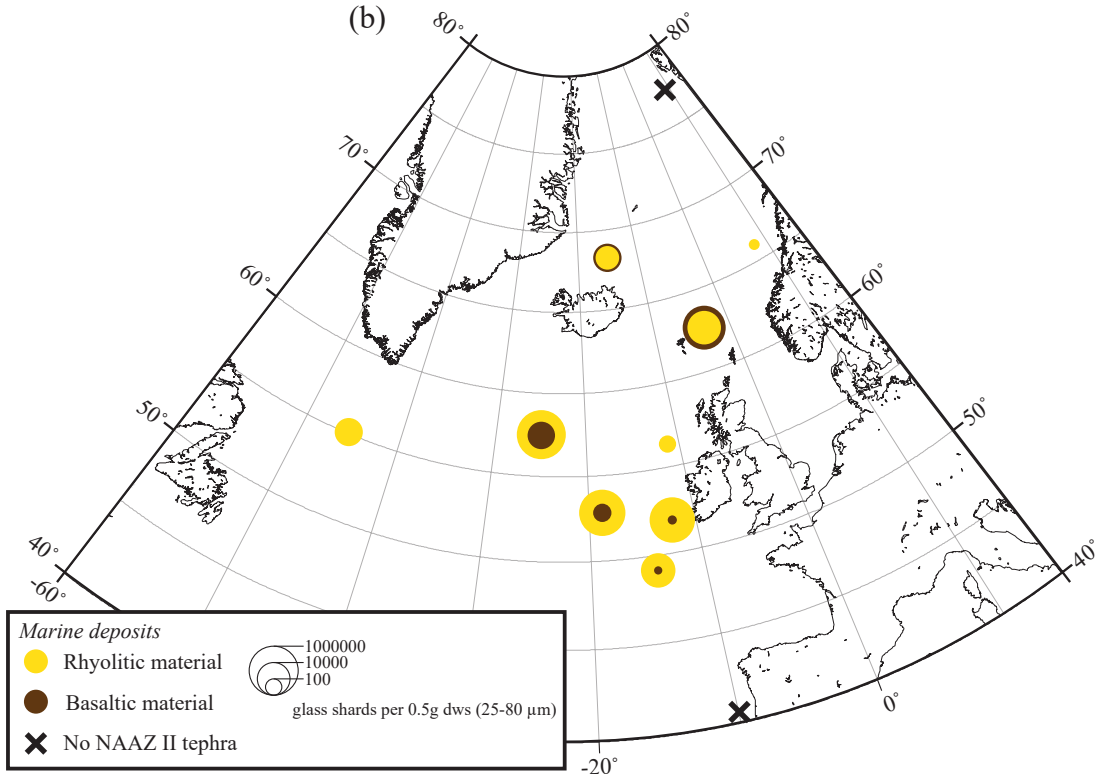
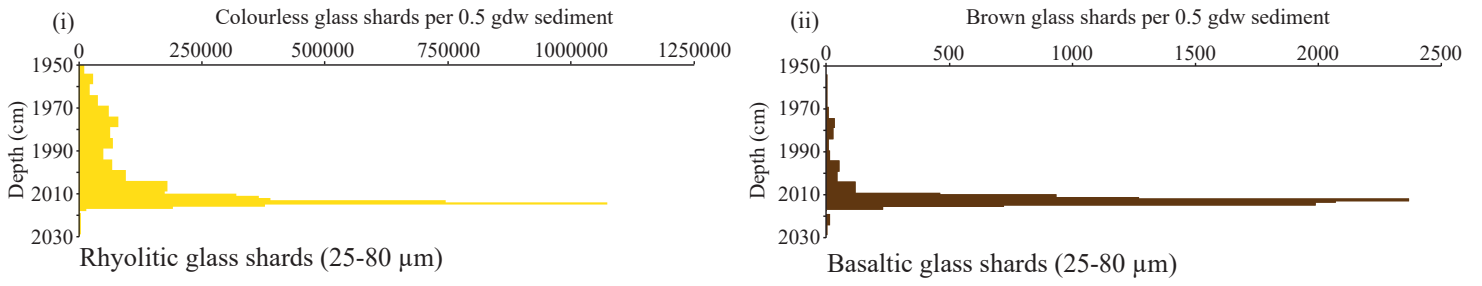
**II-RHY-1:**

- Kvamme et al. (1989)
- Austin et al. (2004)
- - - Wastegård et al. (2006)
- - - Brendryen et al. (2011)

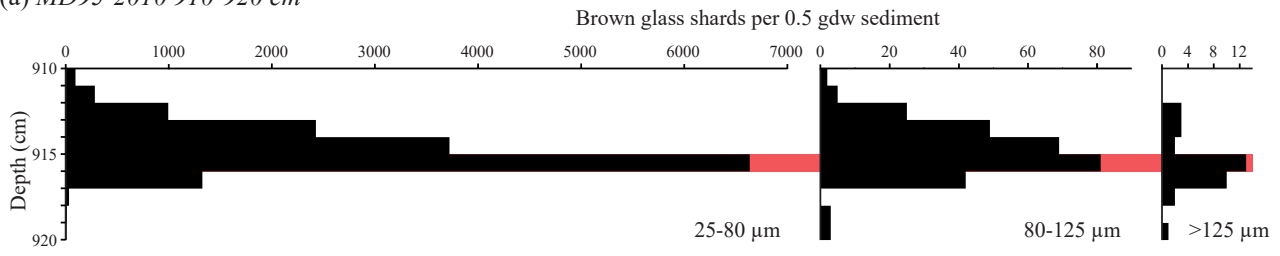
**Kvamme et al. (1989):**

- II-THOL-1
- II-THOL-3
- - - II-THOL-2
- - - II-TAB-1

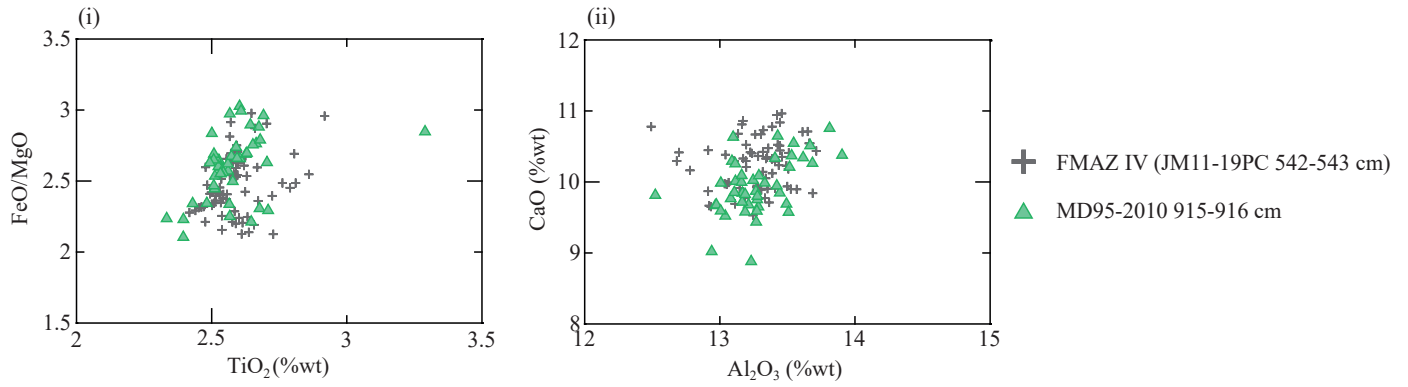
(a) MD99-2251 NAAZ II tephrostratigraphy (Gardar Drift)



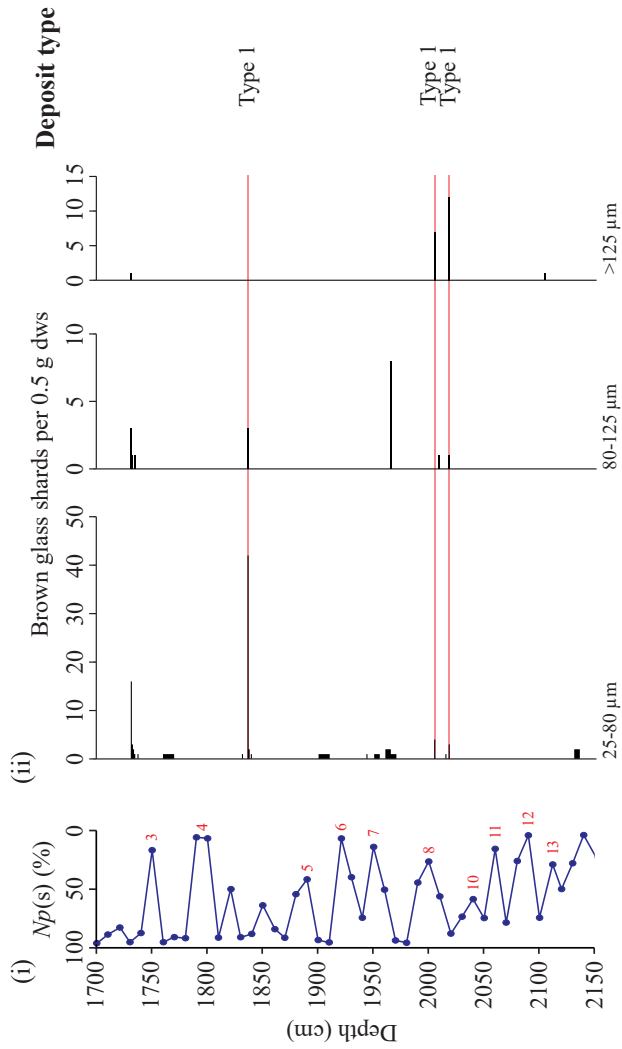
(a) MD95-2010 910-920 cm



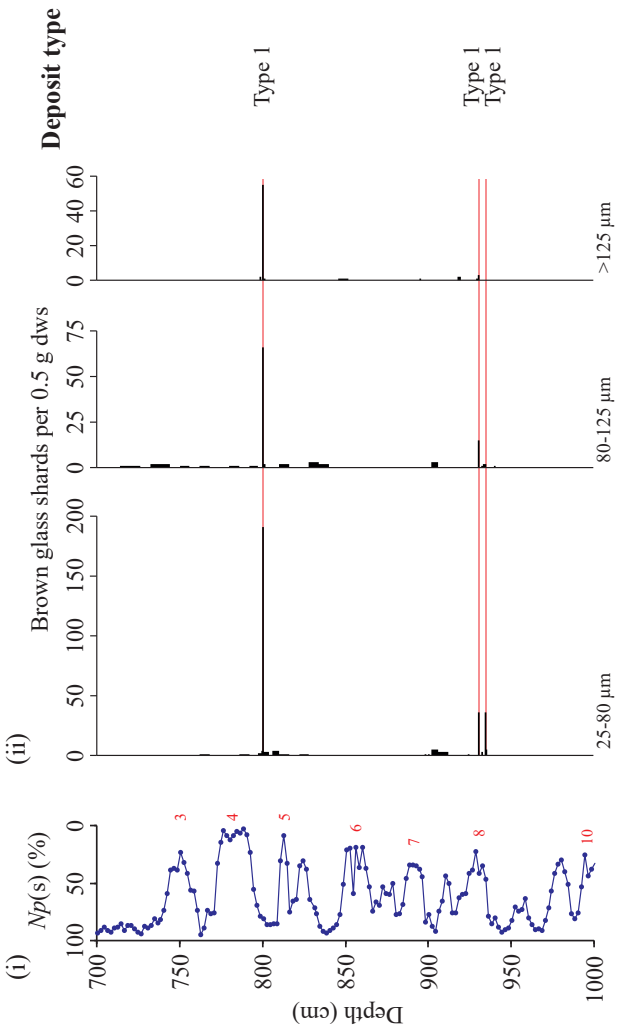
(b)



(a) MD04-2822 (Rockall Trough)



(b) MD04-2829CQ (Rosemary Bank)



(c) MD04-2822 and MD04-2829CQ characterisations

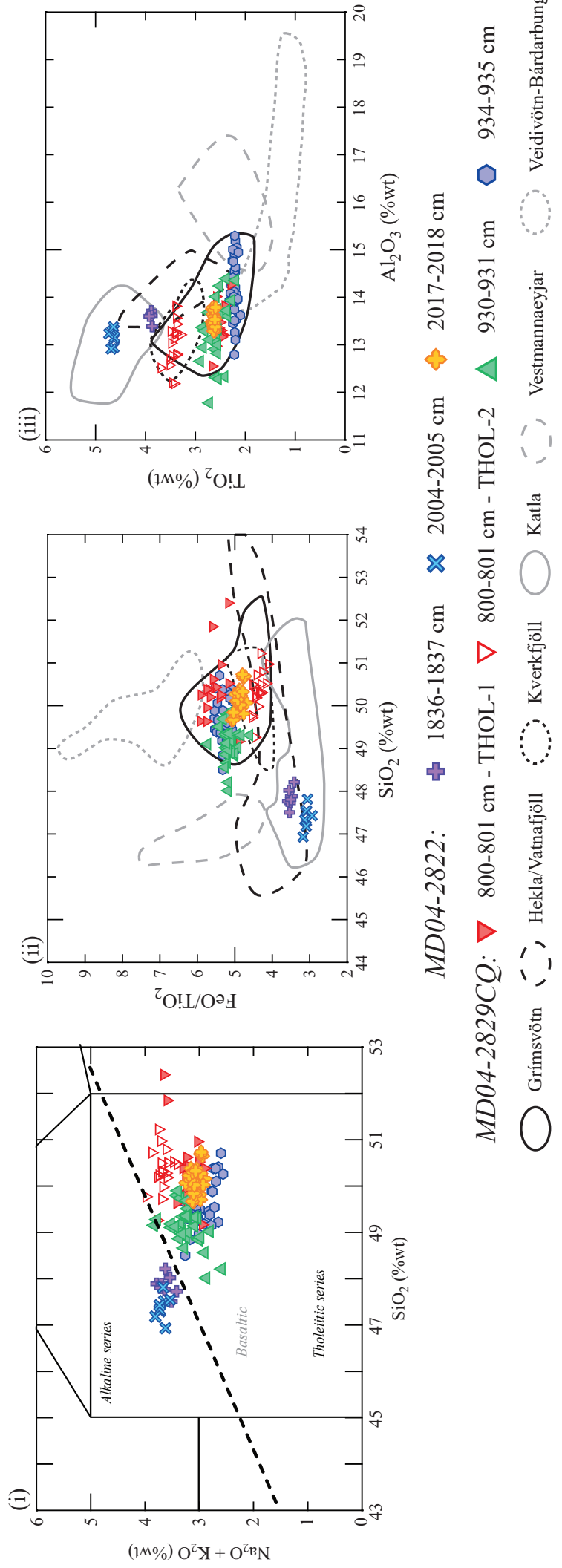
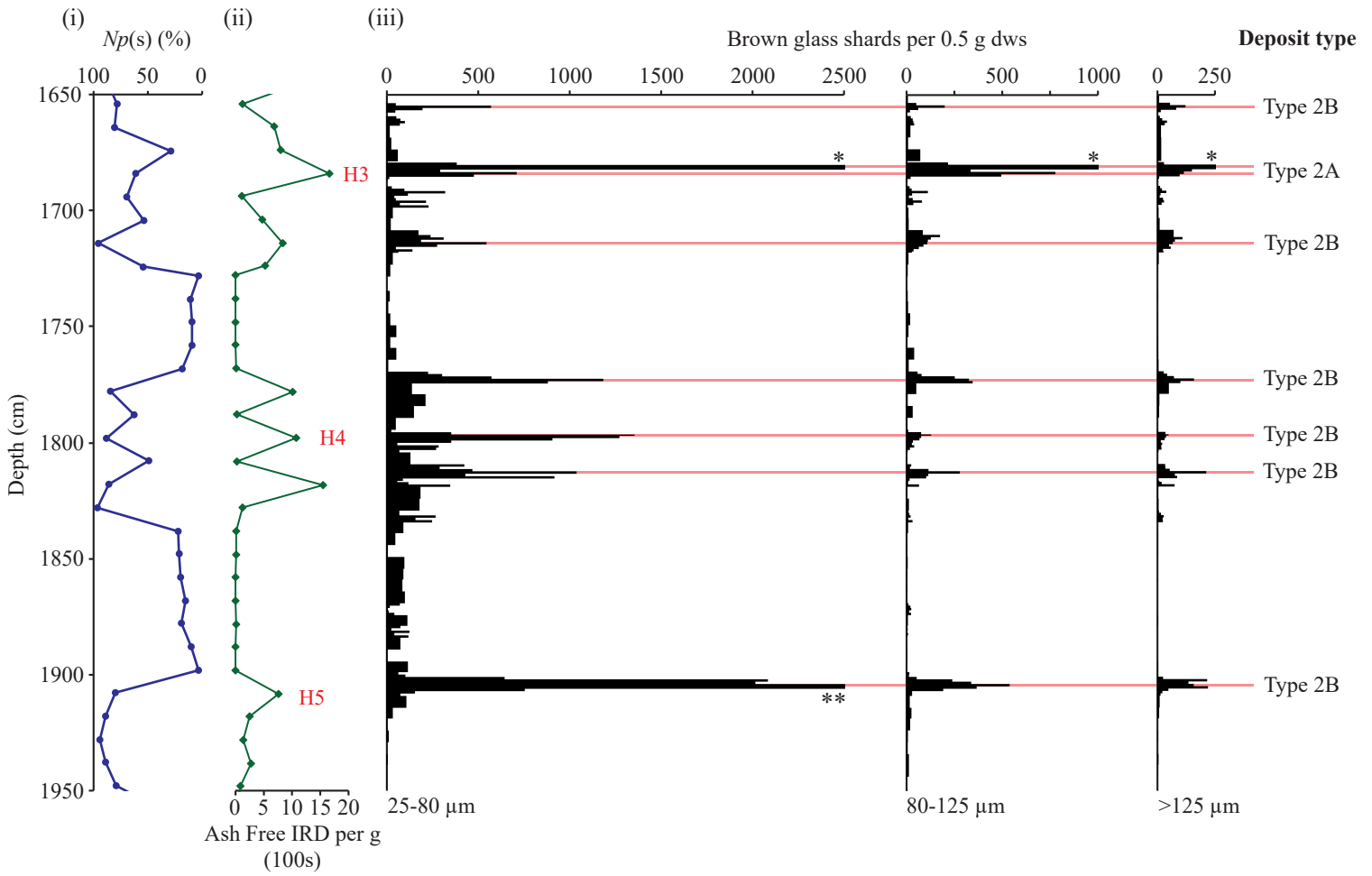
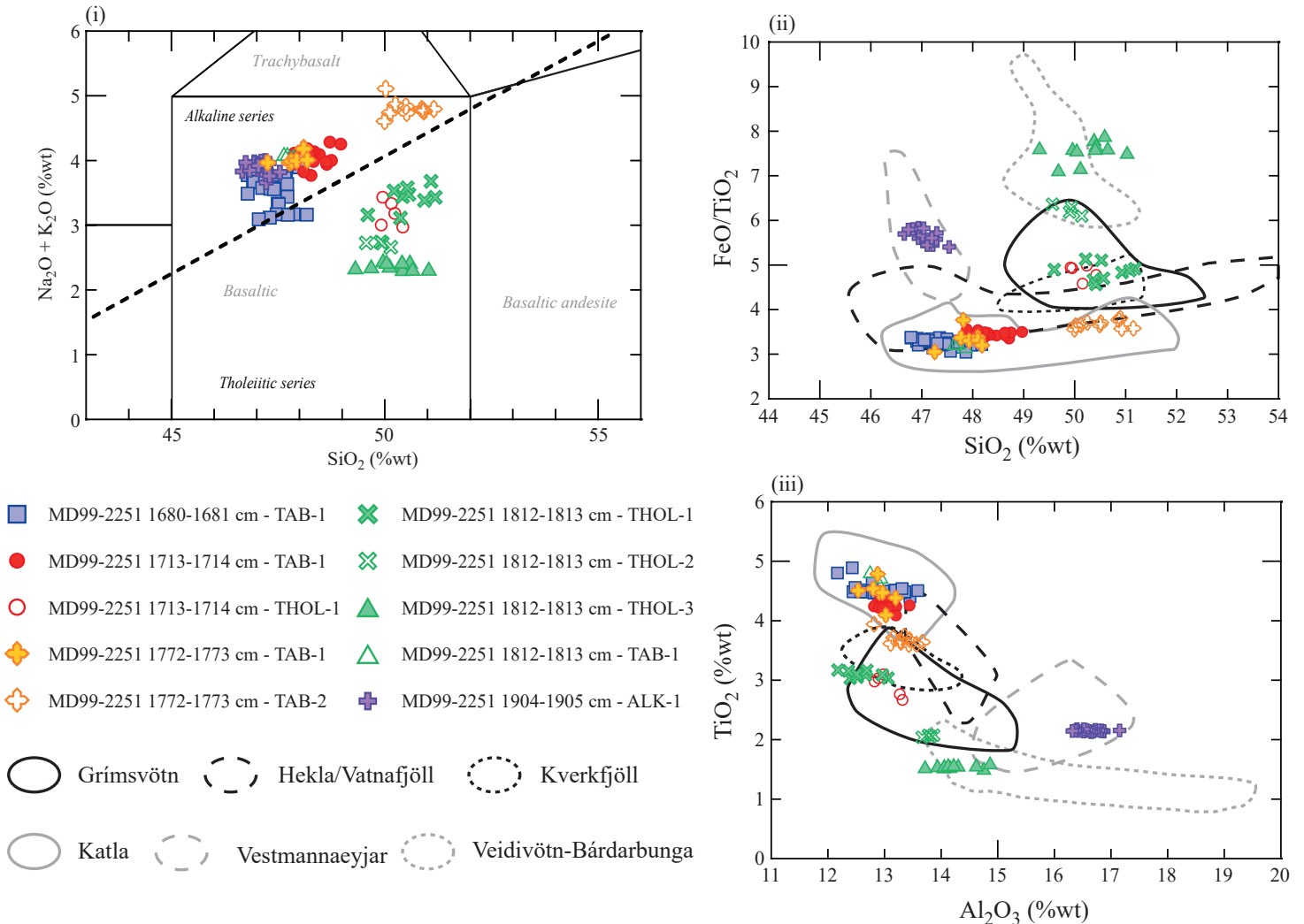


Figure 6

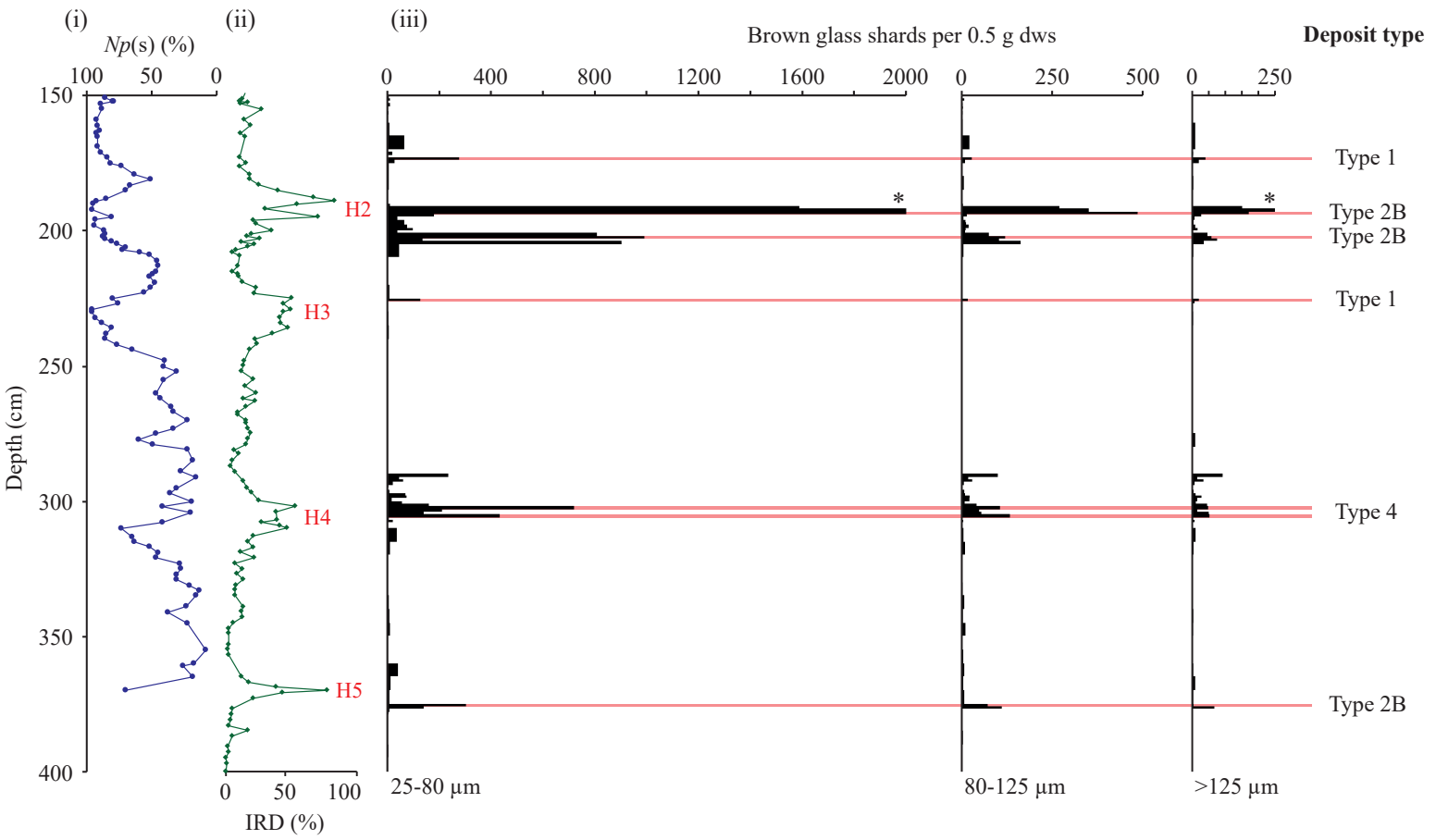
(a) MD99-2251 tephr stratigraphy (Gardar Drift)



(b) MD99-2251 isochronous horizon and significant geochemical populations



(a) GIK23415-9 tepthrostratigraphy (Northern North Atlantic)



(b) GIK23415-9 isochronous horizons and significant geochemical populations

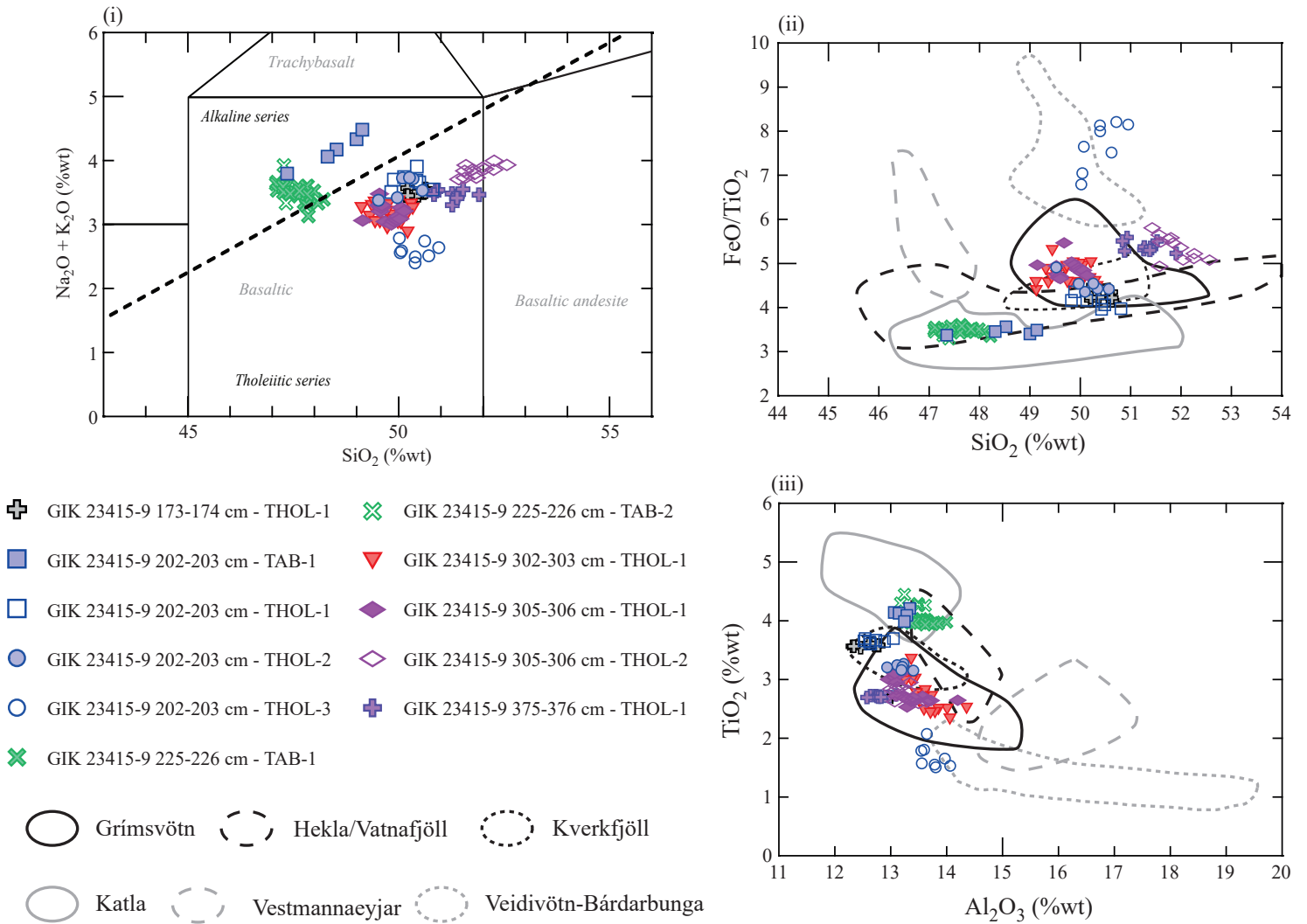




Figure 9

